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# 75997

## Armed Services Technical Information Agency

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# OF

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Date

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*Richard E. Reedy*

OFFICE SECURITY ADVISOR

PROGRESS REPORT ON  
HIGH ALTITUDE PLASTIC BALLOONS  
CONTRACT NONR-710 (01)  
DECEMBER 13, 1951 to JUNE 15, 1952  
VOLUME IV  
CONFIDENTIAL SECURITY INFORMATION  
Copy No. 4 01

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PROGRESS REPORT ON  
RESEARCH AND DEVELOPMENT  
IN THE FIELD OF  
HIGH ALTITUDE PLASTIC BALLOONS

CONDUCTED UNDER  
CONTRACT NONR-710(01), NR 320 159  
FOR PERIOD DECEMBER 13, 1951 to JUNE 15, 1952  
WITH THE  
OFFICE OF NAVAL RESEARCH

AND SPONSORED JOINTLY  
BY THE ARMY, NAVY AND AIR FORCE

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DEPARTMENT OF PHYSICS  
UNIVERSITY OF MINNESOTA  
MINNEAPOLIS 14, MINNESOTA

Copy No. 481

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PROGRESS REPORT ON CONTRACT # 710 (01)

From Initiation of Contract to June 15, 1952

VOLUME IV

CONFIDENTIAL SECURITY INFORMATION

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## SECTION VIII

Page 8-1

### METEOROLOGY

The work of the meteorological group has so far been largely of an exploratory nature. Four principal studies have been initiated: (1) plastic balloon failure in the "jet stream", (2) comparison of the character of the flow in the stratosphere with the flow in the troposphere, (3) a case study of the vorticity change as measured by the constant altitude balloon, (4) the computation of balloon height and temperature.

In addition, this group prepares the balloon trajectory charts and analysis of wind data from the balloon flights, and provides a number of minor services, generally in interpreting meteorological data. Forecasts are made for flight planning purposes and computations of the infrared flux has been made for selected rays.

#### 1. Plastic Balloon Failure in the Jet Stream

It had been observed in earlier General Mills flight data that there is a relatively high incidence of plastic balloon failure in the vicinity of the tropopause. Since this is the coldest region through which the balloon ascends, the brittleness of the balloon fabric at these low temperatures has been suggested as an important cause of failure. Pilot reports of turbulence associated with high winds at the tropopause led Dr. R. A. Craig to advance the opinion that turbulence may be a further contributing factor.

All available data on previous plastic balloon flights are, therefore, being examined with reference to the following meteorological elements: wind velocity, vertical wind shear, and atmospheric stability. This analysis is made as the meteorological data become available, and, at present, is only completed for irregular intervals in the periods 1948-1949 and 1951-1952. The meteorological data have been studied for a total of 38 flights made by General Mills and also by the Moby Dick project. Only five balloon failures

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at the tropopause were noted in this group. Four cases of failure occurred under conditions of higher than average wind velocity (70-100 kt) and relatively large values of the vertical wind shear. One failure could be considered a borderline case with a wind velocity of 50 kt. In this group of 38, however, there were also four cases of successful flights through wind streams of equally high wind velocities and vertical wind shear. No obvious relationship was observed between atmospheric stability or temperature and balloon failure. In one of the cases of balloon failure it was observed that the horizontal wind shear was abnormally large, a value of  $2 \times 10^{-4} \text{ sec}^{-1}$  obtained from measured winds.<sup>1</sup> The further possibility exists, therefore, that the horizontal wind shear may be important in isolating the dangerous meteorological situations.

The erratic nature of this failure is also born out by recent Minnesota flights. Since the balloons in this group are not of identical manufacture, it seems quite likely that differences in balloon construction and materials may be the decisive factor as to whether the "jet stream" will cause balloon failure. The limited number of failures available for study prevents final conclusions on this point.

2. Comparison of the Character of the Flow in the Stratosphere with the Flow in the Troposphere

As a means of interpreting the information obtained on the stratospheric flow, various characteristics of the flow in the stratosphere are being compared with the characteristics of the flow in the troposphere. The trajectories of the early Moby Dick flights in the wind season 1951-1952, were used in a preliminary study in an effort to determine to what elevations one could successfully extrapolate the methods of conventional synoptic analysis. Spe-

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<sup>1</sup>Palmen and Nagler point out that  $10^{-4} \text{ sec}^{-1}$  is a large value of the horizontal shear, Journ. Met., 1948, vol. 5, p. 64.

cifically, 100 mb charts were constructed for the flights which floated near 55,000 ft. and 50 mb charts for the flights at approximately 70,000 ft. From these limited cases it appeared that the broad features of the flow in the 100 mb chart were quite similar to those at 200 mb and the flow described by the balloon's trajectory corresponded, within the limits of accuracy of the trajectory, to the flow indicated by the 100 mb chart. The construction of the 50 mb charts was not successful. The upper air data is quite limited at these elevations and from the available data it appears that considerable changes in flow pattern are possible in the layer 100 to 50 mb.

The velocity fluctuations or accelerations are a quantitative and descriptive characteristic of the flow. The mean velocity over varying time intervals is being computed from the trajectories at constant altitude. Similar computations have been made for the General Mills 300 mb flights so as to provide the tropospheric comparison. The compilation of data and the analysis have not been completed but certain features are evident. Velocity changes over adjacent 5 minute periods are of about equal magnitude in the troposphere and in the stratosphere at 80,000 ft. Velocity changes over adjacent periods of 1 to 2 hours are also very nearly equal in the stratosphere and in the troposphere. In some stratospheric flights, notably Minnesota flight 12, the fluctuations over 20 to 40 minute periods are about an order of magnitude greater than have been observed in the troposphere.<sup>2</sup>

The complicated pattern made by the trajectory of Flight 12 (see the latter part of this section), suggests the above result merely by inspection. The winds at floating level were quite light and since no tropospheric data

<sup>2</sup>The magnitude of the velocity fluctuations in the troposphere were found to be consistent with the observations of Durst, 1948: The fine structure of wind in the free air, QJRMS, vol. 74. Also Durst and Gilbert, 1950: Constant-height balloons--Calculations of geostrophic departures and ensuing discussion, QJRMS, vol. 76, 75-88.

under similar conditions were available, some doubt must remain that this condition is unique in the stratosphere. One further remark in this connection may be in order; both the cusp in the trajectory of Flight 12 and the sudden shift in direction of the trajectory of Flight 8 occur with anticyclonic motion of the trajectory. If the wind is anywhere approaching geostrophic balance \*\* means that the air decelerated as it was directed toward low pressure. One must conclude therefore that in this case, either the Coriolis acceleration is relatively unimportant, which would be true in the flow about very small pressure systems, or that the pressure field is moving rapidly with respect to the wind velocity. Since the complicated trajectory patterns, illustrated by Flight 12, are generally associated with light winds, it may be that these are typical observations in the layer of very light winds which is found between the lower westerly and upper easterly currents throughout the summer half of the year.

### 3. A Case Study of the Vorticity Change

The measurement of atmospheric accelerations by the constant altitude balloon may make possible a more complete description of the atmospheric flow processes. For example, the relationship of the accelerations to be divergence may be expressed by the vorticity equation. The usefulness of the measurement of the accelerations in this connection, is being examined by means of a case study of the vorticity change for a twenty-four hour period corresponding to General Mills Flight 578. The three dimensional field of motion has been analyzed in the layer 400 mb to 200 mb; the vertical velocity has been computed by the adiabatic method outlined by Panofsky (1946). From these values, cross sections of the horizontal divergence were prepared following Fleagle, (1948). This information and the measurements of the balloon trajectory will then be used in the computation of the vorticity change and independent values of the divergence will be computed from the vorticity equation.

The vorticity equation has been derived in vector form following the suggestion of Hsu (1951), the equation obtained is

$$\frac{d}{dt} (\zeta + f) + (\zeta + f) \nabla \cdot \vec{v} - (\nabla \times \vec{v} + 2\vec{\Omega}) \cdot \nabla_2 w = k \cdot \nabla_2 p \times \nabla_2 \alpha,$$

where the symbols have their conventional meteorological significance:

$(\zeta + f)$  is the vertical component of the absolute vorticity;

$\vec{v}$  the velocity vector;

$\vec{\Omega}$  the angular rotation of the earth;

$w$  the vertical component of the velocity;

$k$  the unit vector in the zenith direction;

$p$  and  $\alpha$  are pressure and specific volume of the air, respectively;

$\nabla_2$  and  $\nabla$  are the horizontal and three dimensional del operators.

The advantage of this derivation is the limited number of assumptions required;

the principal assumptions retained here are:

1. That the frictional forces may be neglected;
2. That the gravitational acceleration does not vary in the tangent plane;
3. That the earth is spherical.

The evaluation of vorticity change will be made in two ways; by measurement of the vorticity difference at the beginning and end of the trajectory and by measuring individual terms when expressed in natural coordinates, as suggested by Hsu (1951). It is hoped thereby to determine which features of the accelerations are important in producing the divergence and to provide a better estimate of the accuracy and type of measurement required for a complete computation.

#### 4. Computation of Height and Temperature in the Stratosphere

The balloon elevation is being computed from the photographs taken from the balloon by comparing distances on the film with map distances. The basic limitation to the accuracy of this procedure is the measurement of true or map distance. With a map scale of 1:50,000 the error of an appropriate dis-

tance seems to be at best 0.2% and possibly nearer 0.5% on the average. Somewhat greater accuracy may be possible by using the map only to identify clearly marked section boundaries and assuming the distance between to be exactly one mile.

The balloon elevation computed from the pictures of Flight 8 is compared with the pressure altitude computed from the standard atmosphere in Figure 1. The interval between photographs was found to vary slightly over the period of the flight and the time scale used here was interpolated by comparing the theodolite and photo trajectories. The time altitude curve in Figure 1 is relatively smooth so that it seems likely that large changes in ascent rate can be detected by this method. The oscillation of the values of elevation computed from the photographs about the values determined by the pressure reports, however, suggests that camera sway is not negligible.

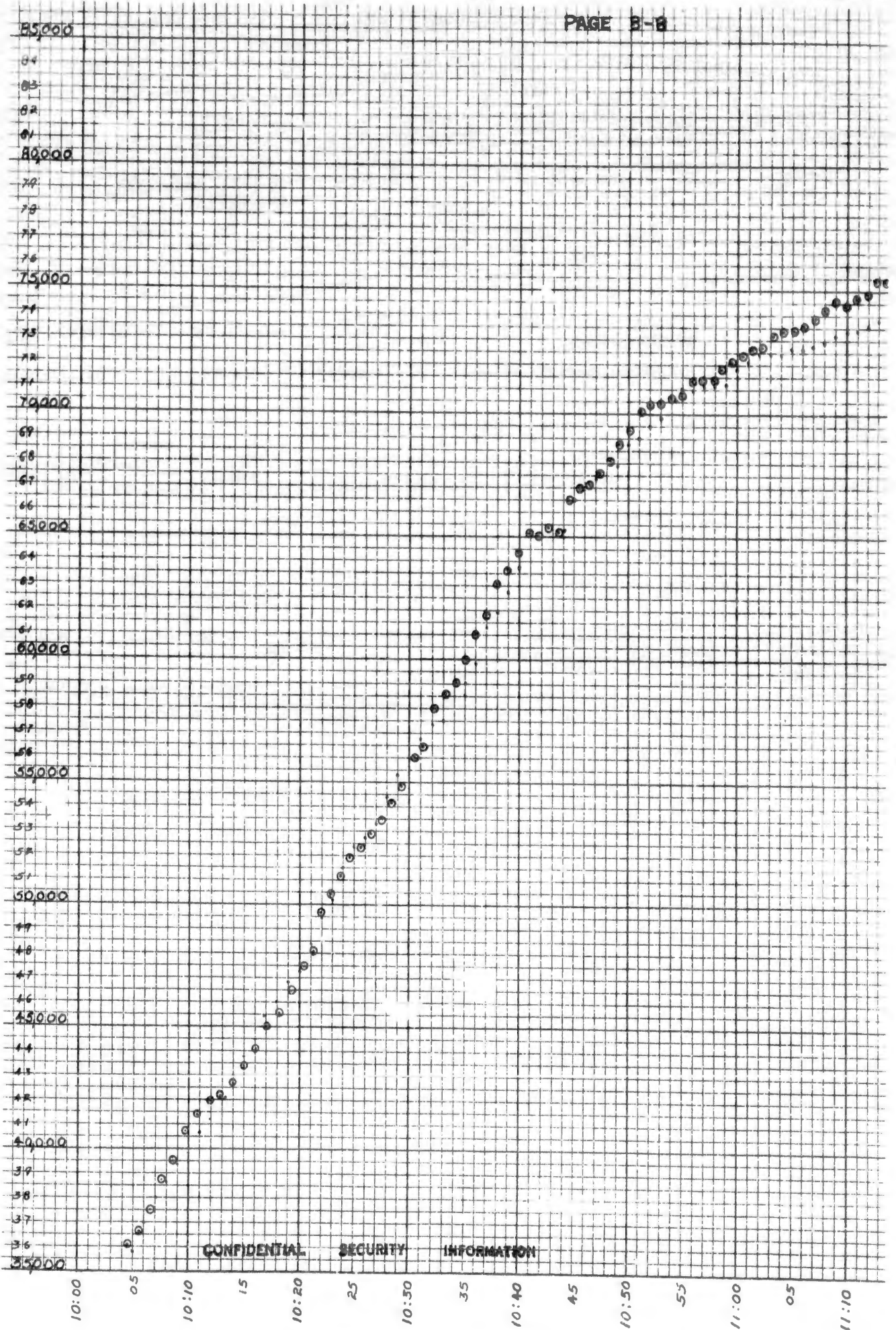
The mean temperature of any layer may also be computed from the elevation and the pressure observations by means of the hydrostatic equation. The temperature distribution resulting from this computation for Flight 8, is shown in Figure 2 where the temperature reported by St. Cloud, the nearest radiosonde observation station, is given for comparison. It is probably a safe assumption that the St. Cloud temperature at 200 mb is correct to within a few degrees since both the height computation and temperature are consistent with the nearby radiosonde reports.

The results of the temperature computation were found to be particularly dependent upon a good time coordination of the various observations. A sample calculation of the effect of the probable error in timing showed that a variation in 10 degrees could be expected from this error alone. In view of the difficulties of instrumentation and measurement encountered in a satisfactory height and temperature determination by this method, it seems desirable to have a further elevation check (for example by double theodolite) during the balloon ascent.

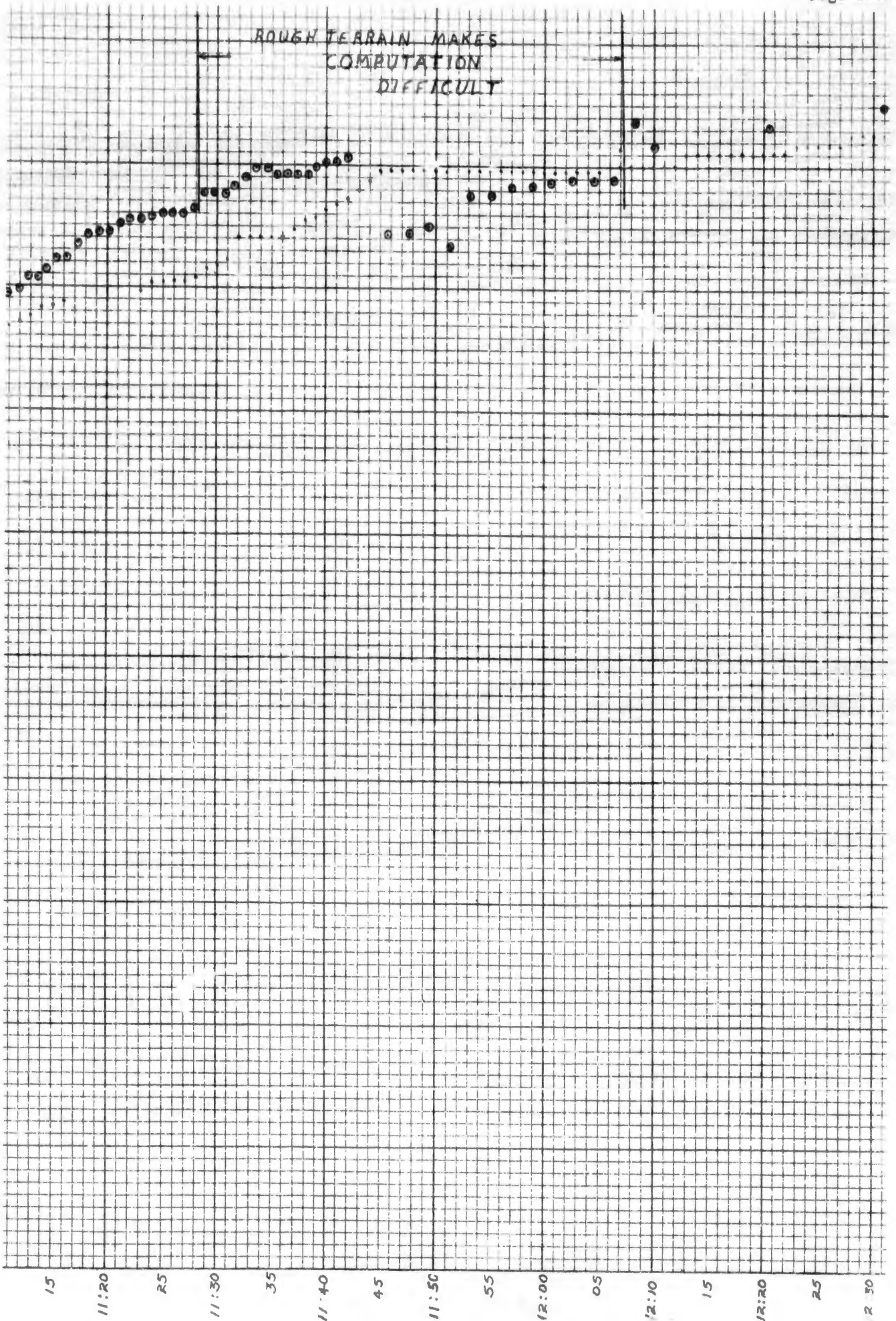
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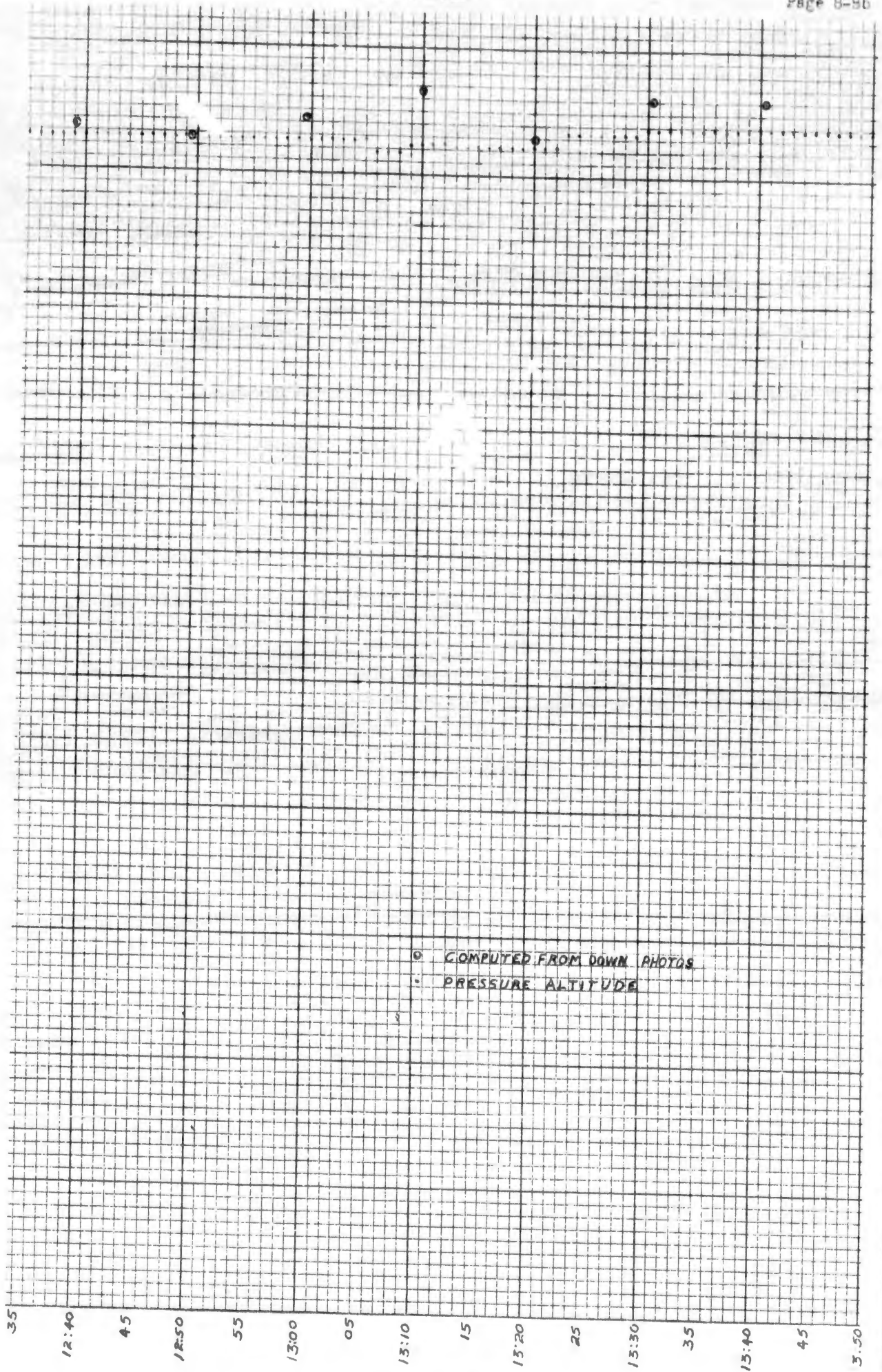
- Durst, C. S., 1948: The fine structure of the wind in the free air, QJRMS, vol. 74.
- , and G. H. Gilbert, 1950: Constant-height balloons--Calculations of geostrophic departures and discussion ensuing, QJRMS, vol. 76, 75-88.
- Fleagle, R. G., 1948: Quantitative factors influencing pressure change, Journal of Meteorology, vol. 5, 281-292.
- Hsu, E. H., 1951: A general equation of horizontal mass divergence in the atmosphere, Journal of Meteorology, vol. 8, 395-397.
- Palmen, E., and K. M. Nagler, 1948: An analysis of the wind and temperature distribution in the free atmosphere over North America in a case of approximately westerly flow, Journal of Meteorology, vol. 5, 64.
- Panofsky, H. A., 1946: Methods of computing vertical motion in the atmosphere, Journal of Meteorology, 3:271-275.

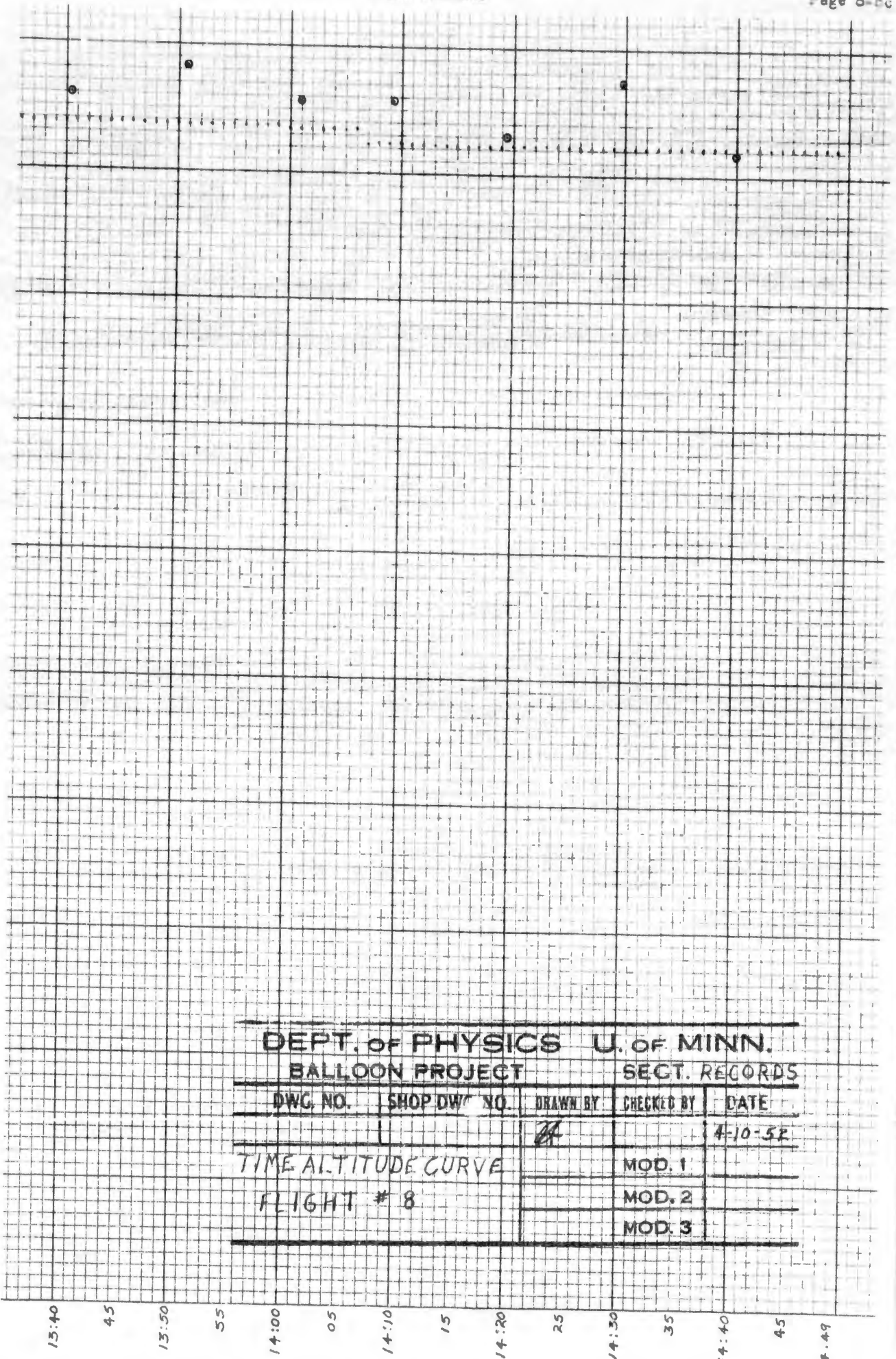
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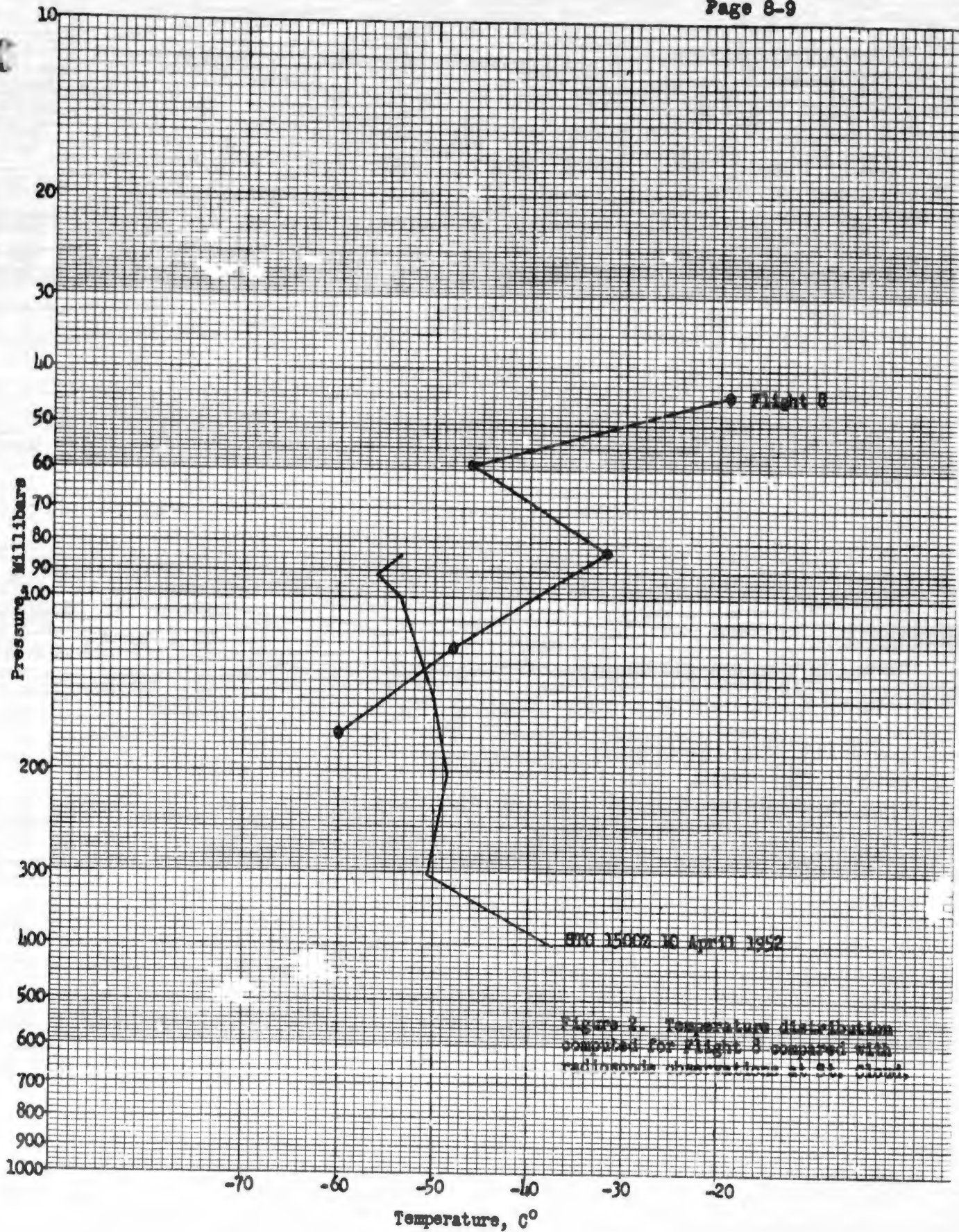
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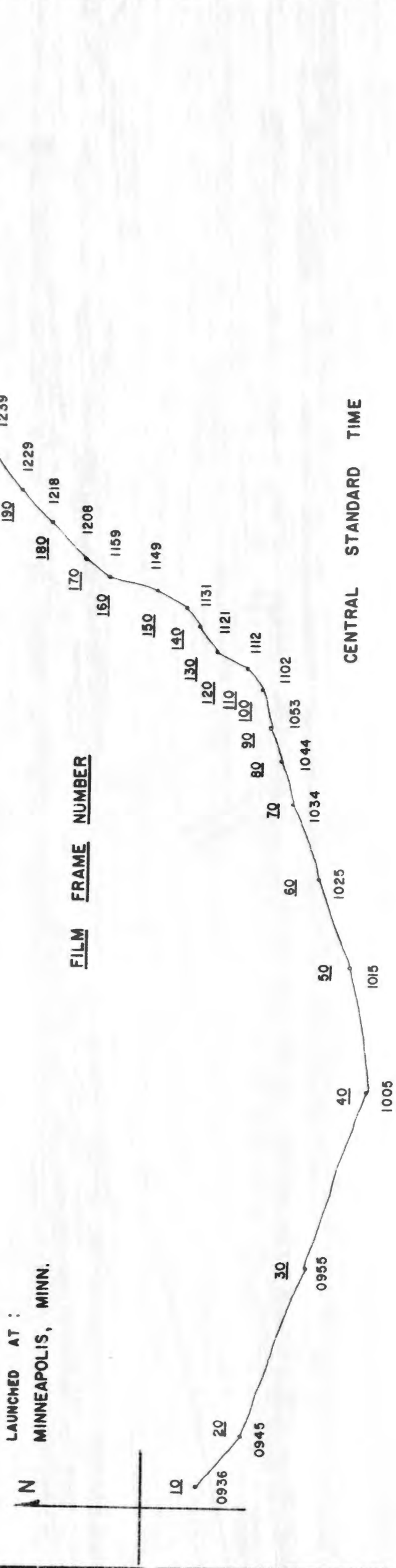


DEPT. OF PHYSICS		U. OF MINN.		
BALLOON PROJECT		SECT. RECORDS		
DWG. NO.	SHOP DWG. NO.	DRAWN BY	CHECKED BY	DATE
		<i>[Signature]</i>		4-10-52
TIME ALTITUDE CURVE			MOD. 1	
FLIGHT # 8			MOD. 2	
			MOD. 3	



870 1500Z 10 April 1952

Figure 2. Temperature distribution computed for Flight 8 compared with standard atmosphere at 20,000 feet

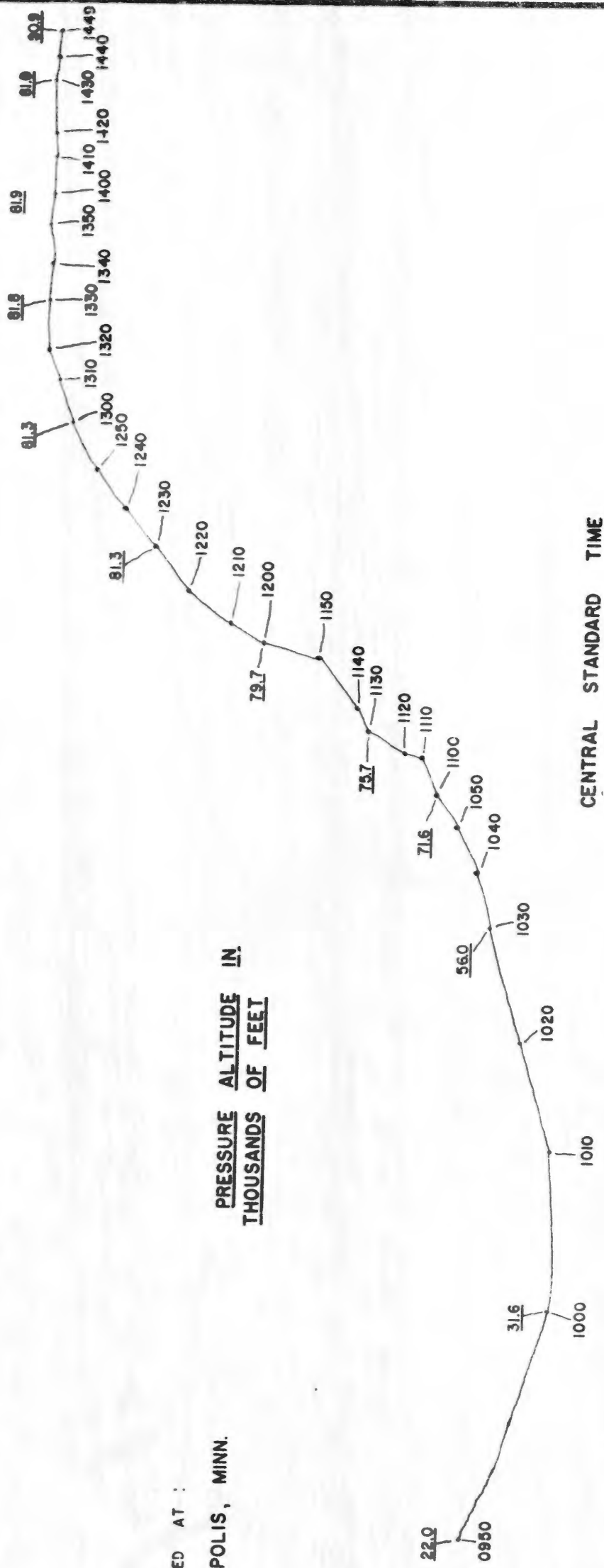


FLIGHT NO. 8  
 Launched 0926 CST  
 10 APRIL, 1952  
 TRAJECTORY FROM DOWN PICTURE  
 SCALE 1:5 x 10<sup>5</sup>



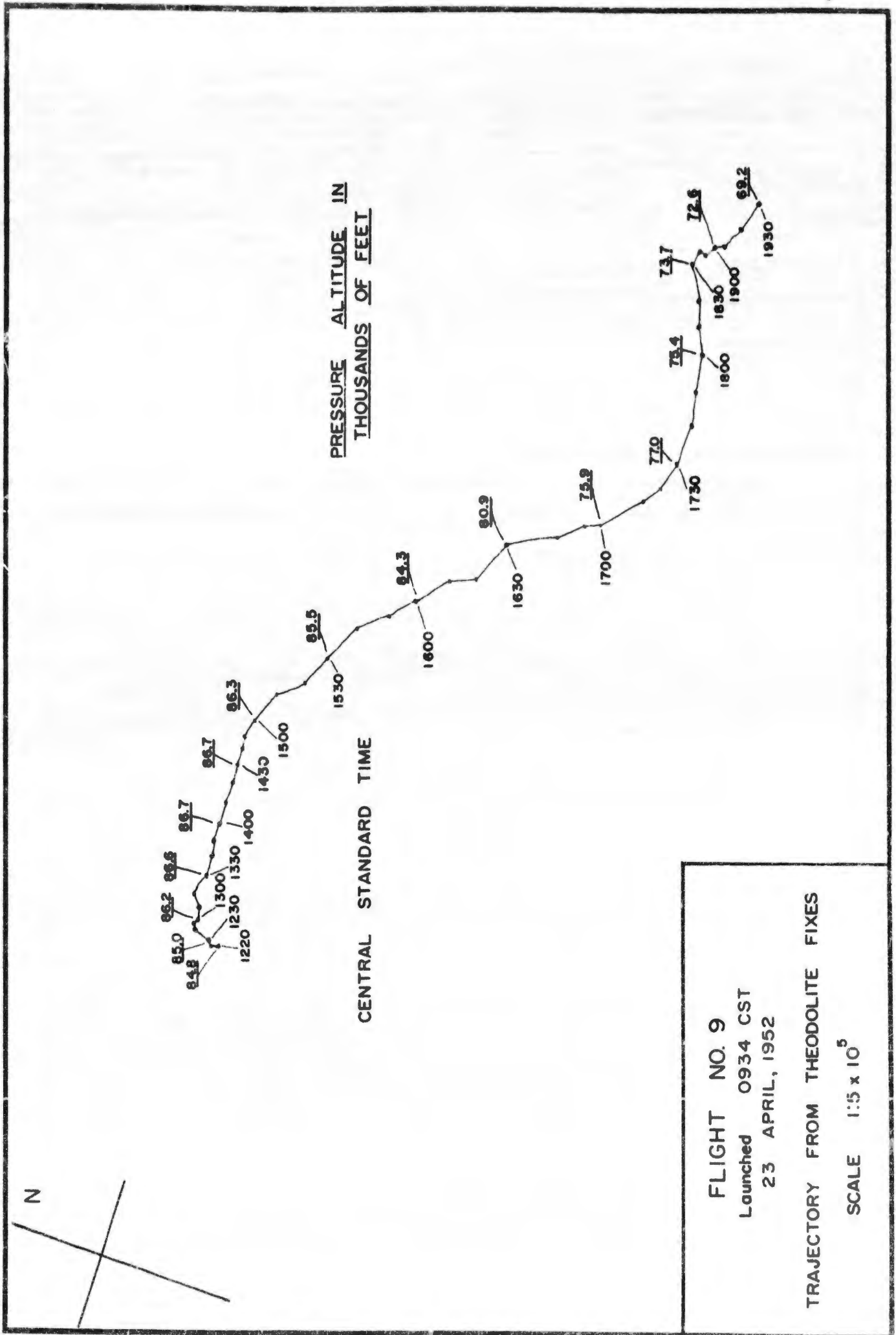
LAUNCHED AT :  
MINNEAPOLIS, MINN.

PRESSURE ALTITUDE IN  
THOUSANDS OF FEET

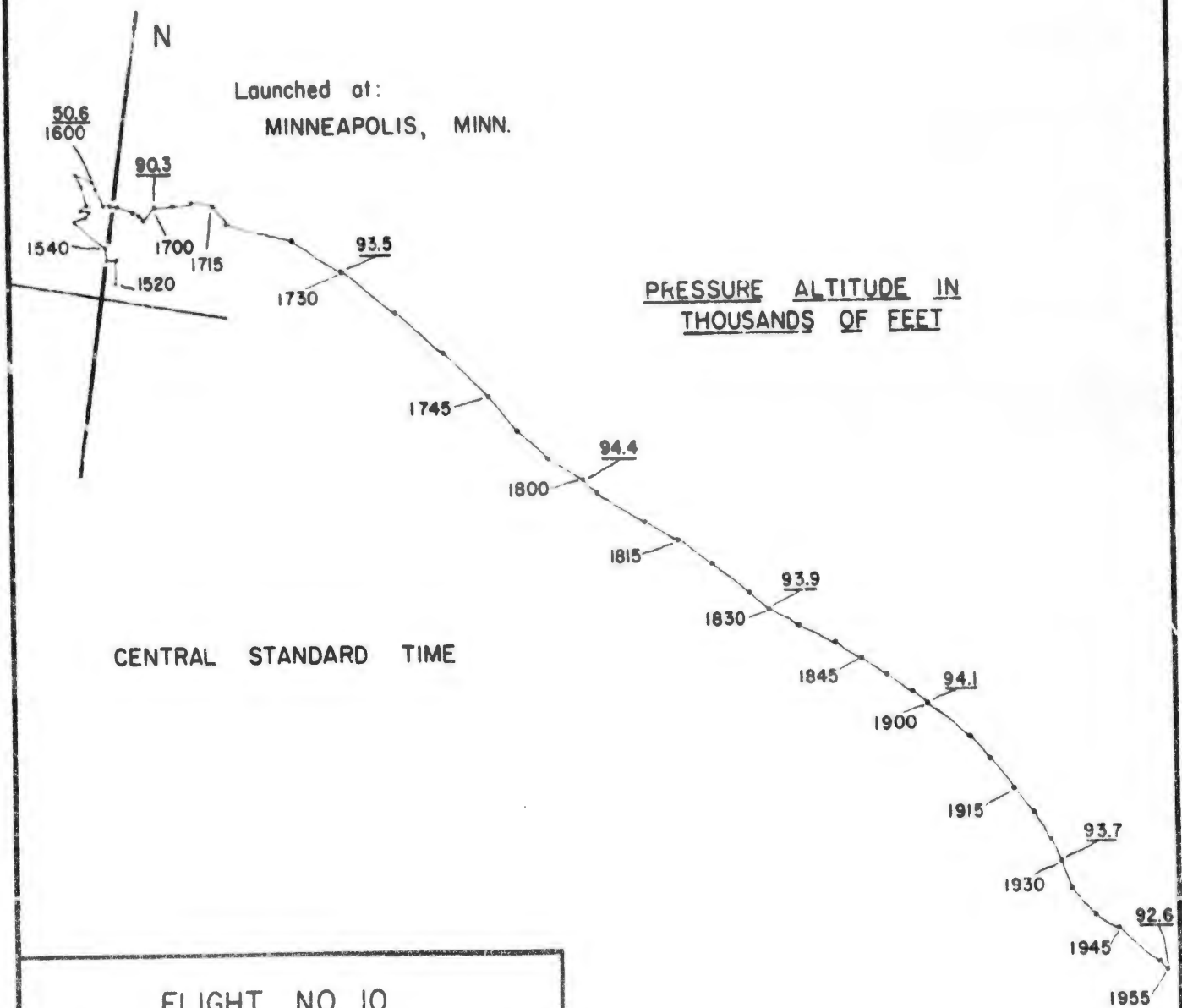


CENTRAL STANDARD TIME

FLIGHT NO. 8  
 Launched 0926 CST  
 10 APRIL, 1952  
 TRAJECTORY FROM THEODOLITE FIXES  
 Scale : 1 : 5 x 10<sup>5</sup>

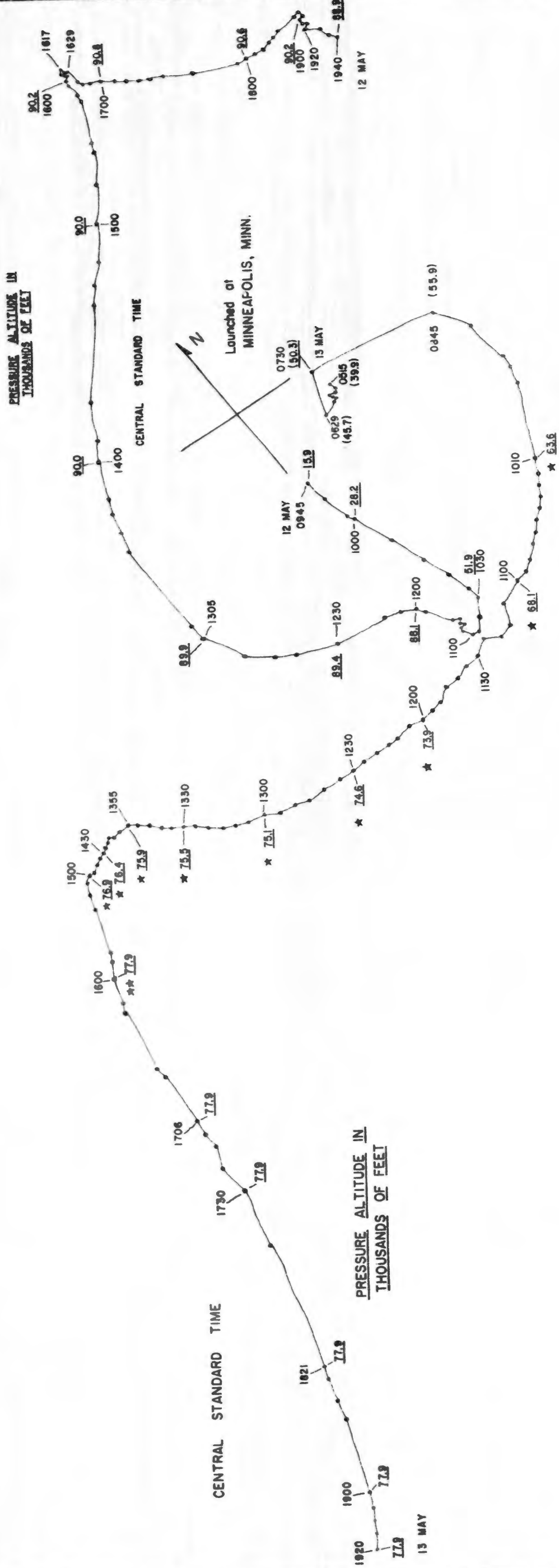


FLIGHT NO. 9  
 Launched 0934 CST  
 23 APRIL, 1952  
 TRAJECTORY FROM THEODOLITE FIXES  
 SCALE 1:5 x 10<sup>5</sup>



FLIGHT NO. 10  
Launched at 1506 CST  
29 APRIL, 1952  
TRAJECTORY FROM THEODOLITE FIXES  
SCALE  $1:5 \times 10^5$

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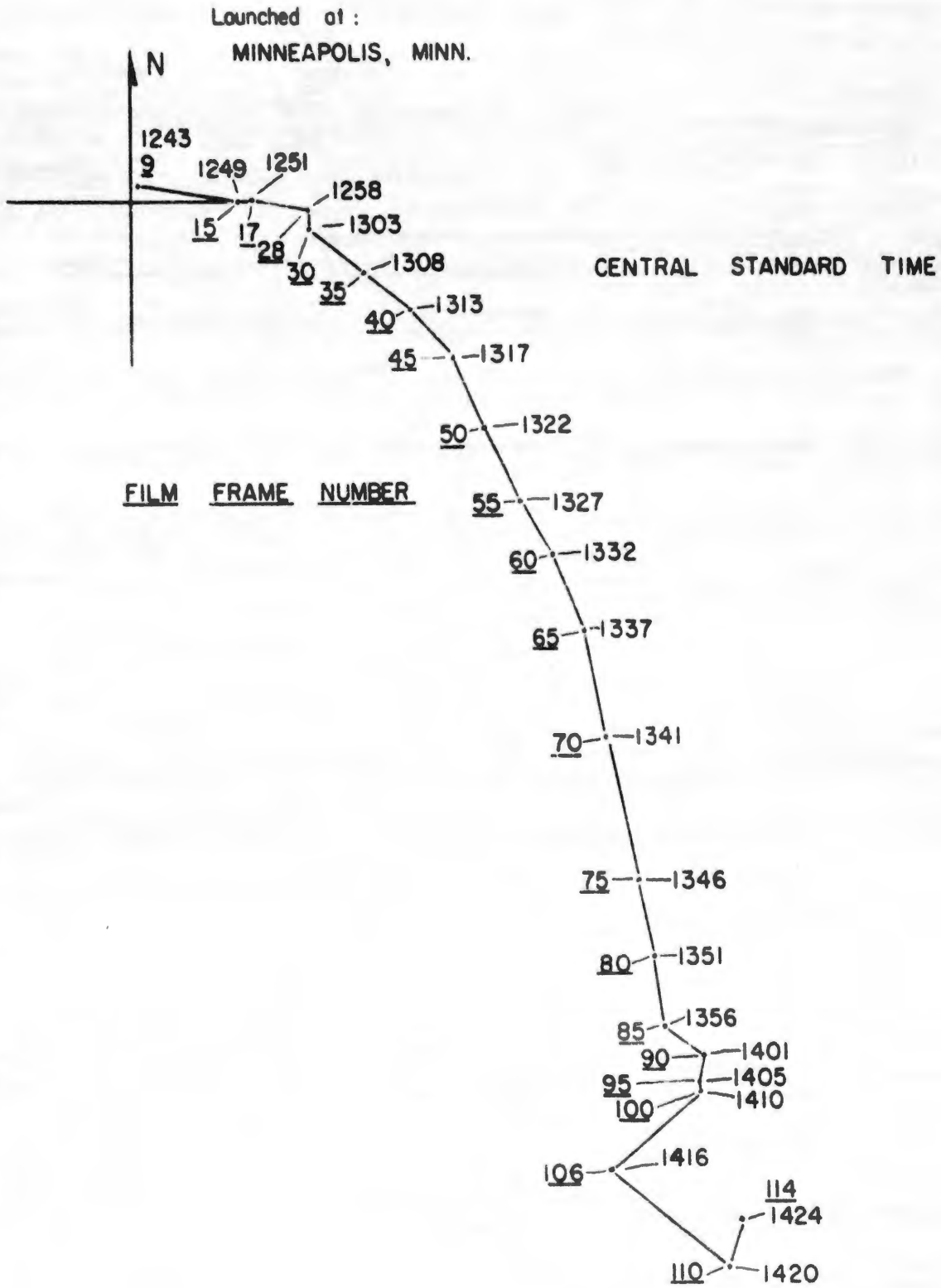
FLIGHT NO. 12  
 Launched 0929 CST  
 12 & 13 MAY, 1952

CODE

( ) ELEVATION FROM SUBTENDED ANGLE  
 \* ELEVATION FROM DOUBLE THEODOLITE  
 \*\* NO OBSERVATIONS AFTER 1600

SCALE 1:5 x 10<sup>6</sup>

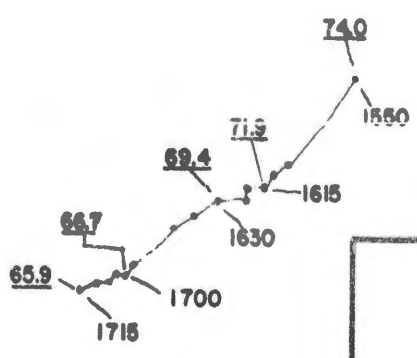
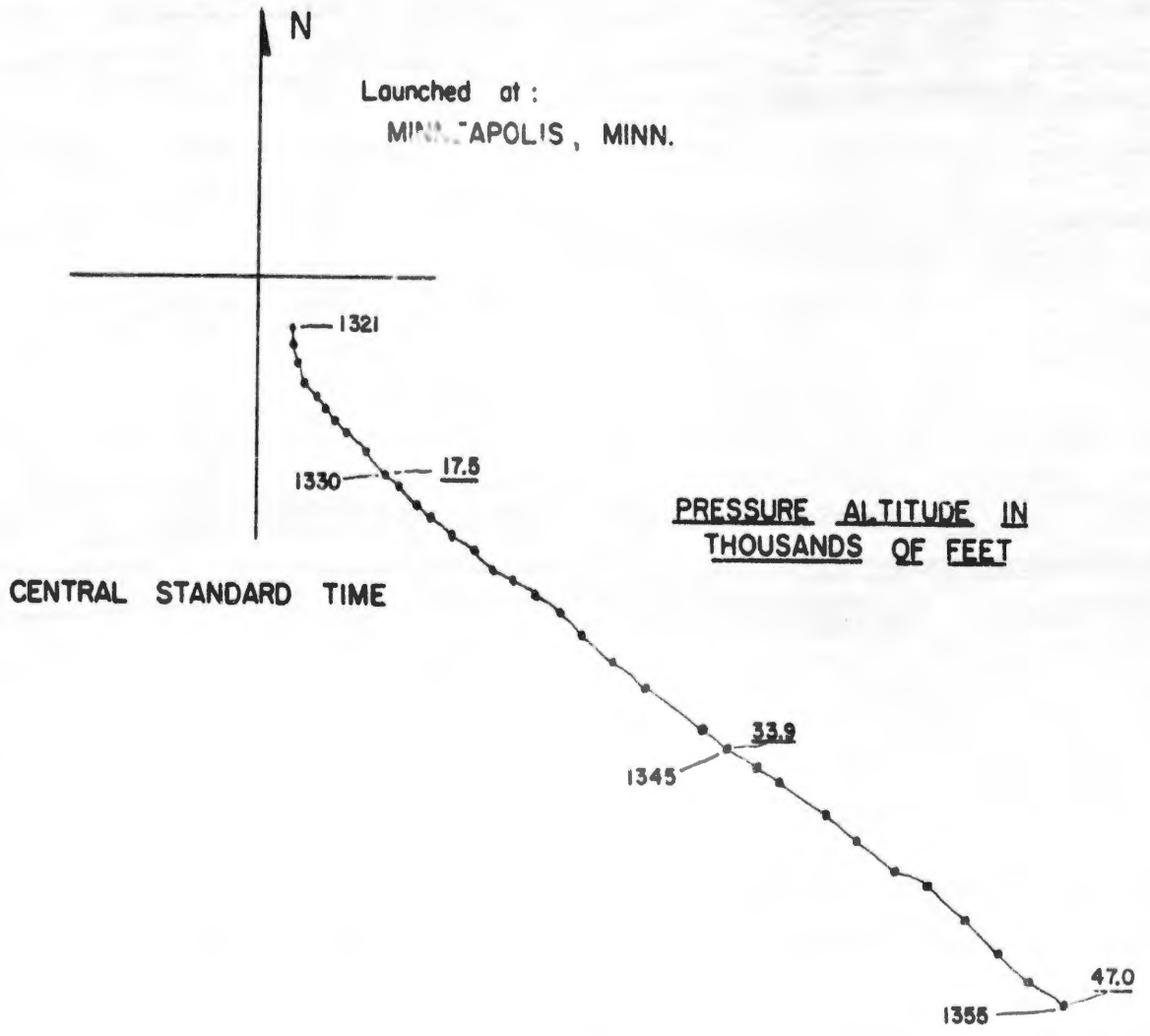
CONFIDENTIAL SECURITY INFORMATION



FLIGHT NO.17  
 Launched 1234 CST  
 5 JUNE, 1952

TRAJECTORY FROM DOWN PICTURES

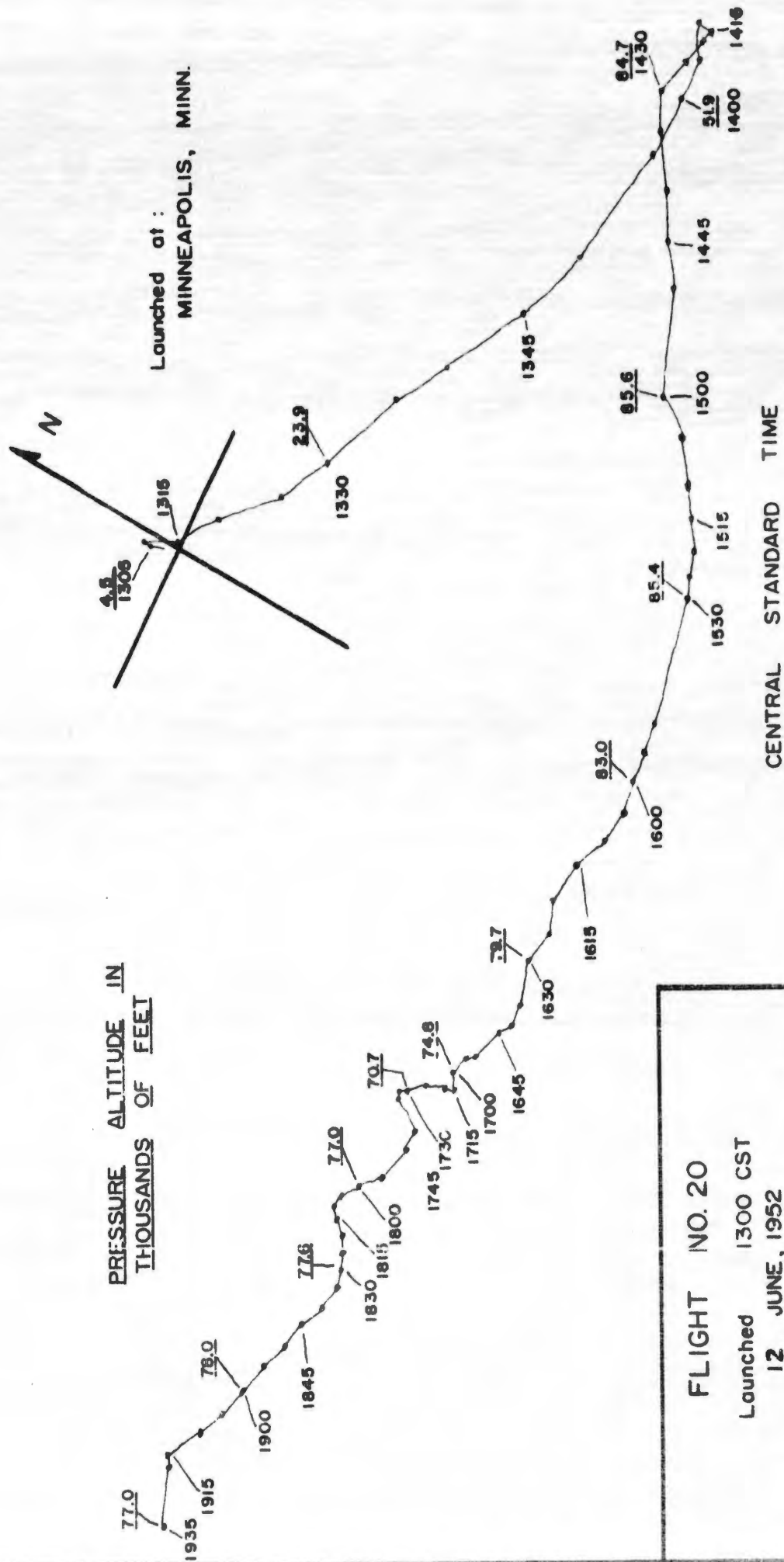
SCALE 1:5 x 10<sup>5</sup>



**FLIGHT NO. 19**  
 Launched 1313 CST  
 10 JUNE, 1952

TRAJECTORY FROM THEODOLITE FIXES

SCALE  $1:5 \times 10^5$



FLIGHT NO. 20  
 Launched 1300 CST  
 12 JUNE, 1952  
 TRAJECTORY FROM THEODOLITE FIXES  
 SCALE 1:5 x 10<sup>5</sup>

-----  
THEORETICAL CEILING  
85,000 FT.

TIME - PRESSURE - ALTITUDE

FLIGHT NO. 4

MARCH 1, 1952

GROSS LOAD 409 LB.

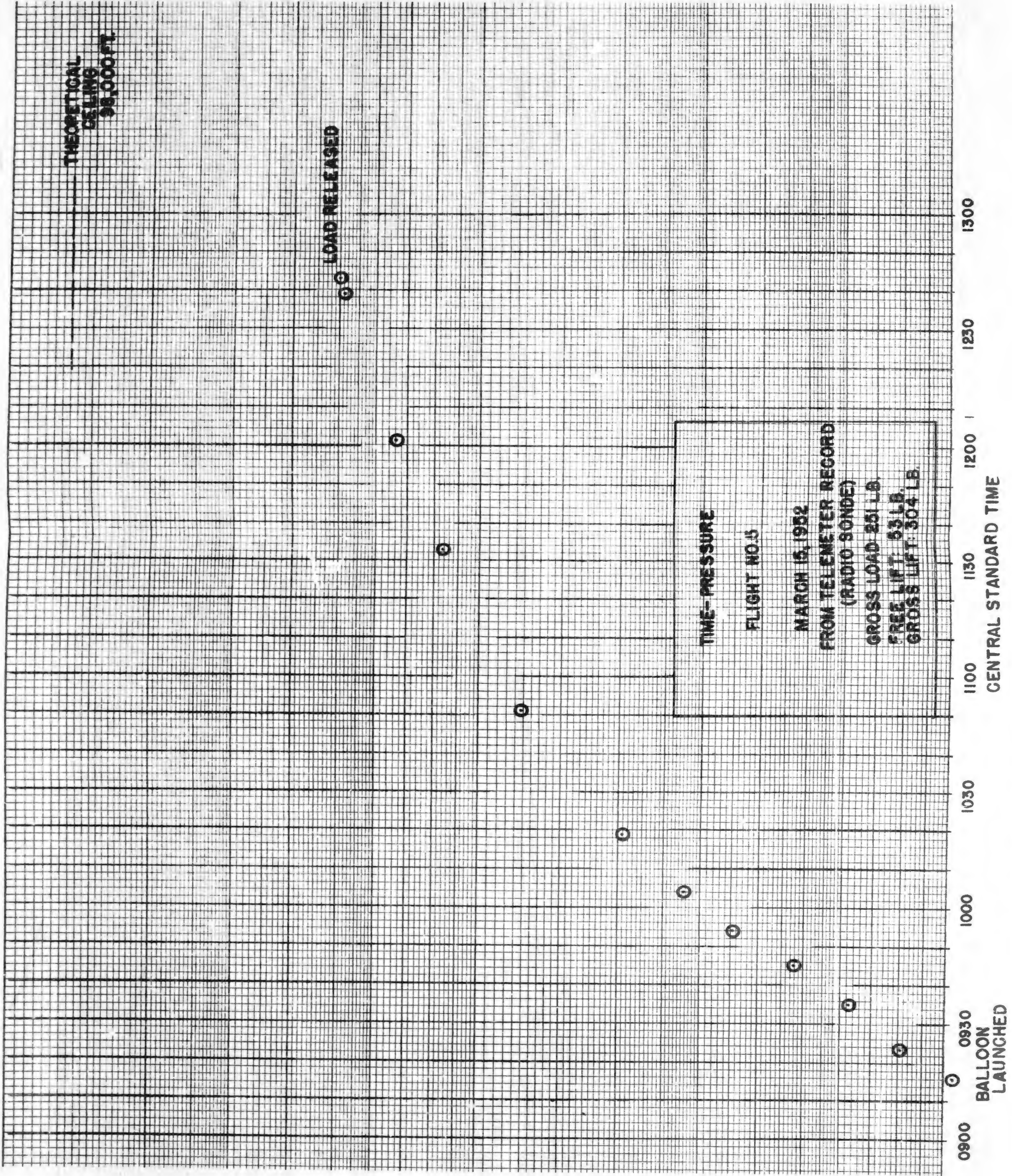
FREE LIFT 96 LB.

GROSS LIFT 505 LB.

DATA FROM  
SINGLE COUNTER TELEMETER

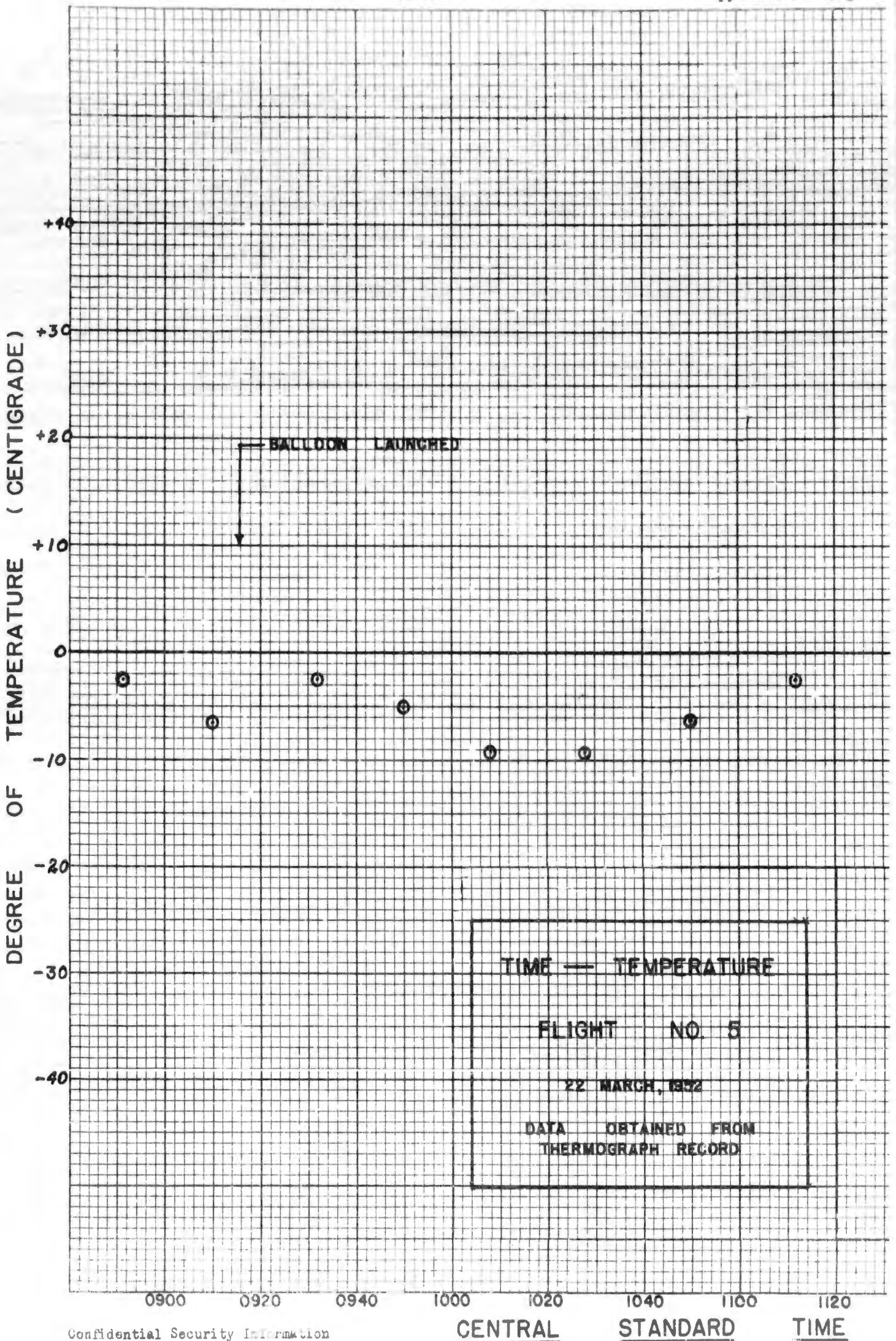
BALLOON  
LAUNCHED

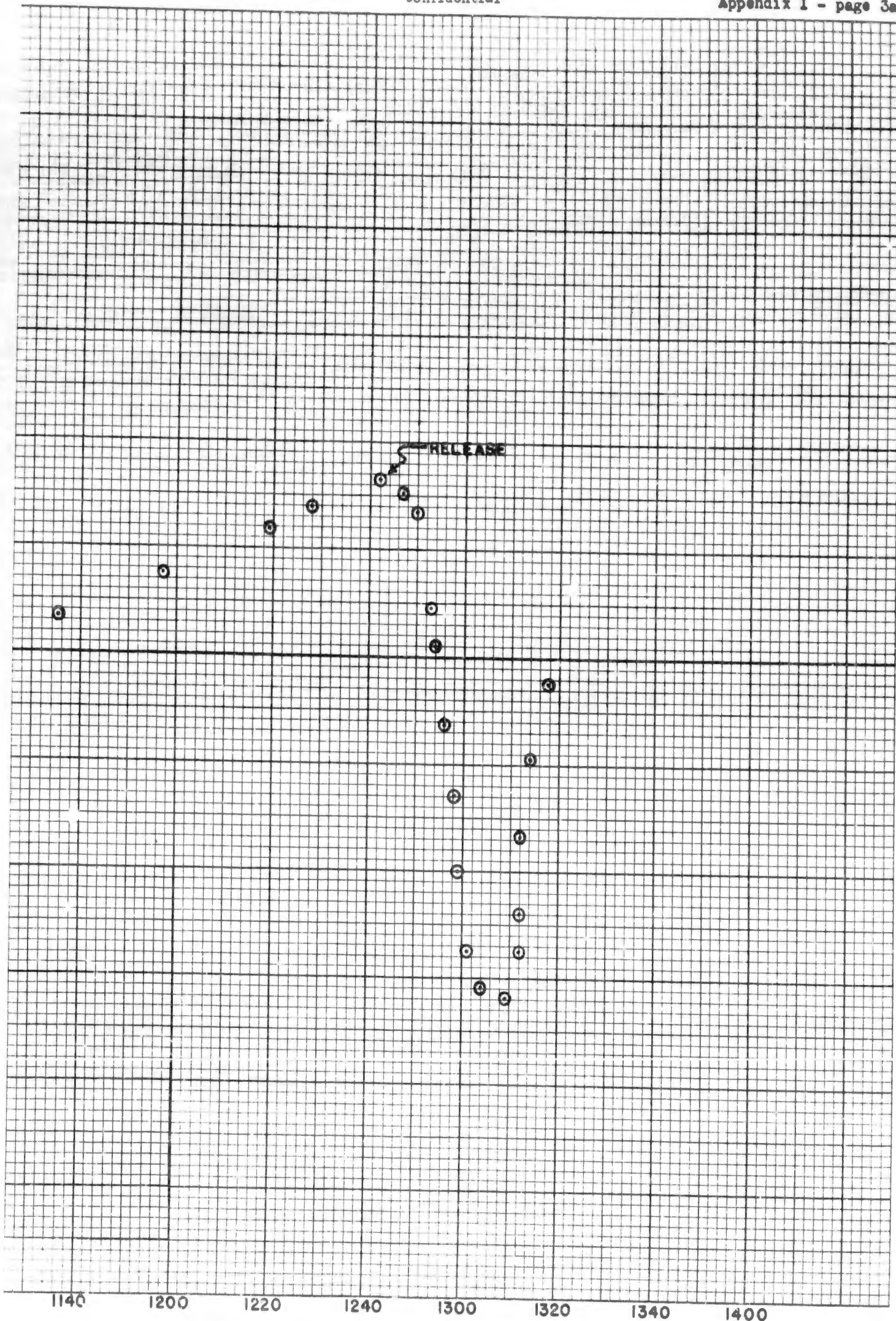
0900 0920 0940 1000 1020 1040 1100 1120 1140  
CENTRAL STANDARD TIME

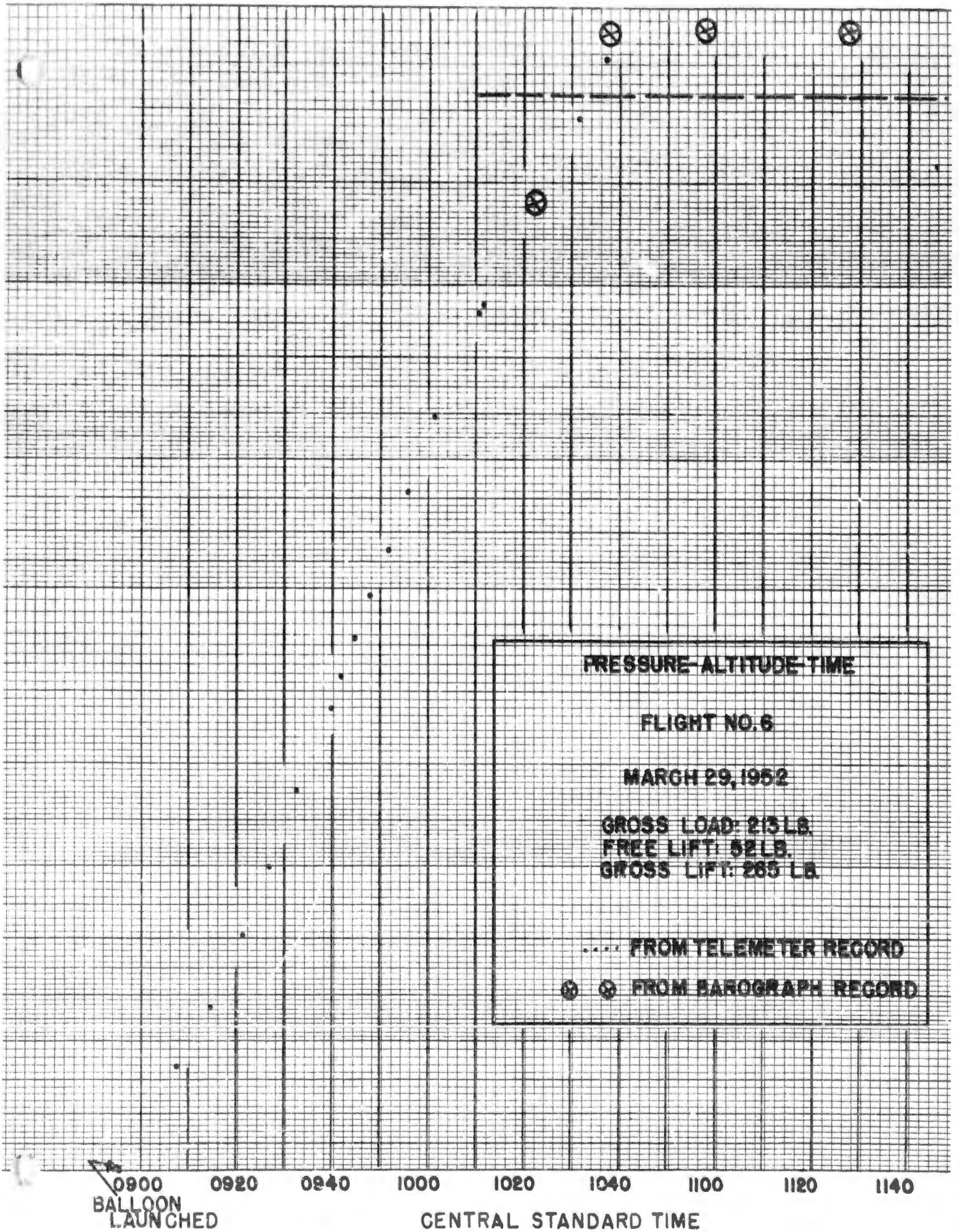


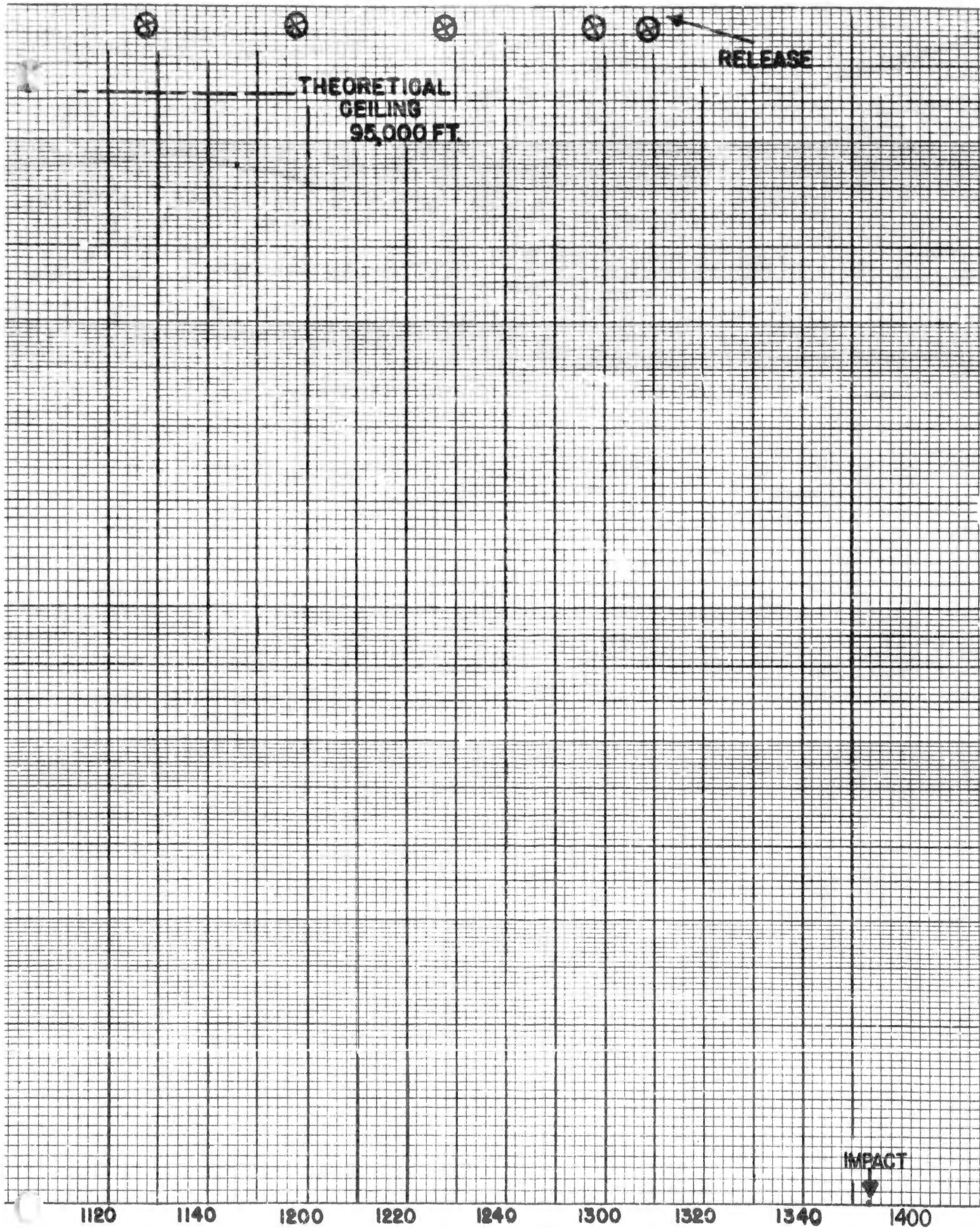
TIME-PRESSURE  
 FLIGHT NO. 5  
 MARCH 15, 1952  
 FROM TELEMETER RECORD  
 (RADIO SONDE)  
 GROSS LOAD: 251 LB.  
 FREE LIFT: 53 LB.  
 GROSS LIFT: 304 LB.

0900 0930  
 BALLOON LAUNCHED  
 1000 1030 1100 1130 1200 1230 1300  
 CENTRAL STANDARD TIME

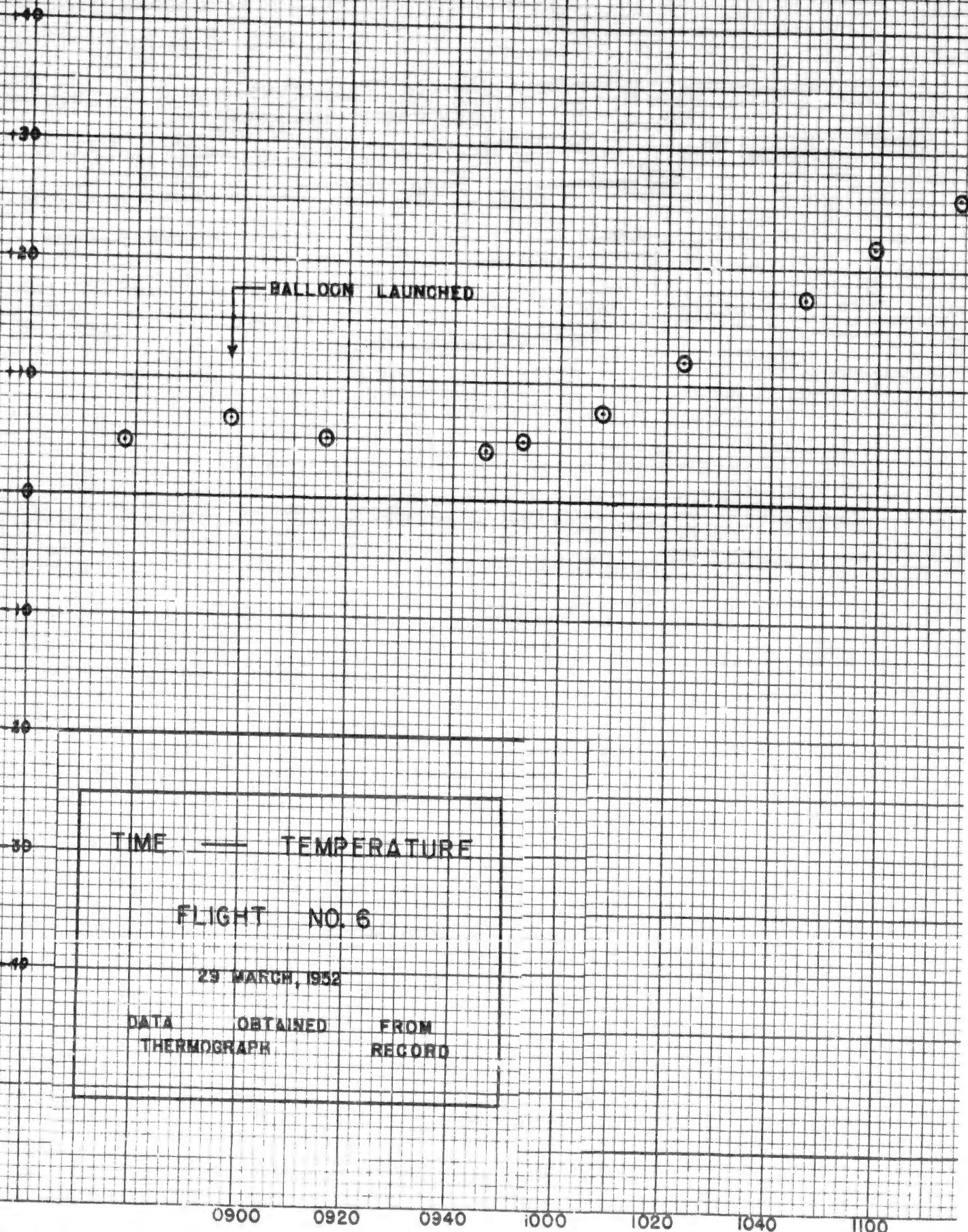




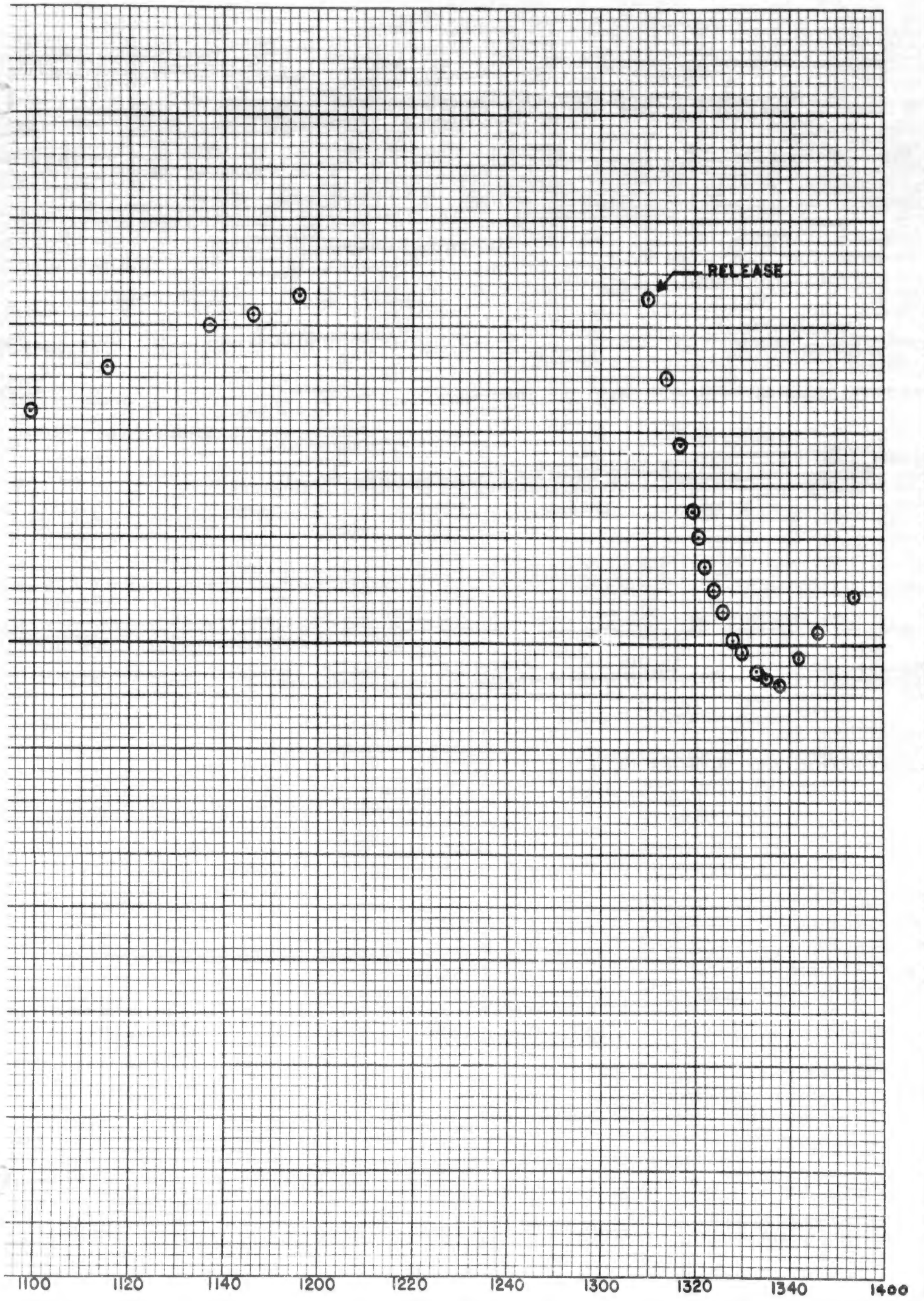




DEGREE OF TEMPERATURE (CENTIGRADE)



TIME — TEMPERATURE  
FLIGHT NO. 6  
29 MARCH, 1952  
DATA OBTAINED FROM THERMOGRAPH RECORD



ALTITUDE IN THOUSAND FEET

TIME - PRESSURE - ALTITUDE

FLIGHT NO. 8

APRIL 10 - 1

GROSS LOAD - 345 LBS.

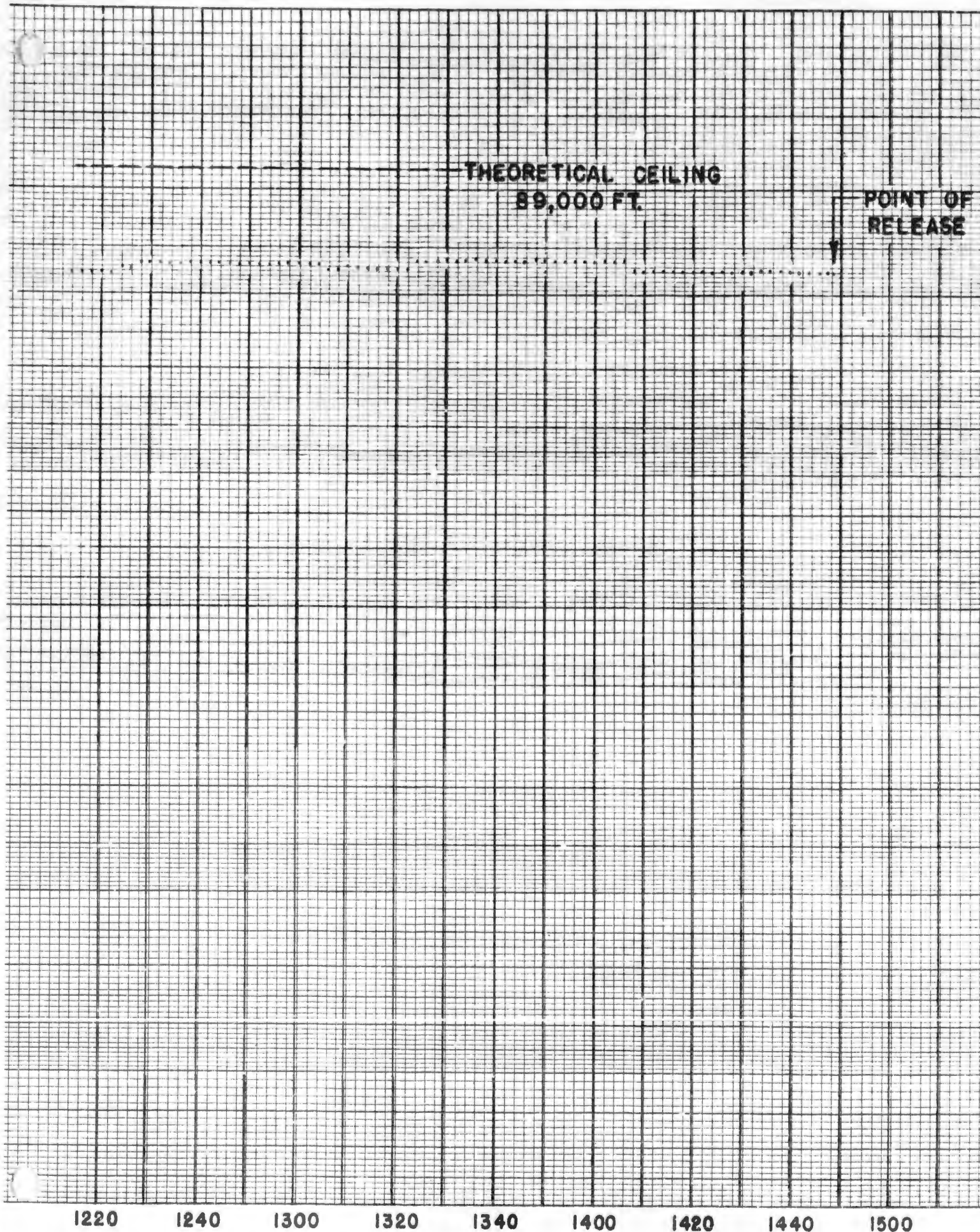
FREE LIFT - 68 LBS.

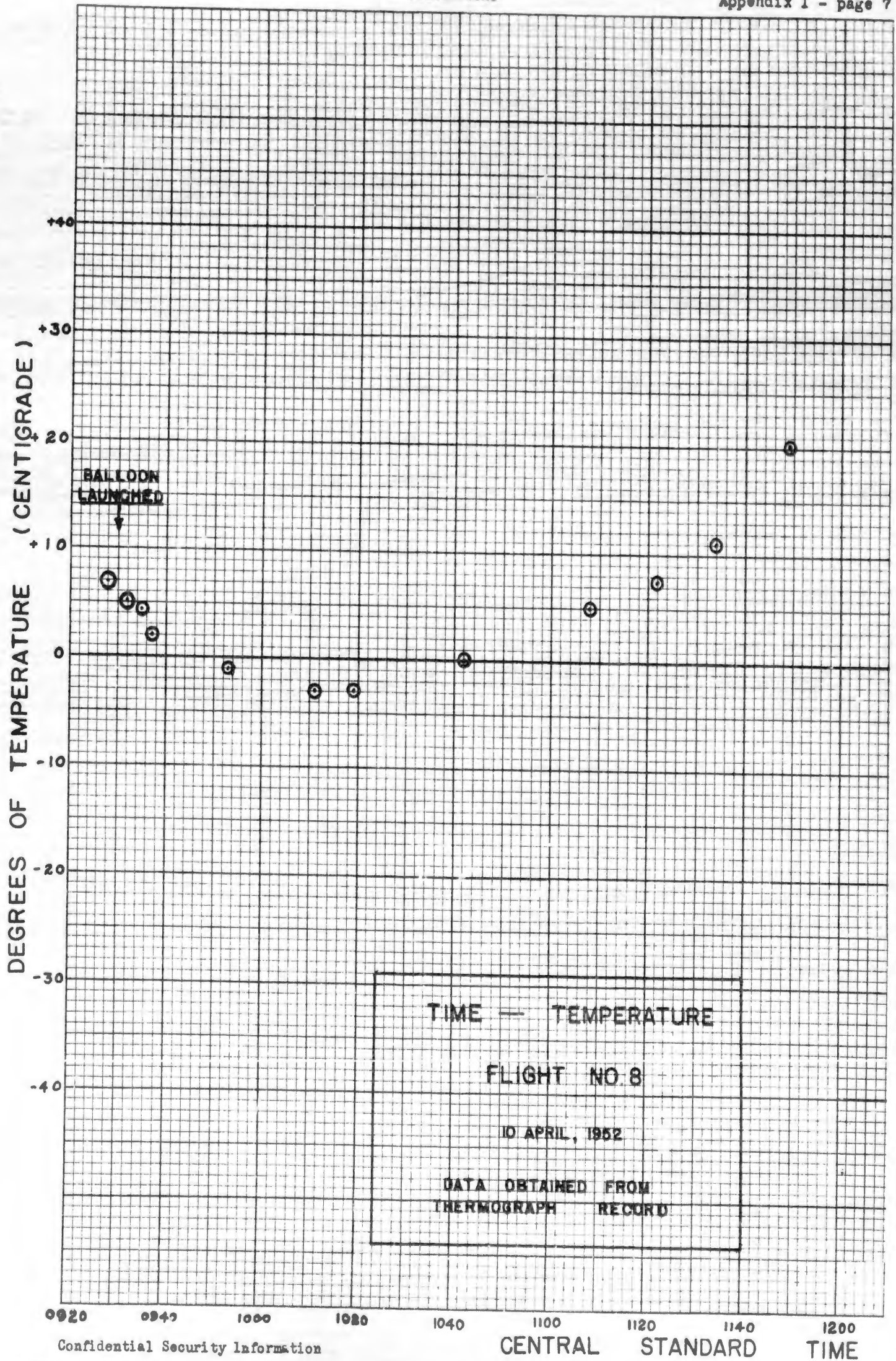
GROSS LIFT - 413 LBS.

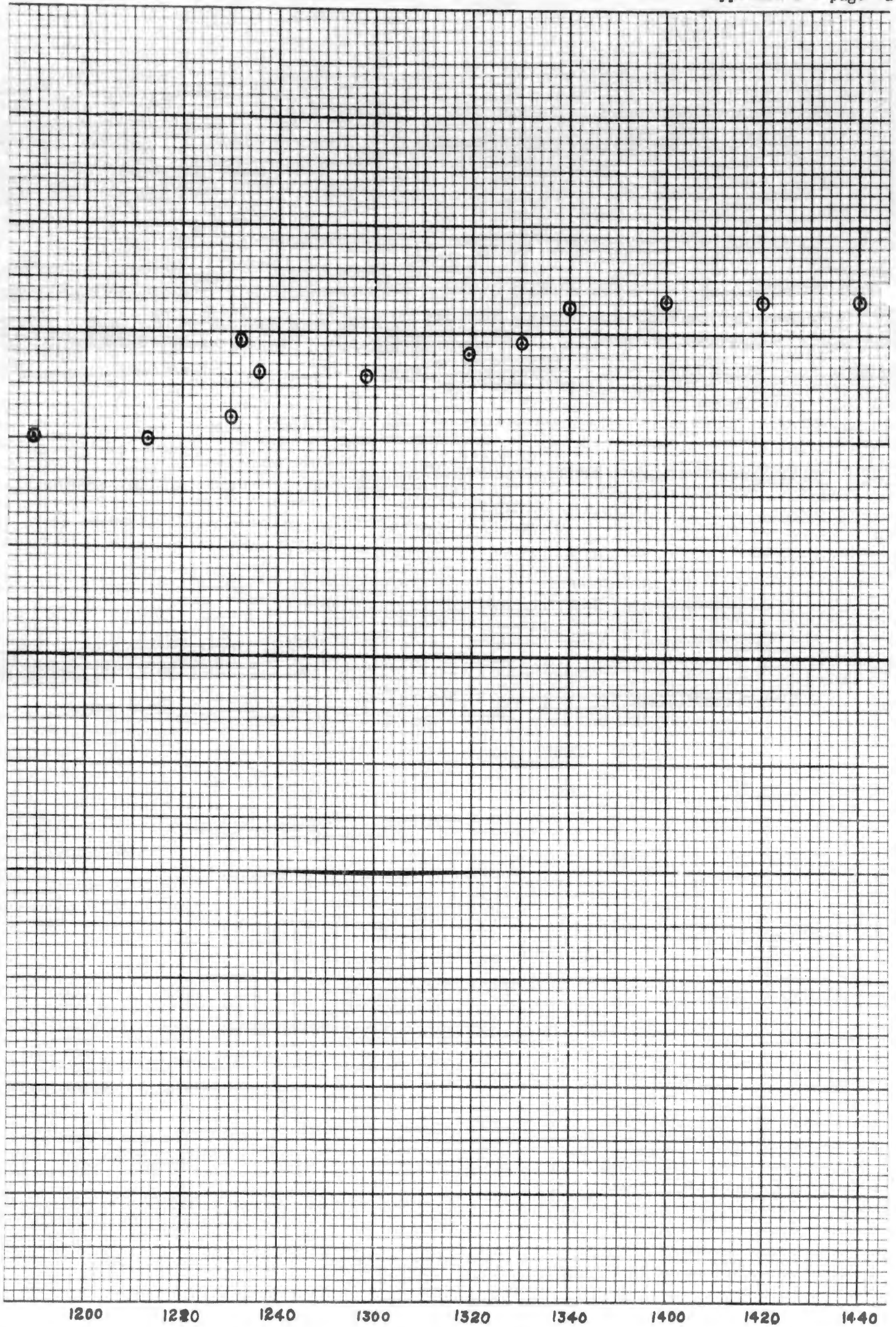
DATA FROM TELEMETER RECORD

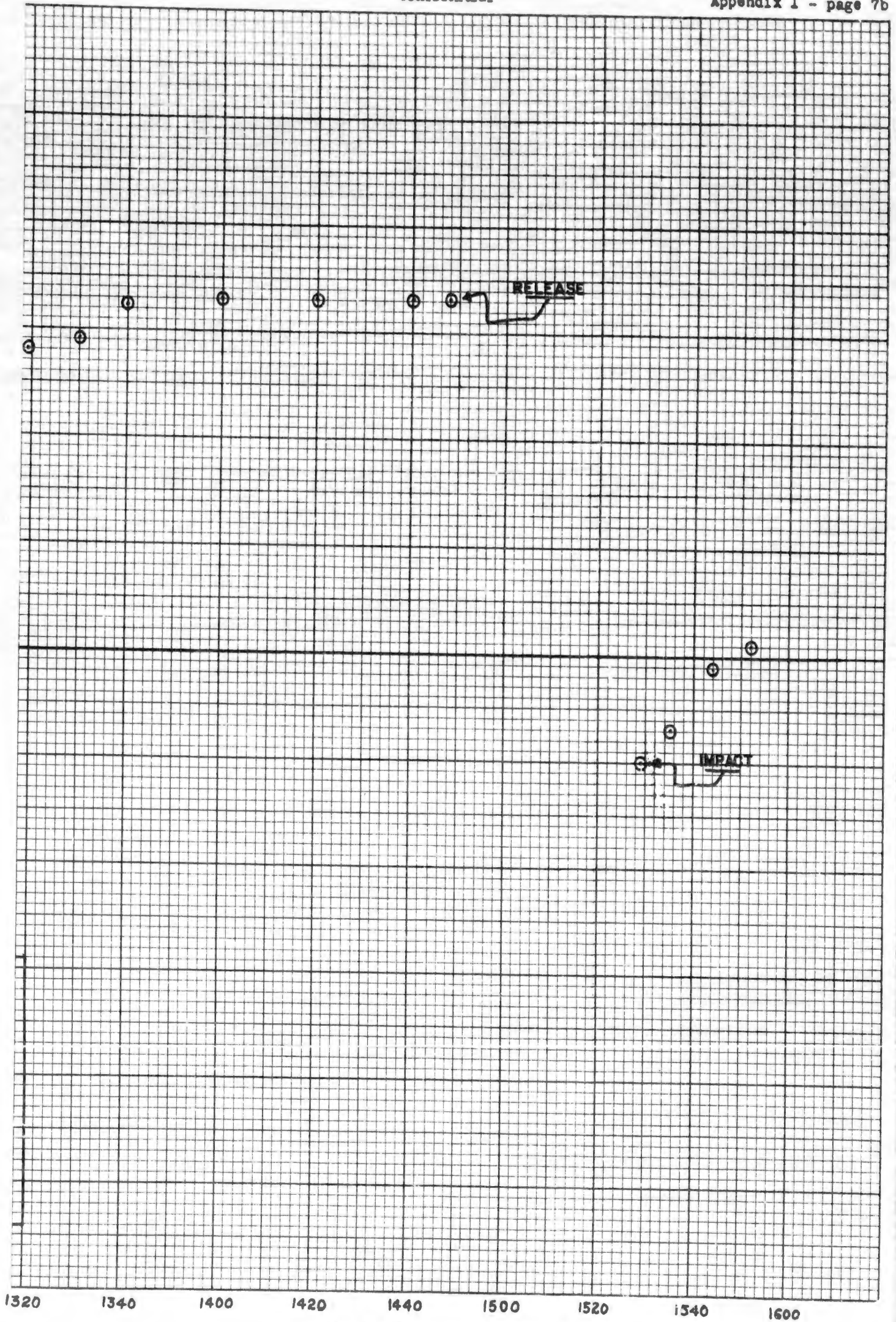
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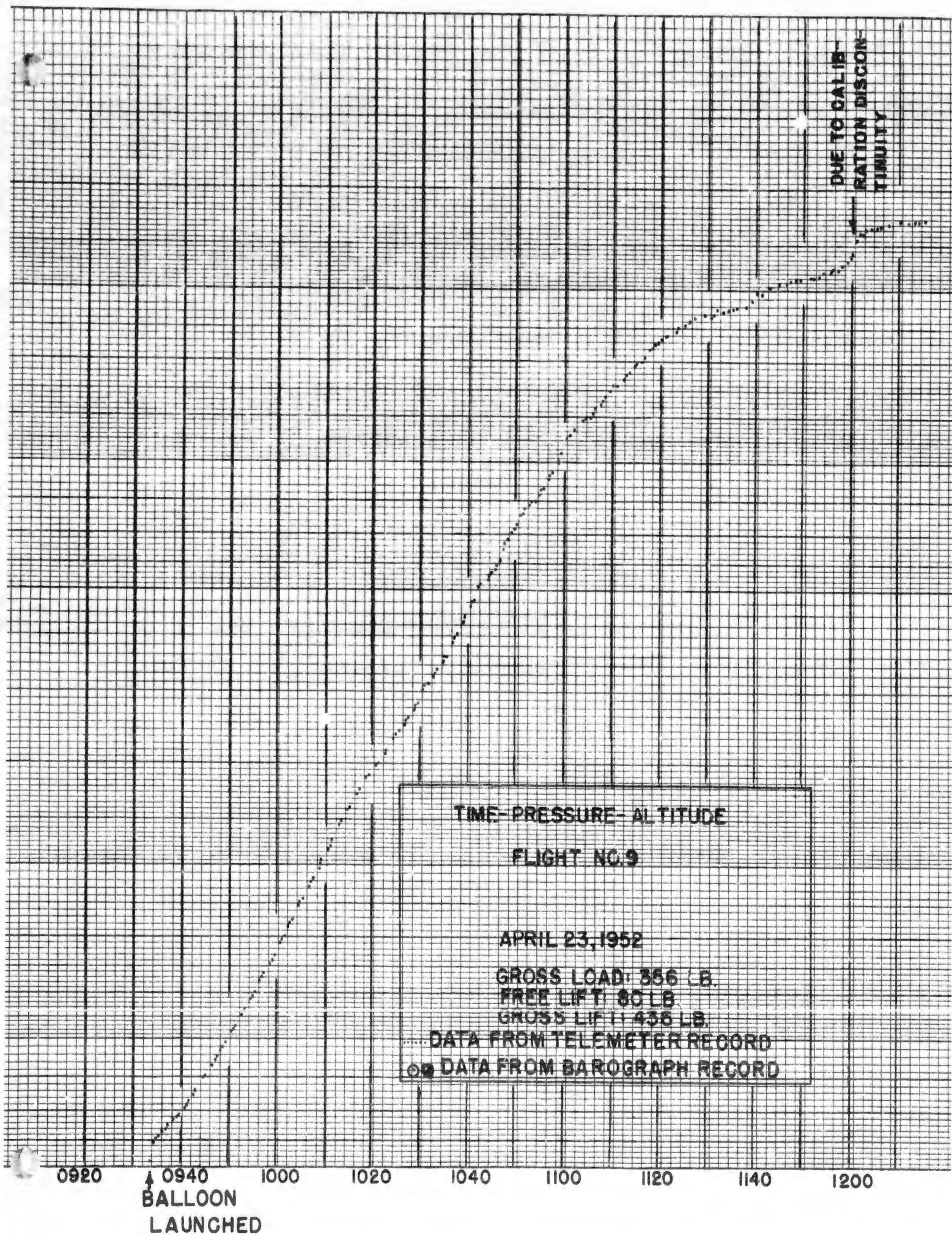
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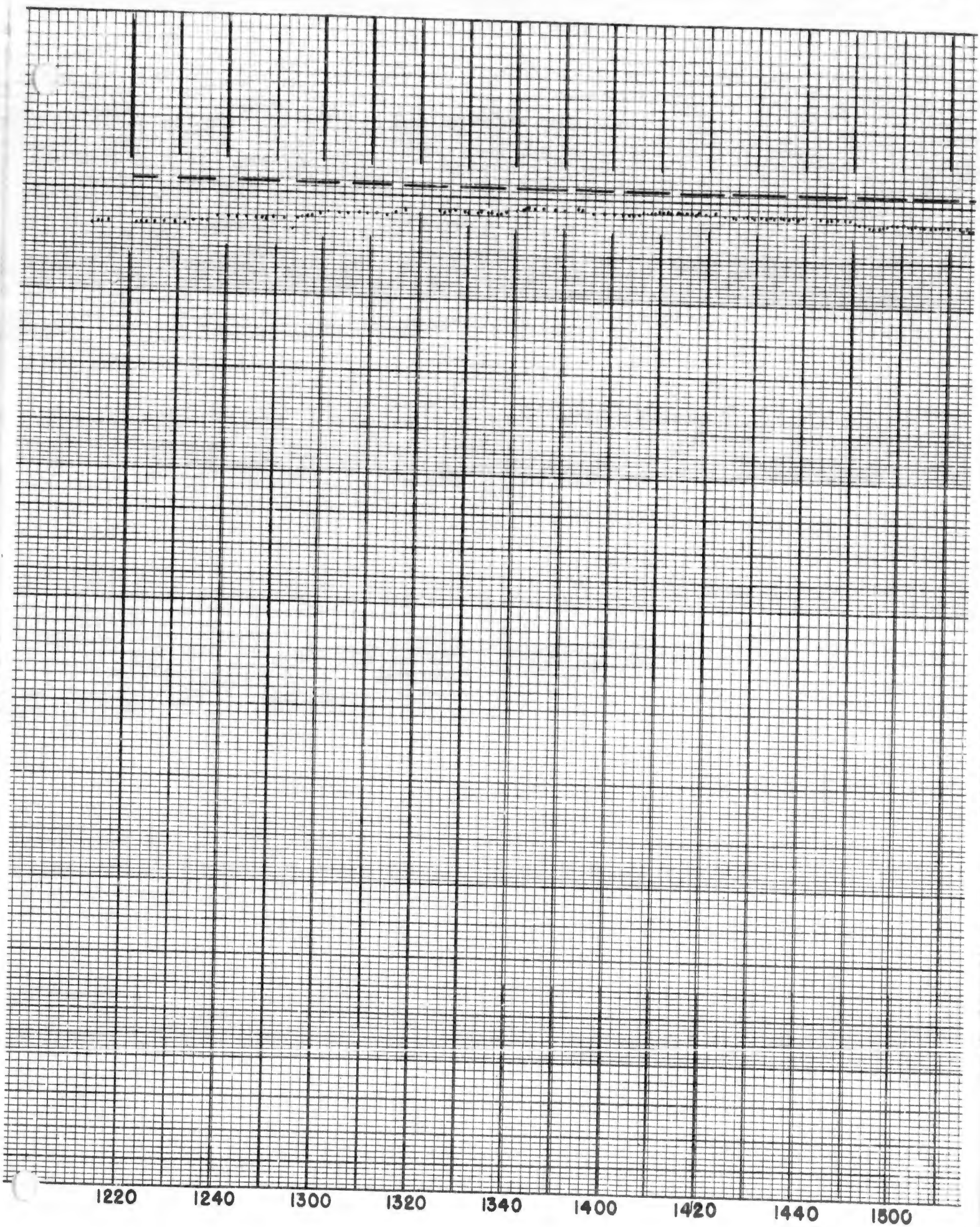


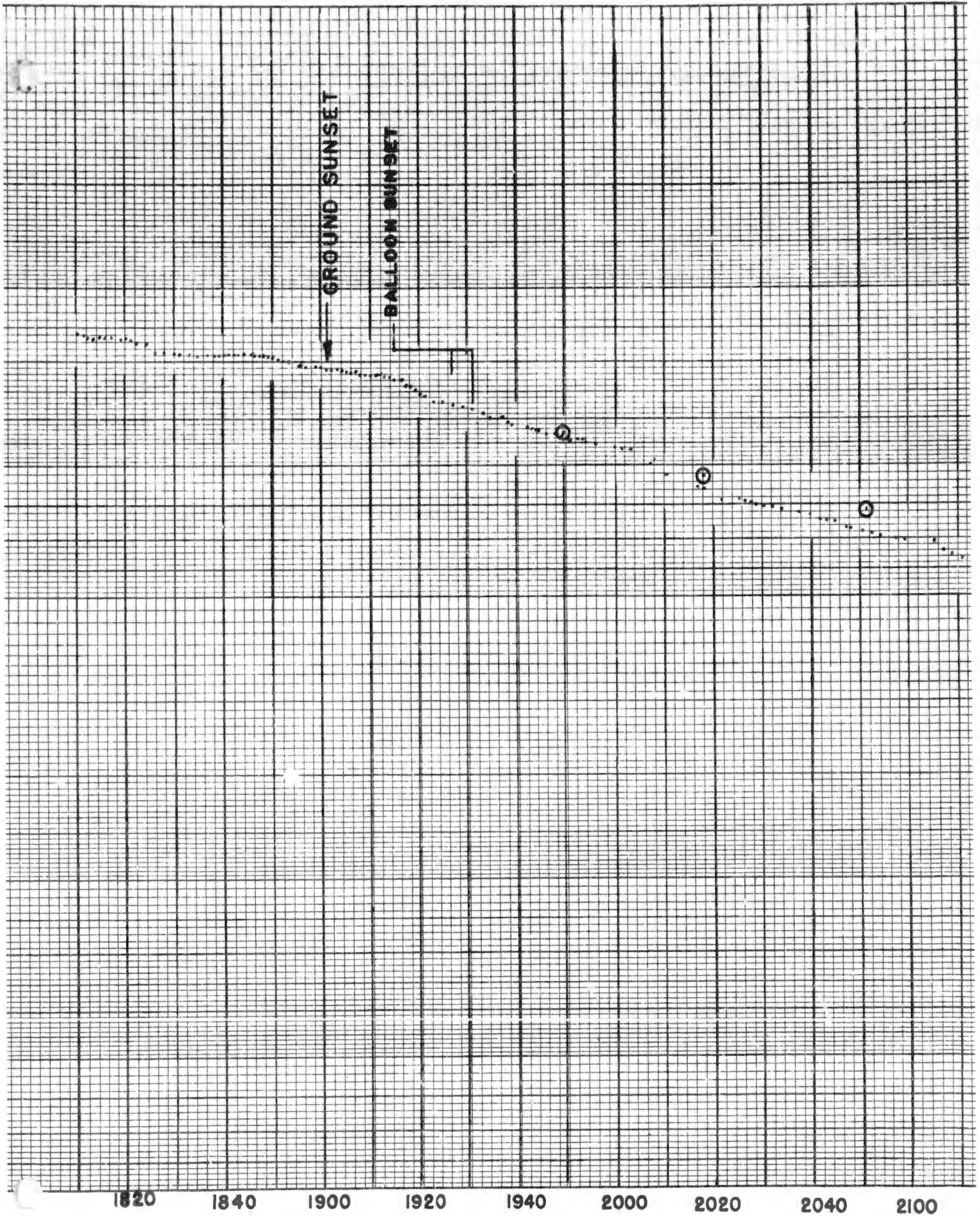


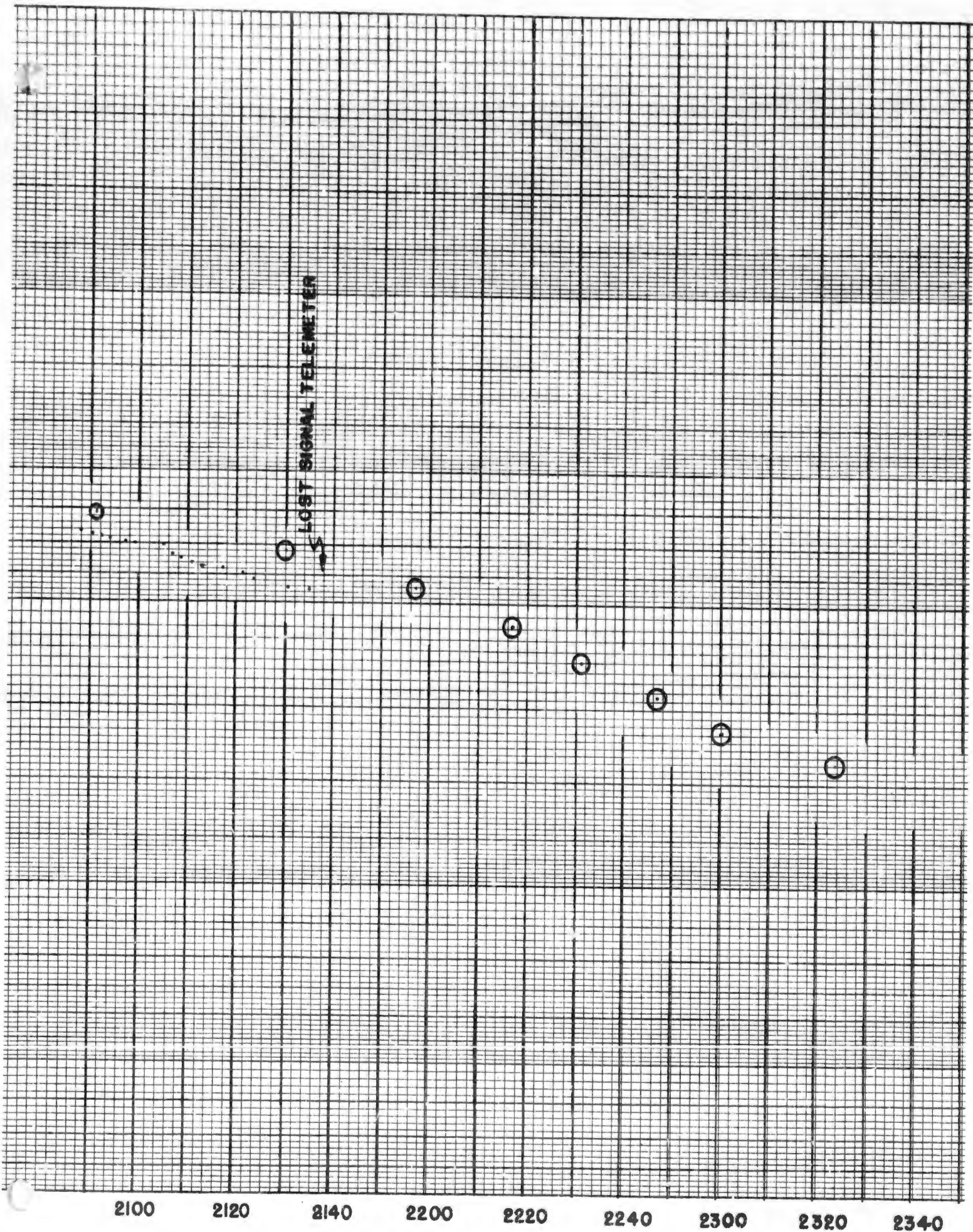




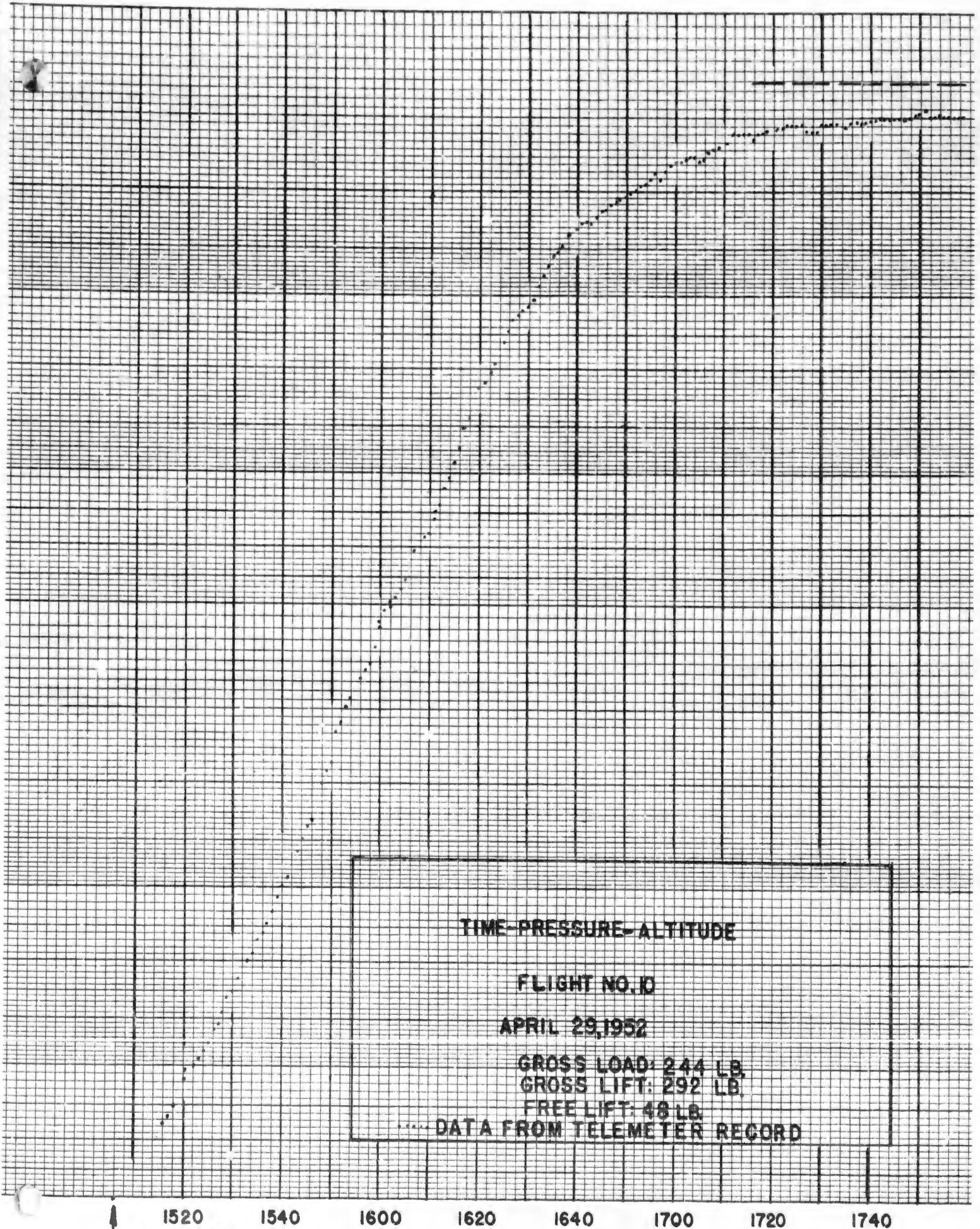




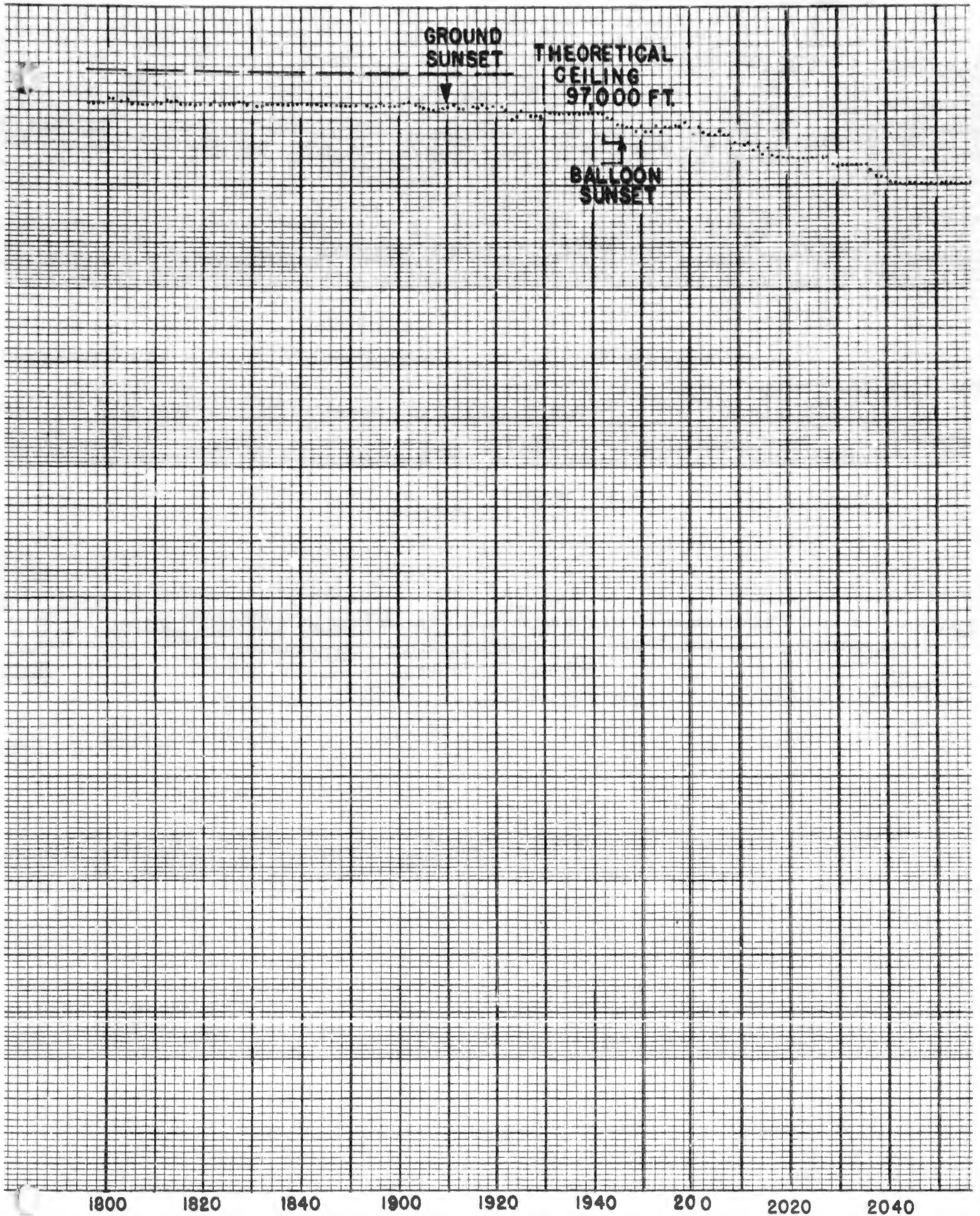


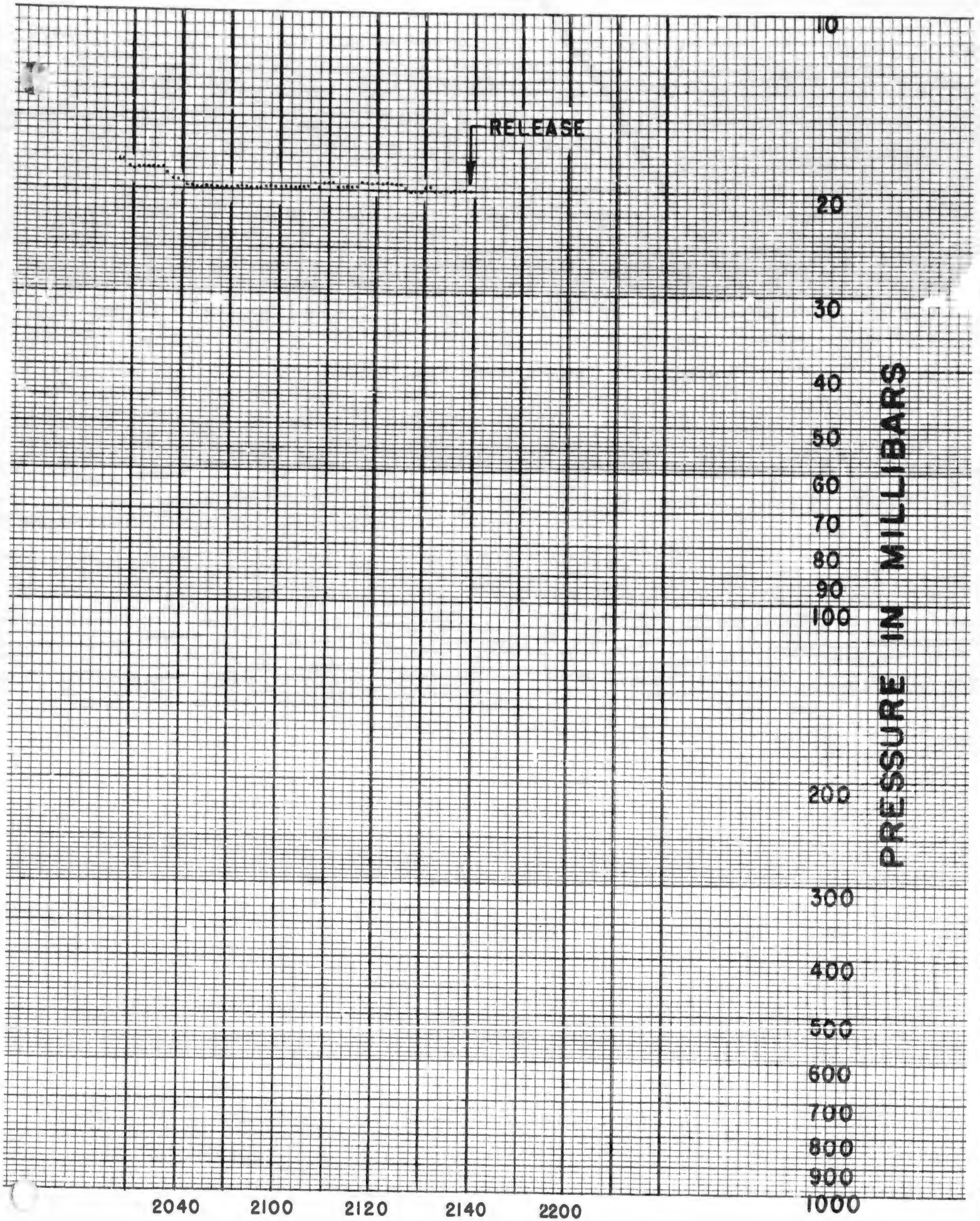


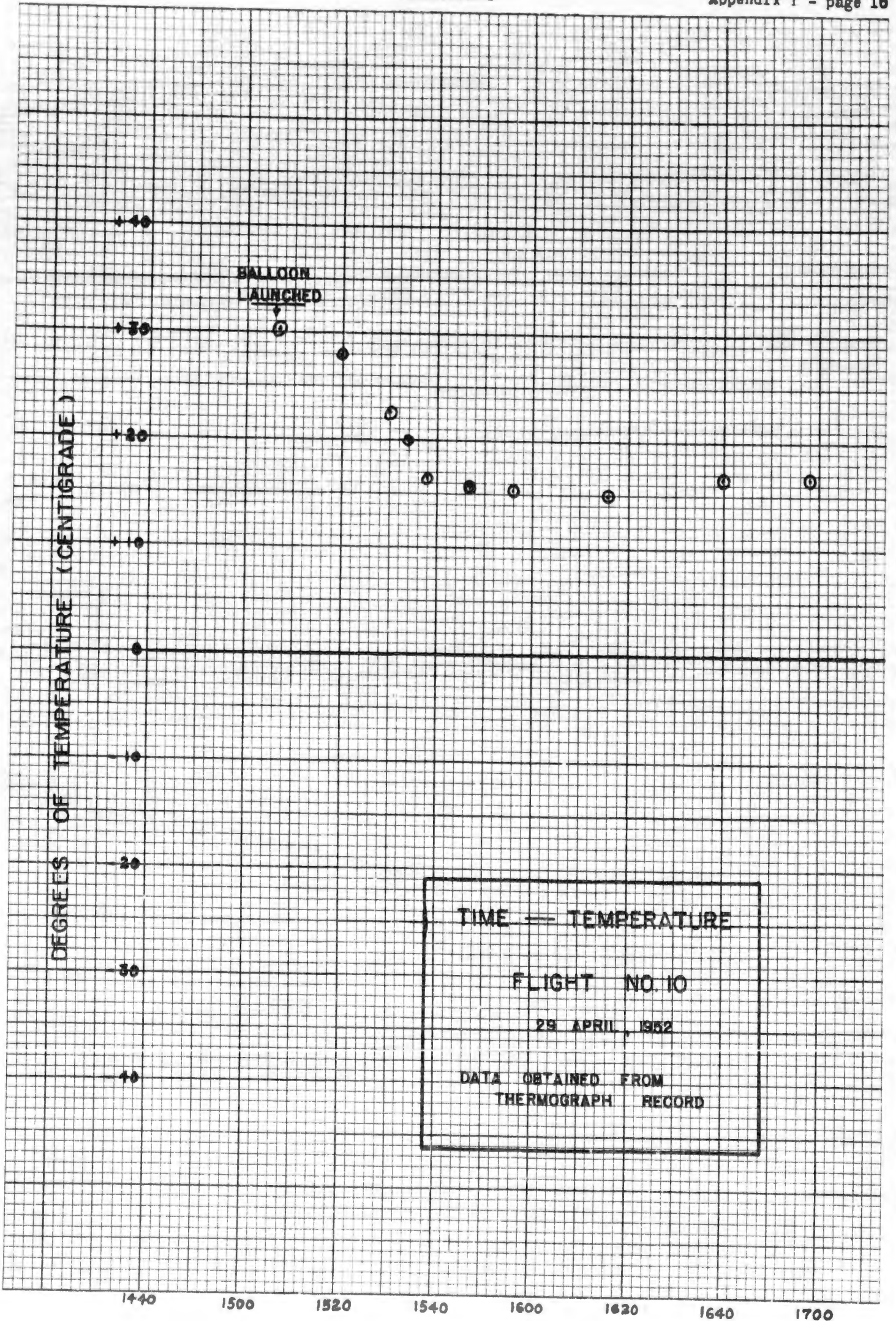


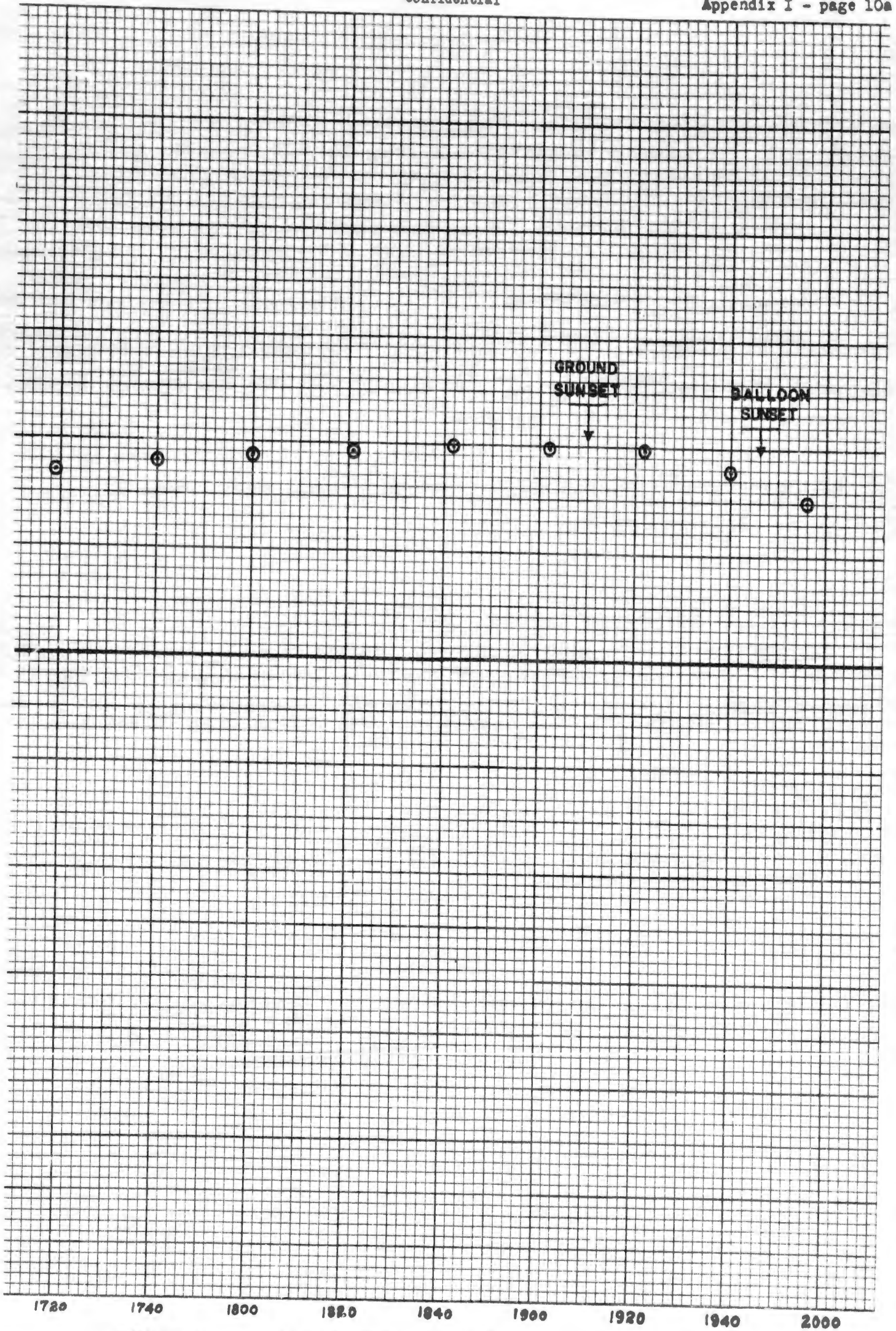


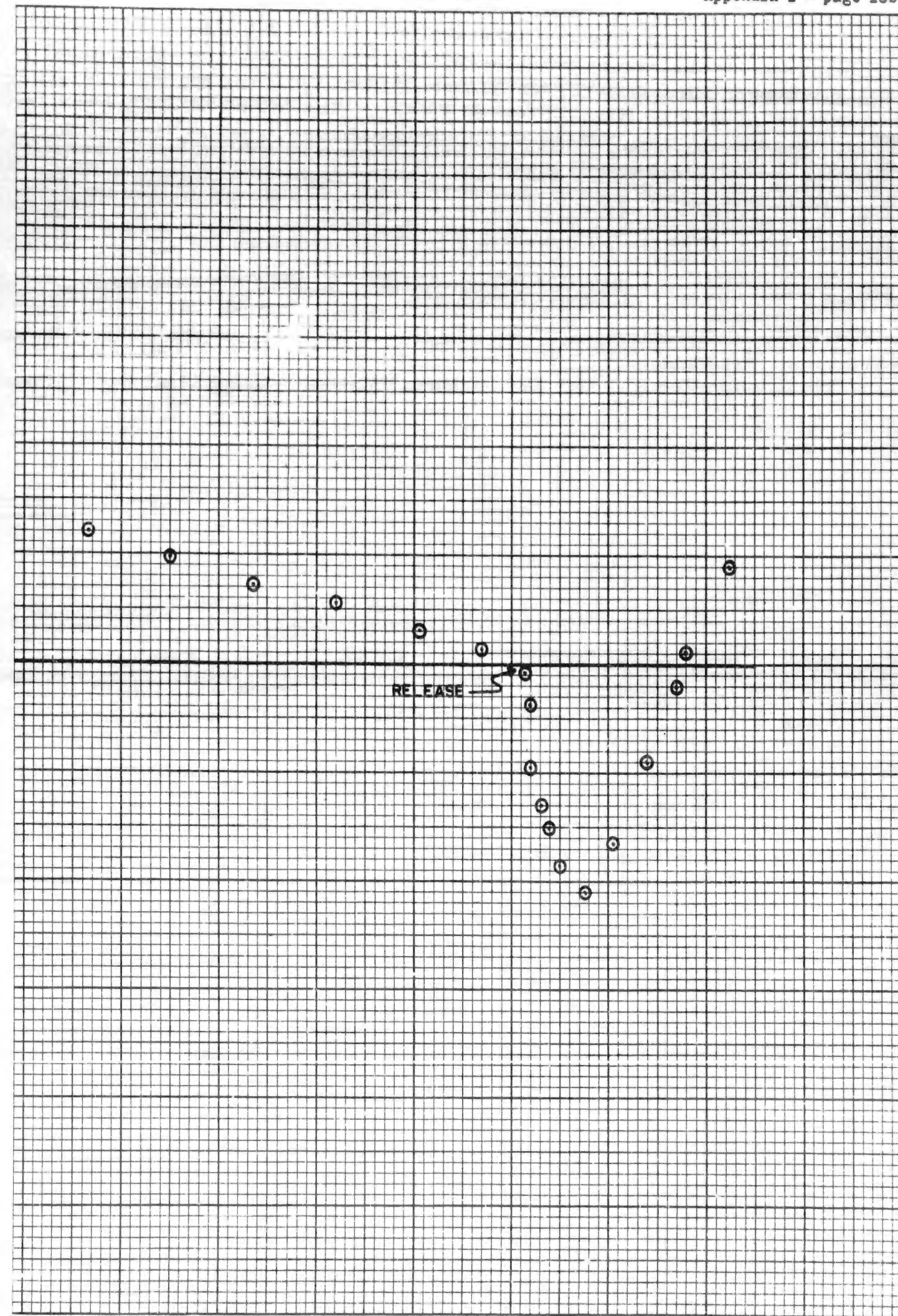
↑  
BALLOON  
LAUNCHED  
15:06

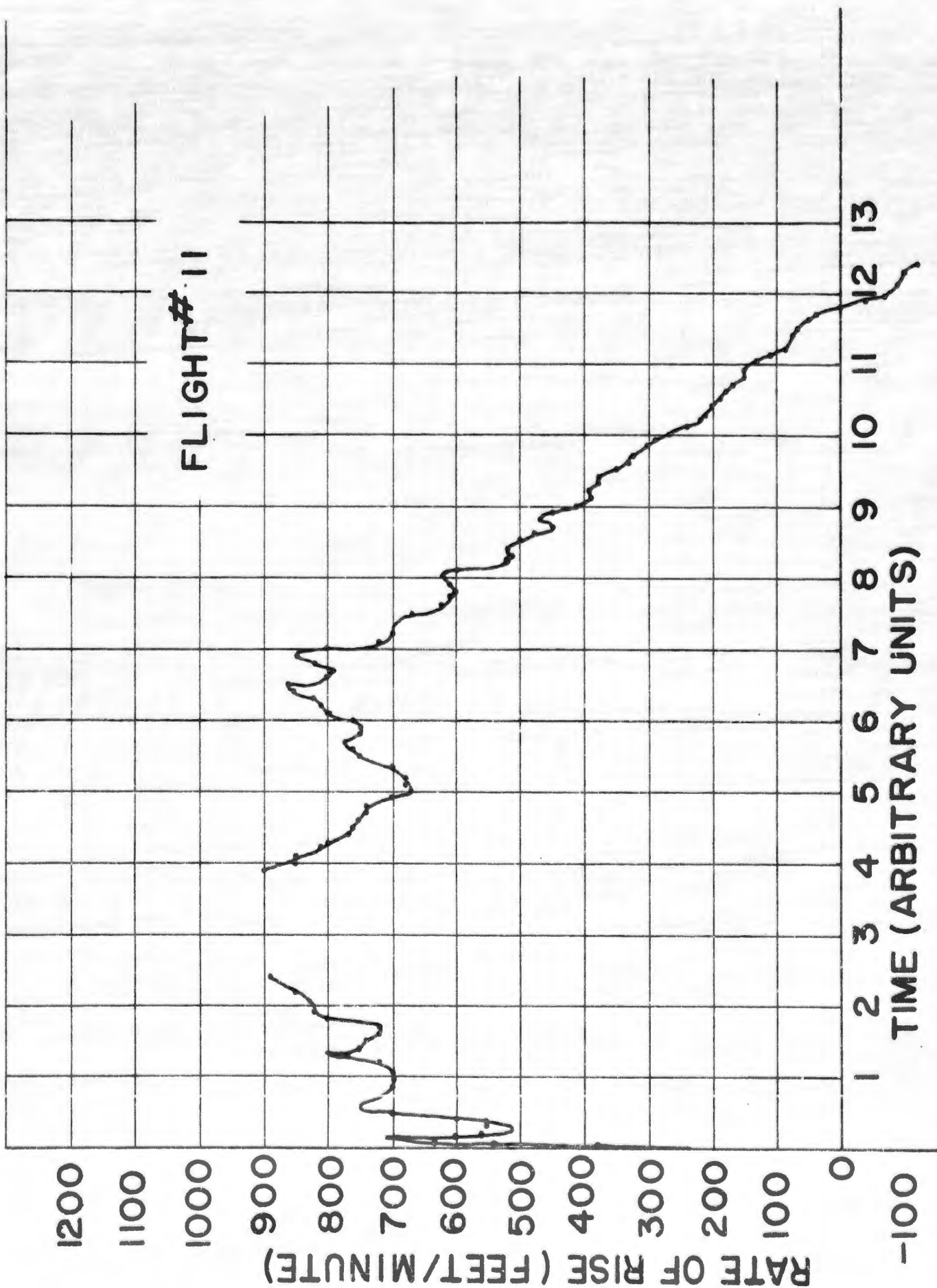


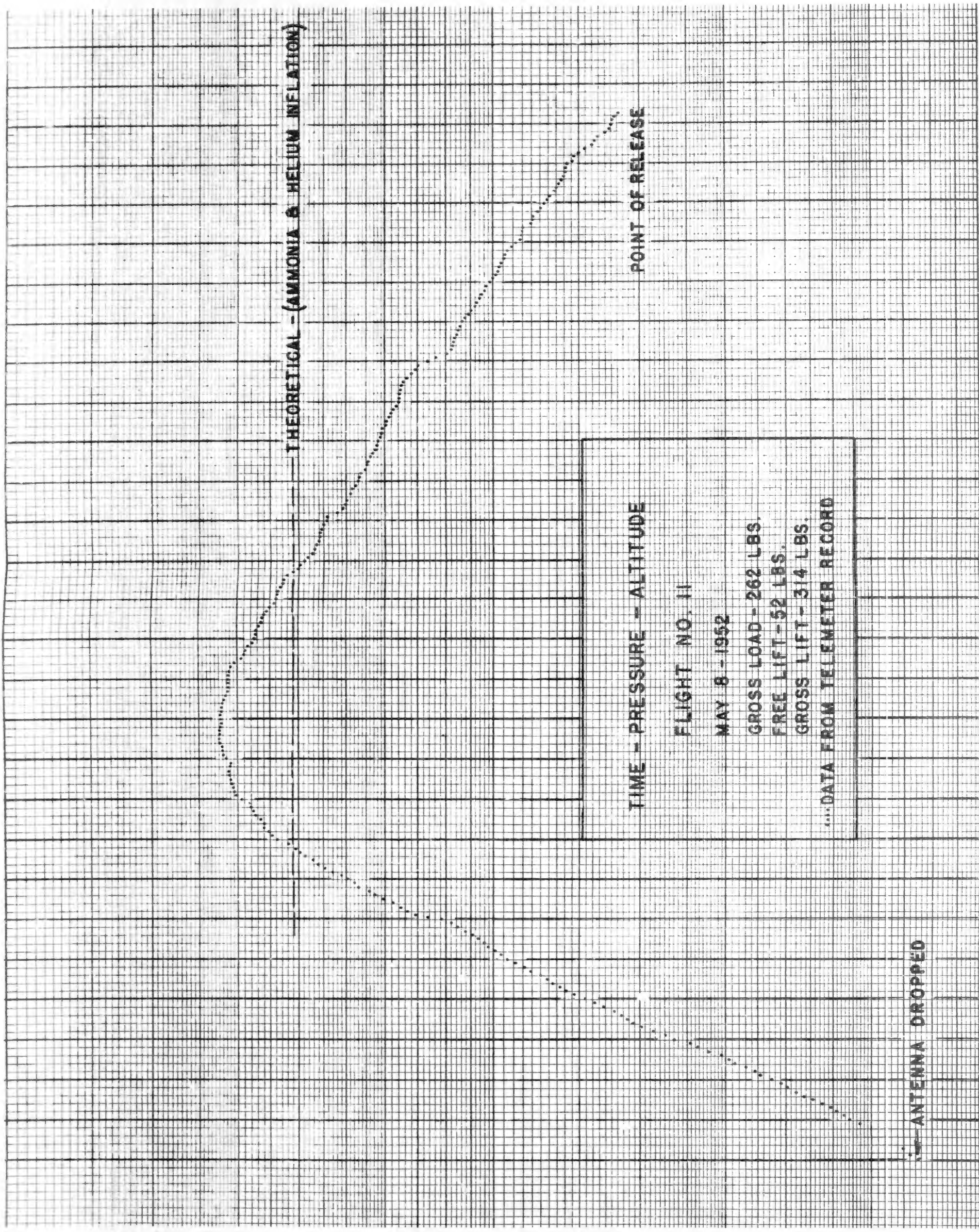


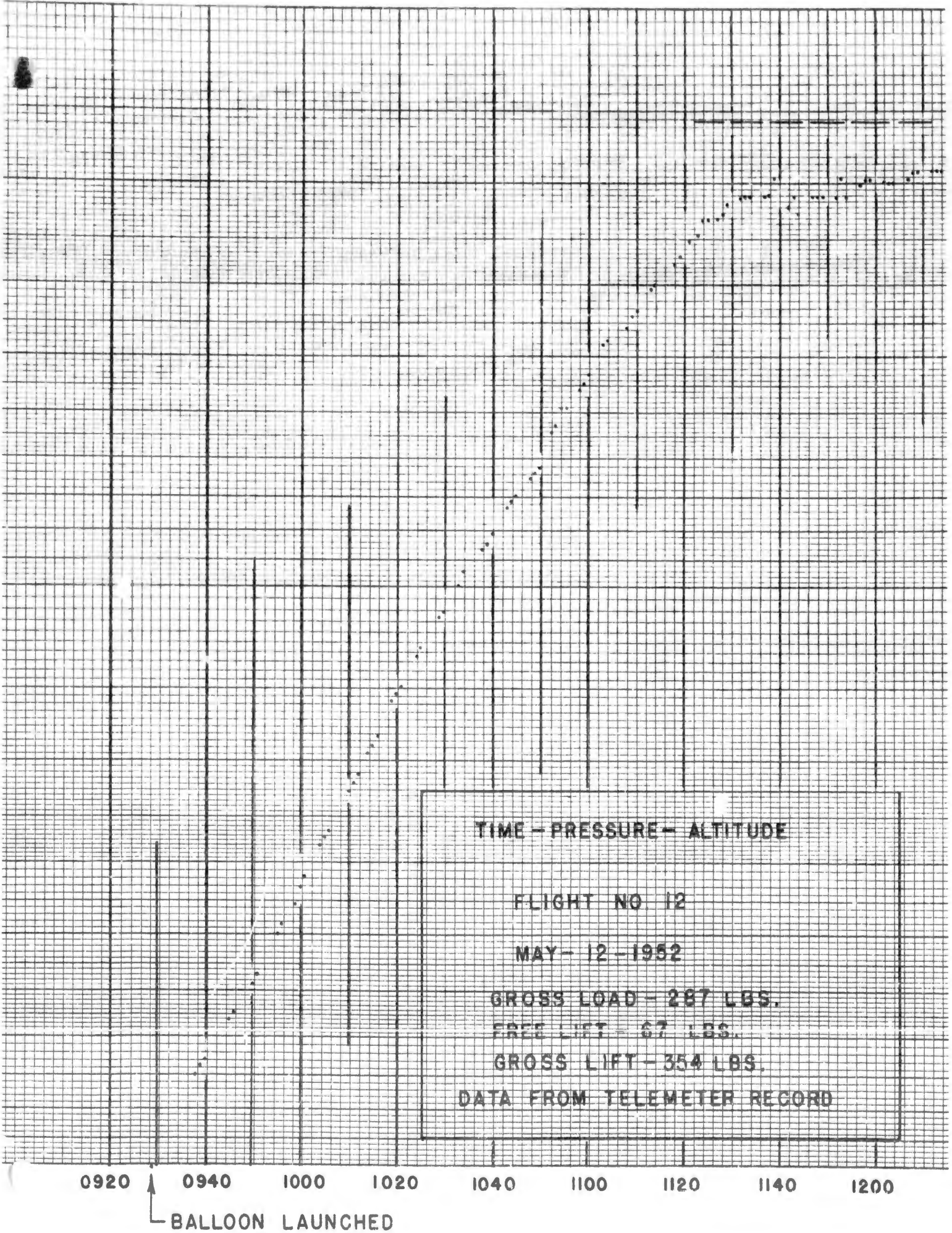


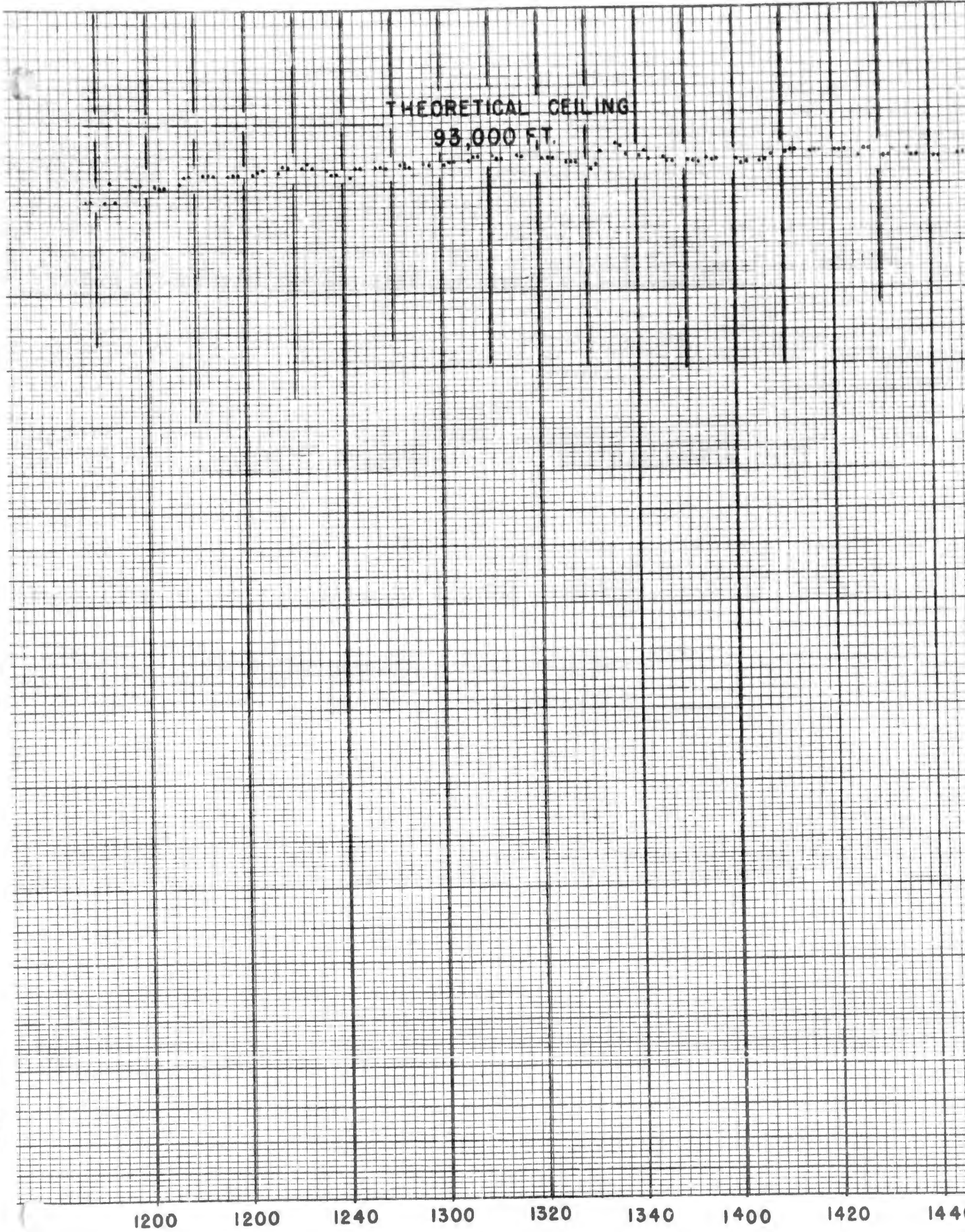


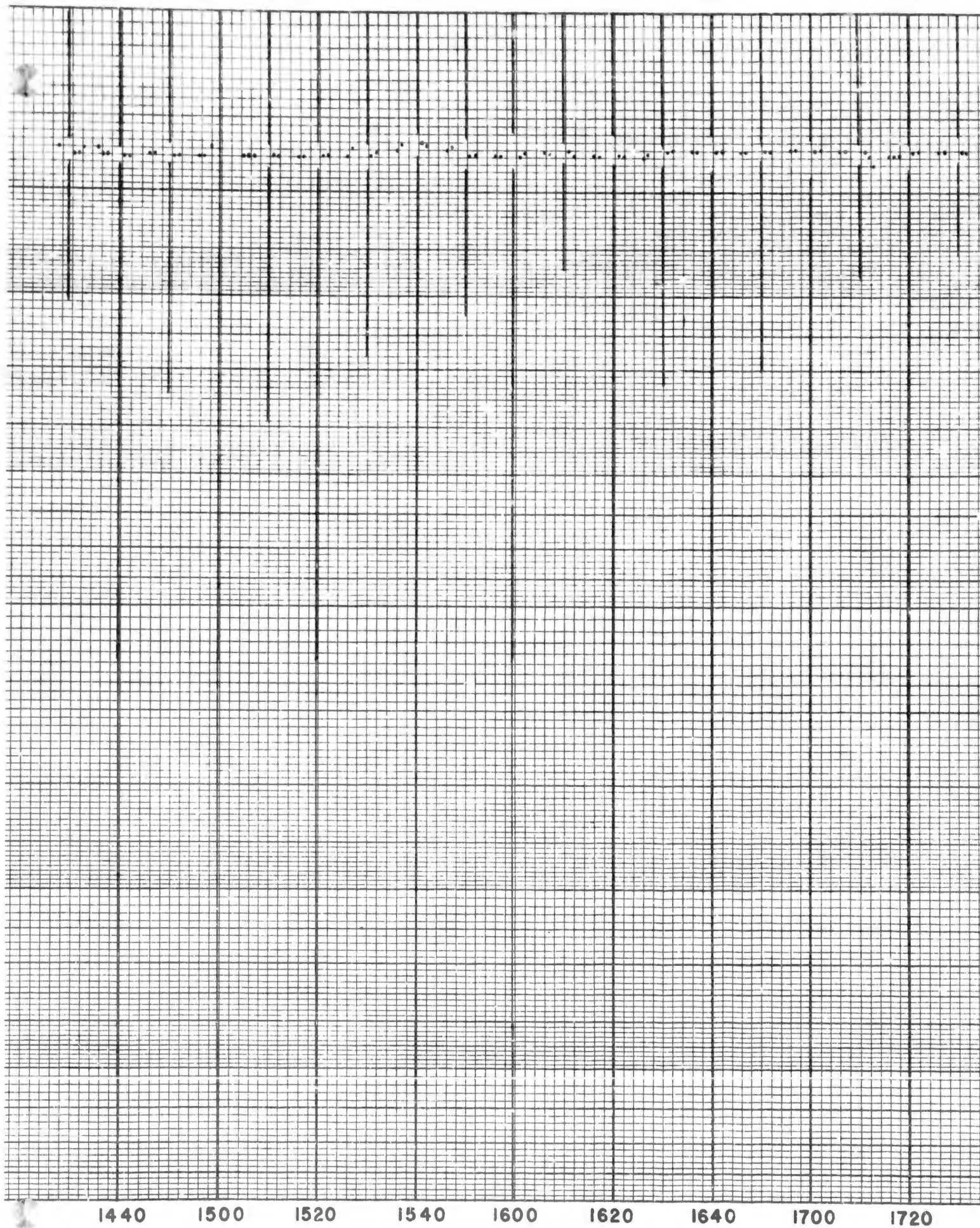




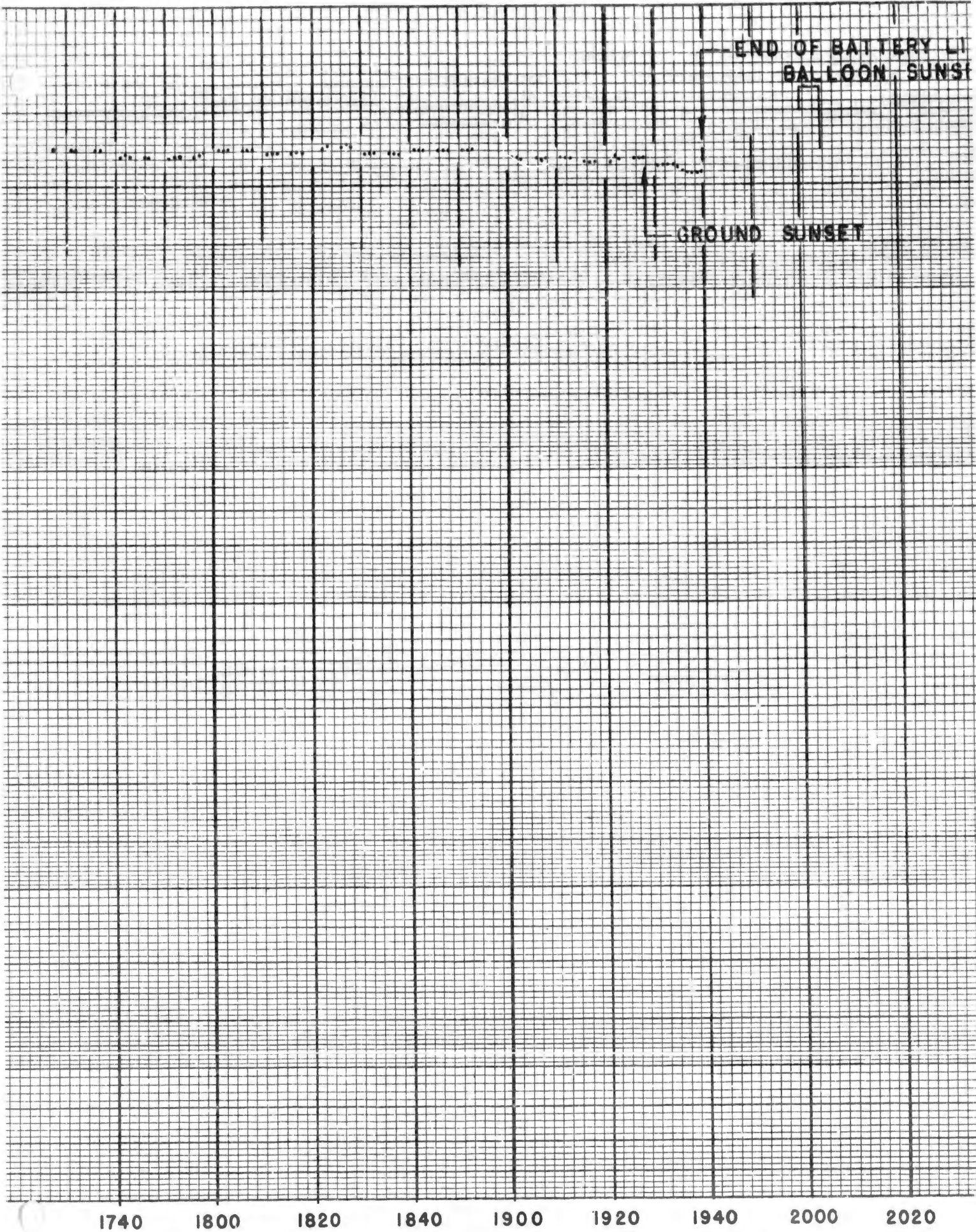




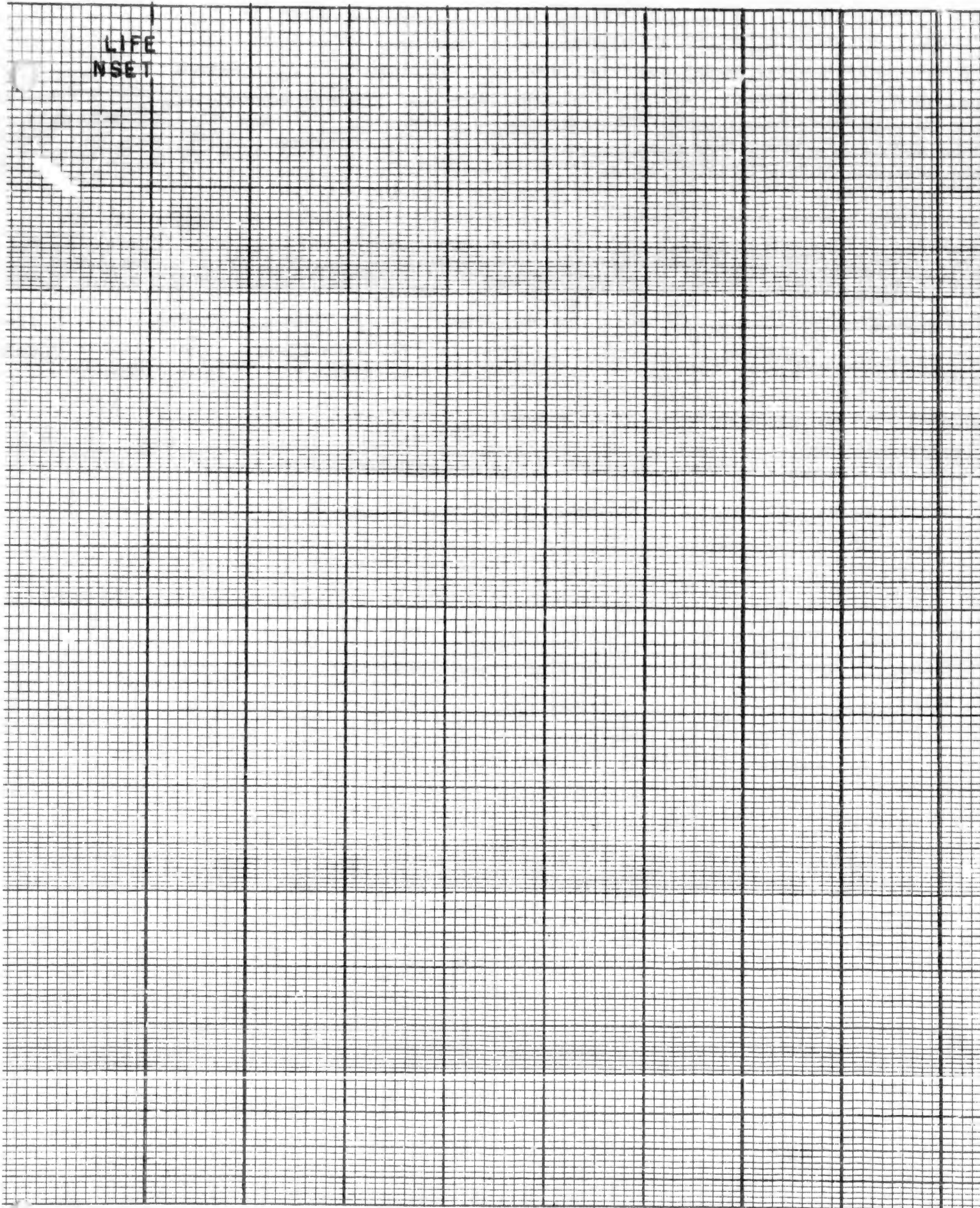




1440 1500 1520 1540 1600 1620 1640 1700 1720



LIFE  
NSET

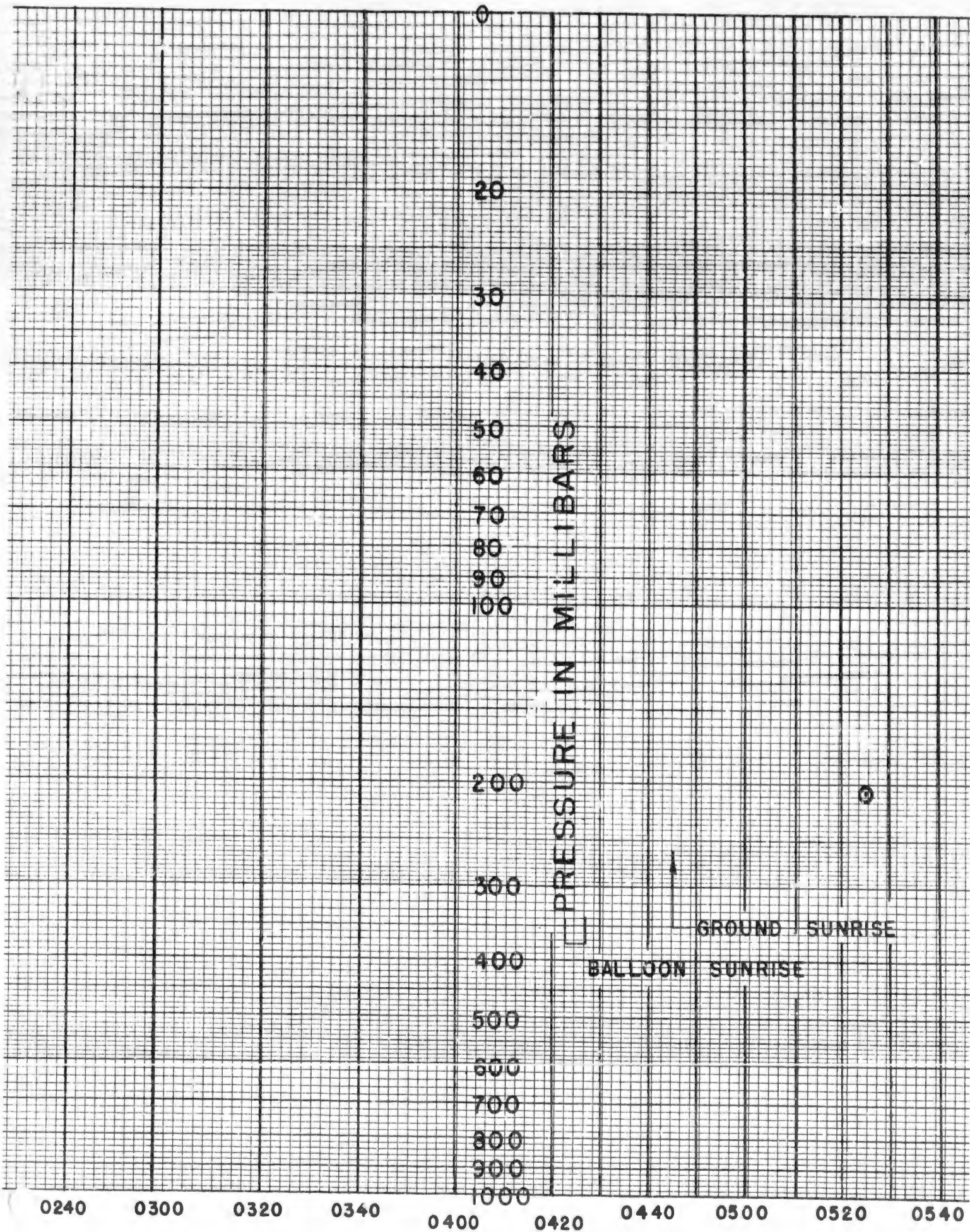


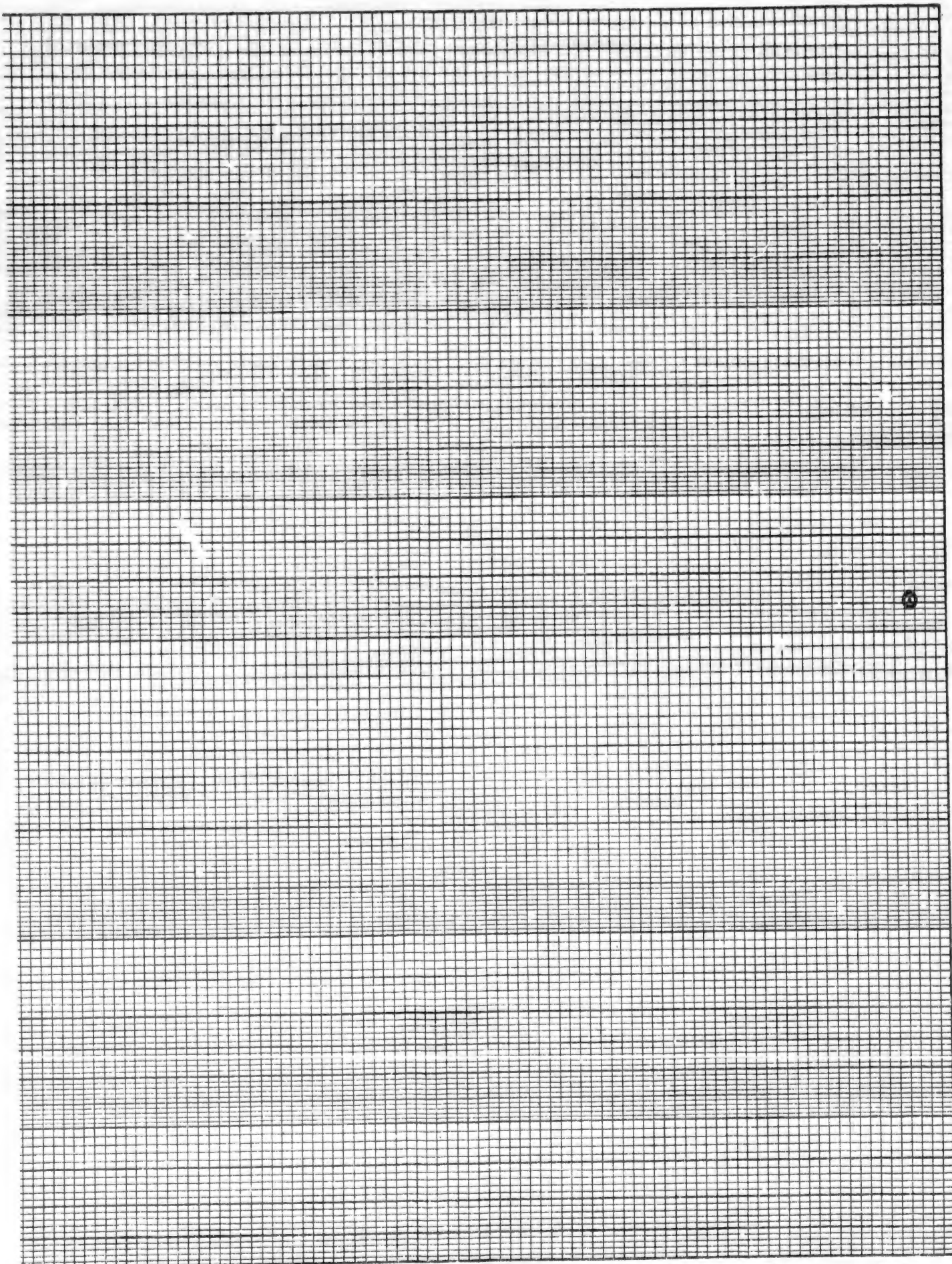
2040 2100 2120 2140 2200 2220 2240 2300 2320

ALTITUDE DATA FOR 13 MAY  
DETERMINED FROM  
OO SUBTENDED ANGLES AND  
...DOUBLE THEODOLITE READINGS

2340 2400 0020 0040 0100 0120 0140 0200 0220 0240

→  
MAY 13-1952





0600

0620

0640

0700

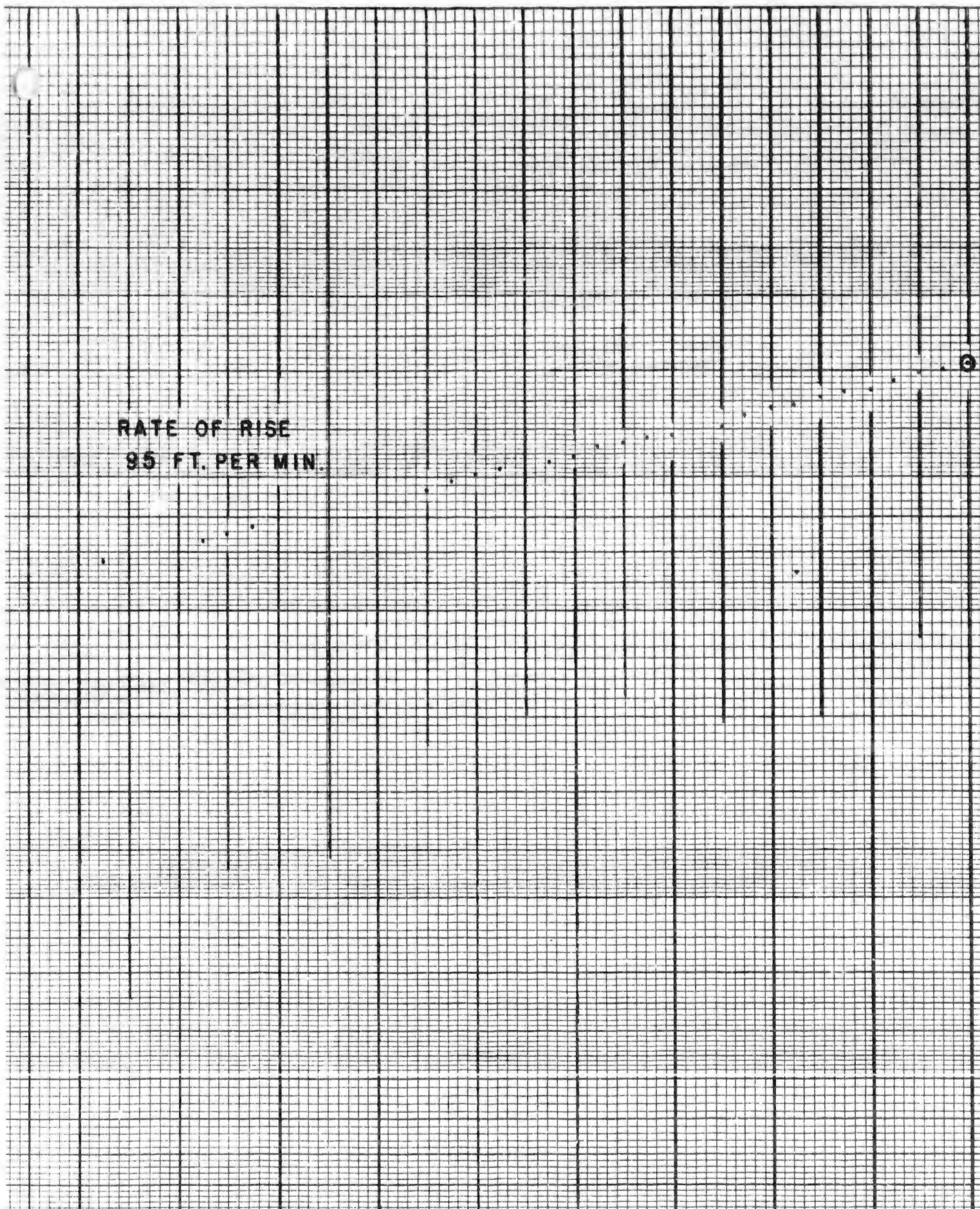
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0740

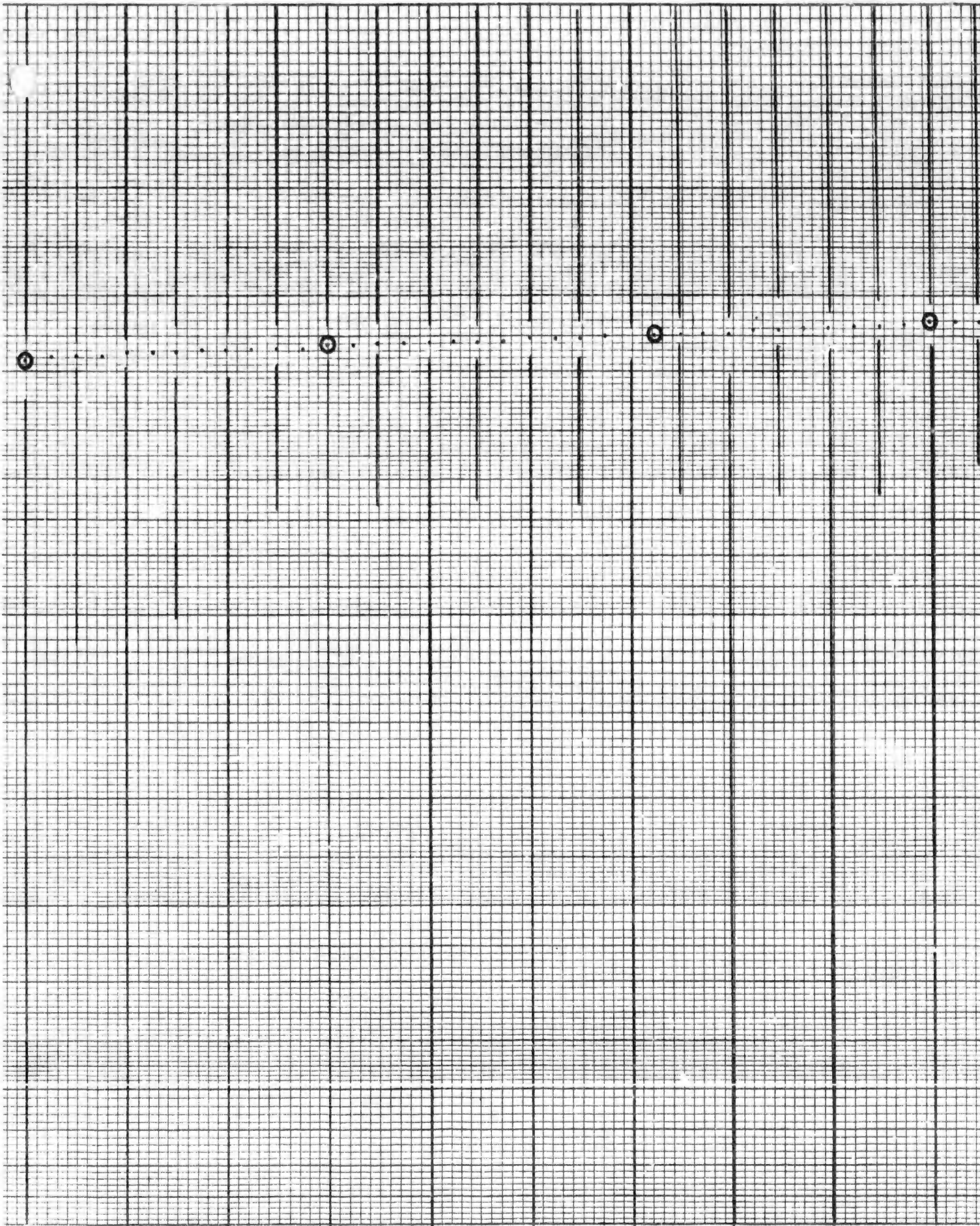
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0820

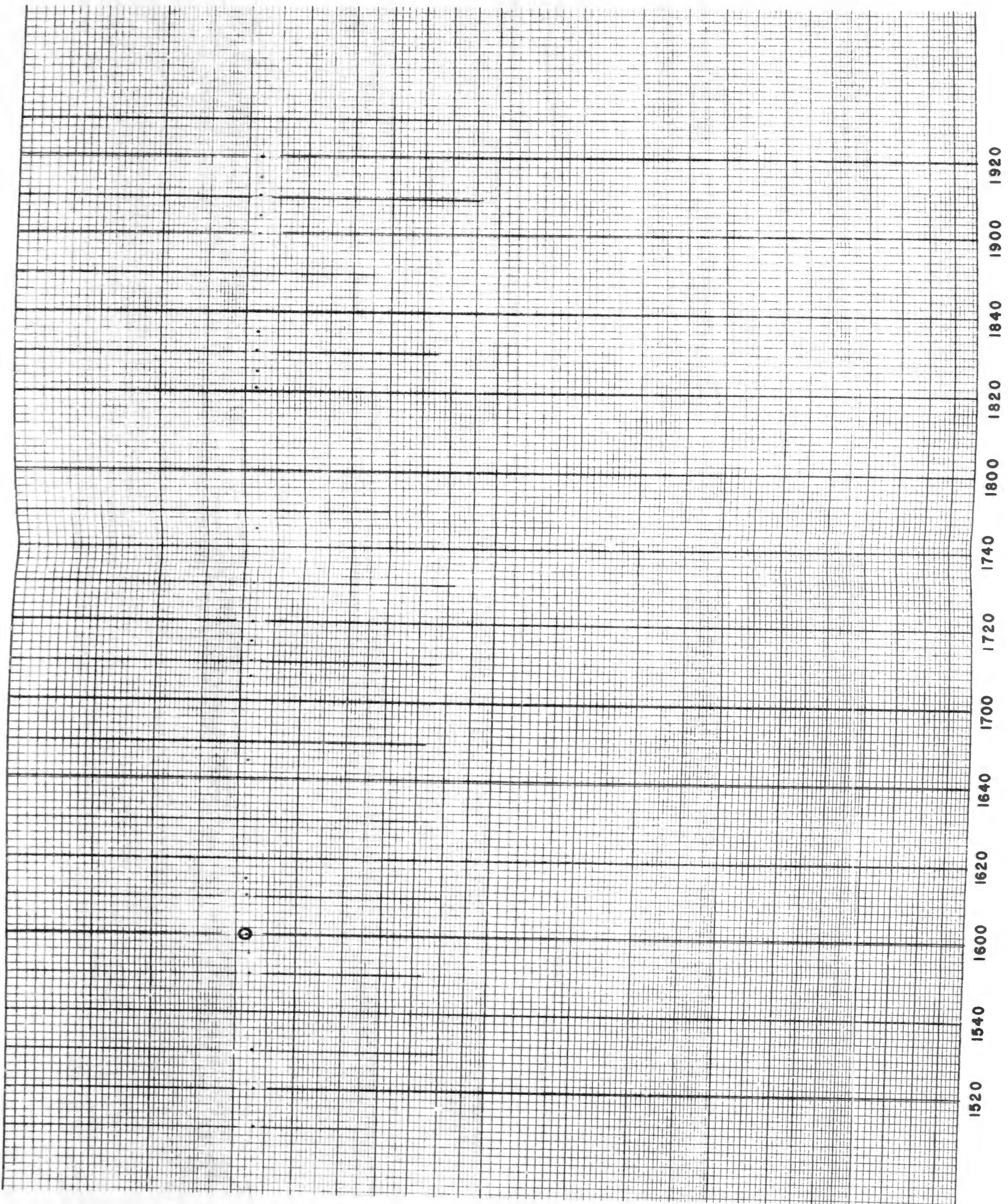
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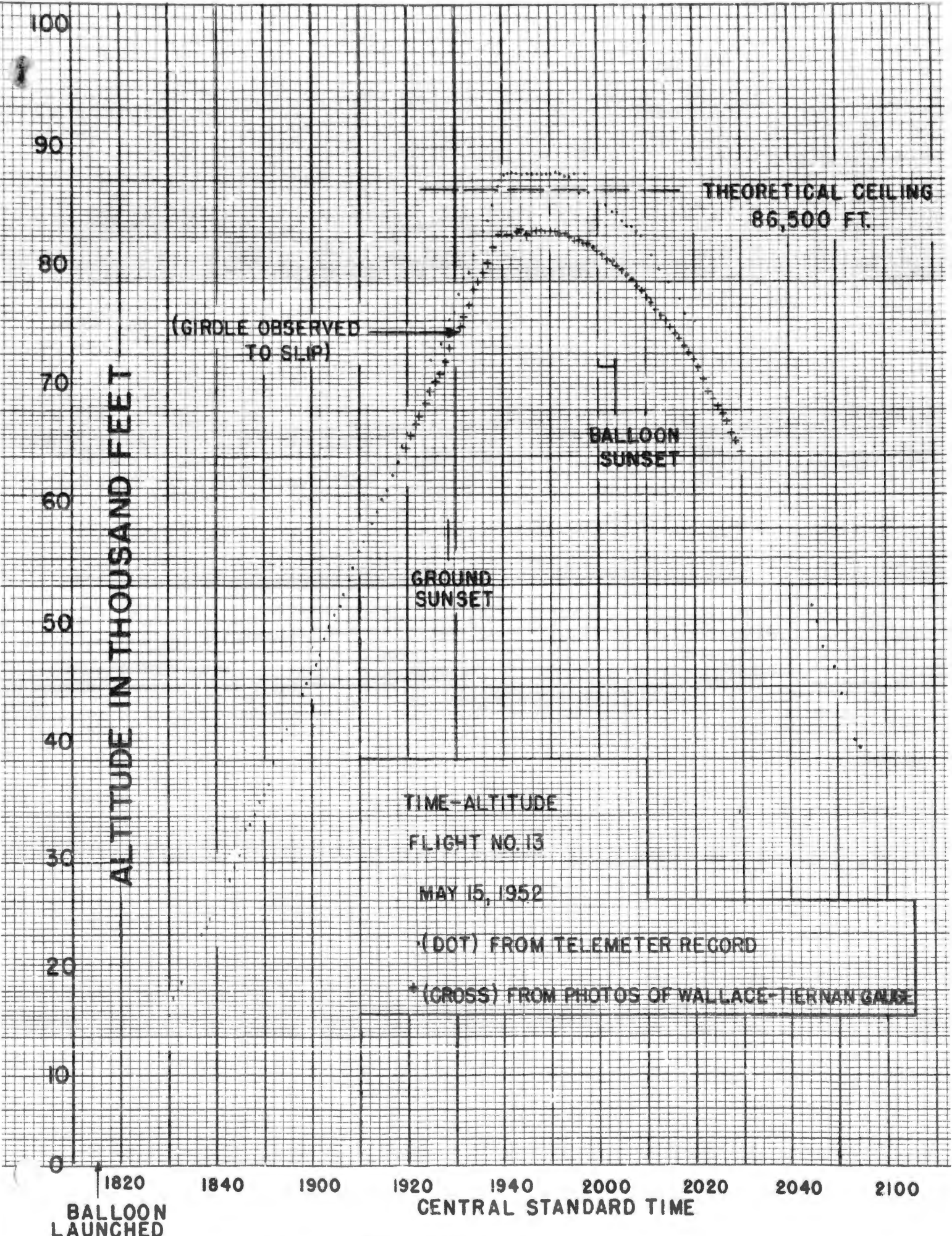


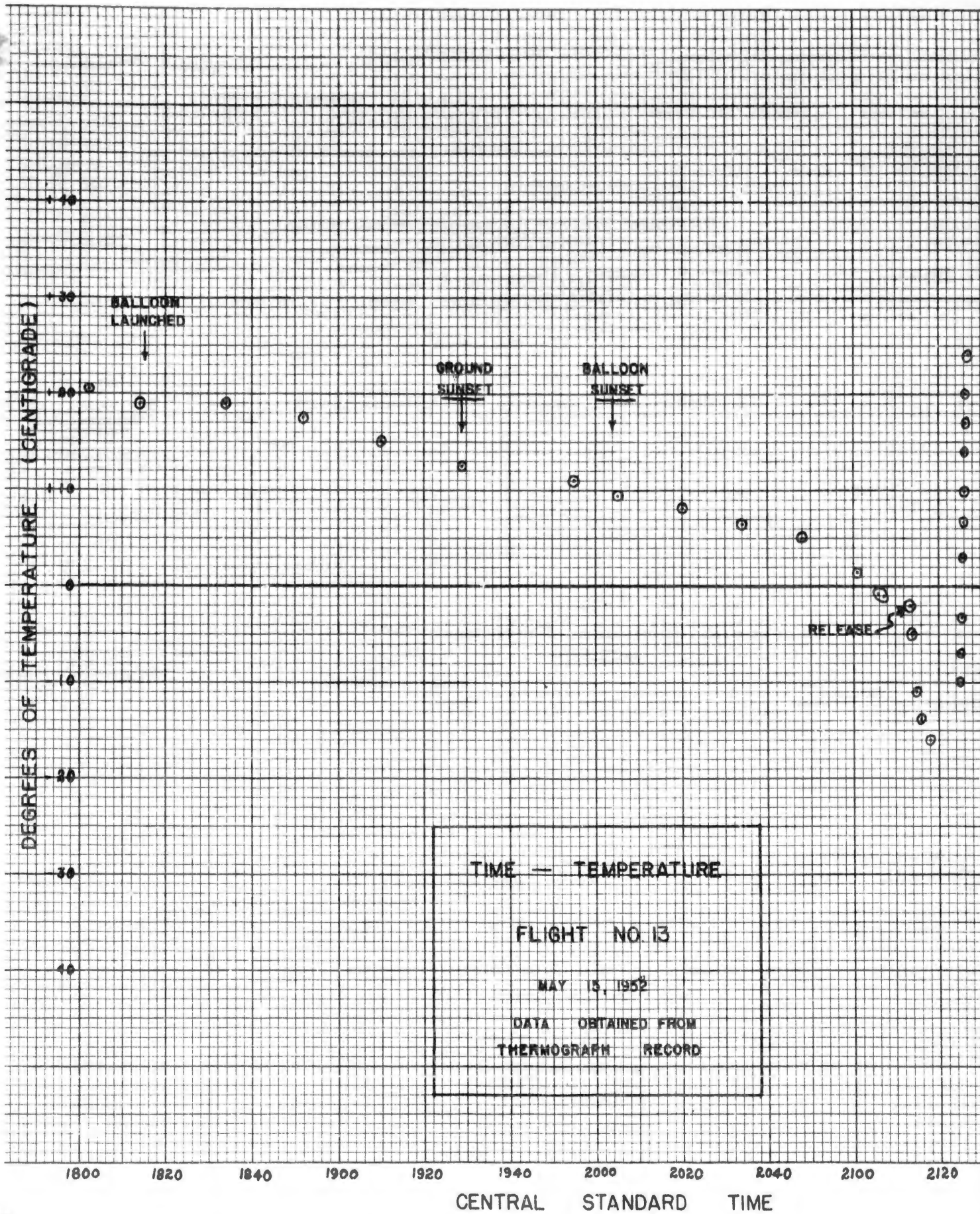
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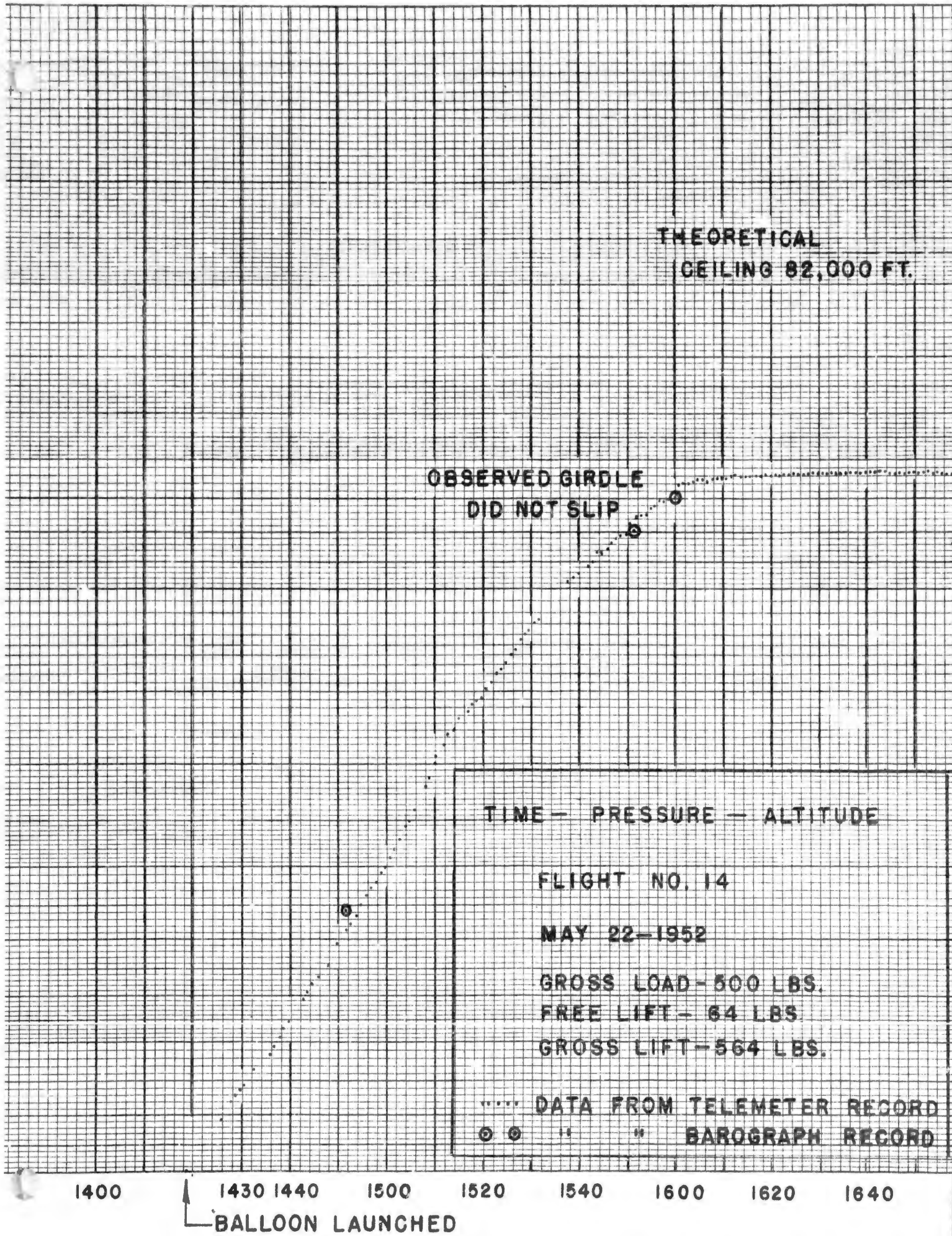


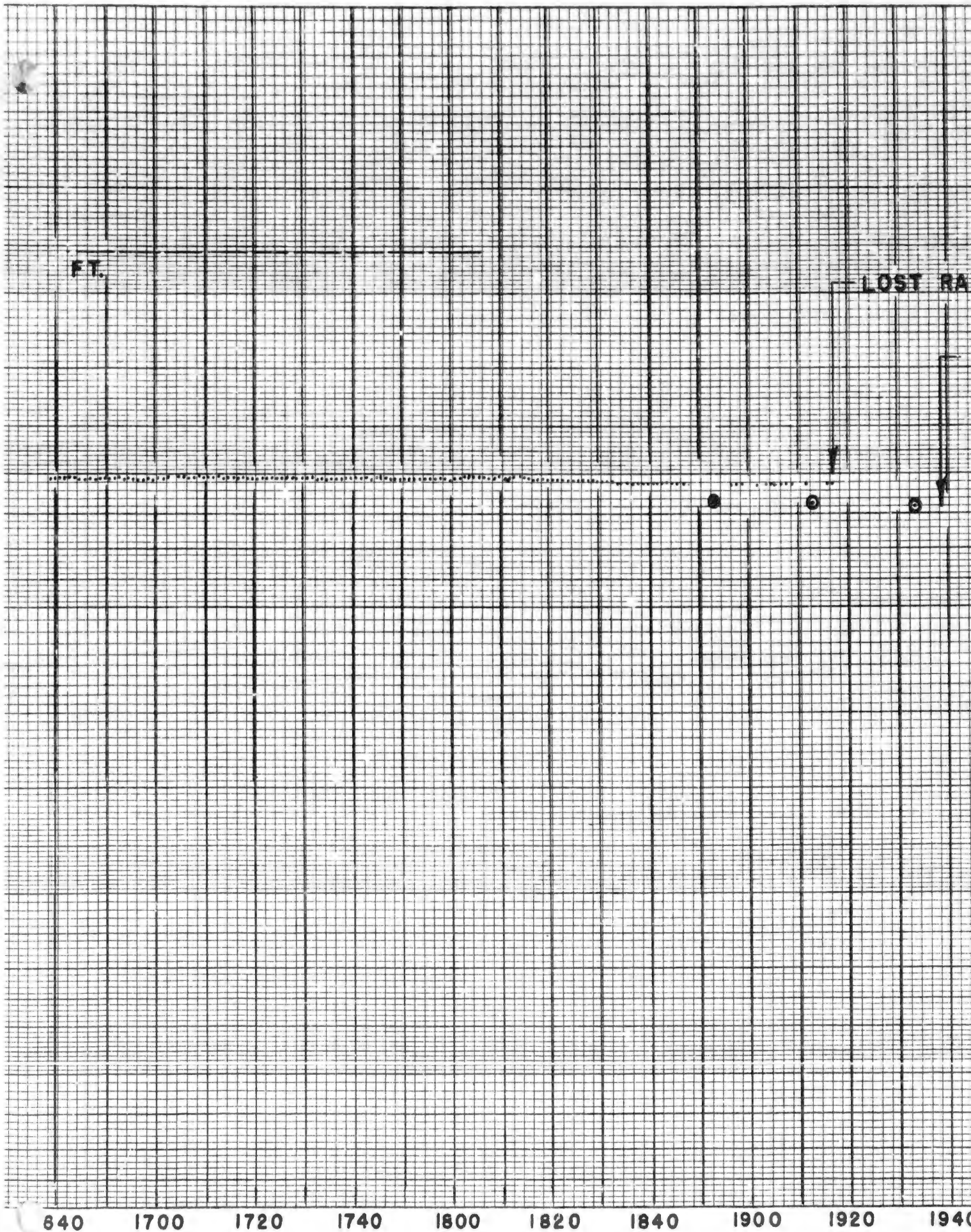
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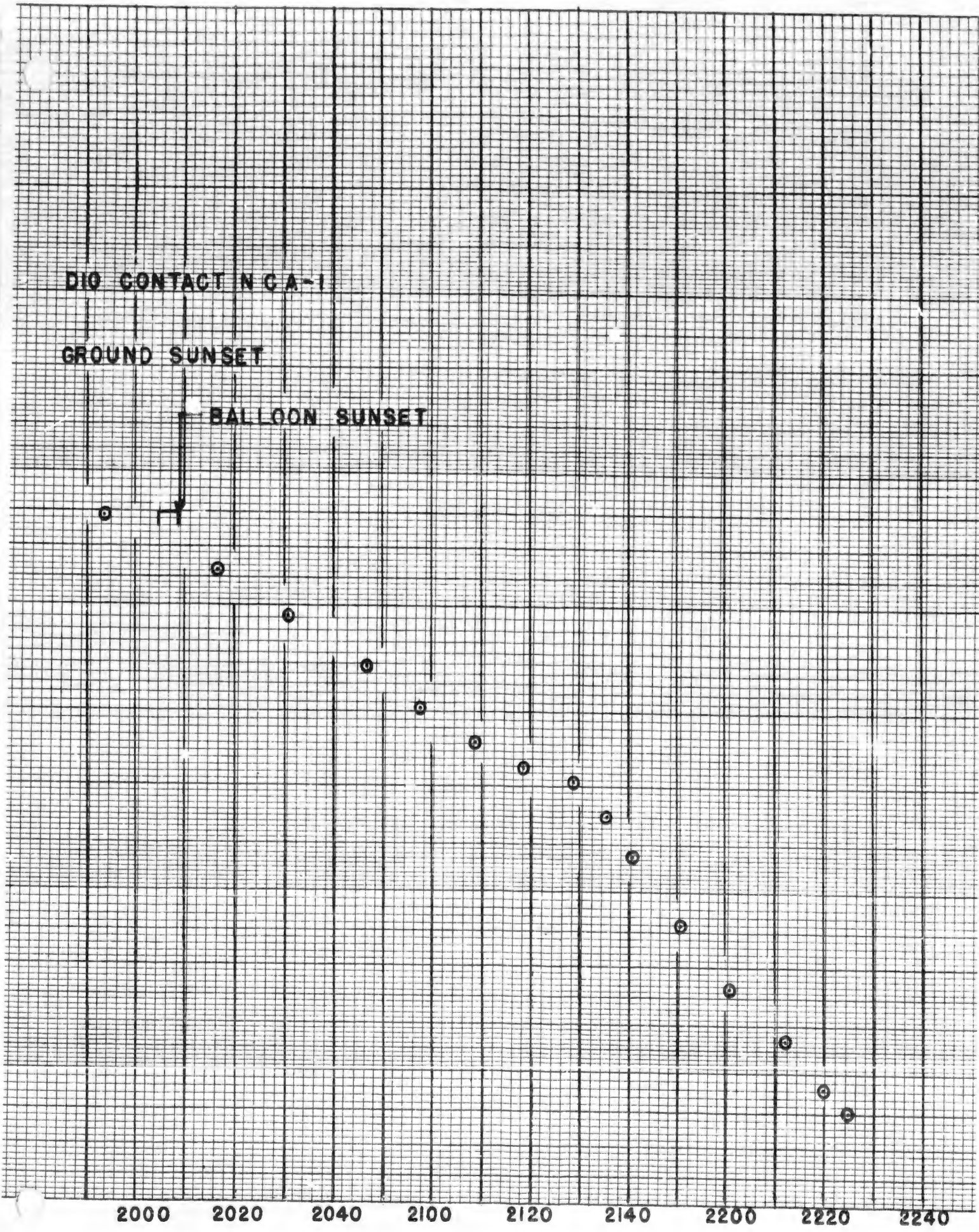


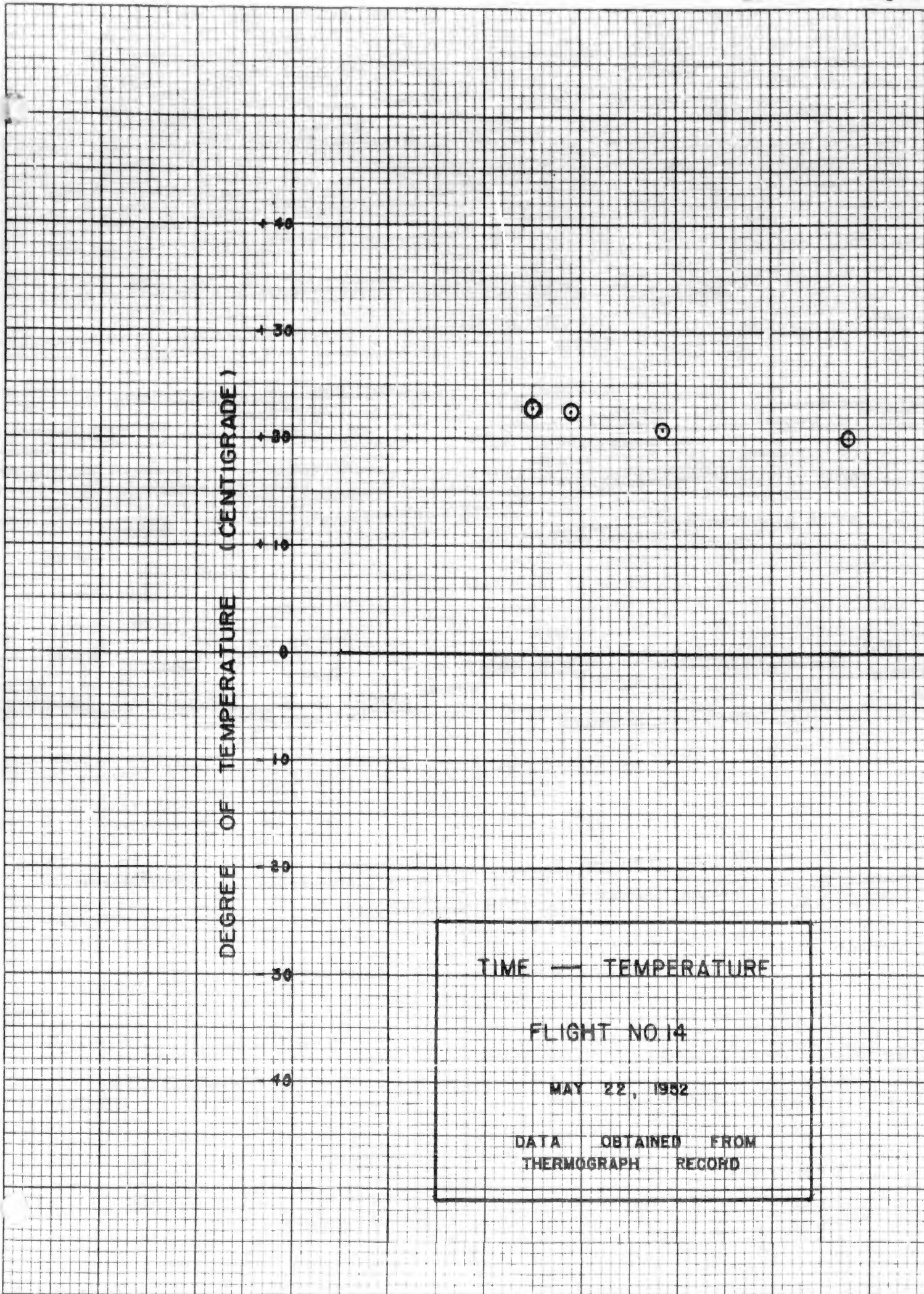


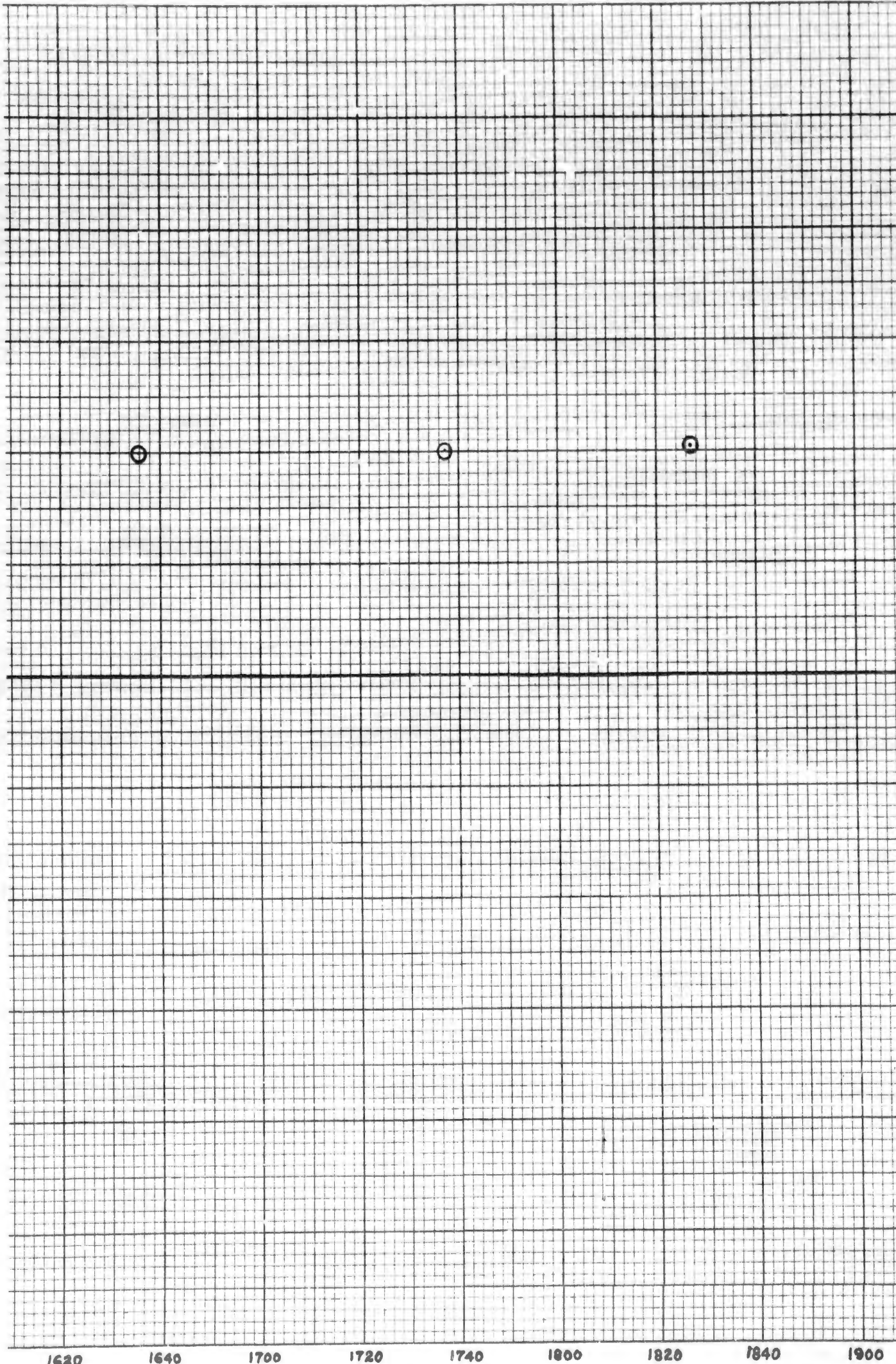


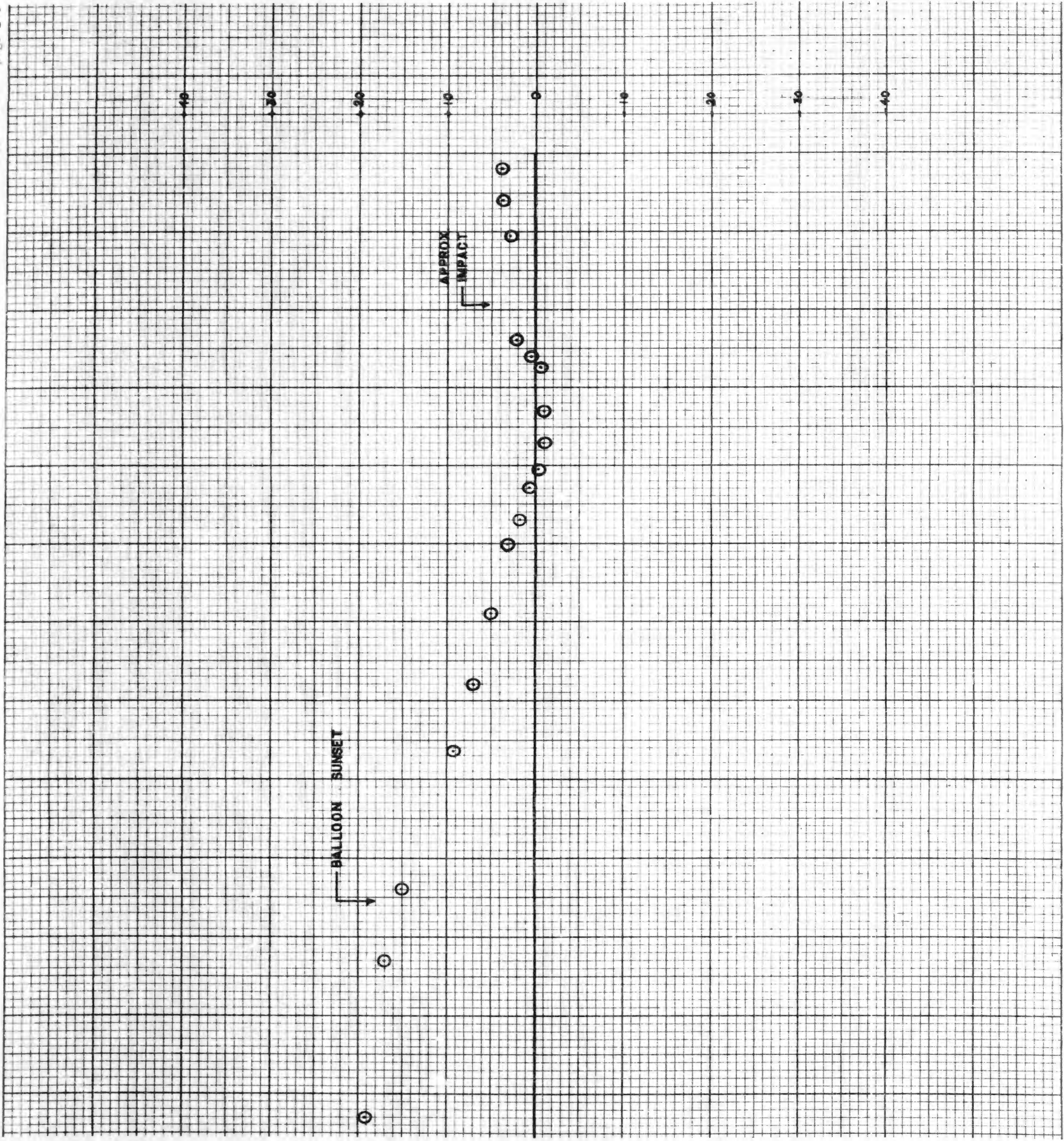


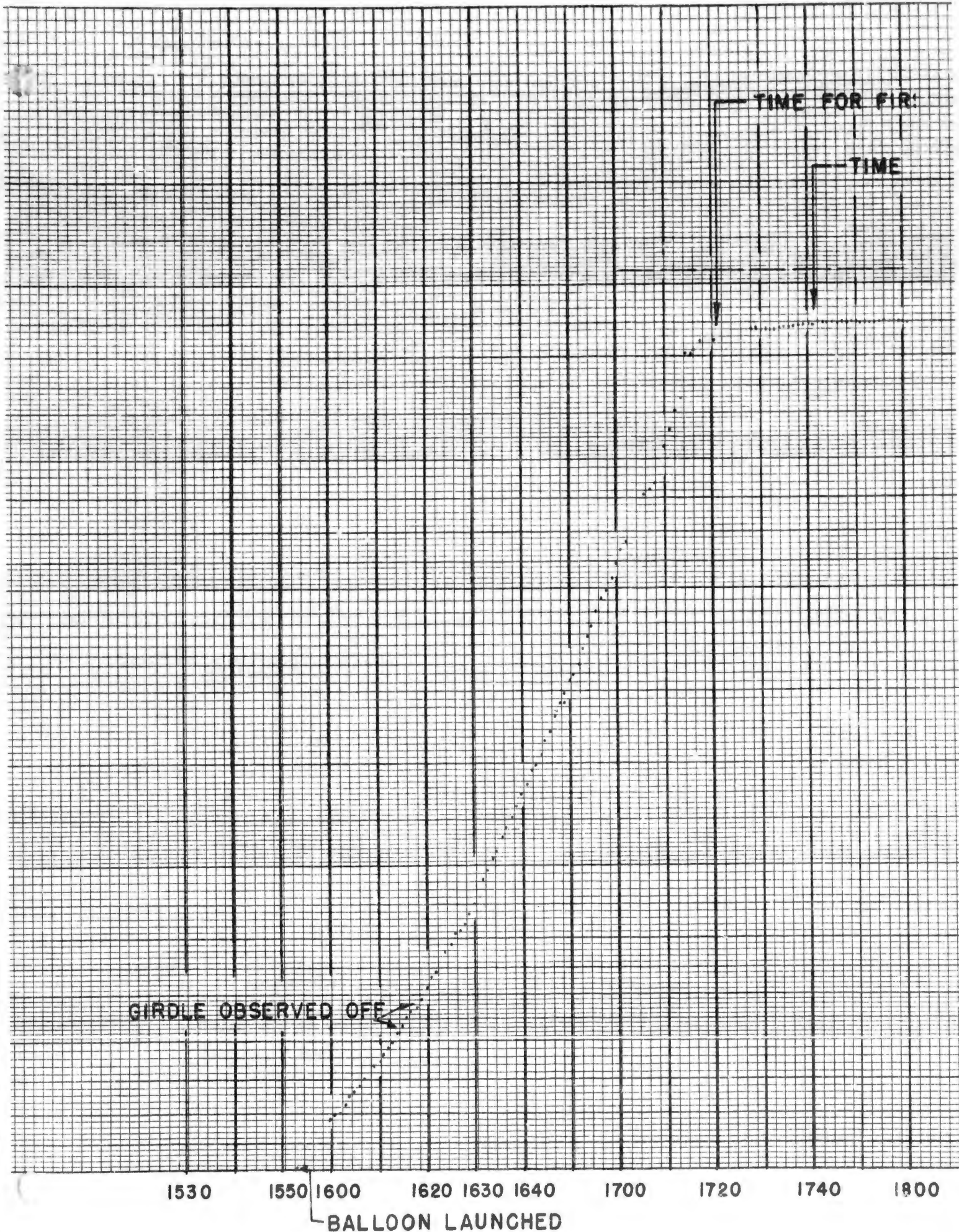












FOR FIRST 6\* BALLAST DROP

TIME FOR NO. 2 - 6\* BALLAST DROP

TIME FOR NO. 3 - 6\* BALLAST DROP

THEORETICAL CEILING  
80,000 FT.

GROUND  
SUNSET

BALLOON  
SUNSET

TIME - PRESSURE - ALTITUDE

FLIGHT NO. 15

MAY 26 - 1952

GROSS LOAD - 275 LBS.

FREE LIFT - 84 LBS.

GROSS LIFT - 359 LBS.

.... ALTITUDE DATA FROM TELEMETER RECORD

⊕ " " " "B" BAROGRAPH RECORD

1740

1800

1820

1840

1900

1920

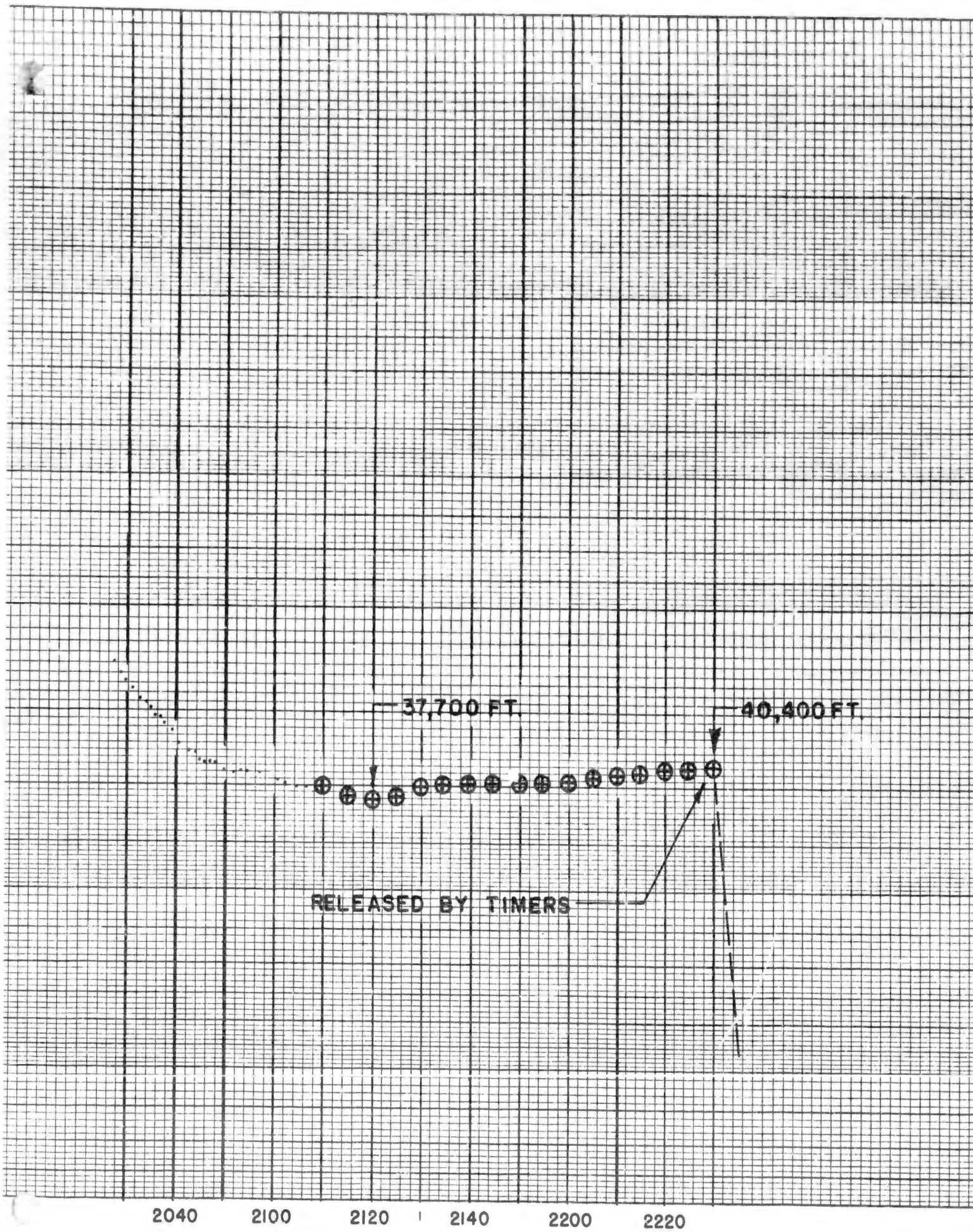
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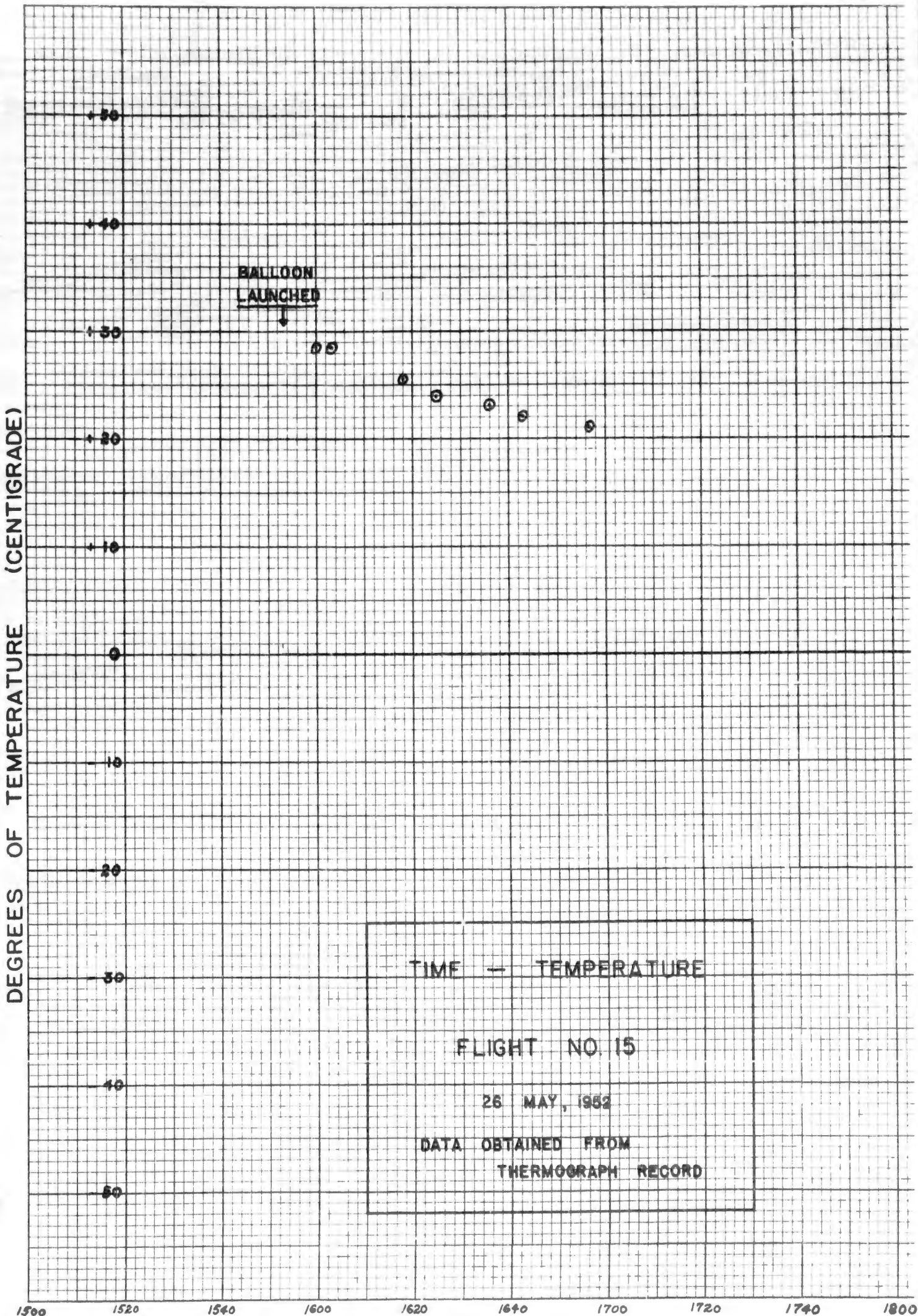
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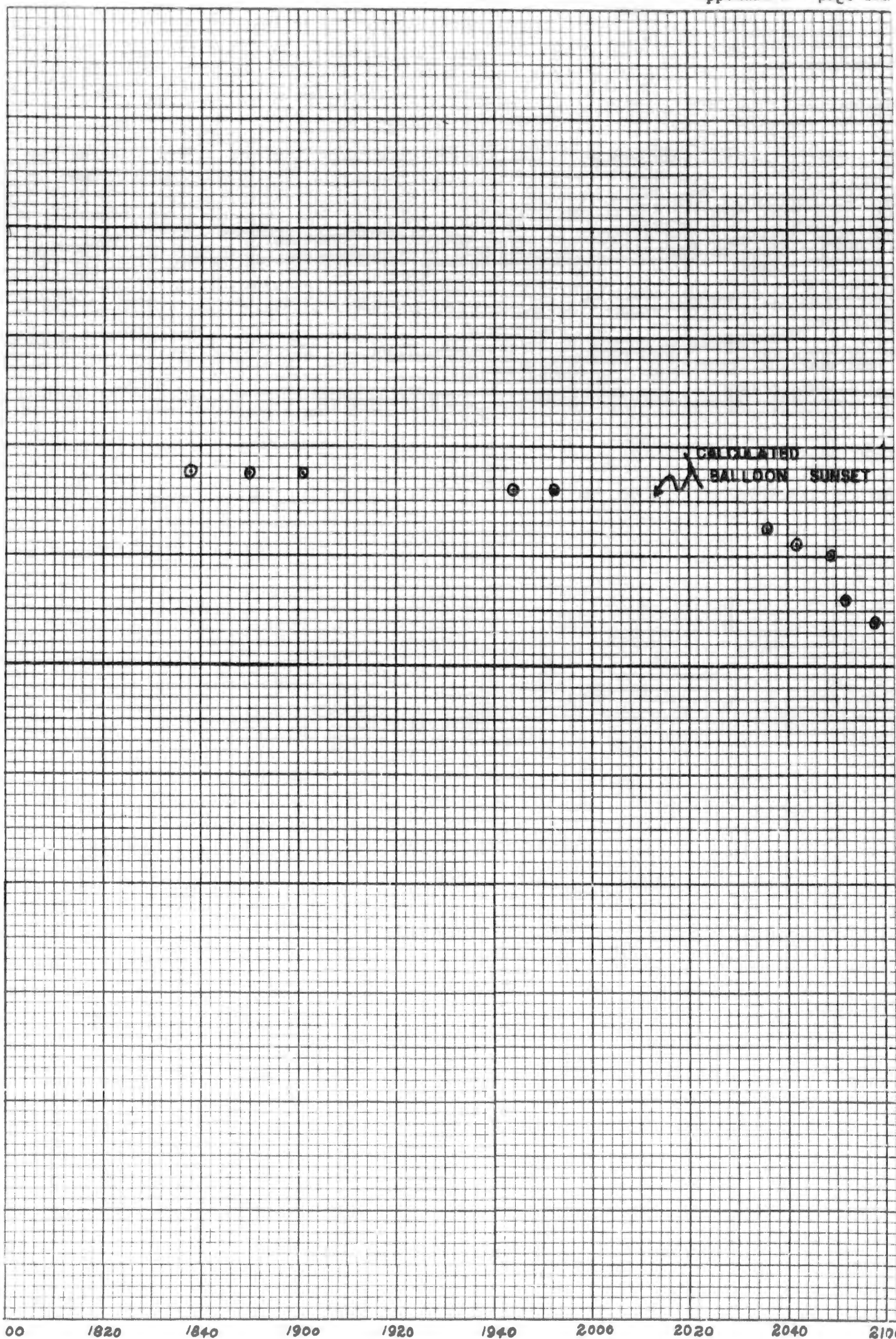
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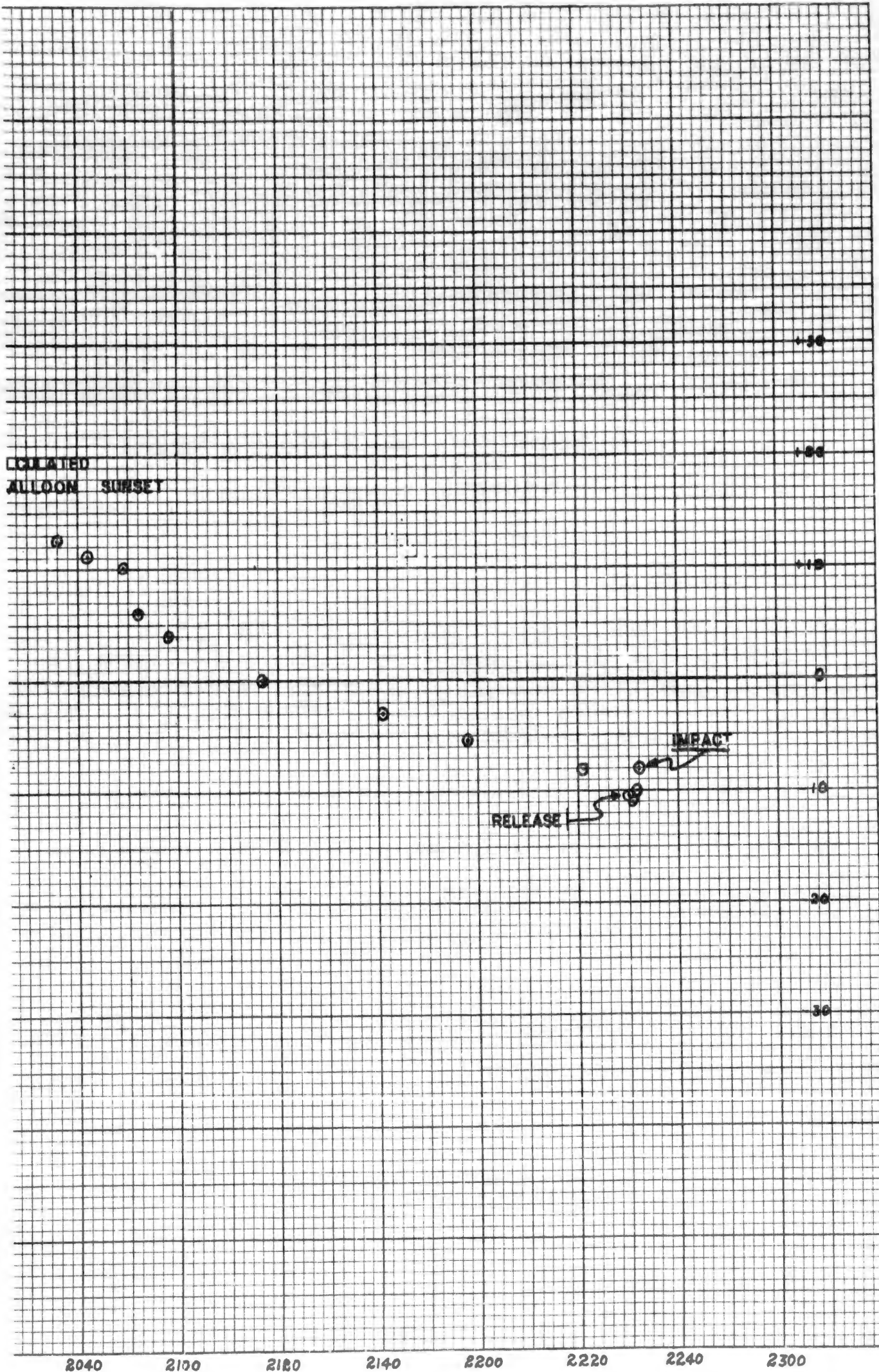
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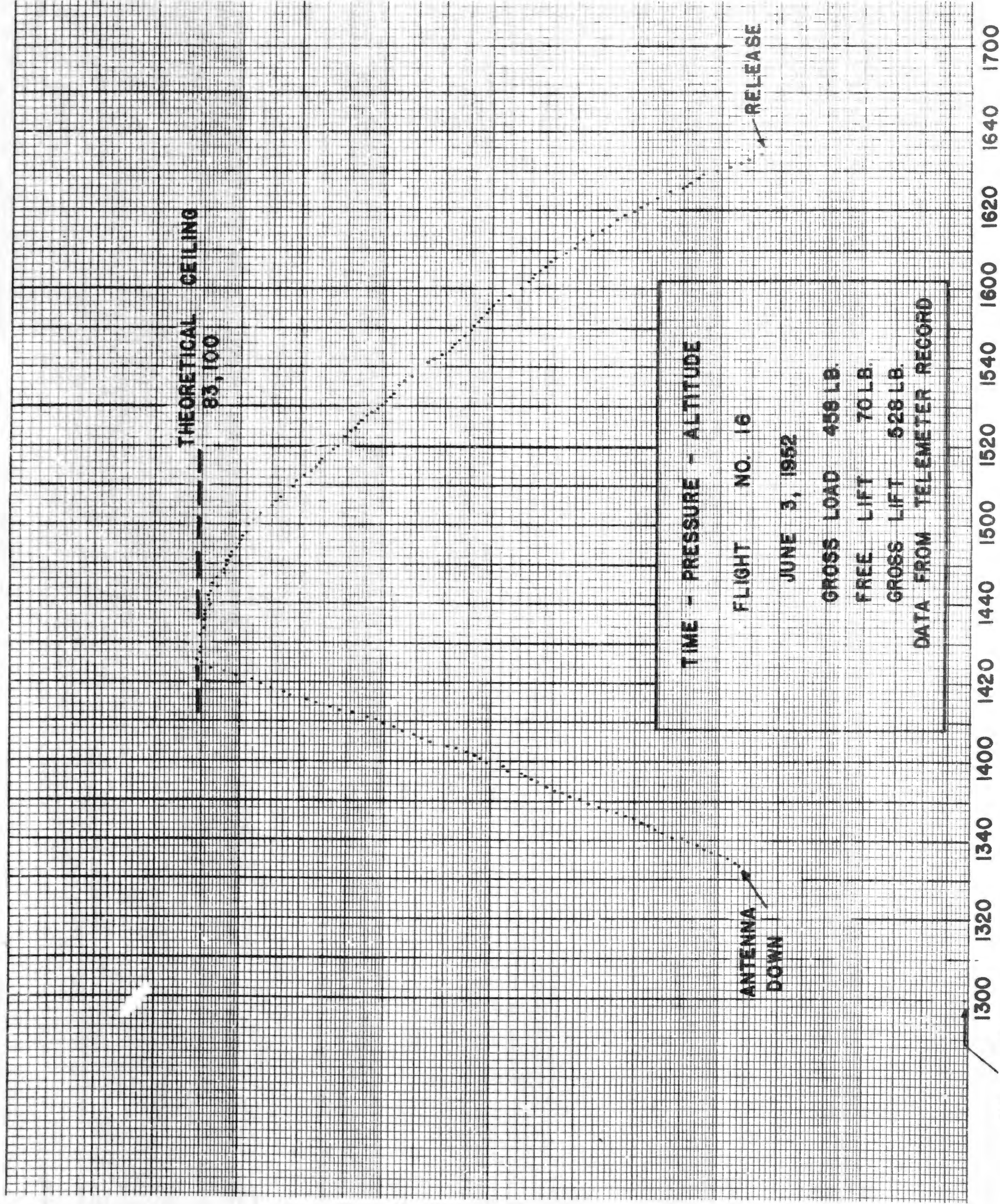
CENTRAL STANDARD TIME







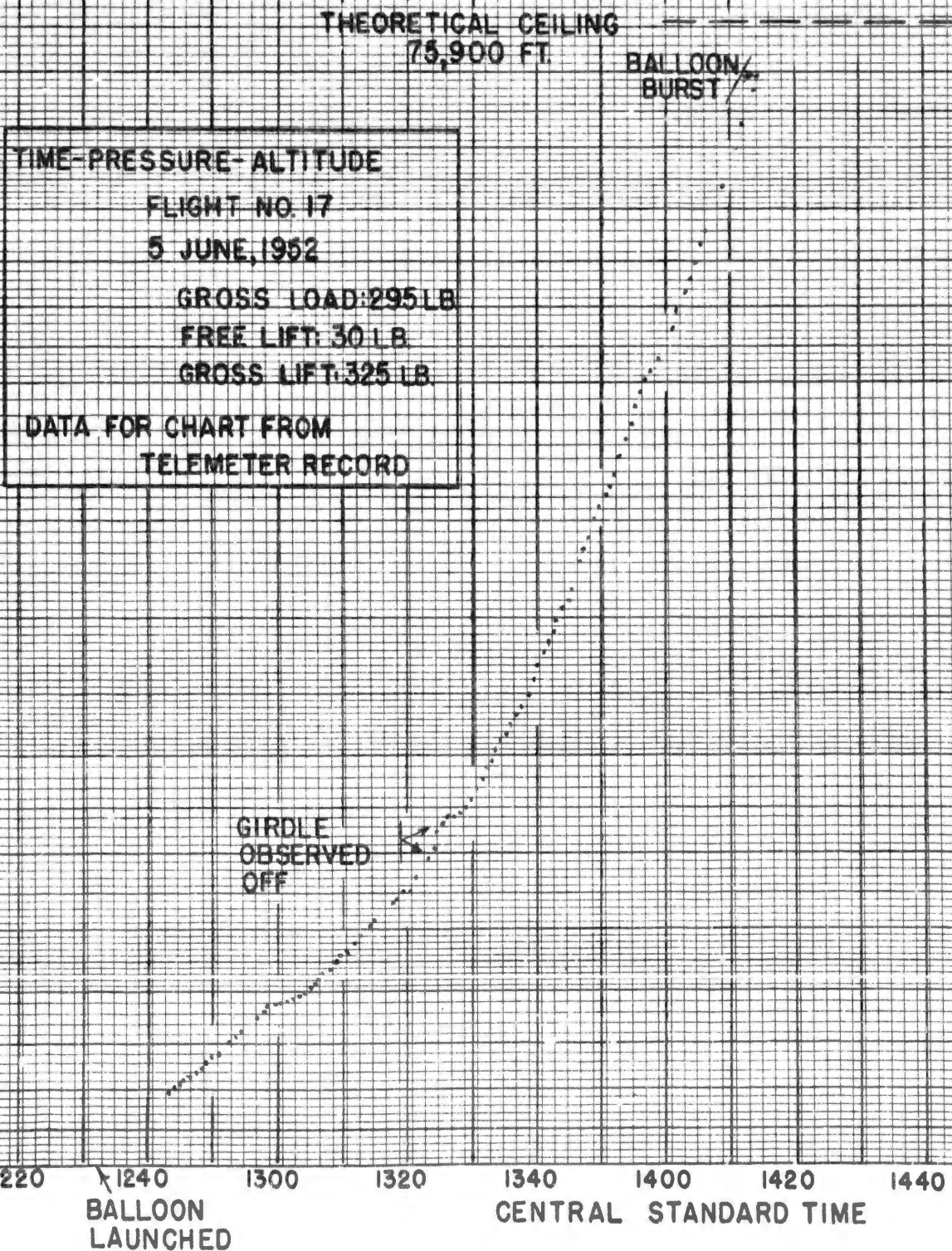


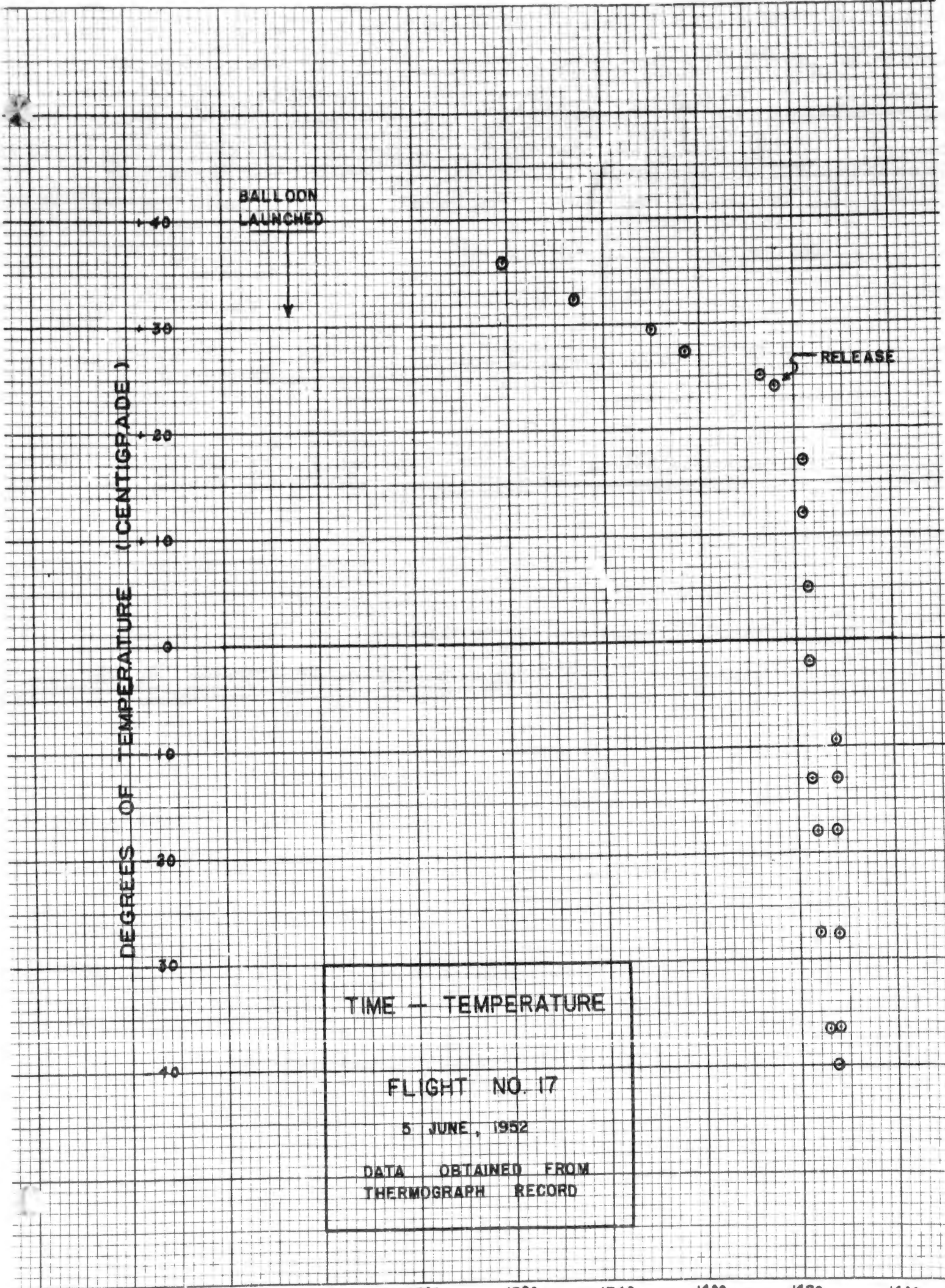


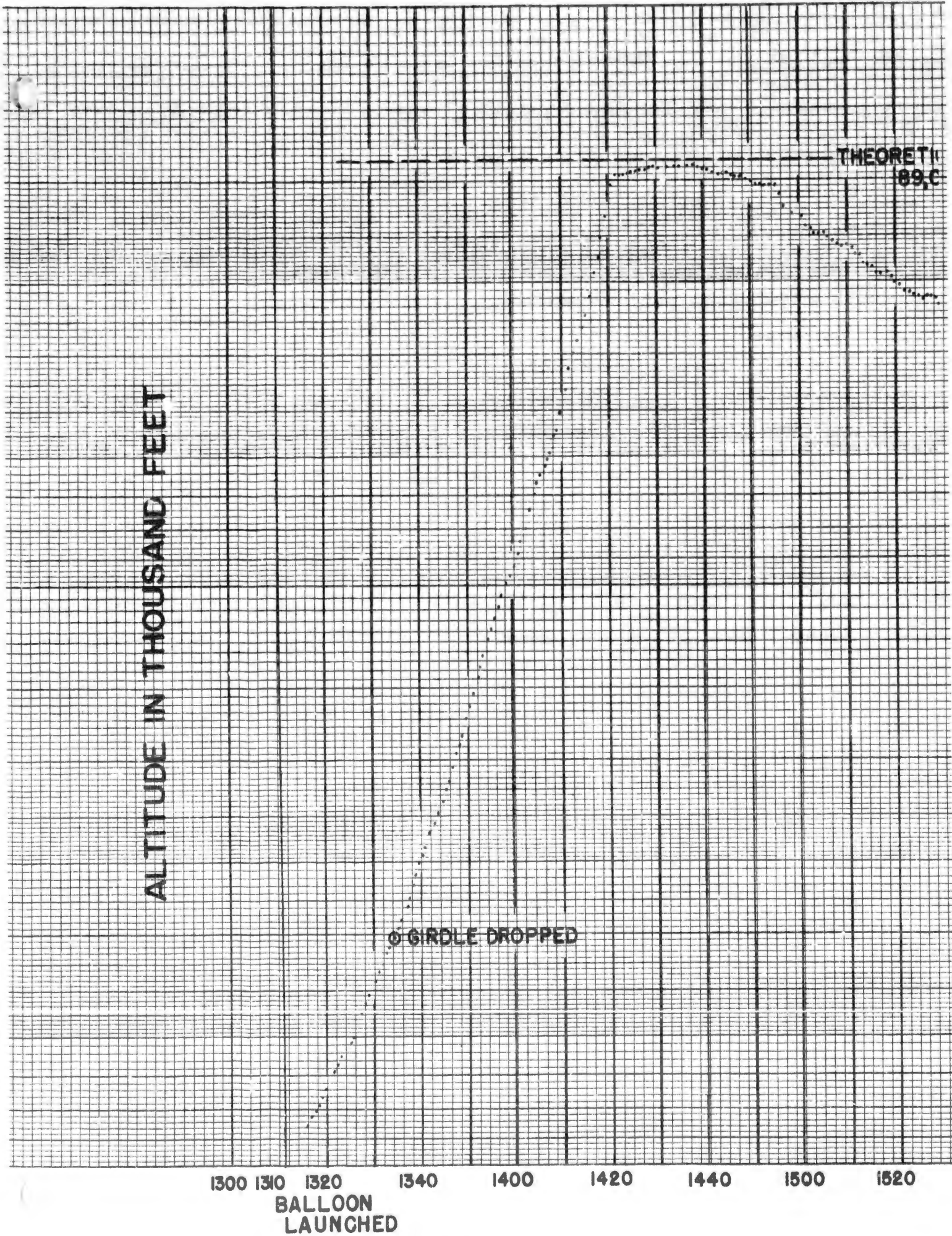
BALLOON LAUNCHED

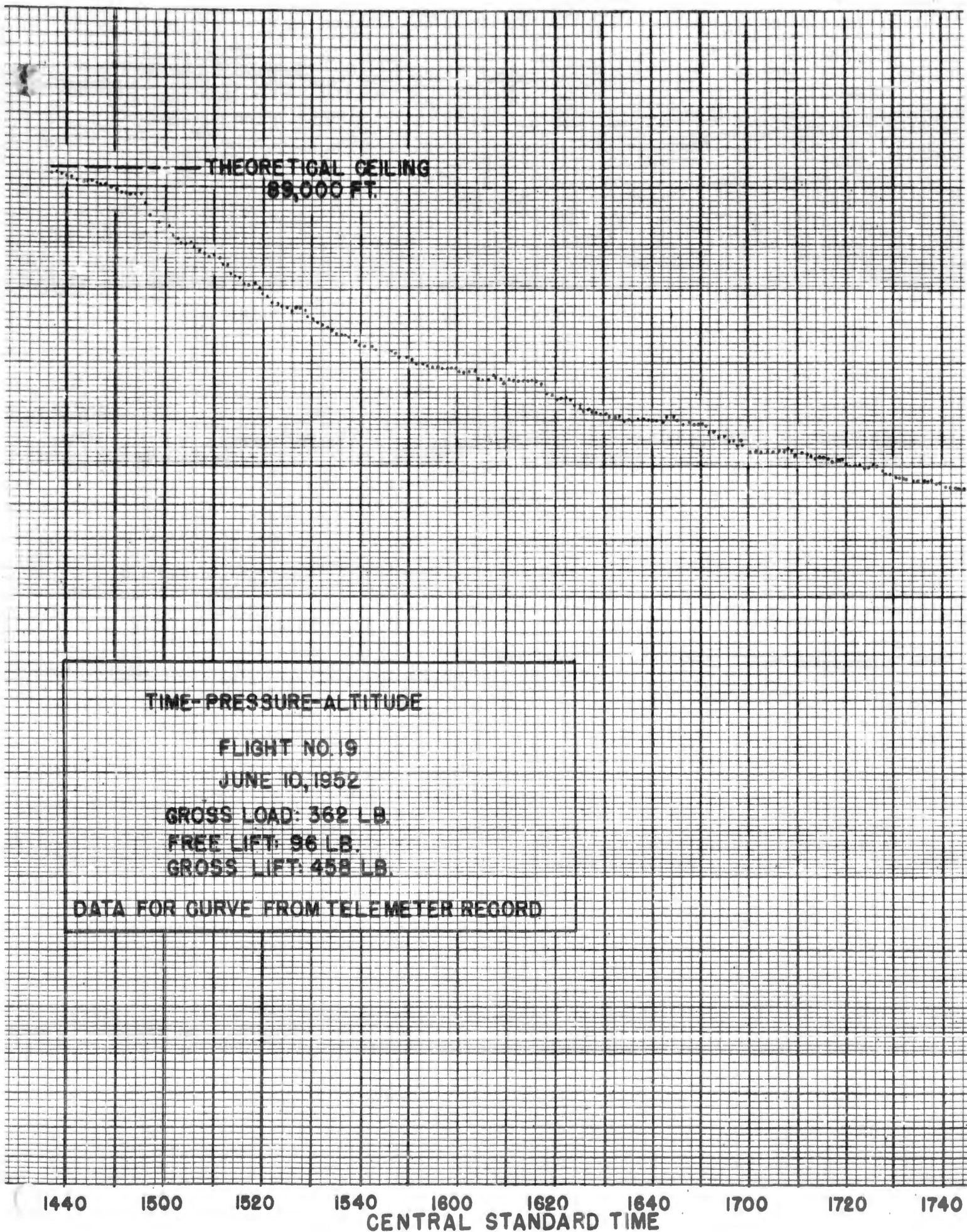
CENTRAL STANDARD TIME

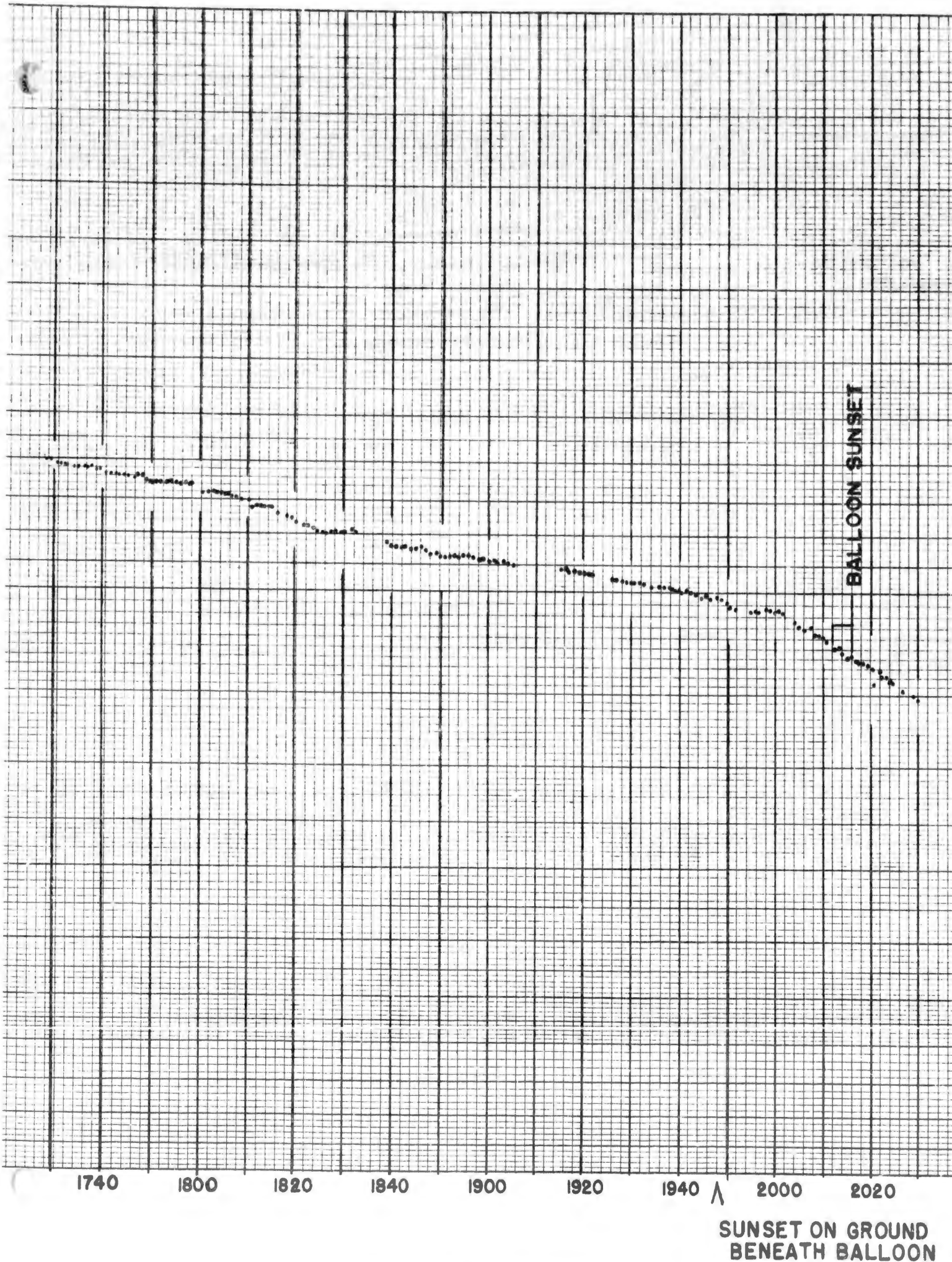
CONFIDENTIAL SECURITY INFORMATION

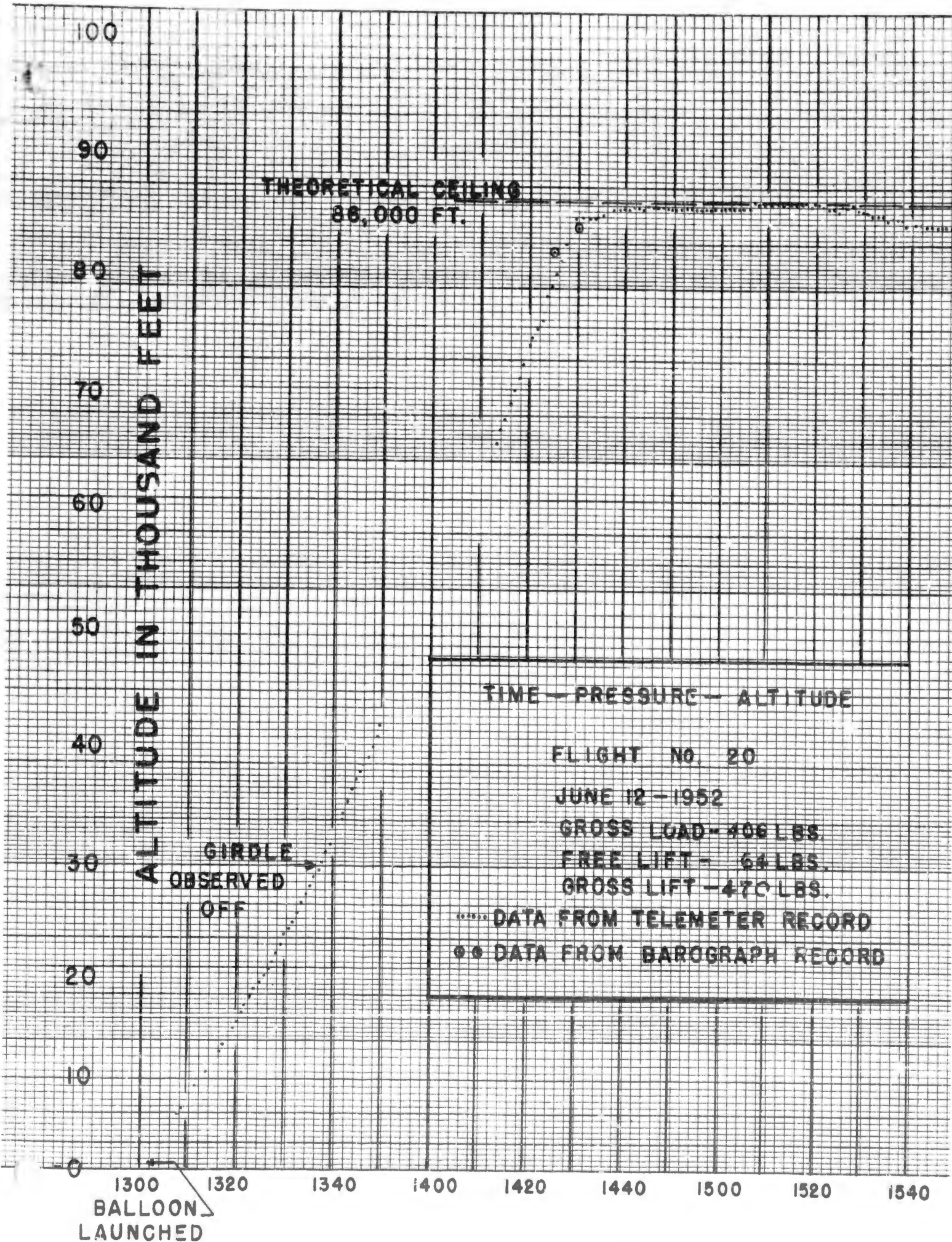


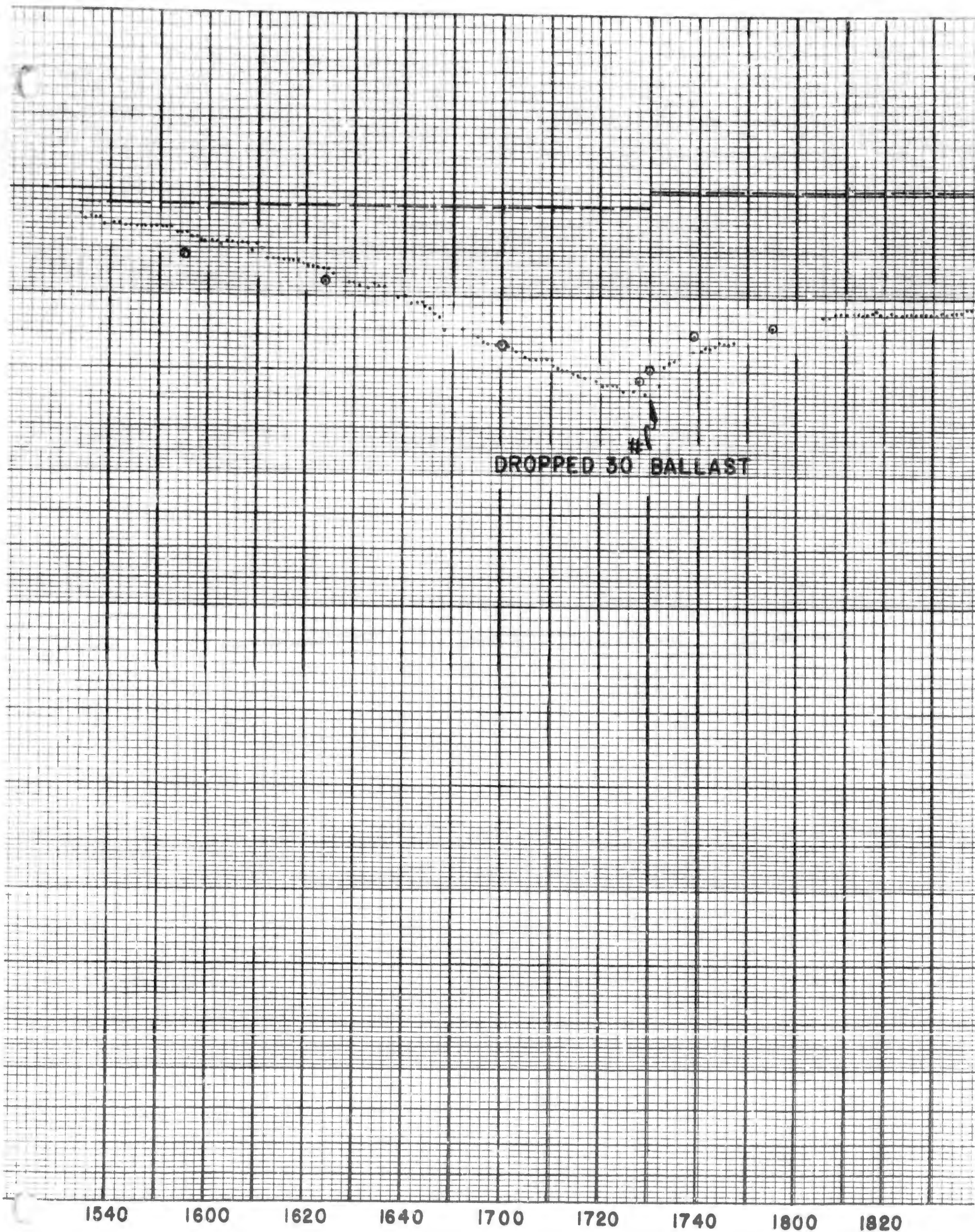


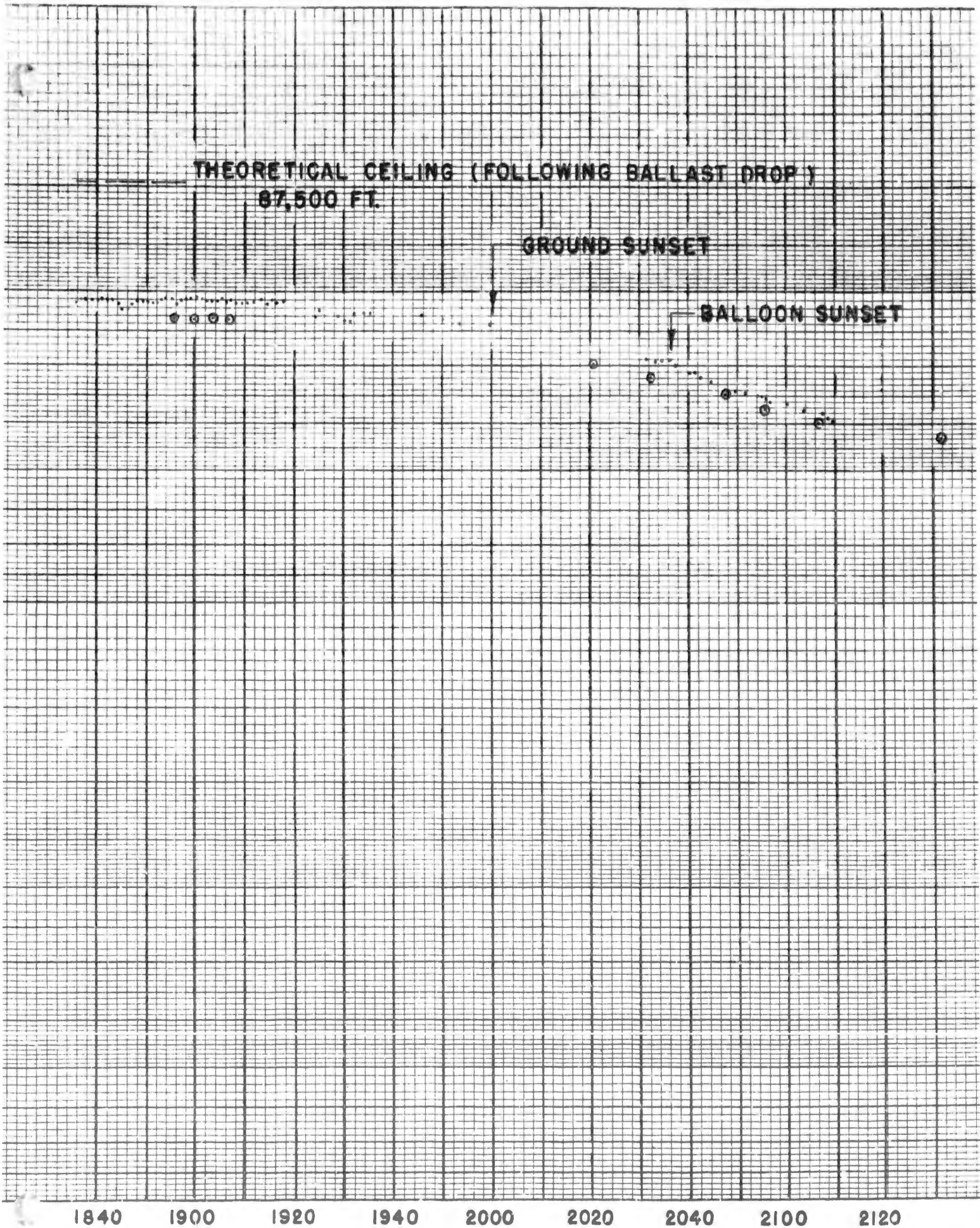


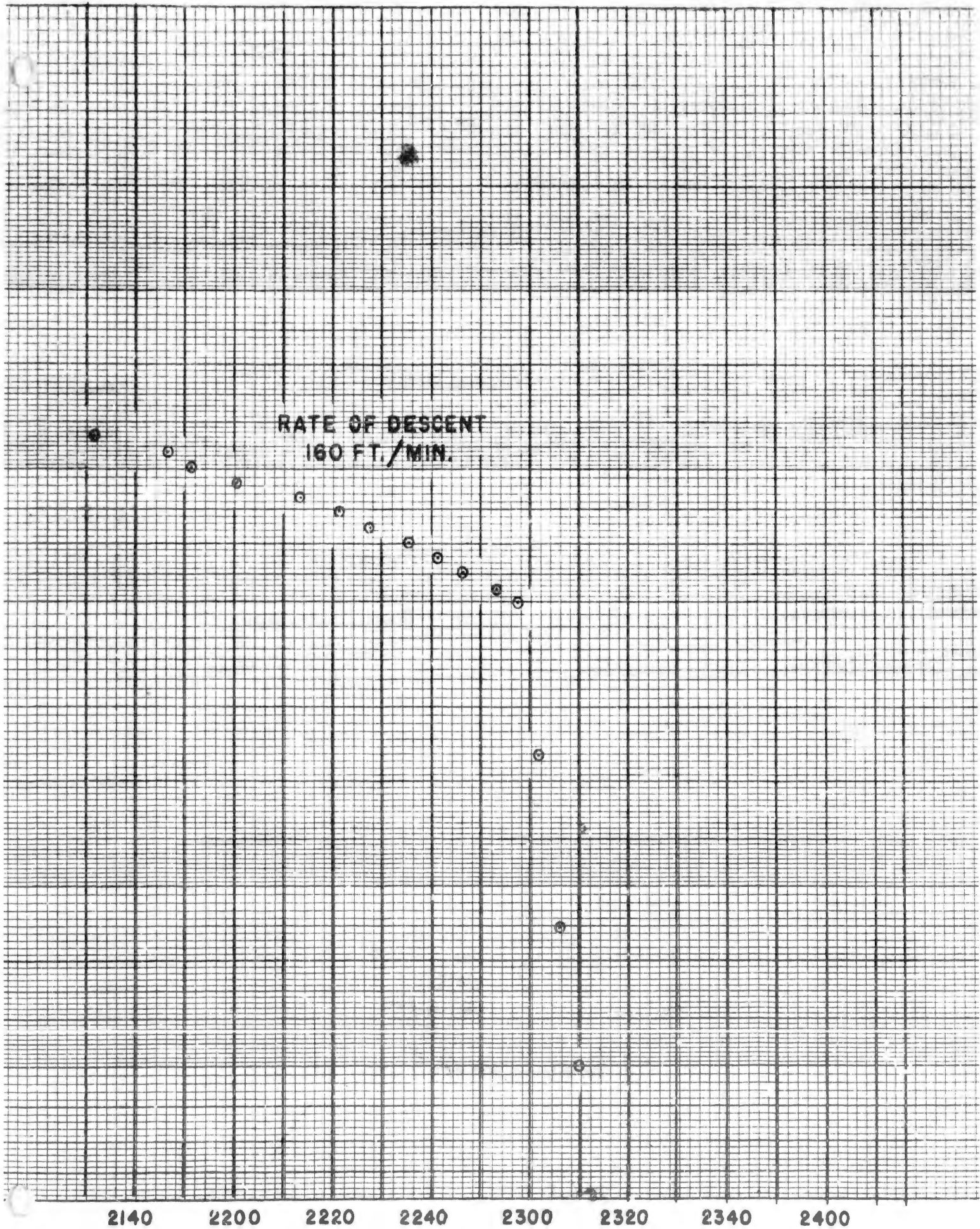


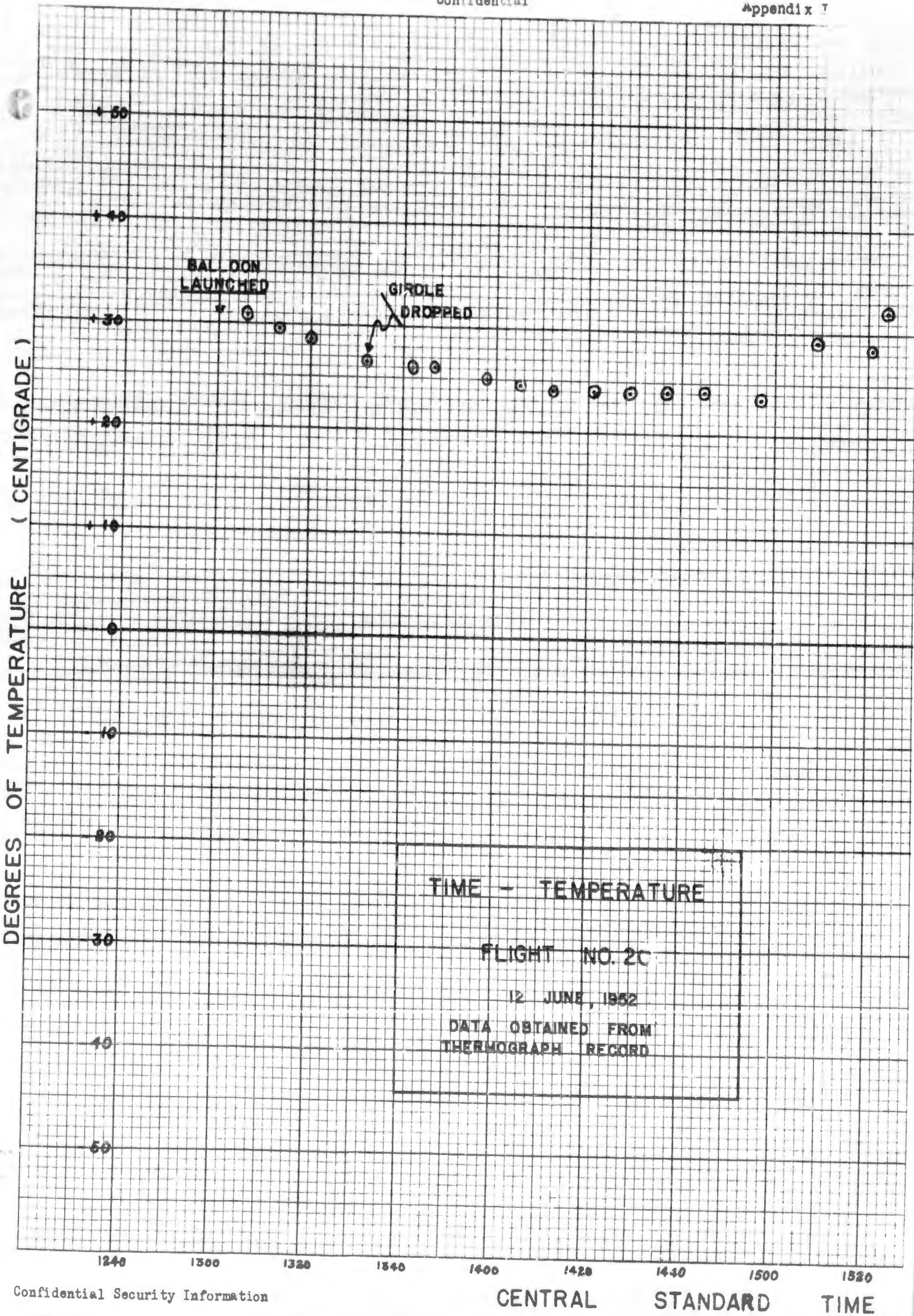


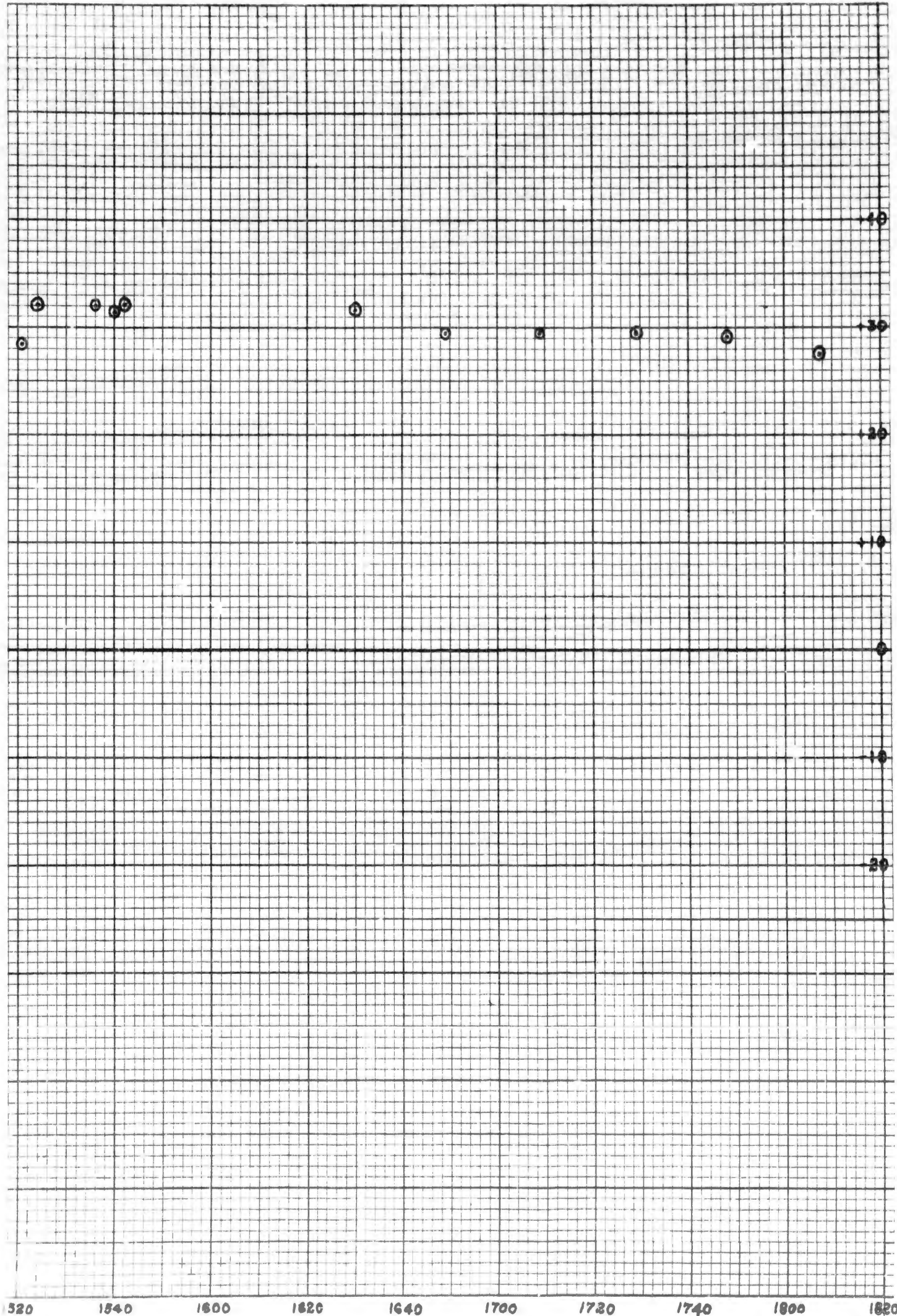


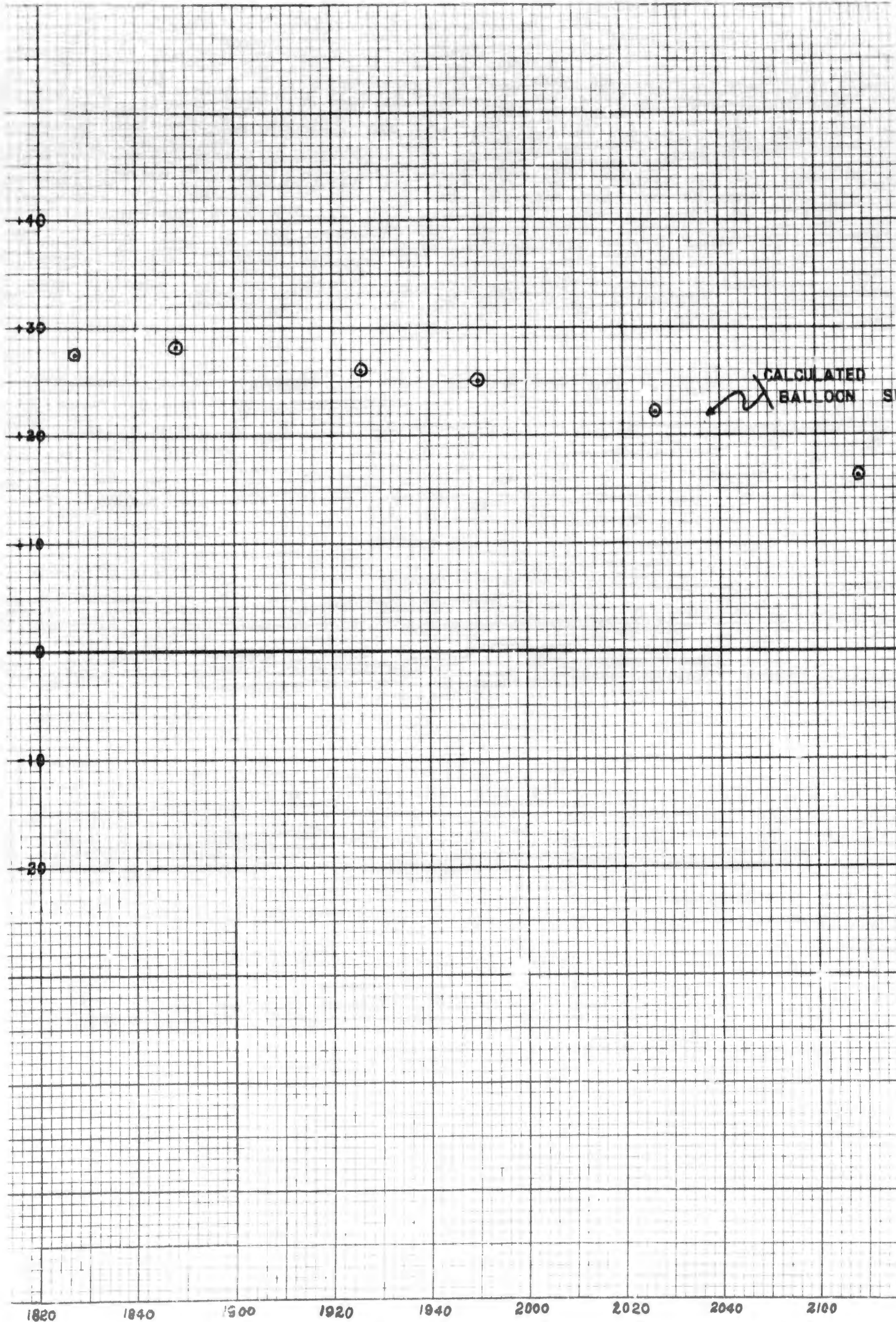


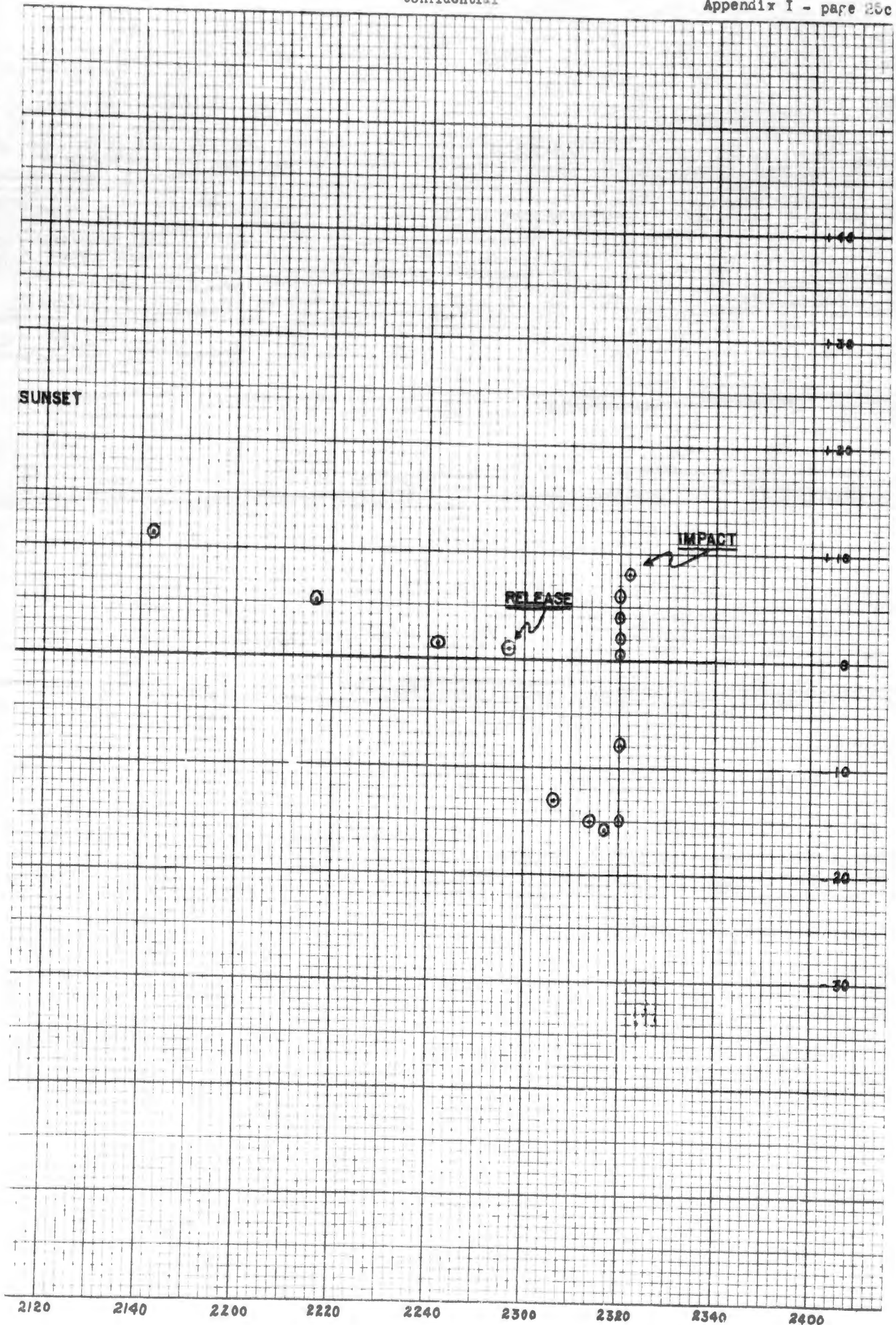












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**APPENDIX X**

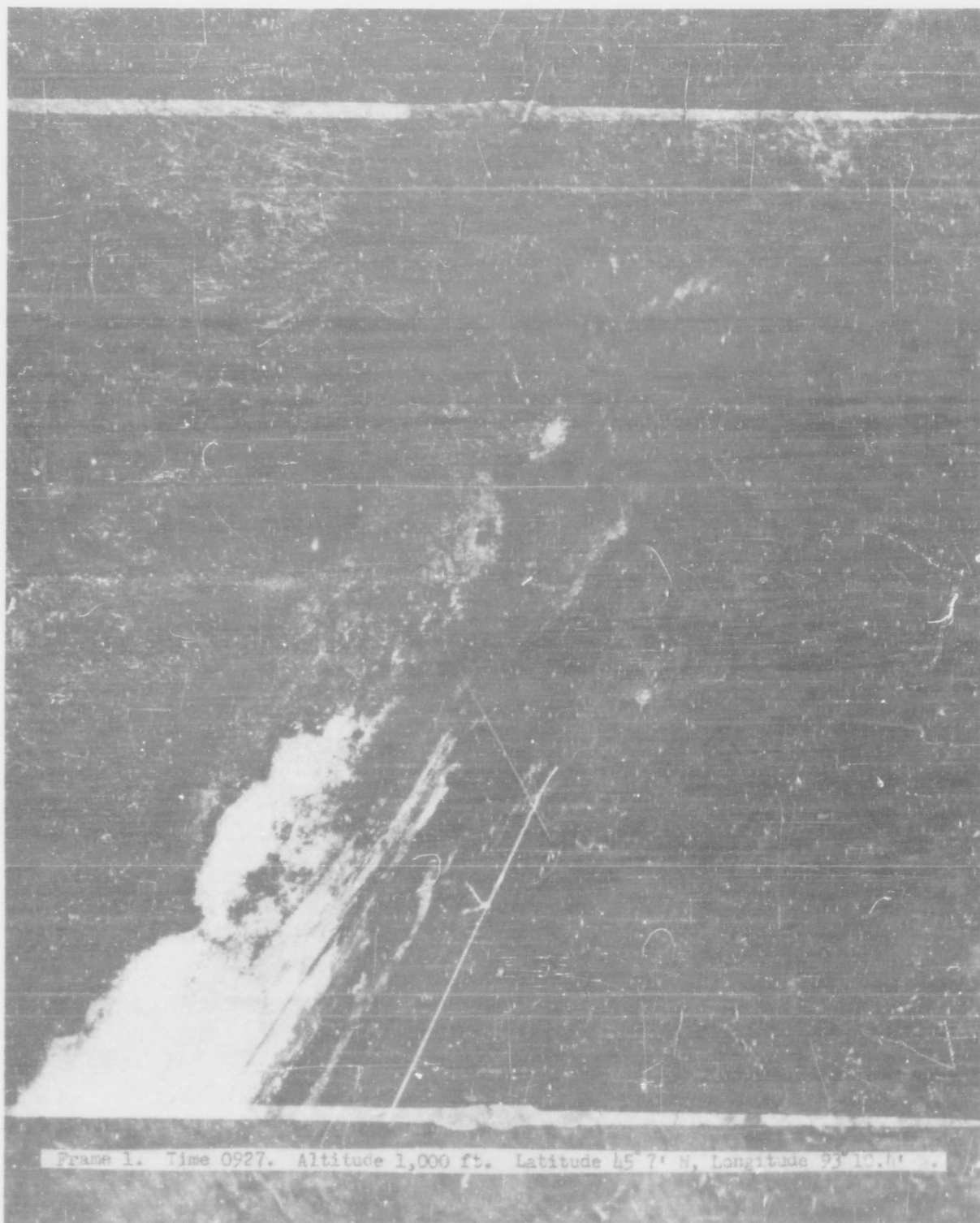
**DOWN PHOTOGRAPHS FROM FLIGHT 8**

These photographs are pictures taken looking down from the gondola at the ground in Flight 8. The complete set of pictures has been used to obtain an accurate trajectory for the flight. The pictures taken during the ascent of the balloon are being used to evaluate the atmospheric temperature.

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Frame 1. Time 0927. Altitude 1,000 ft. Latitude 45° 7' N, Longitude 93° 10.0'

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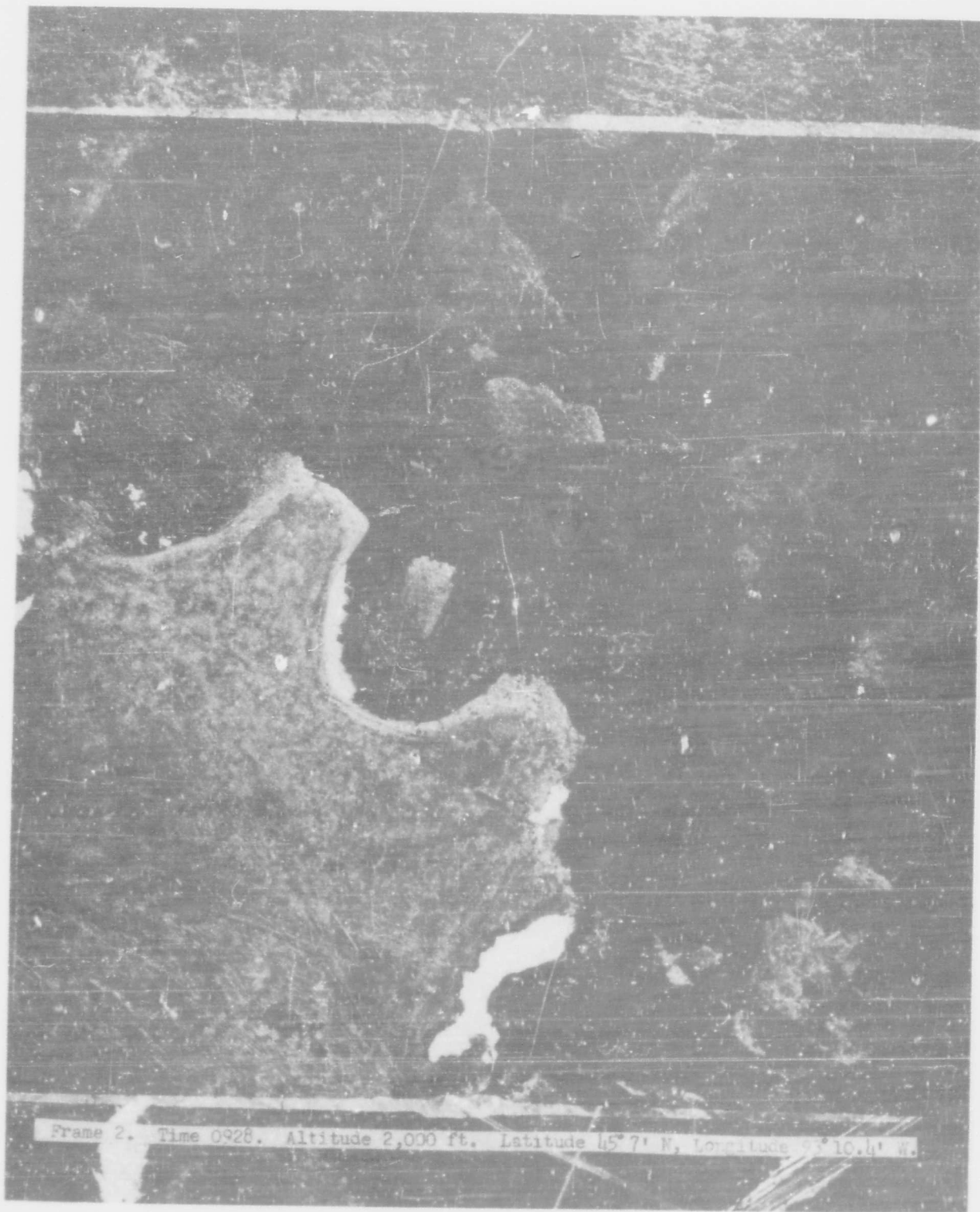


Frame 7. Time 0933. Altitude 6,000 ft. Latitude 45° 5.5' N, Longitude 93° 10.3' W.

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Unclassified Security Information



Frame 11. Time 0937. Altitude 10,000 ft. Latitude 45° 4.2' N, Longitude 93° 9.4' W.

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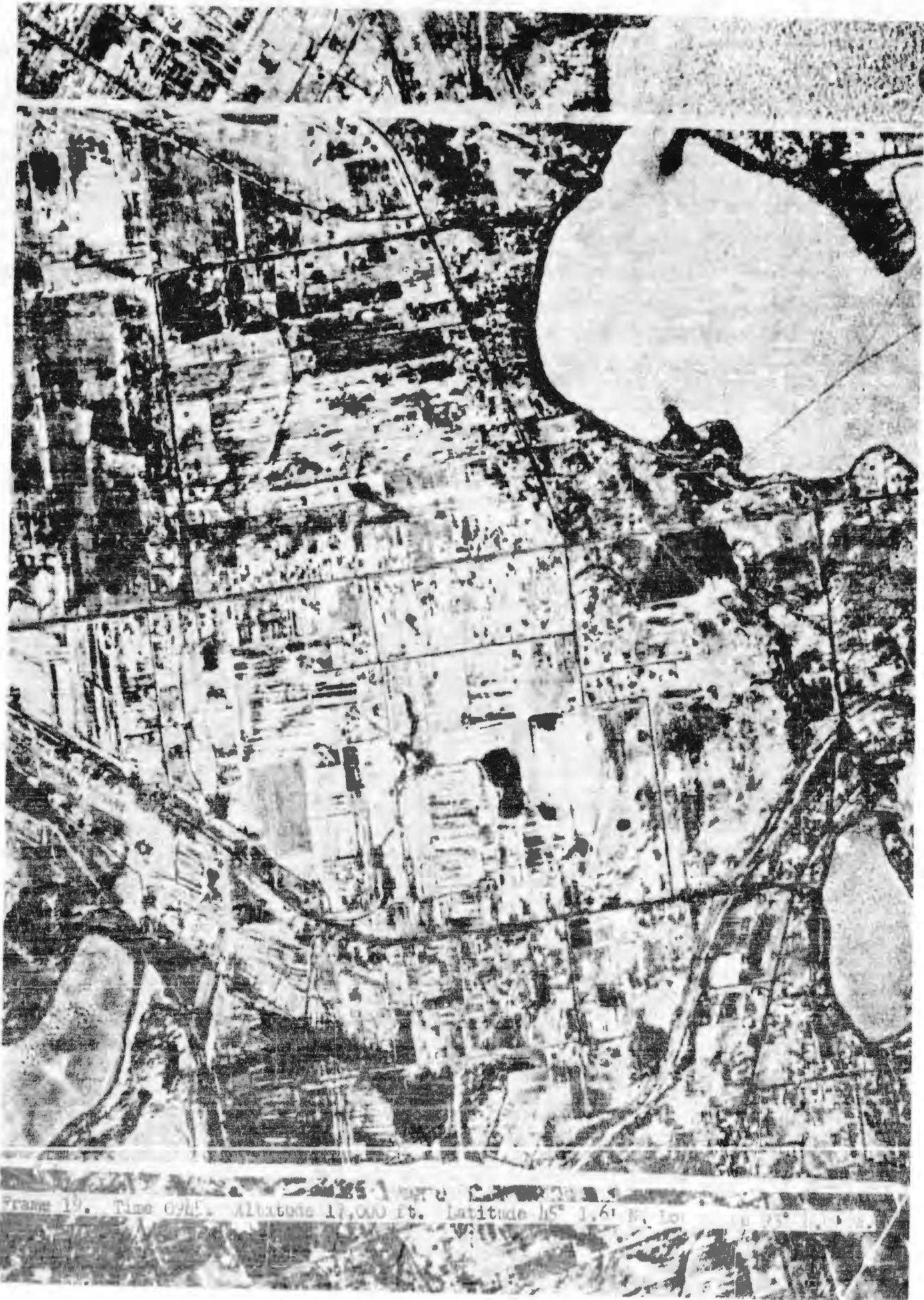


Frame 7. Time 0933. Altitude 6,000 ft. Latitude 45° 5.5' N, Longitude 93° 10.3' W.

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Frame 15. Time 09h1. Altitude 11,000 ft. Latitude 45° 2.5' N, Longitude 93° 7.0'



Frame 19. Time 0945. Altitude 17,000 ft. Latitude 45° 1.61' N, Lon 122° 45' 10.12' W

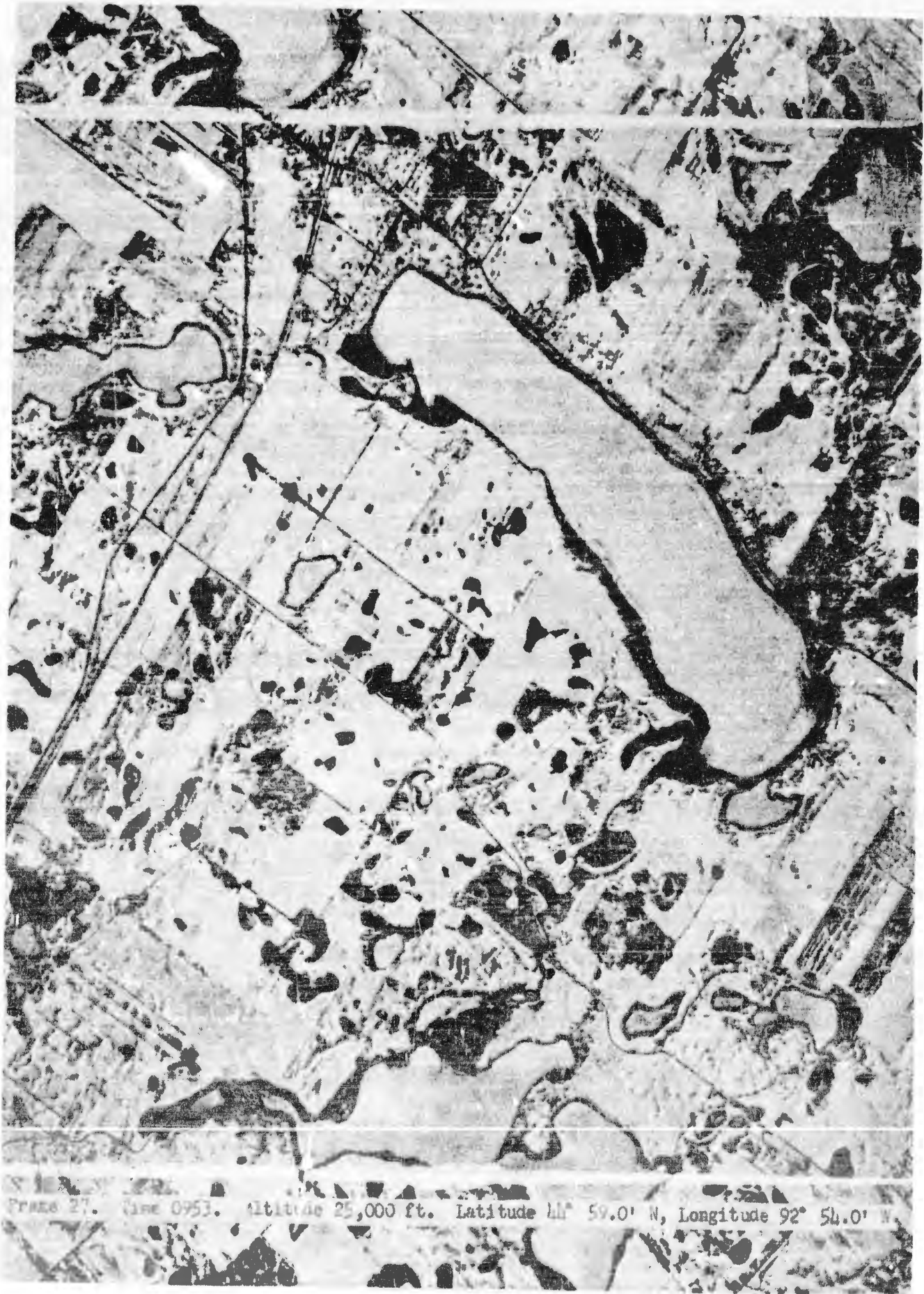
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Frame 23. Time 0919. Altitude 21,000 ft. Latitude 15° 0.6' N. Longitude 92° 59.8' W.

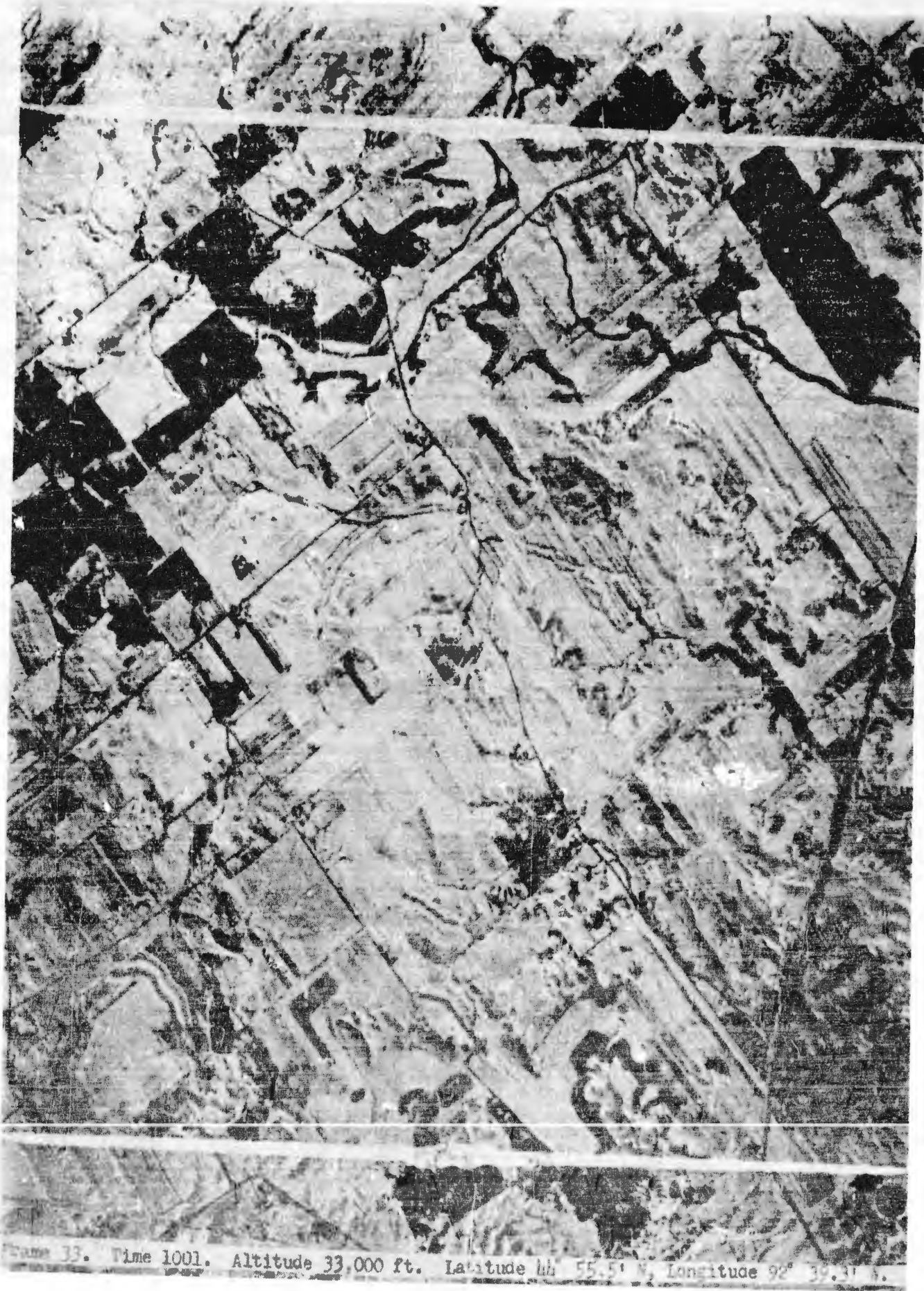
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FRAMES 27. Time 0953. Altitude 25,000 ft. Latitude 44° 59.0' N, Longitude 92° 54.0' W.



Frame 11. Time 0957. Altitude 29,000 ft. Latitude 44° 58.7' N, Longitude 92° 45.5' W



Frame 33. Time 1001. Altitude 33,000 ft. Latitude 44° 55.5' N, Longitude 92° 39.3' W.



Frame 39, Time 1005. Altitude 36,000 ft. Latitude 33° 55.3' N, Longitude 92° 33.1'

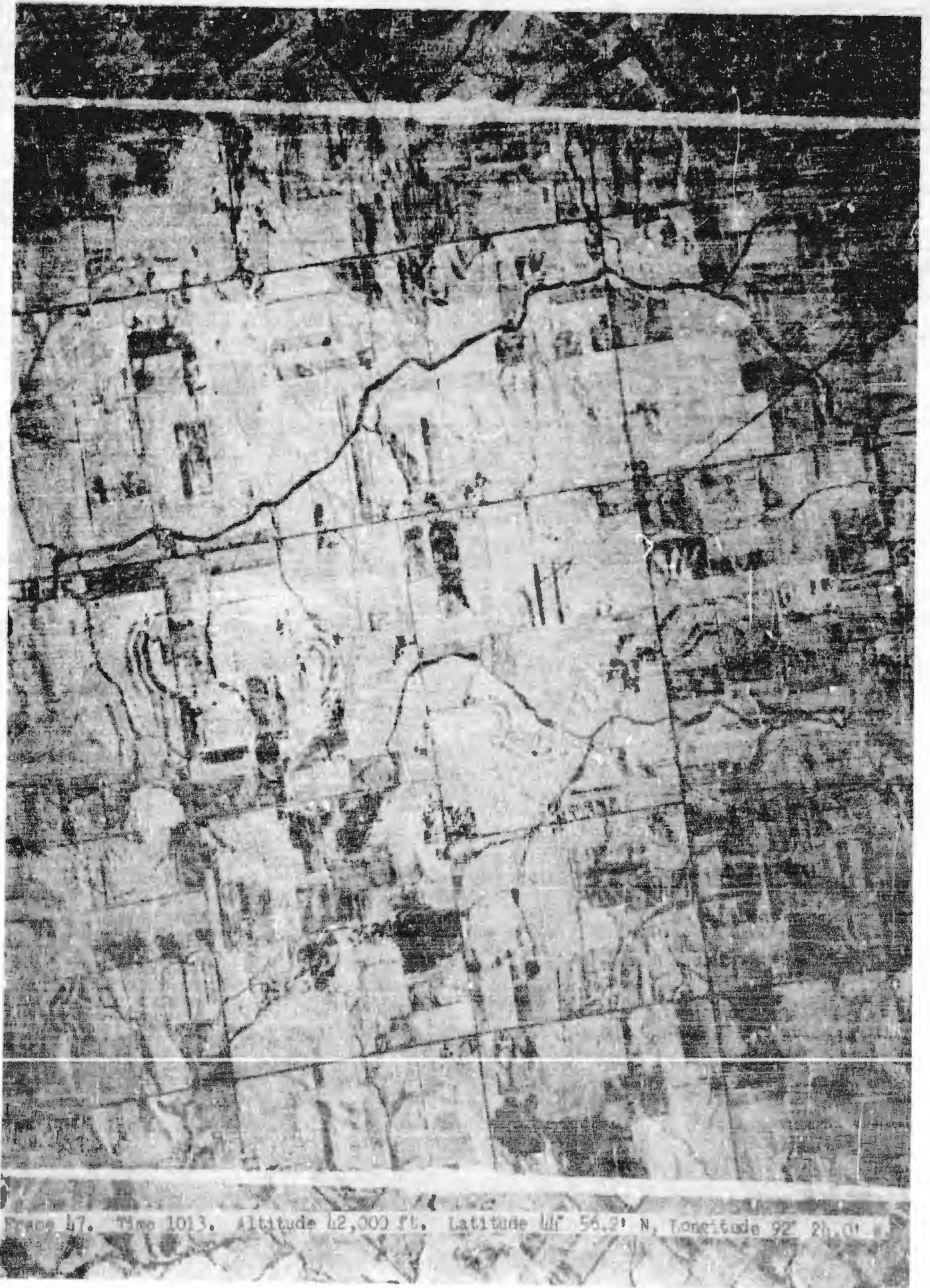
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Frame 13. Time 1009. Altitude 39,000 ft. Latitude 44° 56.0' N, Longitude 92° 29.0' W.

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Frame 17. Time 1013. Altitude 42,000 ft. Latitude 41° 56.9' N, Longitude 92° 21.0'

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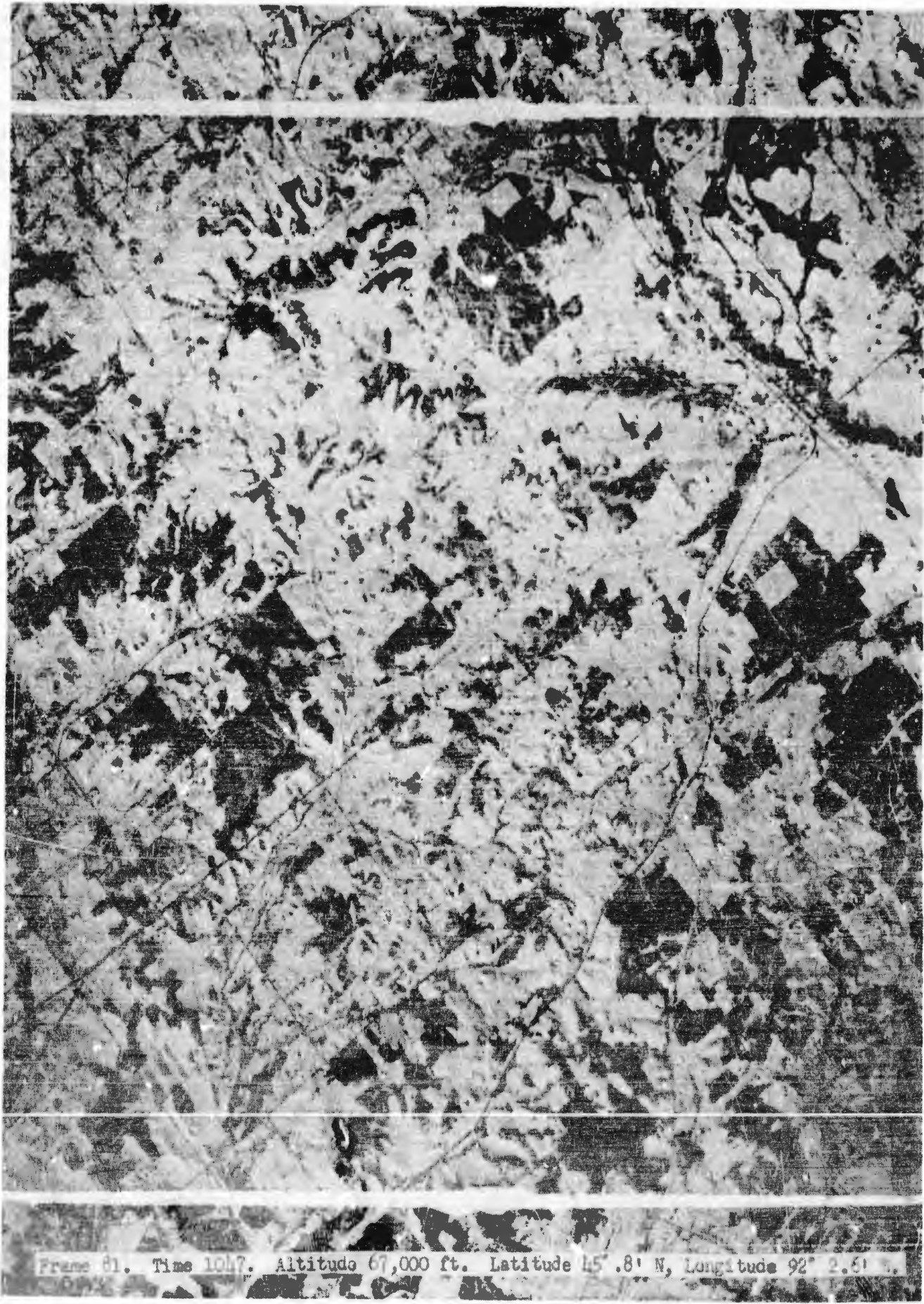
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Frame 72. Time 1036. Altitude 62,000 ft. Latitude 45° N. Longitude 92° 5.6' W.

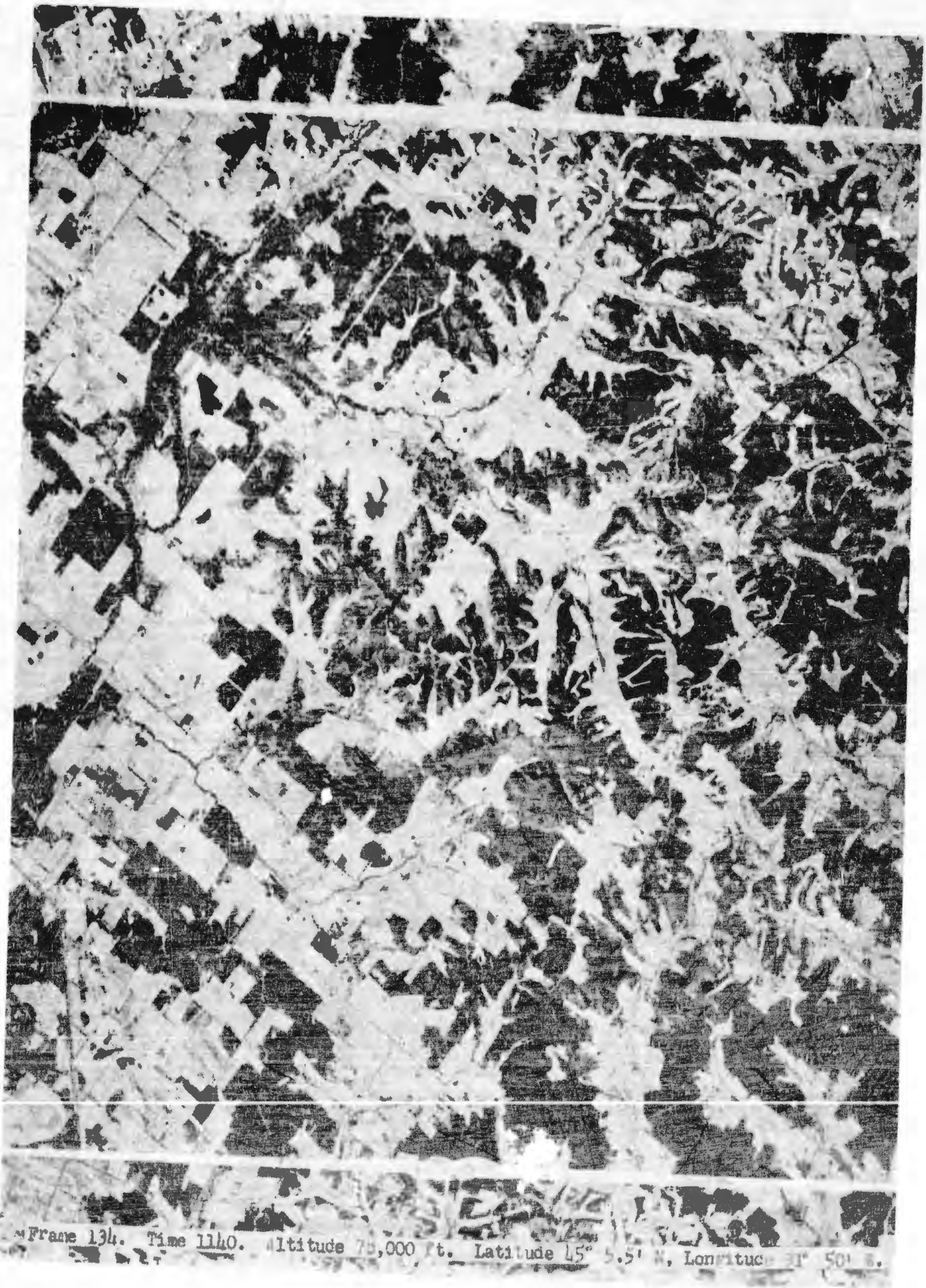
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Frame 81. Time 10h7. Altitude 67,000 ft. Latitude 45° .8' N, Longitude 92° 2.6' W.





Frame 134. Time 1140. Altitude 75,000 ft. Latitude 45° 5.5' N, Longitude 31° 50' E.

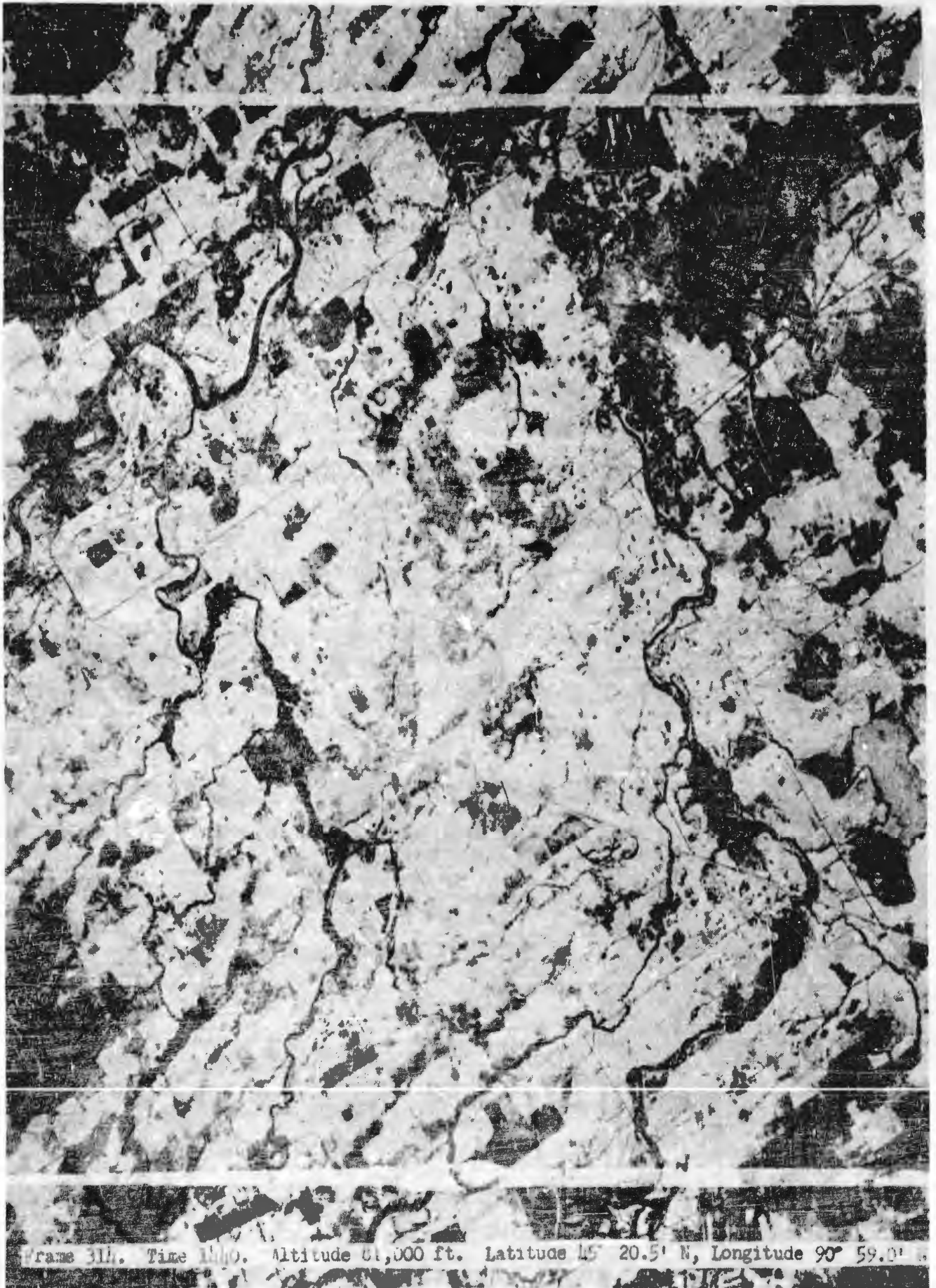


Frame 15h. Time 1200. Altitude 60,000 ft. Latitude 45° 2.2' N, Longitude 91° 46.4' W.



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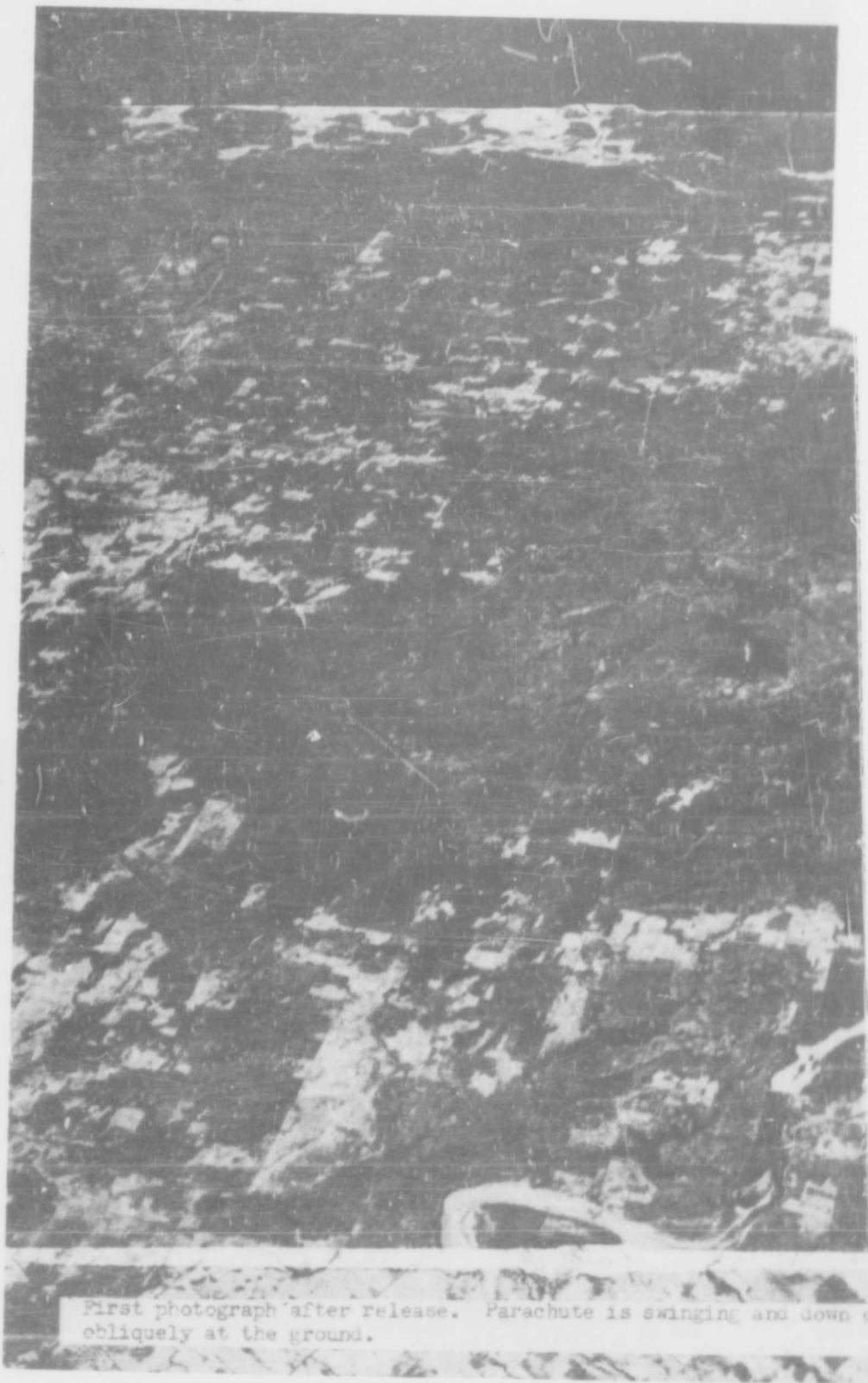


Frame 31h. Time 1440. Altitude 61,000 ft. Latitude 45° 20.5' N, Longitude 90° 59.0' W.

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