

10701403

LAMONT-DOHERTY GEOLOGICAL OBSERVATORY
OF COLUMBIA UNIVERSITY
Palisades, New York 10964

THE INTEGRAL SOLUTION OF THE SOUND FIELD
IN A MULTILAYERED LIQUID-SOLID HALF-SPACE
WITH NUMERICAL COMPUTATIONS FOR LOW-FREQUENCY PROPAGATION
IN THE ARCTIC OCEAN

by

Henry W. Kutschale

CVL-1-70

Technical Report No. 1

Contract N00014-67-1-0108-0015 with the
Office of Naval Research

Work done on behalf of the U.S. Naval Ordnance Laboratory,
White Oak, Silver Spring, Maryland

Reproduction of this document in whole or in part is permitted
for any purpose of the U.S. Government.

Reproduced by the
CLEARINGHOUSE
for Federal Scientific & Technical
Information Springfield Va. 22151

This document has been prepared
by the Clearinghouse and is
available in microfiche

D D C
REGISTERED
APR 22 1970
RESEARCH

Corrections

Page 7, line 3, should read as follows:

in terms of elements of

Page 13, line 4, should read as follows:

$$(p_{zz})_{n-1} = A_{21} \dot{\hat{w}}_0 + A_{22} (\hat{p}_{zz})_0$$

Page 16, line 8, should read as follows:

$$\begin{bmatrix} \dot{\hat{w}}_{n-1} \\ (p_{zz})_{n-1} \end{bmatrix} = a_{n-1} \cdots b_{n-3} a_{n-4} \cdots b_1 \begin{bmatrix} \dot{\hat{w}}_0 \\ 0 \end{bmatrix}$$

Page 19, line 1, should read as follows:

$$A = A_{S_2} A_{S_1} = a_{n-1} a_{n-2} \cdots b_{n-10} \cdots a_1$$

Page 22, line 7, should read as follows:

at depth z at the bottom of the D_1 layer is $p_{D_1}(r, z, k, t) =$

Page 22, line 10, should read as follows:

$$A_{D_1} = a_{D_1} a_{D_1-1} \cdots a_1$$

Page 28, line 2, should read as follows: but from the relation

$$AA^{-1} = I, \quad A_{11}A_{22} - A_{12}A_{21} = 1$$

Page 41, second line, should read as follows:

$$\left\{ \frac{2\sqrt{2}\pi^2 \sqrt{\omega^2}}{\rho_n c^{5/2} \bar{A}_{11} \frac{\partial}{\partial c} \left[\frac{c^2 \bar{A}_{11}}{\text{Im}(\rho_{an})} + \frac{\text{Im}(-\bar{A}_{21})}{\rho_n} \right]} \right\}_{\substack{\omega = \omega_{rn} \\ c = c_{im}}}$$

LAMONT-DOHERTY GEOLOGICAL OBSERVATORY
OF COLUMBIA UNIVERSITY
Palisades, New York 10964

THE INTEGRAL SOLUTION OF THE SOUND FIELD
IN A MULTILAYERED LIQUID-SOLID HALF-SPACE
WITH NUMERICAL COMPUTATIONS FOR LOW-FREQUENCY PROPAGATION
IN THE ARCTIC OCEAN

by

Henry W. Kutschale

CU1-1-70

Technical Report No. 1

Contract N00014-67-A-0108-0016 with the
Office of Naval Research

Work done on behalf of the U.S. Naval Ordnance Laboratory,
White Oak, Silver Spring, Maryland

Reproduction of this document in whole or in part is permitted
for any purpose of the U.S. Government.

February 1970

TABLE OF CONTENTS

	<u>Pages</u>
Abstract	1
Captions for Figures	2
Definition of Symbols	3-5
Introduction	6-9
Formal Solution	9-39
Numerical Computations	40-56
Acknowledgements	56
References	57-58
Appendix	59-64

ABSTRACT

This report develops by matrix methods the integral solution of the wave equation for point sources of harmonic waves in a liquid layer of a multilayered liquid-solid half space in a form convenient for numerical computation on a high-speed digital computer. Only the case is considered of a high-speed liquid bottom underlying the stack of layers. The integral over wave number has singularities in the integrand and is conveniently transformed into the complex plane. By a proper choice of contours, complex poles are displaced to an unused sheet of the two-leaved Riemann surface, and the integral solution for the multilayered system reduces to a sum of normal modes plus the sum of two integrals, one along the real axis and the other along the imaginary axis. Both integrals are evaluated by a Gaussian quadrature formula. Sample computations are presented for low-frequency propagation in the Arctic Ocean sound channel. These are preliminary computations and the ice layer, which averages three meters in thickness, is not included in the layered system. The effects of the ice layer on propagation are currently under investigation.

FIGURES

- Figure 1. Multilayered half-space,
- Figure 2. Contours for integration.
- Figure 3. Variation of sound velocity with depth for Model A.
Table 2 gives additional parameters for this model.
- Figure 4. Computations for Model A. Phase - and group-velocity dispersion and excitation function of pressure dependent only on layering.
- Figure 5. Computations for Model A. Variation with range of the absolute value of pressure for the normal-mode contribution of the sound field. Source depth 150 m. Hydrophone depth 50 m. Source frequency 10 Hz. Source pressure amplitude 1 dyne/cm² re 1m.
- Figure 6. Computations for Model A. Variation with range of absolute value of pressure for the integral contribution of the sound field. Source and detector same as for Figure 5.
- Figure 7. Computations for Model A. Variation with range of absolute value of pressure of the total sound field. Source and detector same as for Figure 5.
- Figure 8. Same as Figure 5 but computations carried to longer range.
- Figure 9. Same as Figure 6 but computations carried to longer range.
- Figure 10. Same as Figure 7 but computations carried to longer range.

DEFINITION OF SYMBOLS

- α_m ; compressional-wave velocity in the m-th layer
 β_m ; shear-wave velocity in the m-th layer
 h_m ; thickness of the m-th layer
 z ; vertical coordinate
 r ; range between source and detector
 t ; time
 ω ; angular frequency
 c ; phase velocity
 k ; wave number
 i ; $\sqrt{-1}$
 I_m ; imaginary part of a complex number $x + iy$
 ϕ_m ; velocity potential in the m-th layer
 $(p)_{zzm}$; normal stress in the m-th layer parallel to z axis
 p_m ; pressure in the m-th liquid layer
 u_m ; horizontal particle displacement in the m-th layer
 w_m ; vertical particle displacement in the m-th layer
 \dot{u}_m ; horizontal particle velocity in the m-th layer
 \dot{w}_m ; vertical particle velocity in the m-th layer
 ρ_m ; density in the m-th layer
 ρ_s ; density at the source
 $k_{\alpha m} = \omega / \alpha_m$
 $k_{\beta m} = \omega / \beta_m$

J_0 ; Bessel function of order 0

Y_0 ; Y Bessel function of order 0

K_0 ; K Bessel function of order 0

$H_0^{(1)}$; Hankel function of the first kind of order 0

$H_0^{(2)}$; Hankel function of the second kind of order 0

$$\beta_{\alpha m} = \sqrt{k_{\alpha m}^2 - k^2}, \quad k < k_{\alpha m}$$

$$\beta_{\alpha m} = -i\sqrt{k^2 - k_{\alpha m}^2}, \quad k > k_{\alpha m}$$

$$\beta_{\beta m} = \sqrt{k_{\beta m}^2 - k^2}, \quad k < k_{\beta m}$$

$$\beta_{\beta m} = -i\sqrt{k^2 - k_{\beta m}^2}, \quad k > k_{\beta m}$$

$$P_m = h_m \beta_{\alpha m}$$

$$Q_m = h_m \beta_{\beta m}$$

$$r_{\alpha m} = \sqrt{\left(\frac{c}{\alpha m}\right)^2 - 1}, \quad c > \alpha m$$

$$r_{\alpha m} = -i\sqrt{1 - \left(\frac{c}{\alpha m}\right)^2}, \quad c < \alpha m$$

$$r_{\beta m} = \sqrt{\left(\frac{c}{\beta m}\right)^2 - 1}, \quad c > \beta m$$

$$r_{\beta m} = -i\sqrt{1 - \left(\frac{c}{\beta m}\right)^2}, \quad c < \beta m$$

$$\bar{P}_m = k h_m r_{\alpha m}$$

$$\bar{Q}_m = k h_m r_{\beta m}$$

$$\gamma_m = 2 \left(\frac{\beta_m}{c} \right)^2$$

INTRODUCTION

This report develops the integral solution of the wave equation for point sources of harmonic waves in a liquid layer of a multilayered liquid-solid half space in a form convenient for numerical computation on a high-speed digital computer. Only the case of a liquid half-space underlying the stack of layers is considered. Furthermore, it is assumed that the speed of sound in this last layer of infinite thickness is greater than the speed of sound in liquid layers immediately above it. The problem of immediate concern is low-frequency propagation from harmonic sources in the central Arctic Ocean at ranges up to ten times the water depth in shallow and intermediate water depths. Solutions are developed for the pressure perturbations detected by a hydrophone at depth or the ice vibrations detected by a seismometer on the ice surface. A modification of the formulation is possible to include harmonic vibration sources on or in the ice, although it appears at the present time that surface sources are of limited application (Greene, 1968).

The solution of the wave equation presented here, based on the Thomson-Haskell matrix method (Thomson, 1950; Haskell, 1953), follows Harkrider (1964) for the solution of the wave equation in an n-layered solid half space. Layer matrices of the type given by Dorman (1962) for computing dispersion in an n-layered liquid - solid half-space are used for the liquid layers. An application of the theorem that the inverse of the

product matrix for the layered system above the source has the same form as the inverse of a layer matrix reduces the integrand of the integral solution to a simple form in terms of product matrices derived in the source-free, plane-wave case. The integral over wave number from zero to infinity has singularities in the integrand and is, therefore, conveniently transformed into the complex $\zeta = k + i\tau$ plane. Performing the contour integration in the ζ plane and noting that branch line integrals corresponding to branch points generated by each layer matrix except the last are zero, we are left with an expression analogous to that of Sorensen (1959) and Leslie and Sorensen (1961) for the two-layer liquid half-space with a high-speed bottom. By the proper choice of contours, complex poles are displaced to an unused sheet of the two-leaved Riemann surface, and the integral solution for the multilayered system reduces to a summation of normal modes plus the sum of two integrals over wave number, one along the real ζ axis and the other along the imaginary ζ axis. Both of these integrals are conveniently computed by a Gaussian quadrature formula.

The physical interpretation of the final solution is straightforward. The normal-mode terms correspond to waves trapped in the Arctic sound channel; that is, refracted and surface reflected (RSR) sounds and reflected sounds incident on the bottom beyond the critical angle. The integral over wave number along the real axis corresponds to waves incident

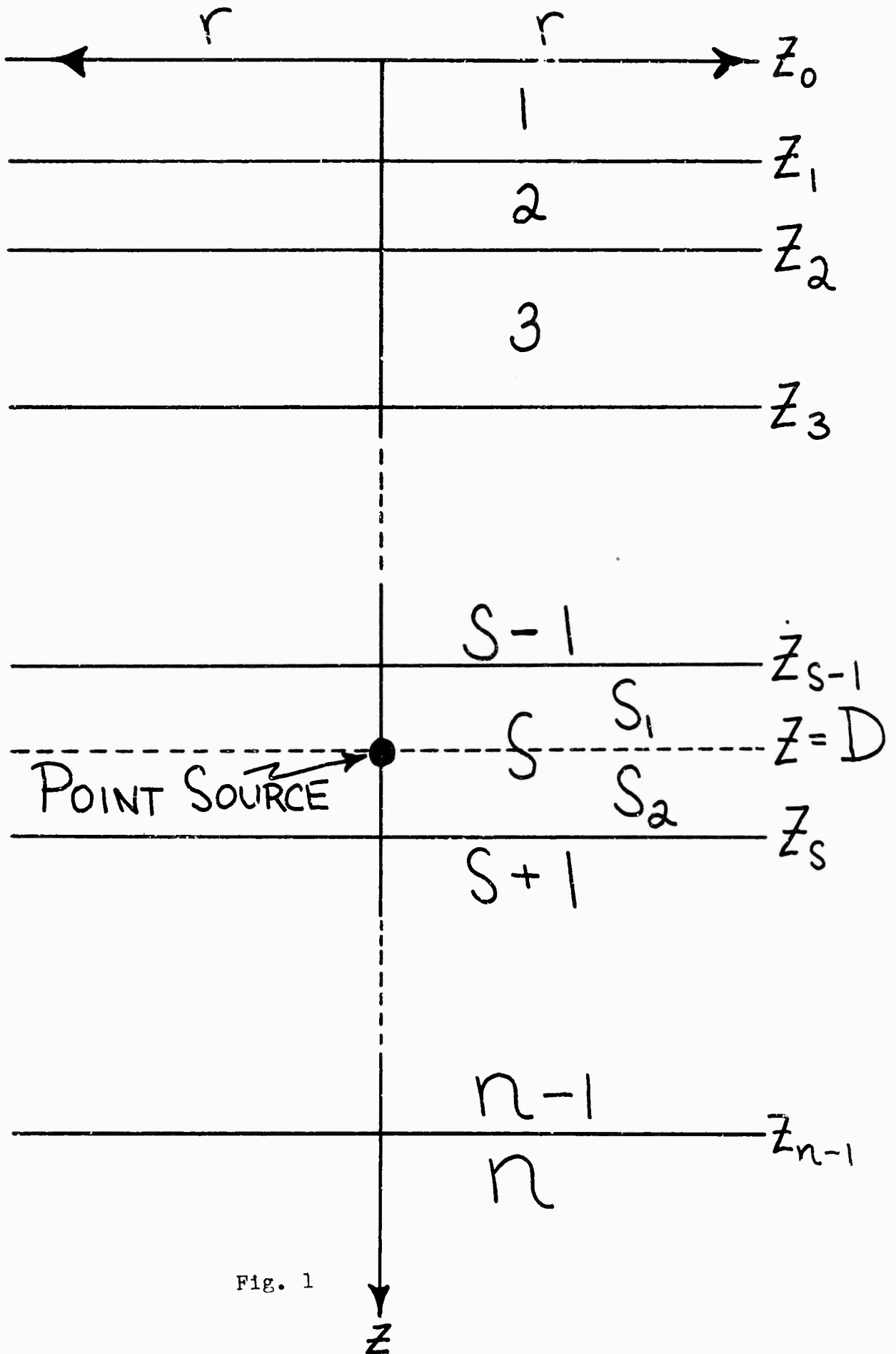


Fig. 1

on the bottom at angles greater than the critical angle of total reflection. Since energy leaves the guide continuously as waves travel down the guide, this term is of importance only to ranges of about ten times the water depth. The contribution to the sound field of the integral along the imaginary axis is of importance only very near the source since the integrand decays exponentially with range and wave number.

The usefulness of the normal mode terms for describing long-range explosive sound transmission in the central Arctic Ocean is shown by Kutschale (1969). In that work the present formulas for the normal modes were extended to explosive sources and the Fourier integral for each mode was evaluated by the principle of stationary phase (Pekeris, 1948). Attenuation by the rough ice boundaries was also included, although omitted here.

FORMAL SOLUTION

Source Free Case

Consider the n -layered interbedded liquid-solid half-space shown in Figure 1 in which the last layer of infinite thickness is liquid and has a higher sound velocity than liquid layers immediately above it; that is, a high-speed bottom. A point source of harmonic waves is located in one of the liquid layers. The velocity potential in the m -th layer, which is

assumed to be liquid, satisfies the wave equation

$$\nabla^2 \phi_m - \frac{1}{a_m^2} \ddot{\phi}_m = 0.$$

In this layer

$$(\phi_{zz})_m = -\rho_m = \rho_m \frac{\partial \phi_m}{\partial t}, \quad \dot{w}_m = \frac{\partial \phi_m}{\partial z}, \quad \dot{u}_m = \frac{\partial \phi_m}{\partial r}.$$

Solutions for ϕ_m , \dot{w}_m and $(\phi_{zz})_m$ are given by

$$\phi_m(r, z, t, k) = \int_0^\infty \hat{\phi}_m(z) J_0(kr) e^{i\omega t} dk$$

$$\dot{w}_m(r, z, t, k) = \int_0^\infty \hat{w}_m(z) J_0(kr) e^{i\omega t} dk$$

$$(\phi_{zz})_m(r, z, t, k) = \int_0^\infty (\hat{\phi}_{zz})_m(z) J_0(kr) e^{i\omega t} dk.$$

Separation of variables yields

$$\hat{\phi}_m(z) e^{i\omega t} = (A_m e^{i\sqrt{k_{dm}^2 - k^2} z} + A_m' e^{-i\sqrt{k_{dm}^2 - k^2} z}) e^{i\omega t}$$

$$\begin{aligned} \hat{w}_m(z) e^{i\omega t} &= \frac{\partial \hat{\phi}_m}{\partial z} e^{i\omega t} = (i\sqrt{k_{dm}^2 - k^2} A_m e^{i\sqrt{k_{dm}^2 - k^2} z} \\ &\quad - i\sqrt{k_{dm}^2 - k^2} A_m' e^{-i\sqrt{k_{dm}^2 - k^2} z}) e^{i\omega t} \end{aligned}$$

and

$$(\hat{p}_{zz})(z)e^{i\omega t} = \rho_m \frac{\partial}{\partial t} \hat{\phi}_m e^{i\omega t} = i\omega\rho_m (A_m e^{i\sqrt{k_{d_m}^2 - k^2}z} + A'_m e^{-i\sqrt{k_{d_m}^2 - k^2}z}) e^{i\omega t}$$

where A_m and A'_m are constants.

Placing the origin at the $(m-1)$ -st interface we get

$$\dot{\hat{w}}_{m-1} = i\sqrt{k_{d_m}^2 - k^2} (A_m - A'_m) \quad (1)$$

$$(\hat{p}_{zz})_{m-1} = i\omega\rho_m (A_m + A'_m) \quad \text{for } z = 0$$

and

$$\dot{\hat{w}}_m = -i\sqrt{k_{d_m}^2 - k^2} (A_m + A'_m) \sin P_m + i\sqrt{k_{d_m}^2 - k^2} (A_m - A'_m) \cos P_m \quad (2)$$

$$(\hat{p}_{zz})_m = i\omega\rho_m (A_m + A'_m) \cos P_m - \omega\rho_m (A_m - A'_m) \sin P_m$$

for $z = h_m$.

Substituting expressions for $[A_m - A_m']$ and $[A_m + A_m']$ from equations (1) in equations (2) yields

$$\hat{\omega}_m = \hat{\omega}_{m-1} \cos P_m - (\hat{p}_{zz})_{m-1} \frac{\sqrt{k_{dm}^2 - k^2}}{i\omega p_m} \sin P_m$$

$$(\hat{p}_{zz})_m = -\hat{\omega}_{m-1} \frac{\omega p_m}{i\sqrt{k_{dm}^2 - k^2}} \sin P_m + (\hat{p}_{zz})_{m-1} \cos P_m$$

or in matrix notation

$$\begin{bmatrix} \hat{\omega}_m \\ (\hat{p}_{zz})_m \end{bmatrix} = \begin{bmatrix} \cos P_m & \frac{i\sqrt{k_{dm}^2 - k^2}}{\omega p_m} \sin P_m \\ \frac{i\omega p_m \sin P_m}{\sqrt{k_{dm}^2 - k^2}} & \cos P_m \end{bmatrix} \cdot \begin{bmatrix} \hat{\omega}_{m-1} \\ (\hat{p}_{zz})_{m-1} \end{bmatrix} =$$

$$\begin{bmatrix} (a_m)_{11} & (a_m)_{12} \\ (a_m)_{21} & (a_m)_{22} \end{bmatrix} \begin{bmatrix} \hat{\omega}_{m-1} \\ (\hat{p}_{zz})_{m-1} \end{bmatrix}$$

In general for an n -layered liquid half-space

$$\begin{bmatrix} \hat{w}_{n-1} \\ (\hat{p}_{zz})_{n-1} \end{bmatrix} = \begin{bmatrix} A_{11} & A_{12} \\ A_{21} & A_{22} \end{bmatrix} \begin{bmatrix} \hat{w}_0 \\ (\hat{p}_{zz})_0 \end{bmatrix} \quad \text{or}$$

$$\hat{w}_{n-1} = A_{11} \hat{w}_0 + A_{12} (\hat{p}_{zz})_0$$

$$(\hat{p}_{n-1})_{zz} = A_{21} \hat{w}_0 + A_{22} (\hat{p}_{zz})_0 \quad \text{where}$$

$$A = \begin{bmatrix} A_{11} & A_{12} \\ A_{21} & A_{22} \end{bmatrix} = a_{n-1} a_{n-2} \cdots a_1$$

$$= \begin{bmatrix} (a_{n-1})_{11} & (a_{n-1})_{12} \\ (a_{n-1})_{21} & (a_{n-1})_{22} \end{bmatrix} \cdots \begin{bmatrix} (a_1)_{11} & (a_1)_{12} \\ (a_1)_{21} & (a_1)_{22} \end{bmatrix}$$

since the vertical particle velocity and pressure are continuous across each interface. For the n -th layer

$$A_n - A_n' = \frac{\hat{w}_{n-1}}{i\sqrt{k_{d_n}^2 - k^2}} = \frac{A_{11} \hat{w}_0}{i\sqrt{k_{d_n}^2 - k^2}} + \frac{A_{12} (\hat{p}_{zz})_0}{i\sqrt{k_{d_n}^2 - k^2}}$$

$$A_n + A'_n = \frac{(\hat{p}_{zz})_{n-1}}{i\omega\rho_n} = \frac{A_{21} \hat{w}_0}{i\omega\rho_n} + \frac{A_{22} (\hat{p}_{zz})_0}{i\omega\rho_n}$$

14.

or in matrix notation

$$\begin{bmatrix} A_n - A'_n \\ A_n + A'_n \end{bmatrix} = E^{-1} \begin{bmatrix} \hat{w}_{n-1} \\ (\hat{p}_{zz})_{n-1} \end{bmatrix}, E^{-1} = \begin{bmatrix} \frac{1}{i(k_{\alpha n}^2 - k^2)} & 0 \\ 0 & \frac{1}{i\omega\rho_n} \end{bmatrix} \quad (3)$$

Applying boundary conditions that the pressure vanishes at the surface and that no upgoing waves travel from infinity, $(p_{zz})_0 = 0$, $A_n = 0$, the relation (3) reduces to

$$\begin{bmatrix} -A'_n \\ A'_n \end{bmatrix} = E^{-1} A \begin{bmatrix} \hat{w}_0 \\ 0 \end{bmatrix}$$

Consider now a solid layer between two liquid layers or at the surface of the laminated halfspace. For this layer we may write (Thomson, 1950; Dorman, 1962)

$$\begin{bmatrix} (a_m)_{11} & (a_m)_{12} & (a_m)_{13} & (a_m)_{14} \\ (a_m)_{21} & (a_m)_{22} & (a_m)_{23} & (a_m)_{24} \\ (a_m)_{31} & (a_m)_{32} & (a_m)_{33} & (a_m)_{34} \\ (a_m)_{41} & (a_m)_{42} & (a_m)_{43} & (a_m)_{44} \end{bmatrix} \begin{bmatrix} \hat{u}_{m-1} \\ \hat{w}_{m-1} \\ (\hat{p}_{zz})_{m-1} \\ 0 \end{bmatrix} = \begin{bmatrix} \hat{u}_m \\ \hat{w}_m \\ (\hat{p}_{zz})_m \\ 0 \end{bmatrix} \quad (4)$$

where the fourth element of the column vector is the tangential stress which is zero at a solid-liquid boundary. The matrix elements are given in the appendix and may be derived following Haskell (1953). Equation (4) yields the three equations.

$$\dot{\hat{u}}_{m-1} = -\frac{(a_m)_{42}}{(a_m)_{41}} \dot{\hat{w}}_{m-1} - \frac{(a_m)_{43}}{(a_m)_{41}} (\hat{p}_{zz})_{m-1} \quad (5)$$

$$\dot{\hat{w}}_m = (a_m)_{21} \dot{\hat{u}}_{m-1} + (a_m)_{22} \dot{\hat{w}}_{m-1} + (a_m)_{23} (\hat{p}_{zz})_{m-1} \quad (6)$$

$$(\hat{p}_{zz})_m = (a_m)_{31} \dot{\hat{u}}_{m-1} + (a_m)_{32} \dot{\hat{w}}_{m-1} + (a_m)_{33} (\hat{p}_{zz})_{m-1} \quad (7)$$

Substituting $\dot{\hat{u}}_{m-1}$ from equation (5) in equations (6) and (7) yields

$$\dot{\hat{w}}_m = \left[(a_m)_{22} - \frac{(a_m)_{21}(a_m)_{42}}{(a_m)_{41}} \right] \dot{\hat{w}}_{m-1} + \left[(a_m)_{23} - \frac{(a_m)_{21}(a_m)_{43}}{(a_m)_{41}} \right] (\hat{p}_{zz})_{m-1}$$

$$\begin{pmatrix} \hat{p}_{zz} \end{pmatrix}_m = \left[(a_m)_{32} - \frac{(a_m)_{31}(a_m)_{42}}{(a_m)_{41}} \right] \dot{\hat{w}}_{m-1} +$$

16.

$$\left[(a_m)_{33} - \frac{(a_m)_{31}(a_m)_{43}}{(a_m)_{41}} \right] \begin{pmatrix} \hat{p}_{zz} \end{pmatrix}_{m-1}$$

or in matrix notation

$$\begin{bmatrix} \dot{\hat{w}}_m \\ \begin{pmatrix} \hat{p}_{zz} \end{pmatrix}_m \end{bmatrix} = \begin{bmatrix} (a_m)_{22} - \frac{(a_m)_{21}(a_m)_{42}}{(a_m)_{41}} & (a_m)_{23} - \frac{(a_m)_{21}(a_m)_{43}}{(a_m)_{41}} \\ (a_m)_{32} - \frac{(a_m)_{31}(a_m)_{42}}{(a_m)_{41}} & (a_m)_{33} - \frac{(a_m)_{31}(a_m)_{43}}{(a_m)_{41}} \end{bmatrix}$$

$$\begin{bmatrix} \dot{\hat{w}}_{m-1} \\ \begin{pmatrix} \hat{p}_{zz} \end{pmatrix}_{m-1} \end{bmatrix} = \begin{bmatrix} (b_m)_{11} & (b_m)_{12} \\ (b_m)_{21} & (b_m)_{22} \end{bmatrix} \begin{bmatrix} \dot{\hat{w}}_{m-1} \\ \begin{pmatrix} \hat{p}_{zz} \end{pmatrix}_{m-1} \end{bmatrix}$$

which is the same as the matrix relation for a liquid layer.

In general, then, for interbedded solids and liquids

$$\begin{bmatrix} \dot{\hat{w}}_{n-1} \\ \begin{pmatrix} \hat{p}_{zz} \end{pmatrix}_{n-1} \end{bmatrix} = a_{n-1} \cdots b_{n-3} a_{n-4} \cdots b_0 \begin{bmatrix} \dot{\hat{w}}_0 \\ 0 \end{bmatrix}$$

$$= A \begin{bmatrix} \dot{\hat{w}}_0 \\ 0 \end{bmatrix}$$

Point Source of Harmonic Waves in a Liquid Layer.

Divide the source layer into two layers as shown in Figure 1. At $z=D$ the pressure is continuous. The vertical particle velocity is continuous everywhere in the plane defined by $z=D$ except at the point source where the liquid above and below the source moves in opposite directions. This may be expressed by writing

$$\delta(\hat{w})_s = 2k.$$

In matrix notation for the source layer

$$\begin{bmatrix} \hat{w}_{s_2}(D) \\ (\hat{p}_{zz})_{s_2}(D) \end{bmatrix} = \begin{bmatrix} \hat{w}_{s_1}(D) \\ (\hat{p}_{zz})_{s_1}(D) \end{bmatrix} + \begin{bmatrix} \delta(\hat{w})_s \\ 0 \end{bmatrix}. \quad (8)$$

For the layers below $z=D$ we have

$$\begin{bmatrix} \hat{w}_{n-1} \\ (\hat{p}_{zz})_{n-1} \end{bmatrix} = A_{s_2} \begin{bmatrix} \hat{w}_{s_2}(D) \\ (\hat{p}_{zz})_{s_2}(D) \end{bmatrix} \quad (9)$$

and for the layers above $z=D$

$$\begin{bmatrix} \hat{w}_{s_1}(D) \\ (\hat{p}_{zz})_{s_1} \end{bmatrix} = A_{s_1} \begin{bmatrix} \hat{w}_0 \\ 0 \end{bmatrix} \quad (10)$$

where $A_{s_2} = a_{n-1} a_{n-2} \cdots a_{s_2}$ and $A_{s_1} = a_{s_1} a_{s_1-1} \cdots a_1$.

Returning now to the relation for the n-th liquid layer

we write from equations (8), (9), and (10)

$$\begin{aligned}
 \begin{bmatrix} -A'_n \\ A'_n \end{bmatrix} &= E^{-1} \begin{bmatrix} \hat{w}_{n-1} \\ (\hat{p}_{zz})_{n-1} \end{bmatrix} = E^{-1} A_{s_2} \begin{bmatrix} \hat{w}_{s_2}(D) \\ (\hat{p}_{zz})_{s_2}(D) \end{bmatrix} \\
 &= E^{-1} A_{s_2} \left\{ \begin{bmatrix} \hat{w}_{s_1}(D) \\ (\hat{p}_{zz})_{s_1}(D) \end{bmatrix} + \begin{bmatrix} \delta(\hat{w})_s \\ 0 \end{bmatrix} \right\} \\
 &= E^{-1} A_{s_2} \left\{ A_{s_1} \begin{bmatrix} \hat{w}_0 \\ 0 \end{bmatrix} + \begin{bmatrix} \delta(\hat{w})_s \\ 0 \end{bmatrix} \right\} \\
 &= E^{-1} \left\{ A_{s_2} A_{s_1} \begin{bmatrix} \hat{w}_0 \\ 0 \end{bmatrix} + A_{s_2} \begin{bmatrix} \delta(\hat{w})_s \\ 0 \end{bmatrix} \right\}. \quad (11)
 \end{aligned}$$

In terms of the inverse matrix $A_{s_1}^{-1}$ of A_{s_1} (11) may be written

$$\begin{bmatrix} -A'_n \\ A'_n \end{bmatrix} = E^{-1} A \left\{ \begin{bmatrix} \hat{w}_0 \\ 0 \end{bmatrix} + A_{s_1}^{-1} \begin{bmatrix} \delta(\hat{w})_s \\ 0 \end{bmatrix} \right\} \quad (12)$$

where $A = A_{S_2} A_{S_1} = a_{n-1} a_{n-2} \cdots b_{n-10} \cdots a_0$

Let

$$\begin{bmatrix} X \\ Y \end{bmatrix} = \begin{bmatrix} \hat{w}_0 \\ 0 \end{bmatrix} + A_{S_1}^{-1} \begin{bmatrix} \delta(\hat{w})_s \\ 0 \end{bmatrix}.$$

Harkrider (1964) has shown that for layered solids the inverse of the product matrix has the same form as the inverse of the layer matrices. Employing an extension of this theorem to interbedded liquids and solids we get

$$A_{S_1}^{-1} = \begin{bmatrix} (A_{S_1})_{22} & -(A_{S_1})_{12} \\ -(A_{S_1})_{21} & (A_{S_1})_{11} \end{bmatrix}.$$

Therefore

$$\begin{aligned} \begin{bmatrix} X \\ Y \end{bmatrix} &= \begin{bmatrix} \hat{w}_0 \\ 0 \end{bmatrix} + \begin{bmatrix} (A_{S_1})_{22} & -(A_{S_1})_{12} \\ -(A_{S_1})_{21} & (A_{S_1})_{11} \end{bmatrix} \begin{bmatrix} \delta(\hat{w})_s \\ 0 \end{bmatrix} \\ &= \begin{bmatrix} \hat{w}_0 \\ 0 \end{bmatrix} + \begin{bmatrix} (A_{S_1})_{22} \delta(\hat{w})_s \\ -(A_{S_1})_{21} \delta(\hat{w})_s \end{bmatrix} \quad \text{or} \end{aligned}$$

$$X = \dot{\hat{w}}_0 + (A_{s_1})_{22} \delta(\dot{\hat{w}})_s$$

$$Y = -(A_{s_1})_{21} \delta(\dot{\hat{w}})_s. \quad (13)$$

We now define $J = E^{-1}A = \begin{bmatrix} J_{11} & J_{12} \\ J_{21} & J_{22} \end{bmatrix}$.

Hence (12) may be written as

$$\begin{bmatrix} -A'_n \\ A'_n \end{bmatrix} = J \begin{bmatrix} X \\ Y \end{bmatrix} = \begin{bmatrix} J_{11}X + J_{12}Y \\ J_{21}X + J_{22}Y \end{bmatrix} \quad \text{or}$$

$$-A'_n = J_{11}X + J_{12}Y$$

$$A'_n = J_{21}X + J_{22}Y. \quad (14)$$

Adding equations (14) we get

$$0 = (J_{11} + J_{21})X + (J_{12} + J_{22})Y \quad \text{and if we}$$

let $T = J_{11} + J_{21}$ and $V = J_{12} + J_{22}$ and solve for X we have

$$X = -\frac{VY}{T}. \quad (15)$$

Carrying out the matrix multiplication $J = E^{-1} A$, expressions for T and V are given by

$$T = \frac{A_{11}}{i\sqrt{k_{d_n}^2 - k^2}} + \frac{A_{21}}{i\omega\rho_n} \quad (16)$$

$$V = \frac{A_{12}}{i\sqrt{k_{d_n}^2 - k^2}} + \frac{A_{22}}{i\omega\rho_n}$$

From the first equation of (13)

$$\dot{\hat{w}}_0 = X - (A_{s_1})_{22} \delta(\dot{\hat{w}})_s \quad \text{and}$$

from equation (15)

$$\dot{\hat{w}}_0 = -\frac{VY}{T} - (A_{s_1})_{22} \delta(\dot{\hat{w}})_s$$

Hence, from the second equation of (13) and equations (16)

$$\dot{\hat{w}}_0 = \frac{\left(\frac{A_{12}}{i\sqrt{k_{d_n}^2 - k^2}} + \frac{A_{22}}{i\omega\rho_n} \right) (A_{s_1})_{21} \delta(\dot{\hat{w}})_s}{\frac{A_{11}}{i\sqrt{k_{d_n}^2 - k^2}} + \frac{A_{21}}{i\omega\rho_n}}$$

$$- (A_{s_1})_{22} \delta(\dot{\hat{w}})_s$$

Therefore the integral solution for the surface verticle particle

velocity is $\dot{w}_0(r, z, k, t) =$

$$e^{i\omega t} \left\{ \frac{\left[\frac{A_{12}}{i\sqrt{k_{dn}^2 - k^2}} + \frac{A_{22}}{i\omega\rho_n} \right] (A_{s_1})_{21} - \left[\frac{A_{11}}{i\sqrt{k_{dn}^2 - k^2}} + \frac{A_{21}}{i\omega\rho_n} \right] (A_{s_1})_{22}}{\frac{A_{11}}{i\sqrt{k_{dn}^2 - k^2}} + \frac{A_{21}}{i\omega\rho_n}} \right\} 2J_0(kr)kdk$$

and from the formula

$$\hat{p}_{zzD_1} = (A_{D_1})_{21} \hat{w}_0 = -p_{D_1} \quad \text{the integral solution for pressure}$$

at depth D_1 is $p_{D_1}(r, z, k, t) =$

$$- e^{i\omega t} \left\{ \frac{\left[\frac{A_{12}}{i\sqrt{k_{dn}^2 - k^2}} + \frac{A_{22}}{i\omega\rho_n} \right] (A_{s_1})_{21}}{\frac{A_{11}}{i\sqrt{k_{dn}^2 - k^2}} + \frac{A_{21}}{i\omega\rho_n}} \right\}$$

$$\frac{\left[\frac{A_{11}}{i\sqrt{k_{dn}^2 - k^2}} + \frac{A_{21}}{i\omega\rho_n} \right] (A_{s_1})_{22}}{\frac{A_{11}}{i\sqrt{k_{dn}^2 - k^2}} + \frac{A_{21}}{i\omega\rho_n}} \left\} (A_{D_1})_{21} 2J_0(kr)kdk$$

where

$$A_{D_1} = a_{D_1} a_{D_1-1} \dots a_0$$

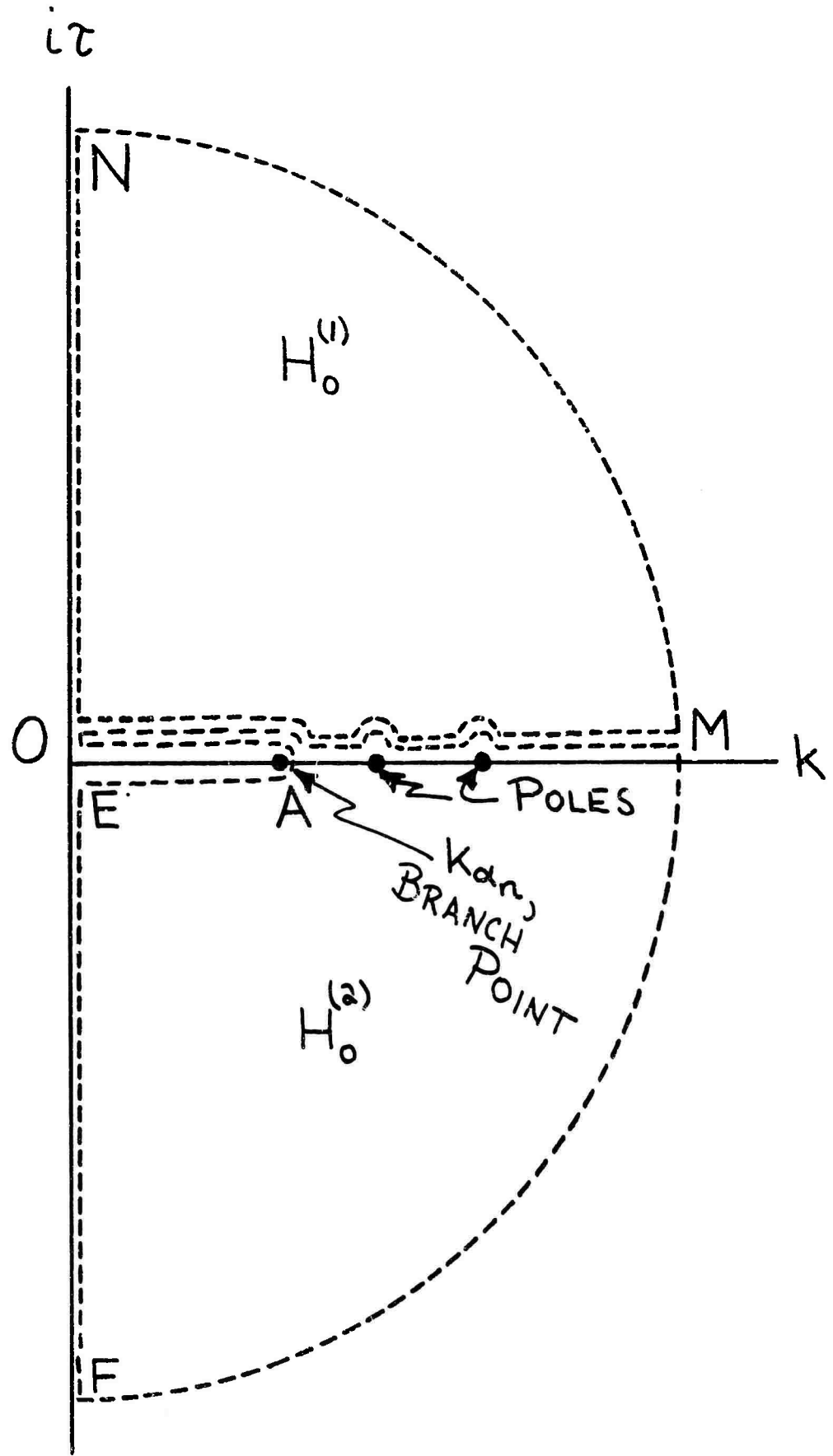


Fig. 2

Evaluation of the Integral Solutions.

The integral solutions have singularities corresponding to zeros of

$$\frac{A_{11}}{i\sqrt{k_{\alpha n}^2 - k^2}} + \frac{A_{21}}{i\omega\rho n}$$

These are simple poles for complex wave number and, therefore, the integral solutions are conveniently evaluated by contour integration in the complex $\beta = k + i\zeta$ plane. The contours are shown in Figure 2. First the transformation

$$2J_0(\beta r) = H_0^{(1)}(\beta r) + H_0^{(2)}(\beta r) \quad \text{is made.}$$

In addition to simple poles at

$$\frac{A_{11}}{i\sqrt{k_{\alpha n}^2 - \beta^2}} + \frac{A_{21}}{i\omega\rho n}$$

branch points occur at

$$\sqrt{k_{\alpha n}^2 - k^2} = 0$$

for the liquid layers and

$$\sqrt{k_{\alpha m}^2 - k^2} = 0 \quad \text{and} \quad \sqrt{k_{\beta m}^2 - k^2} = 0$$

for each solid layer. But since the integrands are even

functions of

$$\sqrt{k_{\alpha m}^2 - k^2} \quad \text{and} \quad \sqrt{k_{\beta m}^2 - k^2}$$

branch line integrals corresponding to branch cuts made to these branch points cancel and only the integral corresponding to the branch cut for $\sqrt{k_{\alpha n}^2 - k^2} = 0$ must be considered.

The Riemann surface has two sheets. To satisfy the vanishing of \dot{w}_0 as ξ approaches infinity, we must remain on the sheet of the Riemann surface where the real part of $i\beta_n$ is greater than zero. For the contours shown in Figure 2 complex poles of

$$\frac{A_{11}}{i\sqrt{k_{dn}^2 - k^2}} + \frac{A_{21}}{i\omega\beta_n} = 0$$

have been displaced to an unused Riemann sheet and all poles are on the real k -axis. See Ewing, Jardetzky and Press (1957, pages 135 to 137) for the proof for a two-layer

liquid half space. Along the branch cut $i\sqrt{k_{dn}^2 - \xi^2}$

is pure imaginary and along the real axis from k_{dn} to

infinity $i(-i\sqrt{\xi^2 - k_{dn}^2})$ is real.

In the first quadrant $\text{Im}[i\sqrt{k_{dn}^2 - \xi^2}]$ is positive and

in the fourth quadrant negative. The integral for \dot{w}_0 (omitting the $e^{i\omega t}$ term) is now transformed to

$$\int_0^M I_1 d\xi + \int_{C(M,N)} I_1 d\xi + \int_N^0 I_1 d\xi = 0 \quad (17)$$

and

$$\int_F^E I_2 d\xi + \int_{C(F,M)} I_2 d\xi + \int_M^0 I_2 d\xi + \int_{OAE} I_2 d\xi =$$

$$2\pi i \sum \text{RES}(I_2) \quad (18)$$

where

$$I_1 = \left\{ \frac{\left[\frac{A_{12}}{i\sqrt{k_{dn}^2 - f^2}} + \frac{A_{22}}{i\omega\rho_n} \right] (A_{s1})_{21} - \left[\frac{A_{11}}{i\sqrt{k_{dn}^2 - f^2}} + \frac{A_{21}}{i\omega\rho_n} \right] (A_{s1})_{22}}{\frac{A_{11}}{i\sqrt{k_{dn}^2 - f^2}} + \frac{A_{21}}{i\omega\rho_n}} \right\}$$

$H_0^{(1)}(fr) f$

$$I_2 = \left\{ \frac{\left[\frac{A_{12}}{i\sqrt{k_{dn}^2 - f^2}} + \frac{A_{22}}{i\omega\rho_n} \right] (A_{s1})_{21} - \left[\frac{A_{11}}{i\sqrt{k_{dn}^2 - f^2}} + \frac{A_{21}}{i\omega\rho_n} \right] (A_{s1})_{22}}{\frac{A_{11}}{i\sqrt{k_{dn}^2 - f^2}} + \frac{A_{21}}{i\omega\rho_n}} \right\}$$

$H_0^{(2)}(fr) f$

Since $H_0^{(1)}(fr)$ approaches zero as f approaches infinity in the first quadrant, (17) is

$$\int_0^\infty I_1 df + \int_{i\infty}^0 I_1 df = 0 \quad \text{or} \quad \int_0^\infty I_1 df = - \int_{i\infty}^0 I_1 df \quad (19)$$

Likewise, since $H_0^{(2)}(fr)$ approaches zero as f approaches infinity in the fourth quadrant (18) is

$$\int_0^{-i\infty} I_2 df + \int_\infty^0 I_2 df + \int_{\text{OAE}} I_2 df = 2\pi i \sum \text{RES}(I_2) \quad \text{or}$$

$$\int_0^{\infty} I_2 dt = \int_0^{-i\infty} I_2 dt + \int_{AOE} I_2 dt = \quad 27.$$

$$- 2\pi i \sum \text{RES}(I_2). \quad (20)$$

Using the identity $H_0^{(1)}(zr) = -H_0^{(2)}(-zr)$ (19) may be written

$$\text{as } \int_0^{\infty} I_1 dk = \int_0^{\infty} F(i\sqrt{k_{\alpha_n}^2 + z^2}) H_0^{(2)}(-izr) z dz$$

and (20) may be written as

$$\int_0^{\infty} I_2 dk = - \int_0^{\infty} F(-i\sqrt{k_{\alpha_n}^2 + z^2}) H_0^{(2)}(-izr) z dz$$

$$+ \int_0^{k_{\alpha_n}} [F(i\sqrt{k_{\alpha_n}^2 - k^2}) - F(-i\sqrt{k_{\alpha_n}^2 - k^2})] H_0^{(2)}(kr) k dk$$

$$- 2\pi i \sum \text{RES}(I_2).$$

$$\text{Hence } \dot{w}_0 = \int_0^{\infty} I_1 dk + \int_0^{\infty} I_2 dk =$$

$$\int_0^{\infty} [F(i\sqrt{k_{\alpha_n}^2 + z^2}) - F(-i\sqrt{k_{\alpha_n}^2 + z^2})] H_0^{(2)}(-izr) z dz$$

$$+ \int_0^{k_{\alpha_n}} [F(i\sqrt{k_{\alpha_n}^2 - k^2}) - F(-i\sqrt{k_{\alpha_n}^2 - k^2})] H_0^{(2)}(kr) k dk$$

$$- 2\pi i \sum \text{RES}(I_2).$$

In the first integral

$$F(i\sqrt{k_{\alpha n}^2 + z^2}) - F(-i\sqrt{k_{\alpha n}^2 + z^2}) = \frac{2(A_{11}A_{22} - A_{12}A_{21})(A_{S_1})_{21}}{\omega \rho_n \sqrt{k_{\alpha n}^2 + z^2} \left(\frac{A_{11}^2}{k_{\alpha n}^2 + z^2} - \frac{A_{21}^2}{\omega^2 \rho_n^2} \right)}$$

but $A_{11}A_{22} - A_{12}A_{21} = 1$

and therefore

$$F(i\sqrt{k_{\alpha n}^2 + z^2}) - F(-i\sqrt{k_{\alpha n}^2 + z^2}) = \frac{2\omega \rho_n \sqrt{k_{\alpha n}^2 + z^2} (A_{S_1})_{21}}{\omega^2 \rho_n^2 A_{11}^2 - (k_{\alpha n}^2 + z^2) A_{21}^2}$$

Likewise, in the second integral

$$F(i\sqrt{k_{\alpha n}^2 - k^2}) - F(-i\sqrt{k_{\alpha n}^2 - k^2}) = \frac{2\omega \rho_n \sqrt{k_{\alpha n}^2 - k^2} (A_{S_1})_{21}}{\omega^2 \rho_n^2 A_{11}^2 - (k_{\alpha n}^2 - k^2) A_{21}^2}$$

Therefore, the integral solution reduces to

$$\dot{w}_0 = \int_0^{\infty} \frac{2\omega \rho_n \sqrt{k_{\alpha n}^2 + z^2} (A_{S_1})_{21} H_0^{(2)}(-izr) z dz}{\omega^2 \rho_n^2 A_{11}^2 - (k_{\alpha n}^2 + z^2) A_{21}^2} + \int_0^{k_{\alpha n}} \frac{2\omega \rho_n \sqrt{k_{\alpha n}^2 - k^2} (A_{S_1})_{21} H_0^{(2)}(kr) k dk}{\omega^2 \rho_n^2 A_{11}^2 - (k_{\alpha n}^2 - k^2) A_{21}^2}$$

$$- 2\pi i \int \text{RES}(I_2)$$

From the formulas $H_0^{(2)}(-izr) = \frac{2i}{\pi} K_0(zr)$ and

$$H_0^{(2)}(kr) = J_0(kr) - iY_0(kr)$$

and multiplying \dot{w}_0 by $\frac{P_0}{i\omega\rho_s}$ for a constant pressure

source of pressure amplitude P_0 at unit distance from the source

$$\dot{w}_0 = \left\{ \int_0^\infty \frac{4P_0\omega\rho_n\sqrt{k_{dn}^2+z^2}(A_{s1})_{21}K_0(zr)zdz}{\pi\omega\rho_s[\omega^2\rho_n^2A_{11}^2 - (k_{dn}^2+z^2)A_{21}^2]} \right.$$

$$+ \left. \int_0^{k_{dn}} \frac{2P_0\omega\rho_n\sqrt{k_{dn}^2-k^2}(A_{s1})_{21}[J_0(kr)-iY_0(kr)]kdk}{i\omega\rho_s[\omega^2\rho_n^2A_{11}^2 - (k_{dn}^2-k^2)A_{21}^2]} \right.$$

$$- 2\pi i \left[\text{RES} \left(\frac{P_0 I_2}{i\omega\rho_s} \right) \right].$$

For the normal mode contribution, $-2\pi i \left[\text{RES} \left(\frac{P_0 I_2}{i\omega\rho_s} \right) \right]$, it is

convenient to rewrite the integral for w_0 in the form

$$\dot{w}_0 = \left\{ \frac{\int_0^\infty \left[\frac{c\bar{A}_{12}}{ikr_{dn}} + \frac{\bar{A}_{22}}{i\omega\rho_n} \right] (A_{s1})_{21} - \left[\frac{c\bar{A}_{11}}{ikr_{dn}} + \frac{\bar{A}_{21}}{i\omega\rho_n} \right] (A_{s1})_{22}}{\left[\frac{c\bar{A}_{11}}{ikr_{dn}} + \frac{\bar{A}_{21}}{i\omega\rho_n} \right]} \right\}$$

$$2 J_0(kr) k dk$$

where the layer matrices \bar{a}_m are defined in the Appendix.

There are, as before, simple poles of I_2 along the real axis at

$$\frac{c\bar{A}_{11}}{ikr_{\alpha n}} + \frac{\bar{A}_{21}}{i\omega\rho_n} = 0 \quad \text{for } \omega = \omega_\ell, k = k_\ell, \ell = 1, 2, \dots$$

This is the period equation for the laminated half space.

From the period equation we write

$$\left(\frac{i\omega\rho_n c}{ikr_{\alpha n}} \right)_{\substack{\omega = \omega_\ell \\ k = k_\ell}} = \left(\frac{-\bar{A}_{21}}{\bar{A}_{11}} \right)_{\substack{\omega = \omega_\ell \\ k = k_\ell}}$$

and hence

$$-2\pi i \int \text{RES} \left(\frac{P_0 I_2}{i\omega\rho_s} \right) =$$

$$-2\pi i \int \left\{ \frac{P_0 \left[\bar{A}_{22} - \frac{\bar{A}_{21}\bar{A}_{12}}{\bar{A}_{11}} \right] (\bar{A}_{s_1})_{21} H_0^{(2)}(kr) k}{i\omega\rho_s \frac{\partial}{\partial k} \left[\frac{c\bar{A}_{11}}{ikr_{\alpha n}} + \frac{\bar{A}_{21}}{i\omega\rho_n} \right]} \right\} \left[\frac{1}{i\omega\rho_n} \right]_{\substack{\omega = \omega_\ell \\ k = k_\ell}}$$

But again $\bar{A}_{11}\bar{A}_{22} - \bar{A}_{21}\bar{A}_{12} = 1$, and therefore

$$-2\pi i \int \text{RES} \left(\frac{P_0 I_2}{i\omega\rho_s} \right) =$$

$$2\pi i \int \left\{ \frac{P_0 \omega (\bar{A}_{s_1})_{21} [J_0(kr) - iY_0(kr)]}{c^3 \rho_s \rho_n \bar{A}_{11} \frac{\partial}{\partial c} \left[\frac{c^2 \bar{A}_{11}}{r_{\alpha n}} + \frac{\bar{A}_{21}}{\rho_n} \right]} \right\}_{\substack{\omega = \omega_\ell \\ c = c_\ell}}$$

In summary, the expressions for vertical particle velocity at the surface, vertical particle displacement at the surface, and pressure in a liquid layer are:

$$\dot{w}_0 = \int_0^{\infty} \frac{4P_0 \rho_n \sqrt{k_{dn}^2 + z^2} (A_{s1})_{21} K_0(zr) z dz}{\pi \rho_s [\omega^2 \rho_n^2 A_{11}^2 - (k_{dn}^2 + z^2) A_{21}^2]} +$$

$$\int_0^{k_{dn}} \frac{2P_0 \rho_n \sqrt{k_{dn}^2 - k^2} (A_{s1})_{21} [J_0(kr) - iY_0(kr)] k dk}{i \rho_s [\omega^2 \rho_n^2 A_{11}^2 - (k_{dn}^2 - k^2) A_{21}^2]}$$

$$+ 2\pi \int_l \left\{ \frac{P_0 \omega (\bar{A}_{s1})_{21} [J_0(kr) - iY_0(kr)]}{c^3 \rho_s \rho_n \bar{A}_{11} \frac{\partial}{\partial c} \left[\frac{c^2 \bar{A}_{11}}{r_{dn}} + \frac{\bar{A}_{21}}{\rho_n} \right]} \right\}_{w=w_l}$$

$c = c_l$

$$w_0 = \frac{\dot{w}_0}{i\omega} = \int_0^{\infty} \frac{4P_0 \rho_n \sqrt{k_{dn}^2 + z^2} (A_{s1})_{21} K_0(zr) z dz}{i \pi \rho_s \omega [\omega^2 \rho_n^2 A_{11}^2 - (k_{dn}^2 + z^2) A_{21}^2]}$$

$$- \int_0^{k_{dn}} \frac{2P_0 \rho_n \sqrt{k_{dn}^2 - k^2} (A_{s1})_{21} [J_0(kr) - iY_0(kr)] k dk}{\omega \rho_s [\omega^2 \rho_n^2 A_{11}^2 - (k_{dn}^2 - k^2) A_{21}^2]}$$

$$+ 2\pi \int \left\{ \frac{P_0 \omega (\bar{A}_{S_1})_{21} [J_0(kr) - iY_0(kr)]}{i\omega c^3 \rho_s \rho_n \bar{A}_{11} \frac{\partial}{\partial c} \left[\frac{c^2 \bar{A}_{11}}{\Gamma_{\alpha_n}} + \frac{\bar{A}_{21}}{\rho_n} \right]} \right\}_{\substack{\omega = \omega_0 \\ c = c_0}} \quad 32.$$

$$b_{D_1} = - \int_0^\infty \frac{4P_0 \rho_n \sqrt{k_{\alpha_n}^2 + z^2} (A_{S_1})_{21} (A_{D_1})_{21} K_0(zr) z dz}{\pi \rho_s [\omega^2 \rho_n^2 A_{11}^2 - (k_{\alpha_n}^2 + z^2) A_{21}^2]}$$

$$+ i \int_0^{k_{\alpha_n}} \frac{2P_0 \rho_n \sqrt{k_{\alpha_n}^2 - k^2} (A_{S_1})_{21} (A_{D_1})_{21} [J_0(kr) - iY_0(kr)] k dk}{\rho_s [\omega^2 \rho_n^2 A_{11}^2 - (k_{\alpha_n}^2 - k^2) A_{21}^2]}$$

$$- 2\pi \int \left\{ \frac{P_0 \omega (\bar{A}_{S_1})_{21} (\bar{A}_{D_1})_{21} [J_0(kr) - iY_0(kr)]}{c^4 \rho_s \rho_n \bar{A}_{11} \frac{\partial}{\partial c} \left[\frac{c^2 \bar{A}_{11}}{\Gamma_{\alpha_n}} + \frac{\bar{A}_{21}}{\rho_n} \right]} \right\}_{\substack{\omega = \omega_0 \\ c = c_0}}$$

where for the normal mode terms for b_{D_1} the relation $b_{D_1} = (\bar{A}_{D_1})_{21} \left(\frac{i\omega_0}{c} \right)$ has been used.

For programming it is convenient to write:

$$\dot{w}_0 = -W_1 + W_2 - iW_3 + iW_4 - iW_5 \quad \text{and for the absolute}$$

value of vertical particle velocity, a quantity conveniently measured in transmission experiments,

$$|\dot{w}_0| = \sqrt{(-W_1 + W_2)^2 + (-W_3 + W_4 - W_5)^2}$$

where

$$W_1 = \int_0^{k_{dn}} \frac{2P_0 \rho_n \sqrt{k_{dn}^2 - k^2} \operatorname{Im}[-(A_{s_1})_{21}] J_0(kr) k dk}{\rho_s \{ \omega^2 \rho_n^2 A_{11}^2 + (k_{dn}^2 - k^2) [\operatorname{Im}(-A_{21})]^2 \}}$$

$$W_2 = 2\pi \int \left\{ \frac{P_0 \omega \operatorname{Im}[-(\bar{A}_{s_1})_{21}] J_0(kr)}{c^3 \rho_s \rho_n \bar{A}_{11} \frac{\partial}{\partial c} \left[\frac{c^2 \bar{A}_{11}}{\operatorname{Im}(\Gamma_{dn})} + \frac{\operatorname{Im}(-\bar{A}_{21})}{\rho_n} \right]} \right\} \Bigg|_{\substack{w=\omega l \\ c=c l}}$$

$$W_3 = \int_0^{\infty} \frac{4P_0 \rho_n \sqrt{k_{dn}^2 + \zeta^2} \operatorname{Im}[-(A_{s_1})_{21}] K_0(\zeta r) \zeta d\zeta}{\pi \rho_s \{ \omega^2 \rho_n^2 A_{11}^2 + (k_{dn}^2 + \zeta^2) [\operatorname{Im}(-A_{21})]^2 \}}$$

$$W_4 = \int_0^{k_{dn}} \frac{2P_0 \rho_n \sqrt{k_{dn}^2 - k^2} \operatorname{Im}[-(A_{s_1})_{21}] Y_0(kr) k dk}{\rho_s \{ \omega^2 \rho_n^2 A_{11}^2 + (k_{dn}^2 - k^2) [\operatorname{Im}(-A_{21})]^2 \}}$$

$$W_5 = 2\pi \int \left\{ \frac{P_0 \omega \operatorname{Im} [-(\bar{A}_{s_1})_{21}] Y_0(kr)}{c^3 \rho_s \rho_n \bar{A}_{11} \frac{\partial}{\partial c} \left[\frac{c^2 \bar{A}_{11}}{\operatorname{Im}(\Gamma_{\alpha_n})} + \frac{\operatorname{Im}(-\bar{A}_{21})}{\rho_n} \right]} \right\}_{\substack{w=u_2 \\ c=c_0}}$$

Likewise for vertical particle displacement

$$w_0 = -W_6 + W_7 - W_8 + i W_9 - i W_{10}$$

$$|w_0| = \sqrt{(-W_6 + W_7 - W_8)^2 + (W_9 - W_{10})^2}$$

$$W_6 = \frac{\int_0^\infty 4 P_0 \rho_n \sqrt{k_{\alpha_n}^2 + c^2} \operatorname{Im} [-(A_{s_1})_{21}] K_0(zr) z dz}{\omega \pi \rho_s \omega \left\{ \omega^2 \rho_n^2 A_{11}^2 + (k_{\alpha_n}^2 + c^2) [\operatorname{Im}(-A_{21})]^2 \right\}}$$

$$W_7 = \frac{\int_0^{k_{\alpha_n}} 2 P_0 \rho_n \sqrt{k_{\alpha_n}^2 - k^2} \operatorname{Im} [-(A_{s_1})_{21}] Y_0(kr) k dk}{\omega \rho_s \left\{ \omega^2 \rho_n^2 A_{11}^2 + (k_{\alpha_n}^2 - k^2) [\operatorname{Im}(-A_{21})]^2 \right\}}$$

$$W_8 = 2\pi \int \left\{ \frac{P_0 \operatorname{Im} [-(\bar{A}_{s_1})_{21}] Y_0(kr)}{c^3 \rho_s \rho_n \bar{A}_{11} \frac{\partial}{\partial c} \left[\frac{c^2 \bar{A}_{11}}{\operatorname{Im}(\Gamma_{\alpha_n})} + \frac{\operatorname{Im}(-\bar{A}_{21})}{\rho_n} \right]} \right\}_{\substack{w=u_2 \\ c=c_0}}$$

$$W_q = \int_0^{k_{\alpha n}} \frac{2P_0 \rho_n \sqrt{k_{\alpha n}^2 - k^2} \operatorname{Im}[-(A_{s1})_{21}] J_0(kr) k dk}{\omega \rho_s \left\{ \omega^2 \rho_n^2 A_{11}^2 + (k_{\alpha n}^2 - k^2) [\operatorname{Im}(-A_{21})]^2 \right\}}$$

$$W_{10} = 2\pi \int_0^{\infty} \left\{ \frac{P_0 \operatorname{Im}[-(\bar{A}_{s1})_{21}] J_0(kr)}{c^3 \rho_s \rho_n \bar{A}_{11} \frac{\partial}{\partial c} \left[\frac{c^2 \bar{A}_{11}}{\operatorname{Im}(\Gamma_{\alpha n})} + \frac{\operatorname{Im}(-\bar{A}_{21})}{\rho_n} \right]} \right\} \frac{W_{10}}{c^2} dk$$

and for pressure

$$p_{D_1} = W_{11} - W_{12} + W_{13} - iW_{14} + iW_{15}$$

$$|p_{D_1}| = \sqrt{(W_{11} - W_{12} + W_{13})^2 + (-W_{14} + W_{15})^2}$$

$$W_{11} =$$

$$\int_0^{\infty} \frac{4P_0 \rho_n \sqrt{k_{\alpha n}^2 + \zeta^2} \operatorname{Im}[-(A_{s1})_{21}] \operatorname{Im}[-(A_{D1})_{21}] K_0(\zeta r) \zeta d\zeta}{\pi \rho_s \left\{ \omega^2 \rho_n^2 A_{11}^2 + (k_{\alpha n}^2 + \zeta^2) [\operatorname{Im}(-A_{21})]^2 \right\}}$$

$$W_{12} =$$

$$\int_0^{k_{dn}} \frac{2 P_0 \rho_n \sqrt{k_{dn}^2 - k^2} \operatorname{Im}[-(A_{S,2,1})] \operatorname{Im}[-(A_{D,2,1})] Y_0(kr) k dk}{\rho_s \left\{ \omega^2 \rho_n^2 A_{11}^2 + (k_{dn}^2 - k^2) [\operatorname{Im}(-A_{2,1})]^2 \right\}}$$

$$W_{13} =$$

$$2\pi \int_0^l \left\{ \frac{P_0 \omega \operatorname{Im}[-(\bar{A}_{S,2,1})] \operatorname{Im}[-(\bar{A}_{D,2,1})] Y_0(kr)}{c^4 \rho_s \rho_n \bar{A}_{11} \frac{\partial}{\partial c} \left[\frac{c^2 \bar{A}_{11}}{\operatorname{Im}(\Gamma_{dn})} + \frac{\operatorname{Im}(-\bar{A}_{2,1})}{\rho_n} \right]} \right\}_{\substack{\omega = \omega_2 \\ c = c_2}}$$

$$W_{14} =$$

$$\int_0^{k_{dn}} \frac{2 P_0 \rho_n \sqrt{k_{dn}^2 - k^2} \operatorname{Im}[-(A_{S,2,1})] \operatorname{Im}[-(A_{D,2,1})] J_0(kr) k dk}{\rho_s \left\{ \omega^2 \rho_n^2 A_{11}^2 + (k_{dn}^2 - k^2) [\operatorname{Im}(-A_{2,1})]^2 \right\}}$$

$$W_{15} =$$

$$2\pi \int_0^l \left\{ \frac{P_0 \omega \operatorname{Im}[-(\bar{A}_{S,2,1})] \operatorname{Im}[-(\bar{A}_{D,2,1})] J_0(kr)}{c^4 \rho_s \rho_n \bar{A}_{11} \frac{\partial}{\partial c} \left[\frac{c^2 \bar{A}_{11}}{\operatorname{Im}(\Gamma_{dn})} + \frac{\operatorname{Im}(-\bar{A}_{2,1})}{\rho_n} \right]} \right\}_{\substack{\omega = \omega_2 \\ c = c_2}}$$

The layer matrices for the integral along the real axis are programmed as

$$\begin{bmatrix} (a_m)_{11} & -\text{Im}[(a_m)_{12}] & (a_m)_{13} & -\text{Im}[(a_m)_{14}] \\ \text{Im}[(a_m)_{21}] & (a_m)_{22} & \text{Im}[(a_m)_{23}] & (a_m)_{24} \\ (a_m)_{31} & -\text{Im}[(a_m)_{32}] & (a_m)_{33} & -\text{Im}[(a_m)_{34}] \\ \text{Im}[(a_m)_{41}] & (a_m)_{42} & \text{Im}[(a_m)_{43}] & (a_m)_{44} \end{bmatrix}$$

for solid layers and

$$\begin{bmatrix} (a_m)_{11} & \text{Im}[(a_m)_{12}] \\ -\text{Im}[(a_m)_{21}] & (a_m)_{22} \end{bmatrix}$$

for liquid layers. For the integral along the imaginary axis the matrices are

$$\begin{bmatrix} (a_m)_{11} & (a_m)_{12} & -\text{Im}[(a_m)_{13}] & -\text{Im}[(a_m)_{14}] \\ (a_m)_{21} & (a_m)_{22} & -\text{Im}[(a_m)_{23}] & -\text{Im}[(a_m)_{24}] \\ \text{Im}[(a_m)_{31}] & \text{Im}[(a_m)_{32}] & (a_m)_{33} & (a_m)_{34} \\ \text{Im}[(a_m)_{41}] & \text{Im}[(a_m)_{42}] & (a_m)_{43} & (a_m)_{44} \end{bmatrix}$$

for solid layers and

$$\begin{bmatrix} (a_m)_{11} & \text{Im}[(a_m)_{12}] \\ -\text{Im}[(a_m)_{21}] & (a_m)_{22} \end{bmatrix}$$

for liquid layers. Likewise, for the normal mode terms the layer matrices are programmed as

$$\begin{bmatrix} (\bar{a}_m)_{11} & -\operatorname{Im}[(\bar{a}_m)_{12}] & (\bar{a}_m)_{13} & -\operatorname{Im}[(\bar{a}_m)_{14}] \\ \operatorname{Im}[(\bar{a}_m)_{21}] & (\bar{a}_m)_{22} & \operatorname{Im}[(\bar{a}_m)_{23}] & (\bar{a}_m)_{24} \\ (\bar{a}_m)_{31} & -\operatorname{Im}[(\bar{a}_m)_{32}] & (\bar{a}_m)_{33} & -\operatorname{Im}[(\bar{a}_m)_{34}] \\ \operatorname{Im}[(\bar{a}_m)_{41}] & (\bar{a}_m)_{42} & \operatorname{Im}[(\bar{a}_m)_{43}] & (\bar{a}_m)_{44} \end{bmatrix}$$

for solid layers and

$$\begin{bmatrix} (\bar{a}_m)_{11} & \operatorname{Im}[(\bar{a}_m)_{12}] \\ -\operatorname{Im}[(\bar{a}_m)_{21}] & (\bar{a}_m)_{22} \end{bmatrix}$$

for liquid layers.

NUMERICAL COMPUTATIONS

Computer programs were written in double precision Fortran IV to compute the pressure in dynes/cm² and the vertical particle displacement in millimicrons. The programs are run on the IBM 360/91. For numerical results presented here, computing time was under six minutes. It appears that practical limits of numerical precision make the present development most useful for the Arctic at frequencies below 50 Hz in water depths to 1 km. At higher frequencies or in greater water depths the integrands may be so oscillatory that it is difficult to achieve the desired accuracy in the numerical integrations without excessive computing time.

Computations are made in two stages. The first program, an extension of Dorman's (1962) PV 7 dispersion program, computes phase- and group-velocity dispersion, the excitation function dependent only on layering of the medium, and the excitation function for the particular source and detector depths. In the first case for the m-th normal mode the excitation function is defined by

$$\left\{ \frac{2\sqrt{2\pi} \sqrt{\omega}}{\rho_n c^{3/2} \bar{A}_{11}} \frac{\partial}{\partial c} \left[\frac{c^2 \bar{A}_{11}}{\text{Im}(\rho_n)} + \frac{\text{Im}(-\bar{A}_{21})}{\rho_n} \right] \right\}_{\substack{\omega = \omega_m \\ c = c_m}}$$

for pressure and

$$\left\{ \frac{2\sqrt{2\pi}\sqrt{\omega}}{\rho_n c^{5/2} \bar{A}_{11}} \frac{\partial}{\partial c} \left[\frac{c^2 \bar{A}_{11}}{\text{Im}(\bar{r}_{\alpha n})} + \frac{\text{Im}(-\bar{A}_{21})}{\rho_n} \right] \right\}$$

$\omega = \omega_m$
 $c = c_m$

for vertical particle velocity. In the second case the excitation function is

$$\left\{ \frac{2\sqrt{2\pi}\sqrt{\omega} \text{Im}[-(\bar{A}_{S1})_{21}] \text{Im}[-(\bar{A}_{D1})_{21}]}{\rho_n \rho_s c^{7/2} \bar{A}_{11}} \frac{\partial}{\partial c} \left[\frac{c^2 \bar{A}_{11}}{\text{Im}(\bar{r}_{\alpha n})} + \frac{\text{Im}(-\bar{A}_{21})}{\rho_n} \right] \right\}$$

$\omega = \omega_m$
 $c = c_m$

for pressure and

$$\left\{ \frac{2\sqrt{2\pi}\sqrt{\omega} \text{Im}[-(\bar{A}_{S1})_{21}]}{\rho_n \rho_s c^{5/2} \bar{A}_{11}} \frac{\partial}{\partial c} \left[\frac{c^2 \bar{A}_{11}}{\text{Im}(\bar{r}_{\alpha n})} + \frac{\text{Im}(-\bar{A}_{21})}{\rho_n} \right] \right\}$$

$\omega = \omega_m$
 $c = c_m$

for vertical particle velocity. These definitions were chosen to be useful also at long ranges where $H_0^{(2)}(kr) =$

$$\sqrt{\frac{2}{\pi kr}} e^{i(\frac{\pi}{4} - kr)}$$

The second program computes the three integrals, the normal mode contribution to the sound field, and writes and plots the absolute value of p_{D1} , or w_0 as a

TABLE 1.

Model Parameters for Solensen's (1959) Computations: Source
 Pressure Amplitude re im, .3048 dyne/cm²; Source Depth, 15.2 m;
 Hydrophone Depth, 15.2m; Range, 3.048 m

Layer Thickness, m	Compressional Velocity, m/sec	Shear Velocity, m/sec	Density gm/cm ³
18.3	1463.0	0	1.03
Infinite	1609.3	0	1.24

TABLE 2.

Integral Real Axis, Real Part

Frequency, Hz	Sorensen's (1959) Value, dynes/cm ²	Our Value, dynes/cm ²
10	.004713	.004685
20	.030170	.030153
40	.040278	.040485
80	.023193	.023121
160	- .021839	- .022201
320	- .074395	- .074011

TABLE 3.

Integral Real Axis, Imaginary Part

Frequency, Hz	Sorensen's Value, dynes/cm ²	Our Value, dynes/cm ²
10	- .002567	- .002551
20	- .022260	- .022240
40	- .061960	- .062274
80	- .050986	- .050641
160	- .047979	- .047750
320	.006299	.008316

TABLE 4.

Integral Imaginary Axis

Frequency, Hz	Sorensen's Value, dynes/cm ²	Our Value, dynes/cm ²
10	.095080	.095146
20	.081938	.082076
40	.041887	.041759
80	.030795	.030937
160	.030214	.030356
320	.010593	.011043

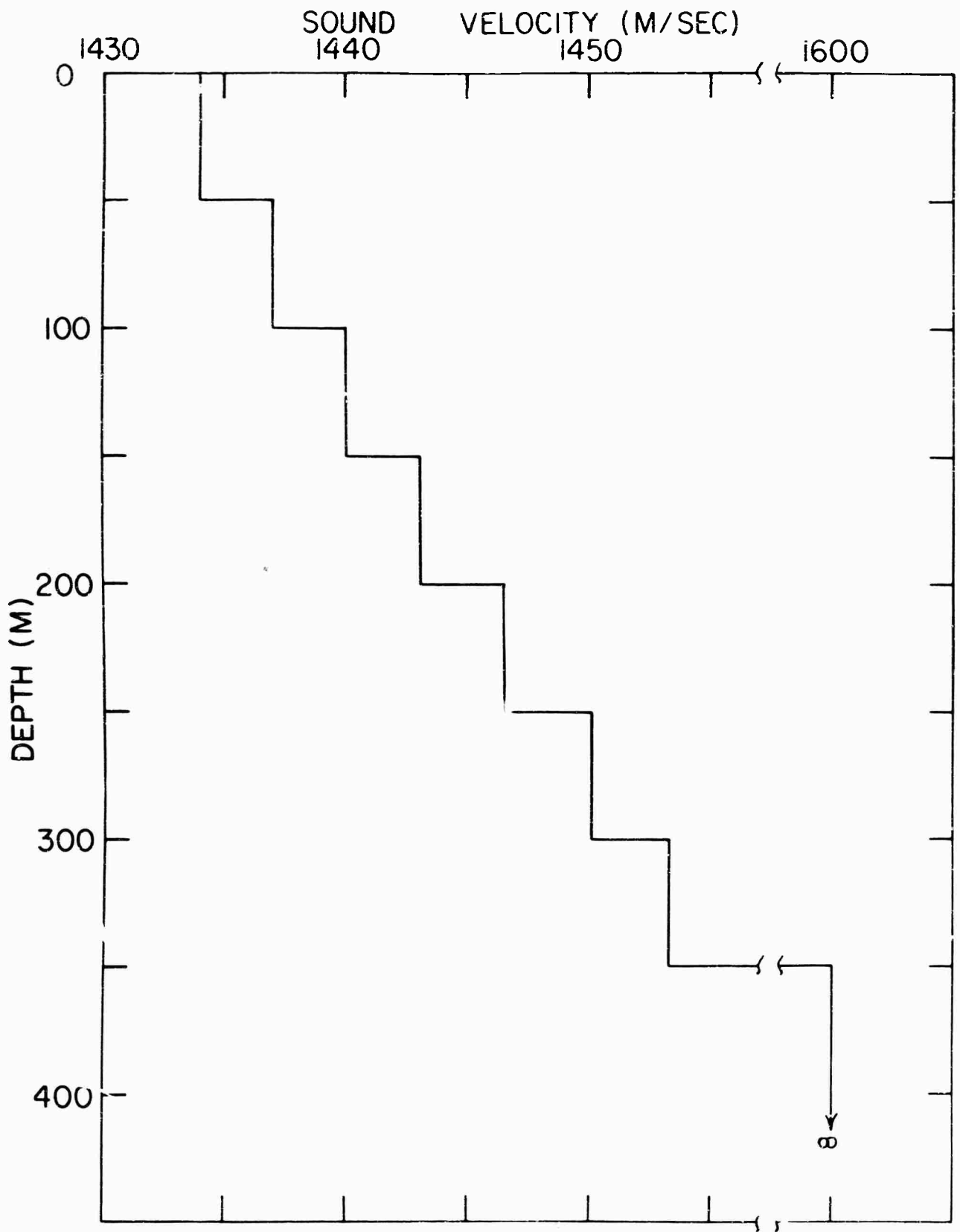


Fig. 3

TABLE 5.
Model A

Layer Thickness, m	Compressional Velocity, m/sec	Shear Velocity, m/sec	Density, gm/cm ³
50	1434.0	0	1.03
50	1437.0	0	1.03
50	1440.0	0	1.03
50	1443.0	0	1.03
50	1446.0	0	1.03
50	1450.0	0	1.03
50	1453.2	0	1.03
Infinite	1600.0	0	1.20

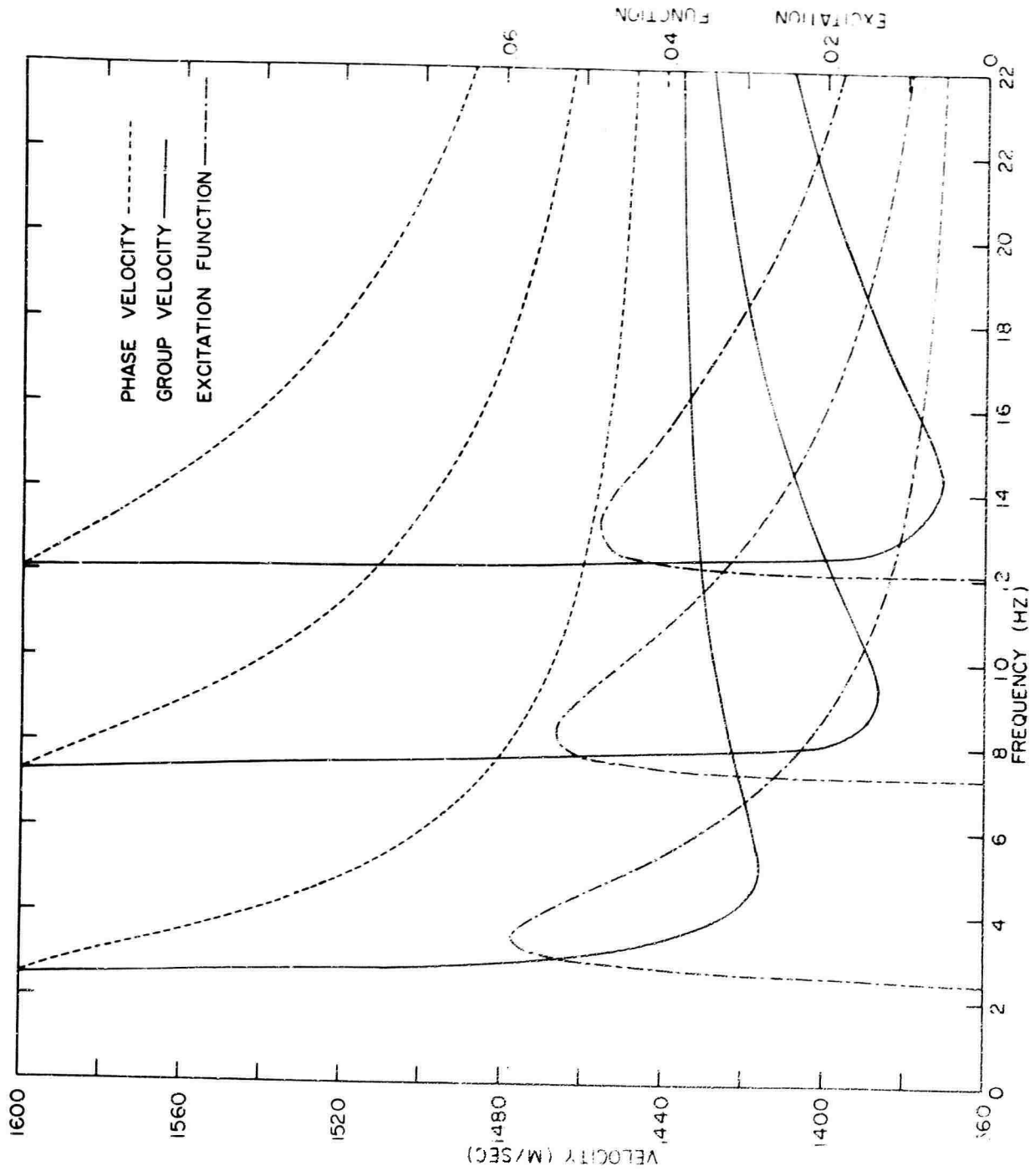


FIG. 4

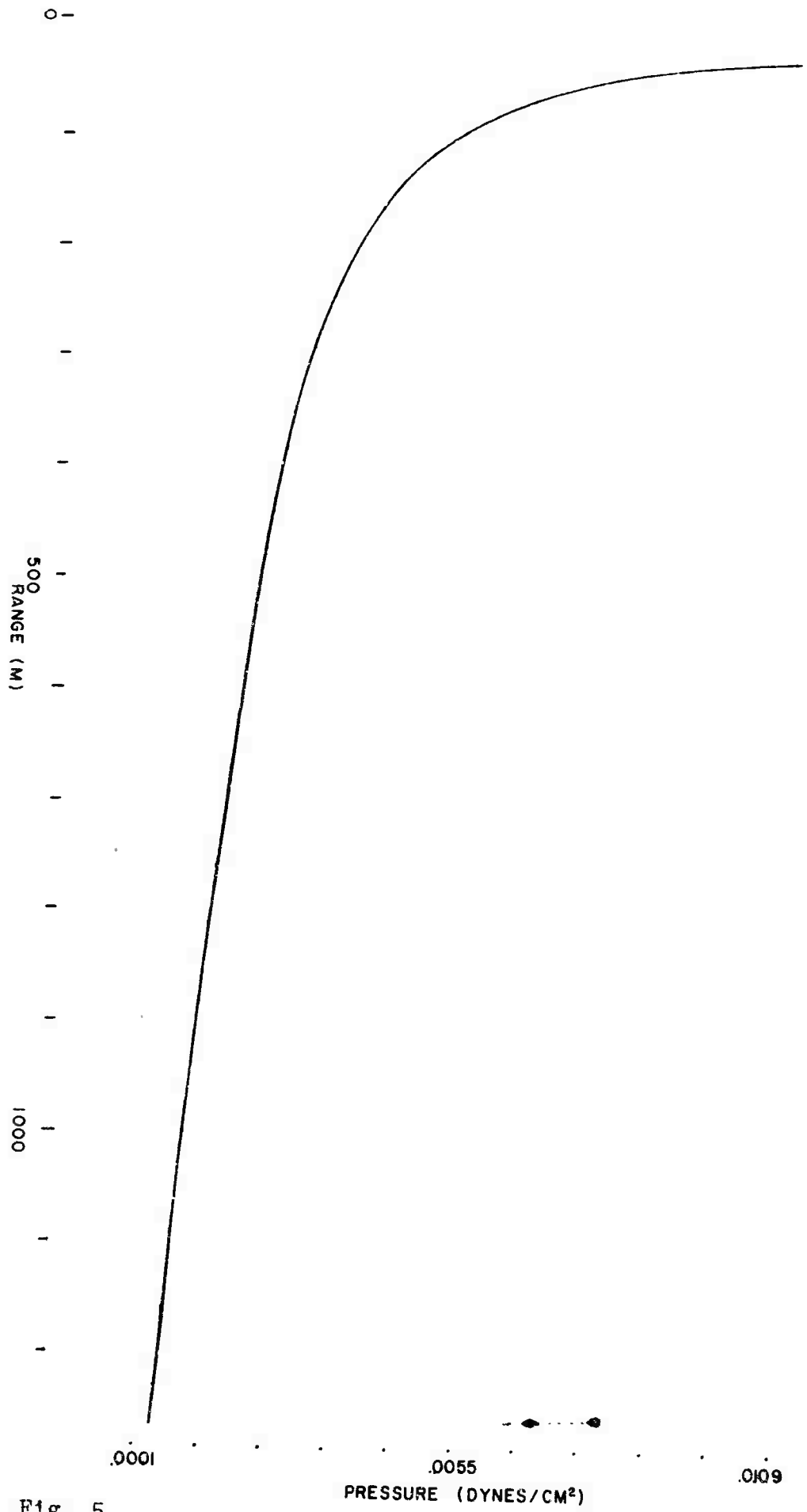


Fig. 5

PRESSURE (DYNES/CM²)

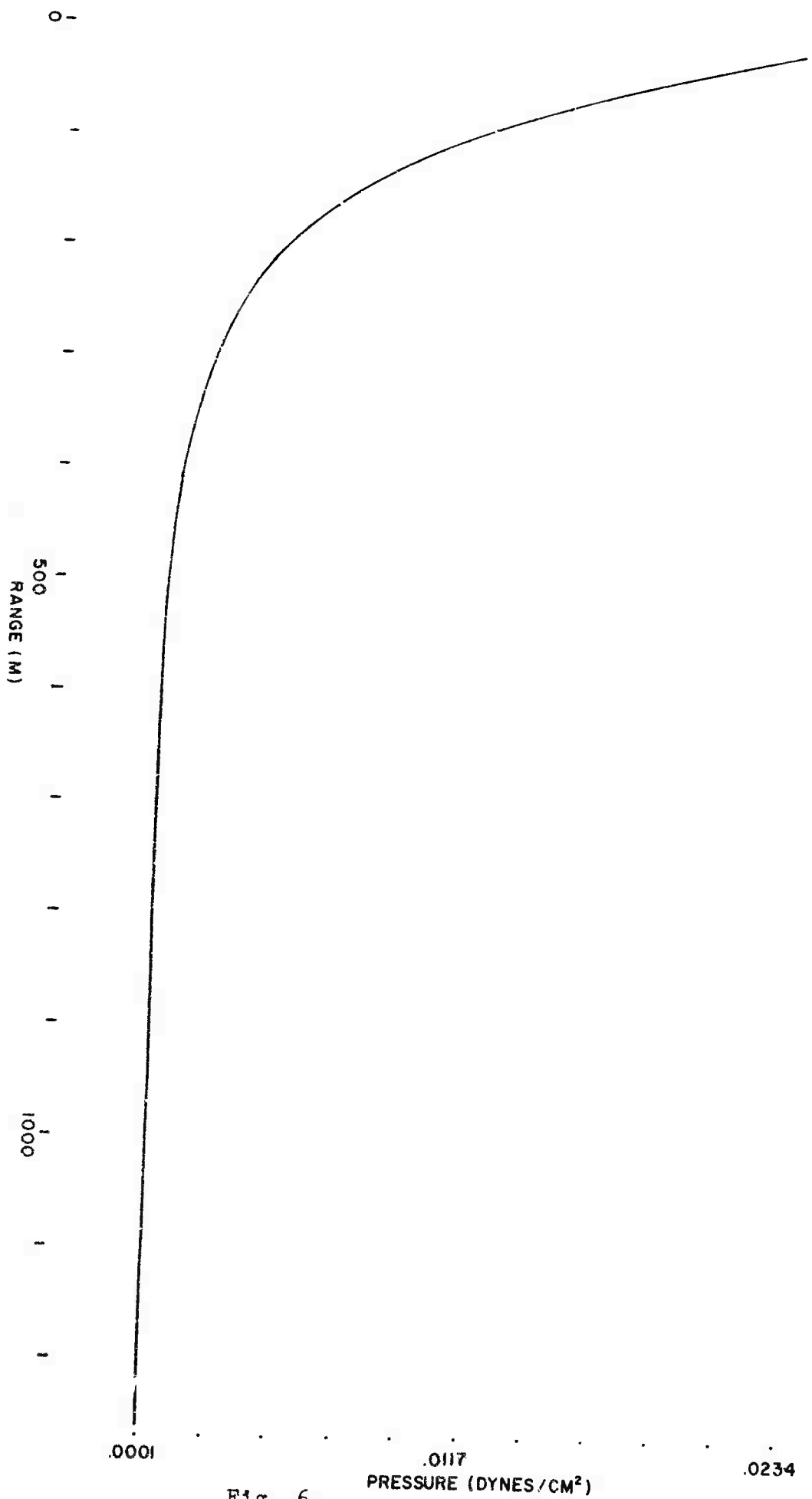


Fig. 6

PRESSURE (DYNES/CM²)

.0234

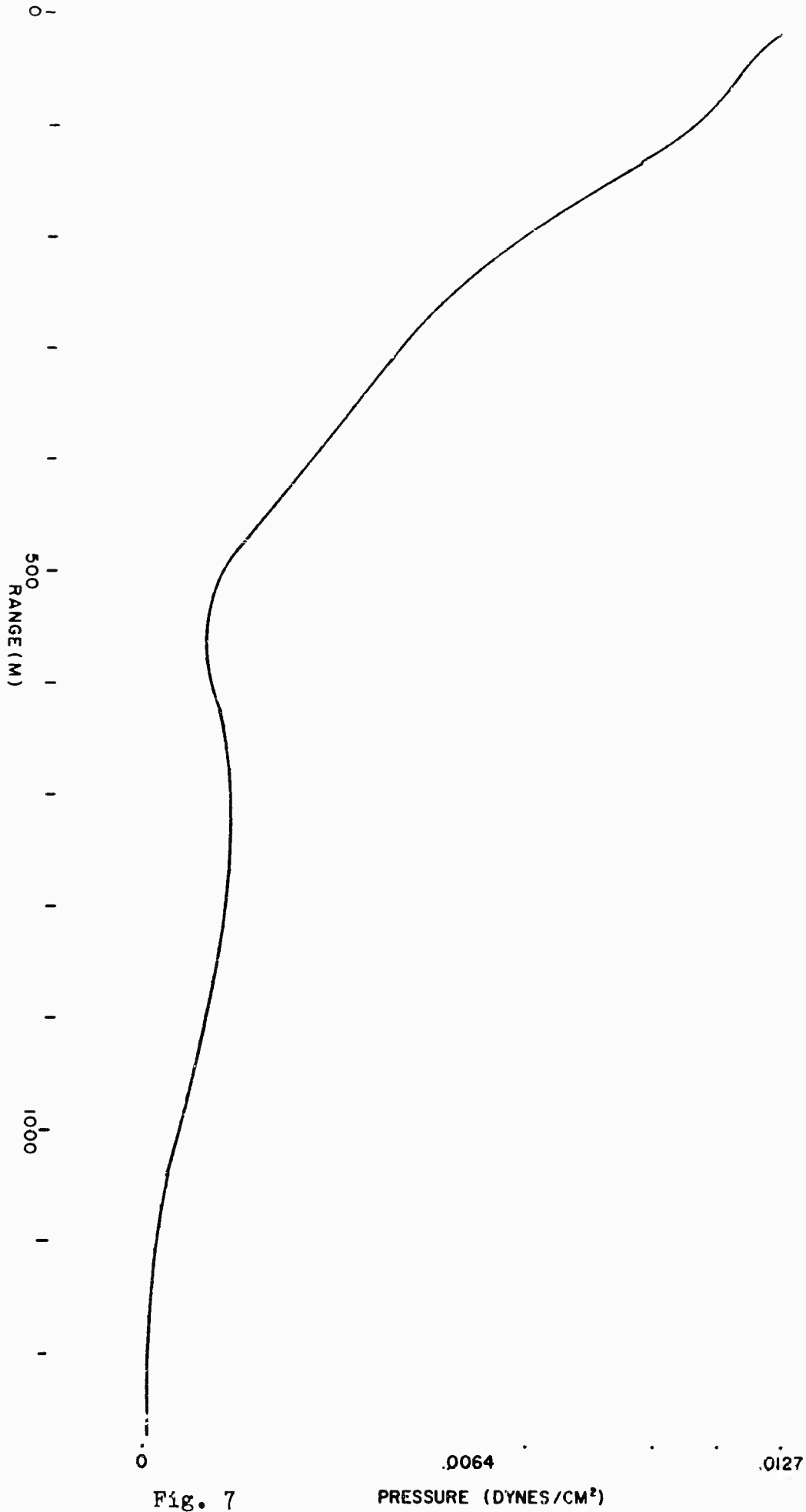


Fig. 7

PRESSURE (DYNES/CM²)

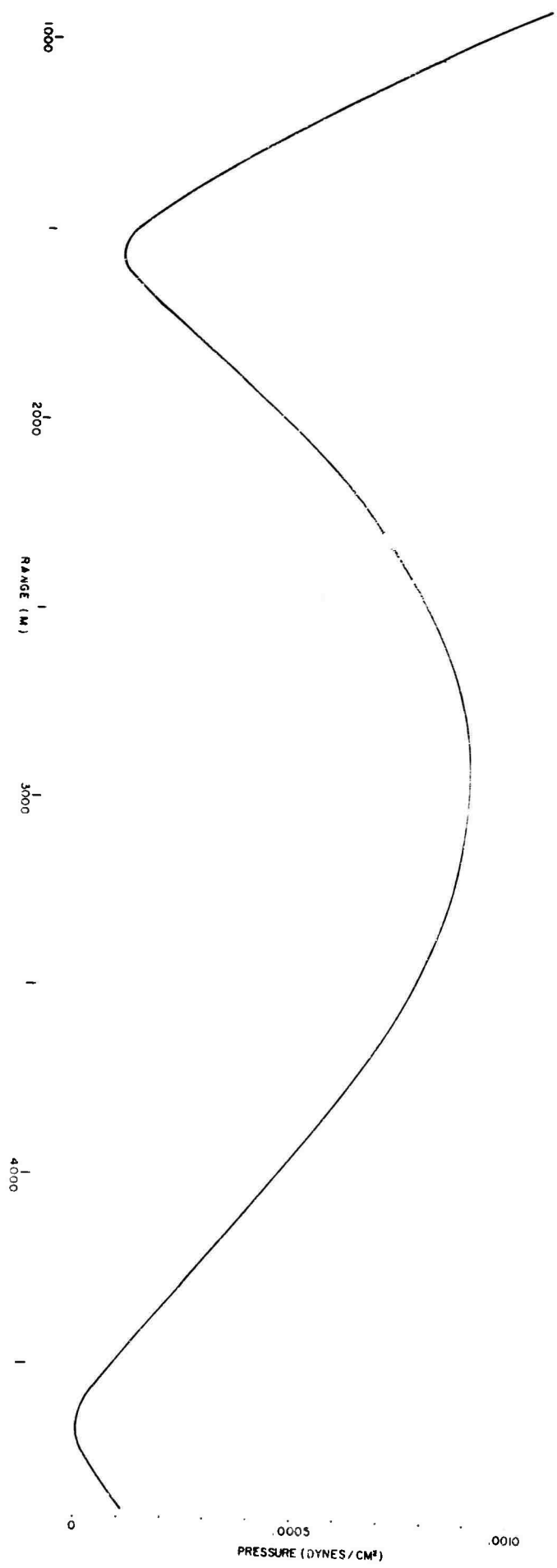


Fig. 8

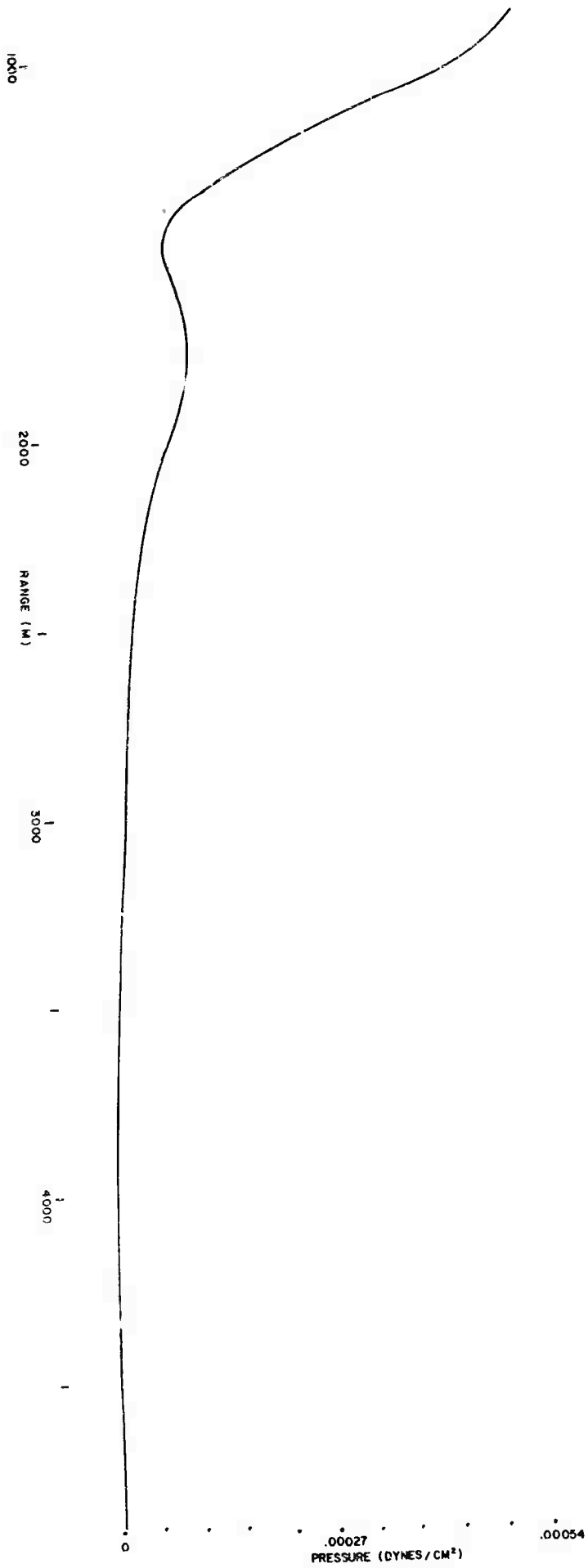


Fig. 9

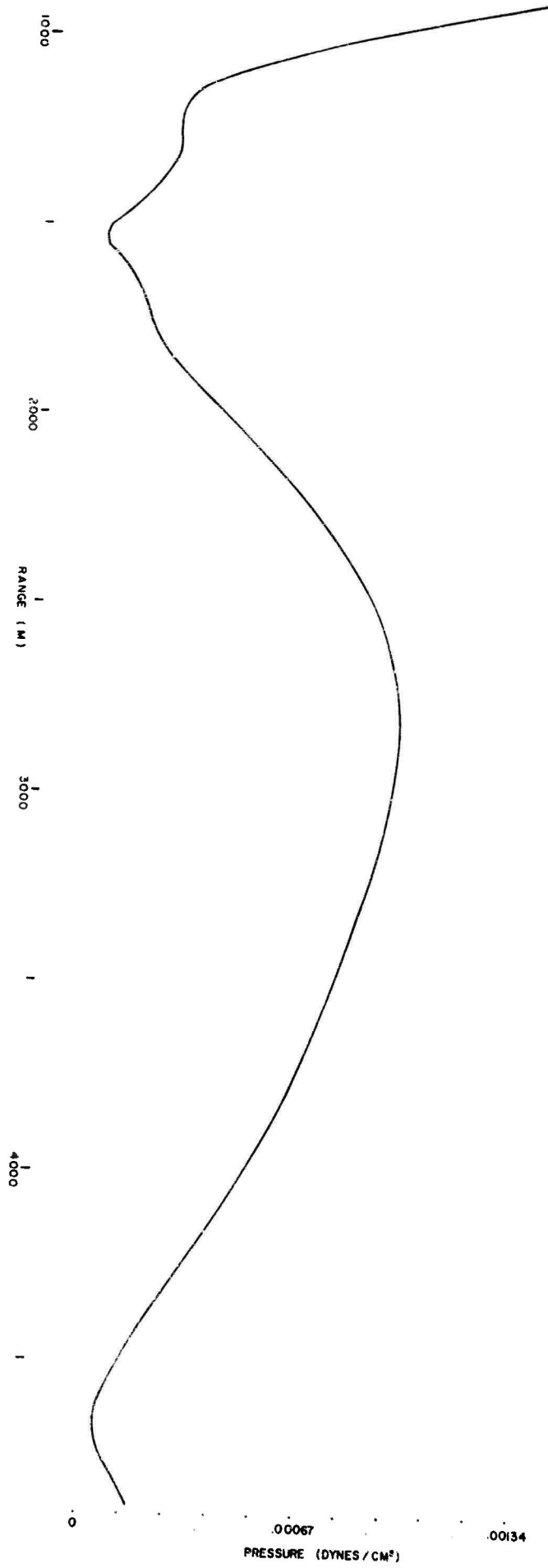


Fig. 10

function of range. In this program provision is also made to compute the three integrals as a function of frequency or detector depth at a specified range. The integrals are evaluated by Gaussian quadrature formulas (see, for example, Hildebrand, 1956). Care must be taken to employ a sufficient number of Gaussian points to obtain the desired accuracy of the integrals. A thirty-two point Gaussian quadrature formula was found to give sufficient accuracy for the computations presented here out to a range of 1500 m, but at longer ranges, higher frequencies, or greater water depths a ninety-six point formula is used. Abscissas and weight factors for the Gaussian integration are given by Davis and Rabinowitz (1956, 1958). The integration along the imaginary axis extends to infinity, but in practice a finite upper limit is chosen which gives sufficient accuracy. This may be done because the integrand is an exponential function of $-kr$ for large k or r . The programs were checked by computations for the same two-layer model of Table 1 used by Sorensen (1959). Tables 2, 3 and 4 show that the two sets of numerical results are in very close agreement.

As an illustration of the programs for the Arctic, computations were made for the layered model A of Figure 3, which closely follows the observed variation of sound velocity with depth. Additional parameters for the model are given in Table 5. This is a preliminary model and the ice sheet, which averages about 3-m in thickness in the central Arctic Ocean, was not included. The effects of the ice sheet on propagation are currently under investigation. Figure 4 shows phase - and group-velocity dispersion and the excitation function for pressure. At a frequency of 10 Hz only two normal modes are excited. The range dependence of the absolute

value of pressure for the normal mode contribution, the integral contribution, and the total pressure are shown in Figures 5 through 10.

ACKNOWLEDGMENTS

Dr. J. Dorman of the Lamont-Doherty Geological Observatory kindly supplied a copy of his PV-7 dispersion program. Computing facilities were provided by the Columbia University Computing Center. This work was supported by the U. S. Naval Ordnance Laboratory and the Office of Naval Research under contract N00014-67-A-0108-0016.

REFERENCES

- Davis, P. and P. Rabinowitz, Abscissas and weights for Gaussian quadratures of high order, J. Res. NBS, 56, 35-37, 1956.
- Davis, P. and P. Rabinowitz, Additional abscissas and weights for Gaussian quadratures of high order: Values for $n = 64$, 80, and 96, J. Res. NBS, 60, 613-614, 1958.
- Dorman, J., Period equation for waves of Rayleigh type on a layered, liquid-solid half-space, Bull. Seismol. Soc. Am., 52, 389-397, 1962.
- Ewing, W. M., W. S. Jardetzky and F. Press, Elastic Waves in Layered Media, McGraw-Hill, New York, 1957.
- Greene, C. R., Arctic operation of seismic transducers, Sea Operations Department, AC Electronics - Defense Research Laboratories of General Motors Corporation, AC-DRL TR 68-53, Santa Barbara, California, 1968.
- Harkrider, D. G., Surface waves in multilayered elastic media I. Rayleigh and Love waves from buried sources in a multilayered elastic half-space, Bull. Seismol. Soc. Am., 54, 627-679, 1964.
- Haskell, N. A., The dispersion of surface waves on multilayered media, Bull. Seismol. Soc. Am., 43, 17-34, 1953.
- Hildebrand, F. B., Introduction to Numerical Analysis, McGraw-Hill, New York, 1956.
- Kutschale, H., Arctic hydroacoustics, Arctic, 22, 246-264, 1969.
- Leslie, C. B., and N. R. Sorensen, Integral solution of the shallow water sound field, J. Acoust. Soc. Am., 33, 323-329, 1961.

- Fekeris, C. L., Theory of propagation of explosive sound in shallow water in Propagation of sound in the ocean, Geol. Soc. Am. Memior, 27, 1-117, 1948.
- Sorensen, N. R., Integral solution of the shallow water sound field, U. S. Naval Ordnance Laboratory, NAVORD Report 6656, White Oak, Maryland, 1959.
- Thomson, W. T., Transmission of elastic waves through a stratified solid medium, J. Appl. Phys., 21, 89-93, 1950.

APPENDIX

Matrix elements for normal modes (see Haskell, 1953, and Dorman, 1962).

Solid layers

$$(\bar{a}_m)_{11} = (\bar{a}_m)_{44} = \gamma_m \cos \bar{P}_m - (\gamma_m - 1) \cos \bar{Q}_m$$

$$(\bar{a}_m)_{12} = (\bar{a}_m)_{34} = i \left[(\gamma_m^{-1} (\gamma_m - 1)) \sin \bar{P}_m + \gamma_m \gamma_{\beta m} \sin \bar{Q}_m \right]$$

$$(\bar{a}_m)_{13} = (\bar{a}_m)_{24} = -(\rho_m c^2)^{-1} (\cos \bar{P}_m - \cos \bar{Q}_m)$$

$$(\bar{a}_m)_{14} = i (\rho_m c^2)^{-1} \left[\gamma_{\alpha m}^{-1} \sin \bar{P}_m + \gamma_{\beta m} \sin \bar{Q}_m \right]$$

$$(\bar{a}_m)_{21} = (\bar{a}_m)_{43} = -i \left[\gamma_{\alpha m} \gamma_m \sin \bar{P}_m + (\gamma_{\beta m}^{-1} (\gamma_m - 1)) \sin \bar{Q}_m \right]$$

$$(\bar{a}_m)_{22} = (\bar{a}_m)_{33} = -(\gamma_m - 1) \cos \bar{P}_m + \gamma_m \cos \bar{Q}_m$$

$$(\bar{a}_m)_{23} = i (\rho_m c^2)^{-1} (\gamma_{\alpha m} \sin \bar{P}_m + \gamma_{\beta m}^{-1} \sin \bar{Q}_m)$$

$$(\bar{a}_m)_{31} = (a_m)_{42} = \rho_m c^2 \delta_m (\delta_m - 1) (\cos \bar{P}_m - \cos \bar{Q}_m)$$

$$(\bar{a}_m)_{32} = i \rho_m c^2 \left[\frac{r_{\alpha m}^{-1} (r_m - 1)^2}{r_m^2 r_{\beta m}} \sin \bar{P}_m + \sin \bar{Q}_m \right]$$

$$(a_m)_{41} = i \rho_m c^2 \left[\delta_m^2 r_{\alpha m} \sin \bar{P}_m + \frac{r_{\beta m}^{-1} (r_m - 1)^2}{r_m} \sin \bar{Q}_m \right]$$

Liquid layers

$$(\bar{a}_m)_{11} = (a_m)_{44} = \cos \bar{P}_m$$

$$(\bar{a}_m)_{12} = \frac{i r_{\alpha m}}{\rho_m c^2} \sin \bar{P}_m$$

$$(\bar{a}_m)_{21} = \frac{i \rho_m c^2}{r_{\alpha m}} \sin \bar{P}_m$$

Matrix elements for integrals along real axis:

Solid layers

$$(a_m)_{11} = (a_m)_{44} = 2 \left(\frac{k}{k_{\beta m}} \right)^2 \cos P_m - \left(2 \left(\frac{k}{k_{\beta m}} \right)^2 - 1 \right) \cos Q_m$$

$$(a_m)_{12} = (a_m)_{34} = i \left[\frac{k \left(2 \left(\frac{k}{k_{\beta m}} \right)^2 - 1 \right) \sin P_m}{\sqrt{k_{\alpha m}^2 - k^2}} + \frac{2k}{k_{\beta m}^2} \sqrt{k_{\beta m}^2 - k^2} \sin Q_m \right]$$

$$(a_m)_{13} = (a_m)_{24} = -\frac{k}{\rho_m \omega h} [\cos P_m - \cos Q_m]$$

$$(a_m)_{14} = i \left[\frac{k^2 \sin P_m}{\rho_m \omega \sqrt{k_{dm}^2 - k^2}} + \frac{1}{\omega \rho_m} \frac{\sqrt{k^2 - k_{\beta m}^2}}{k_{\beta m}} \sin Q_m \right]$$

$$(a_m)_{21} = (a_m)_{43} = -i \left[2 \left(\frac{k}{k_{\beta m}} \right)^2 \frac{\sqrt{k_{dm}^2 - k^2}}{k} \sin P_m + \left(2 \left(\frac{k}{k_{\beta m}} \right)^2 - 1 \right) \frac{k \sin Q_m}{\sqrt{k_{\beta m}^2 - k^2}} \right]$$

$$(a_m)_{22} = (a_m)_{33} = \left[- \left(2 \left(\frac{k}{k_{\beta m}} \right)^2 - 1 \right) \cos P_m + 2 \left(\frac{k}{k_{\beta m}} \right)^2 \cos Q_m \right]$$

$$(a_m)_{23} = \frac{i}{\rho_m \omega} \left[\sqrt{k_{dm}^2 - k^2} \sin P_m + \frac{k^2 \sin Q_m}{\sqrt{k_{\beta m}^2 - k^2}} \right]$$

$$(a_m)_{31} = (a_m)_{42} = \frac{2 \rho_m \omega}{k} \left(\frac{k}{k_{\beta m}} \right)^2 \left(2 \left(\frac{k}{k_{\beta m}} \right)^2 - 1 \right) \cdot \left[\cos P_m - \cos Q_m \right]$$

$$(a_m)_{32} = \frac{i \rho_m \omega}{k} \left[\left(2 \left(\frac{k}{k_{\beta m}} \right)^2 - 1 \right)^2 \frac{k \sin P_m}{\sqrt{k_{dm}^2 - k^2}} + \right]$$

$$2 \left(\frac{k}{k_{\beta m}} \right)^2 \sqrt{k_{\beta m}^2 - k^2} \sin Q_m$$

62.

$$(a_m)_{41} = \frac{i \rho_m \omega}{k} \left[\left(\frac{k}{k_{\beta m}} \right)^2 \sqrt{k_{\alpha m}^2 - k^2} \sin P_m + \left(\frac{k}{k_{\beta m}} \right)^2 - 1 \right] \frac{k \sin Q_m}{\sqrt{k_{\beta m}^2 - k^2}}$$

Liquid layers

$$(a_m)_{11} = (a_m)_{44} = \cos P_m$$

$$(a_m)_{12} = \frac{i \sqrt{k_{\alpha m}^2 - k^2} \sin P_m}{\omega \rho_m}$$

$$(a_m)_{21} = \frac{i \rho_m \omega}{\sqrt{k_{\alpha m}^2 - k^2}} \sin P_m$$

Matrix elements for integral along imaginary axis:

Solid layers

$$P_m = h_m \sqrt{k_{\alpha m}^2 + \tau^2}$$

$$Q_m = h_m \sqrt{k_{\beta m}^2 + \tau^2}$$

$$(a_m)_{11} = (a_m)_{44} = -2 \left(\frac{\tau}{k_{\beta m}} \right)^2 \cos P_m + \left(2 \left(\frac{\tau}{k_{\beta m}} \right)^2 + 1 \right) \cos Q_m$$

$$(a_m)_{12} = (a_m)_{34} = \left(2 \left(\frac{\tau}{k_{\beta m}} \right)^2 + 1 \right) \frac{\tau \sin P_m}{\sqrt{k_{\alpha m}^2 + \tau^2}} -$$

$$\frac{2 \tau}{k_{\beta m}} \sqrt{k_{\beta m}^2 + \tau^2} \sin Q_m$$

$$(a_m)_{13} = (a_m)_{24} = -\frac{i \tau}{\rho_m \omega} \left[\cos P_m - \cos Q_m \right]$$

$$(a_m)_{14} = i \left[\frac{-\tau^2}{\rho_m \omega} \frac{\sin P_m}{\sqrt{k_{\beta m}^2 + \tau^2}} + \frac{1}{\omega \rho_m} \sqrt{k_{\beta m}^2 + \tau^2} \sin Q_m \right]$$

$$(a_m)_{21} = (a_m)_{43} = \left[\frac{2\tau}{k_{\beta m}^2} \sqrt{k_{\beta m}^2 + \tau^2} \sin P_m - \left(2 \left(\frac{\tau}{k_{\beta m}} \right)^2 + 1 \right) \frac{\tau \sin Q_m}{\sqrt{k_{\beta m}^2 + \tau^2}} \right]$$

$$(a_m)_{22} = (a_m)_{33} = \left(2 \left(\frac{\tau}{k_{\beta m}} \right)^2 + 1 \right) \cos P_m - 2 \left(\frac{\tau}{k_{\beta m}} \right)^2 \cos Q_m$$

$$(a_m)_{23} = \frac{i}{\rho_m \omega} \left[\sqrt{k_{\beta m}^2 + \tau^2} \sin P_m - \frac{\tau^2 \sin Q_m}{\sqrt{k_{\beta m}^2 + \tau^2}} \right]$$

$$(a_m)_{31} = (a_m)_{42} = i \left[-\frac{2\rho_m \omega \tau}{k_{\beta m}^2} \left(2 \left(\frac{\tau}{k_{\beta m}} \right)^2 + 1 \right) \cos P_m + \frac{2\rho_m \omega \tau}{k_{\beta m}^2} \left(2 \left(\frac{\tau}{k_{\beta m}} \right)^2 + 1 \right) \cos Q_m \right]$$

$$(a_m)_{32} = i\omega\rho_m \left[\left(2 \left(\frac{z}{k_{\beta m}} \right)^2 + 1 \right)^2 \frac{\sin P_m}{\sqrt{k_{\alpha m}^2 + z^2}} - \right.$$

64.

$$\left. \left(\frac{2z}{k_{\beta m}} \right)^2 \sqrt{k_{\beta m}^2 + z^2} \sin Q_m \right]$$

$$(a_m)_{41} = i\omega\rho_m \left[- \left(\frac{2z}{k_{\beta m}} \right)^2 \sqrt{k_{\alpha m}^2 + z^2} \sin P_m + \right.$$

$$\left. \left(2 \left(\frac{z}{k_{\beta m}} \right)^2 + 1 \right)^2 \frac{\sin Q_m}{\sqrt{k_{\beta m}^2 + k^2}} \right]$$

Liquid layers

$$P_m = h_m \sqrt{k_{\alpha m}^2 + z^2}$$

$$(a_m)_{11} = (a_m)_{44} = \cos P_m$$

$$(a_m)_{12} = \frac{i\sqrt{k_{\alpha m}^2 + z^2} \sin P_m}{\omega\rho_m}$$

$$(a_m)_{21} = \frac{i\rho_m \omega \sin P_m}{\sqrt{k_{\alpha m}^2 + z^2}}$$

DOCUMENT CONTROL DATA - R & D

(Security classification of title, body of abstract and indexing annotation must be entered when the overall report is classified)

1. ORIGINATING ACTIVITY (Corporate author) Lamont-Doherty Geological Observatory of Columbia University		2a. REPORT SECURITY CLASSIFICATION Unclassified	
		2b. GROUP Arctic	
3. REPORT TITLE The Integral Solution of the Sound Field in a Multilayered Liquid-Solid Half-Space with Numerical Computations for Low-Frequency Propagation in the Arctic Ocean.			
4. DESCRIPTIVE NOTES (Type of report and inclusive dates) Technical Report			
5. AUTHOR(S) (First name, middle initial, last name) Henry W. Kutschale			
6. REPORT DATE February 1970		7a. TOTAL NO. OF PAGES 64	7b. NO. OF REFS 13
8a. CONTRACT OR GRANT NO. N00014-67-A-0108-0016		9a. ORIGINATOR'S REPORT NUMBER(S) 1	
b. PROJECT NO. NR 307-320/1-6-69 (415)		9b. OTHER REPORT NO(S) (Any other numbers that may be assigned this report)	
c.			
d.			
10. DISTRIBUTION STATEMENT Reproduction of this document in whole or in part is permitted for any purpose of the U.S. Government.			
11. SUPPLEMENTARY NOTES		12. SPONSORING MILITARY ACTIVITY U.S. Naval Ordnance Laboratory, White Oak, Silver Spring, Maryland Office of Naval Research, Washington,	
13. ABSTRACT This report develops by matrix methods the integral solution of the wave equation for point sources of harmonic waves in a liquid layer of a multilayered liquid-solid half space in a form convenient for numerical computation on a high-speed digital computer. Only the case is considered of a high-speed liquid bottom underlying the stack of layers. The integral over wave number has singularities in the integrand and is conveniently transformed into the complex plane. By a proper choice of contours, complex poles are displaced to an unused sheet of the two-leaved Riemann surface, and the integral solution for the multilayered system reduces to a sum of normal modes plus the sum of two integrals, one along the real axis and the other along the imaginary axis. Both integrals are evaluated by a Gaussian quadrature formula. Sample computations are presented for low-frequency propagation in the Arctic Ocean sound channel. These are preliminary computations and the ice layer, which averages three meters in thickness, is not included in the layered system. The effects of the ice layer on propagation are currently under investigation.			

14 KEY WORDS	LINK A		LINK B		LINK C	
	ROLE	WT	ROLE	WT	ROLE	WT
Arctic						
Hydroacoustics						
Sound field						
Layered media						
Matrix						
Integral solution						
Digital computer						
Numerical computations						