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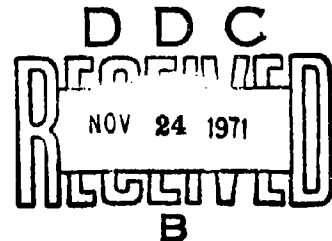
**ATMOSPHERIC AEROSOLS
BETWEEN 700 AND 3000 m ABOVE SEA LEVEL
PART V**

**A Study of the Effects of Atmospheric Fine Structure
Characteristics on the Vertical Distribution of Aerosols**

FINAL TECHNICAL REPORT

by

Reinhold Reiter
Rudolf Sládkovič
and
Walter Carnuth



July 1971

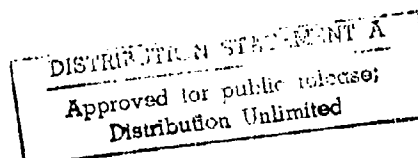
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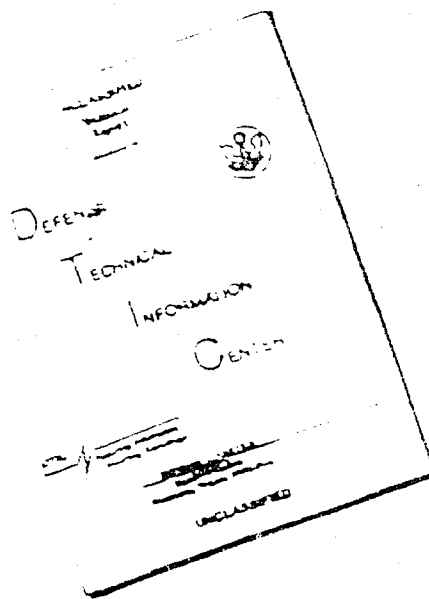
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ABSTRACT

Our research work for the purpose of studying vertical aerosol exchange up to 3000 m a.s.l., which was explained in our previous reports, has consistently been carried on. It comprises the obtaining, to the extent possible, of uninterrupted recording data, on the one hand, and, on the other, the development of appropriate methods of electronic data processing and their application toward the derivation, for practical utilization, of relations between aerological parameters and the incremental exchange coefficient. Following are details on the investigations conducted:

Obtaining of Data: Continuous recording of RaB and RaC at 700 m, 1800 m, and 3000 m a.s.l.; computation of the main vertical exchange coefficients for the strata between these levels, and computation of the ionization rate profile. Recording of air conductivities and of the concentration as well as the particle size spectrum of Aitken nuclei at the same levels. Application of these data for computing the vertical profiles of the Aitken nucleus profiles, from the soundings of air conductivities between the said levels. Performance of more than 400 soundings between 700 m and 1800 m a.s.l., and more than 1000 soundings between 1000 m and 3000 m a.s.l., by means of the cable car telemetry systems, for obtaining the profiles of the following parameters: wind velocity, temperature, humidity, polar air conductivities, and potential gradient. Application of the profiles of aerosol particle concentration as computed from the air conductivity profiles, for computing the profile of the incremental exchange coefficients.

Data Handling: Development of computer programs for automatic computation, from the said recordings, of all essential parameters and their altitude dependence; automatic transfer of the data to parameter data tables.

Scientific Processing of Data: Derivation of significant statistical relations between aerological data such as gradients of temperature, of potential and equivalent potential temperatures, of vapor pressure, on the one hand, and the incremental exchange coefficients, on the other. Detailed analyses to investigate the interrelations between all essential parameters, and to deduce appropriate principles of classification as a basis for future improved statistical data processing.

In the appendix, numerous graphs, copies of original recordings, and parameter data tables illustrate the relations found, up to now, and afford a glimpse into our analysing workshop. The most important computer program is printed out in full.

The narrative portion is concluded by a preview of advanced studies intended for the near future.

I. OBJECTIVES AND BACKGROUND

Our investigations carried out during the period from 1 July 1970 until 30 June 1971 are based upon the preparatory work of the preceding years of operation. The fundamental problem of the current investigations which have been greatly expanded, instrumentation-wise, by the telemetry system aboard the Zugspitze cable car, has been expounded in detail in the preceding report (1) and in a publication (2). It is assumed, therefore, that a reiteration of the final objective of our research may here be dispensed with. Earlier and more recent literature references to be considered have been thoroughly discussed in (1,2), also. New publications by other scientists of any importance have not since been added.

The subject to be presented, however, in this chapter is how does the work performed during the reporting period differ significantly from that of the preceding period up to the middle

of 1970. During the latter it was our objective (a) to shield the cable car telemetry systems against all effects of the weather and to obtain by them homogeneous series of recordings; (b) to develop and apply the physical and mathematical bases for the complete numeric evaluation of the recording runs, up to and including computation of the incremental exchange coefficients; (c) to continually record at the three stations the Aitken nuclei concentrations, natural radioactivity (RaB, RaC), and polar conductivities, and to utilize them in the evaluations; and (d) to derive initial deductions from the results.

The performance of evaluating operations was predominantly based on manual processes. However, certain phases of the computations, e.g. those of Aitken nucleus concentration from the total electric conductivity of the air, and of the incremental exchange coefficients from the vertical gradients of the conductivity data, were already performed electronically with the Hewlett Packard desk computer.

Now, during the reporting period of 1 July 1970 through 30 June 1971, the entire evaluation technique was completely changed. A new computer (Intertechnique Multi 8) with 12 k words (1 word = 9 bit) affords us the possibility to perform the entirety of all computation processes in one single operation, and to print them out in tables. A considerable portion of the time, however, had to be used to develop computer programs, to test them and adopt them to the computation problem in all details. The newly acquired accessory equipment such as fast reader and fast tape perforator afforded new possibilities to optimize the computation work, however, initially required time consuming conversions of the computer programs. Furthermore, it was necessary to transfer our stock of primary data, i.e. the crude data read from the recordings, on punch tapes by means of teletype, and do this both retroactively to the greatest extent possible, and continually up until they were caught up actu-

ally with the daily soundings.

Hence, there was only a relatively short period left for a thorough statistical processing of the data material as the same can only be based upon the final tables from the computer. Nevertheless, it is possible in this report to convey some interesting - tentative - results. These as well as the detailed analyses described below are in this case based upon Zugspitze cable car soundings only. Analog processing of the Wank cable car soundings will be done later. Also the vertical wind profiles recorded by both systems shall be considered at a later date.

It is by no means gratifying to merely process the data obtained statistically, "blindly", as it were. Such a "blind statistics" is even bound - if solely relied on - to lead on to wrong tracks and into incomplete knowledge. It is rather an indispensable necessity to analyse in detail and "individually", such series of soundings as belong together. Only from such practice, i.e. from an undeniably time consuming detailed analysis, grows the foundation of experience upon which a meaningful and successful statistical evaluation may be based. Part of this report will, therefore, deal with the initial results of such detail analyses.

Our long-range objective is - beside the continuation of the soundings, computer work, and detailed analysis - (1) a complete synopsis of the existing interrelations, and (2) solution of the problem to render the disclosed interrelations applicable for practical use.

II. OBTAINING OF DATA

1. Technical Facilities

a) Cable Car Telemetry Systems

During the reporting period the telemetry system in the Wank cable car was working without any troubles. In order to obtain data by it which are comparable in every respect to those obtained by the Zugspitze cable car system, the arrangement of the instruments in the bottom of the Wank cable car was changed. The inlet tube for the measuring air flow was assimilated to that of Zugspitze cable car in length, diameter and position relative to the body of the car.

To begin with, the telemetry system of the Zugspitze cable car required some further improvements.

Due to the large altitudinal difference traversed, condensation would occur occasionally on some of the sensitive electronic parts such as resistors of extremely high ohmic value. We therefore had to resort to installing those parts into sealed cartridges which are kept dry inside through anhydrous CaO. This solved the problem, and even the most severe weather conditions during the run or during waiting period at Zugspitze peak no longer affected the functioning of the unit.

Furthermore, it was possible to completely eliminate a disturbing sensitivity of the cable car equipment as well as the relay station at St. Martin's, to nearby lightning strokes; so, interesting recordings are now available from inside the thunderclouds.

Both systems have repeatedly been tested, calibrated and compared with one another in the laboratory.

b) Recordings of Natural Radioactivity in the Air
(Concentration of RaB and RaC).

Recordings of RaB and RaC concentrations at the 3 stations were made without interruption, evaluated and utilized within the scientific program. This was done, on the one hand, by computing the mean vertical exchange coefficients between 700/1800 m and 1800/3000 m a.s.l., and on the other hand by using the data in computation of the ionization rate. The results of the recordings for the reporting period are compiled in X. Appendix, Section 4.

c) Recording of Aitken Nuclei Concentration at the
Stations.

Aitken nuclei concentration and particle size distribution within the Aitken nuclei range were recorded at the two stations of Wank peak and Garmisch-Partenkirchen without any interruptions or disturbances. All data have been continually analyzed. These (and many other) data are picked up and processed at the latter station by our electronic computer system. Apart from basic research studies on Aitken nuclei and their behavior, these data are primarily used as a basis at 700 m and 1800 m a.s.l., for the conversion of total conductivities into nuclei concentrations.

A new Aitken nuclei counter of the make Environment/one, USA, was installed and put into operation at the Zugspitze station. This unit is smaller and lighter than the GE units, which is essential for operation at Zugspitze peak with the limited space available. But despite considerable expenditure of repair and reequipment efforts, it unfortunately was not yet possible to obtain homogenous recording series over more than a few weeks.

d) Recording of Polar Conductivities and Space Charge Densities at the Stations.

During the reporting period, instruments were put into operation for the recording of polar conductivities and space charge densities at 1800 m a.s.l. (Wank peak) and 700 m (valley floor). From the terminals of our cable car recording range, they furnish data, continuous in time, for direct comparison and link-up. The instruments are functioning satisfactorily, even under bad weather condition. The equipment that had been previously installed at Zugspitze peak has, without interruption, been yielding recorded data which have been evaluated.

e) Numerical Processing of Recorded Data

The methods of numerical processing of recorded data are outlined in (1,2).

2. Volume of Processed Data

From 1 July 1970 through 31 May 1971, we made

413 recording runs by Wank cable car telemetry system
and 1,015 recording runs by Zugspitze cable car telemetry
system.

The manual evaluation operations that were still necessary have all been completed.

During the same period 380 Wank cable car and 900 Zugspitze cable car runs were transferred to punch tapes and mathematically analyzed by computer. The final results are printed out in tabulated form. Based on the fact that each run yields roughly 500 individual data, our file presently comprises 650,000 data.

III. DATA PROCESSING

1. The Physical-Mathematical Fundamentals for Data Processing

The physical-mathematical computation operations are described in every detail in (1), Chapter IV (also cf (2)). They were retained in every detail, hence there is no need for reiteration here.

2. Brief Description of the Overall Procedure of Technical Data Processing

Processing of the data, starting with transfer from the moving cable car up to the print-out of the completed tables is shown in Fig.1.

The sensors and the transmitter on the running cable car transmit the measured values to the receiver in codified form and in the known manner. Thereupon follows the electronic decoding of the data and their analog recording on three xy-recorders. The y-channel of each of the three recorders is fed with the analog pressure signal as parameter of altitude. The x-channels are occupied by the analog data for polar air conductivities, wet and dry temperatures and potential gradient + wind speed. The three record sheets obtained per sounding are then subject to intermediate manual processing: the relative values are read off at the main levels (vertical intervals of 100 or 200 m) and at the secondary levels in between. These secondary levels are determined by the processing scientist, based upon the given aerological fine structures of the soundings. These relative values are then compiled in a work chart (intermediate table) per sounding. Thereupon, they are punched (per sounding) on an 8-channel punch tape by means of a teletypewriter. These tapes are then checked for errors (omissions, wrong signs, logical errors) with the aid of a test program; defective tapes are corrected (rate of errors not in excess of 5%). This test pro-

cess is not reflected in the diagram. After the test, the punched tapes are read via the fast reader, into the core store of the computer. The reading time per sounding is approx. 2 seconds. With the aid of the computer program meanwhile developed, which likewise has to be read into the core store of the computer, the result data are computed. The computation process takes 120 seconds. Thereupon, the result data are punched by means of a fast perforator on a punch tape (approx. 80 seconds per sounding). It is advisable to punch the data of several soundings, one after the other, on a continuous tape of 250 m length. The latter is then fed into another teletypewriter which will print out the result data in the form of clearly arranged tables. They are then furnished to the scientist who will study the mutual correlations of the data, taking into consideration any characteristic atmospheric conditions.

3. The Computer Program as Such

The computer program is described in full in Chapter X, Appendix, 5.

4. Description of Final Data Tables

In Chapter X, Appendix, 3, a number of Final Data Tables are printed, true to original, as supplements to individual runs. They are arranged as follows:

Headings: Z = Zugspitze system (W = Wank system)
Number of run; Date; Hour of start of run.

a. Main levels (Haupt-Niveaus")

Vertical intervals 100 m for Wank system;
Vertical intervals 200 m for Zugspitze system.

b. Special levels, selected by temperature structure
("MP-Temperatur")

Vertical intervals are roughly 30 m as a minimum.

c. Special levels selected by humidity structure -- to
the extent applicable -- ("MP-Feuchtigkeit")

Vertical intervals see b.

d. Special levels, selected by air conductivity struc-
ture -- to the extent applicable -- ("MP-Leitfähigkeit")

Vertical intervals see b,

Meaning of letters designating vertical columns of
 Tables:

Z	: Number of level
H	: Altitude (m.a.s.l.)
P	: Pressure (mb)
T	: Air temperature ($^{\circ}\text{C}$)
TF	: Wet temperature ($^{\circ}\text{C}$)
PT	: Potential temperature ($1/10^{\circ}\text{K}$)
EPT	: Equivalent potential temperature ($1/10^{\circ}\text{K}$)
E	: Water vapor pressure (mb)
S	: Specific humidity (g/kg)
RF	: Relative humidity (mm%)
+L	: Positive polar conductivity (10^{-14} 1/Ohm.m)
-L	: Negative polar conductivity (10^{14} 1/Ohm.m)
+/-	: Ratio of both conductivities
L	: Total conductivity (see above)
DH	: Height difference (m)
TV	: Virtual temperature ($^{\circ}\text{C}$)
G-T	: Gradient of T
G-E	: Gradient of E
G-S	: Gradient of S

G-RF : Gradient of RF
 G-PT : Gradient of PT
 I : Ionization rate (ions/cm³ sec)
 N : Number of Aitken nuclei (cm⁻³)
 A : Incremental exchange coefficient (g cm⁻¹ sec⁻¹)
 D : Diffusion coefficient (cm² sec⁻¹)

Definition of sign of gradients:

The gradients are positive when:

T : decreases with height
PT : increases with height
E : increases with height
S : increases with height
RF : increases with height

The meanings of the letters and gradients should be born in mind when later the Tables and graphs are discussed.

IV. STATISTICAL RESULTS

1. Preliminary Remarks; Kind of Relations Studied; Principles of Selection

Our basic idea in the statistical analyses was to find significant relations between the fine structure characteristics of the meteorological parameters, on the one hand, and the respective value of the incremental vertical exchange coefficient A_1 on the same vertical interval, on the other hand. There are numerous possibilities of doing so. After preliminary tests, the following relations appeared sufficiently promising to be numerically investigated (for legend see above).

- a) $A_1 = f (G-T)$ Inversion
 b) $A_1 = f (G-PT)$

- c) $A_1 = f(G-E)$ Inversions and other barring layers
- d) $A_1 = f(G-PT)$ General relation,

There was no question, from the outset, that all values from the computer tables were not suitable to be "blindly" used for differentiation of the structure of these relations. The definition of the principles of selection, however, is a tentative one and may be varied or supplemented by other, additional ones in subsequent, more exhaustive investigations. In regard to the above defined relations, the following principles of selection were applied in the following:

a) Data of such vertical intervals only were used (no matter whether bounded by main or special levels) containing distinct temperature inversions (G-T negative). These inversion layers were required to have settled both in time and position; cases with a marked vertical movement of the inversion level were excluded, likewise such with a sudden formation or disintegration of the inversion structure. That pair of values was used at that vertical interval where the steepest negative gradient of T was reached.

b) Same definition as a).

c) Again, data of only one vertical interval were used per run, with the following conditions applied: if there was no negative G-T value anywhere that vertical interval having the lowest positive value was selected; if there was a vertical interval with - G-T, the latter was used; if there were several such the highest negative G-T value was used as criterion of selection. At the same time, G-E was required to show a maximum.

d) Almost all existing pairs of values were used, except such, derived from periods or layers where the structure was subject to rapid changes in time, a transformation of the local meteorological conditions was in progress, or where fog or

precipitations were traversed during a run. Likewise, values from layers with distinct inversions have not been included; they belong under b).

In the following part of this we show the distribution of the individual pairs of values as dots in the graphs, on the one hand, and each time a respective statistical evaluation, on the other. The latter was made with our Hewlett Packard computer which, in the meantime, has been equipped with an extended memory. The pairs of values were keyed into the computer. The latter then separated them according to A-values. To this end, each A decade was subdivided into 10 equal intervals. The computer automatically determined the 40 mean values between $A_1 = 0.1$ and $A_1 = 1000$, plus the pertinent criteria of significance

$$\sigma_M = \pm \sqrt{\frac{\sum_{i=1}^n (x_i - \bar{x})^2}{(n-1) 2n}}$$

The results of these statistical investigations are presented in the form of "stepped graphs". In the discussion of the graphs below the above listed criteria of selection should be born in mind, as appropriate.

2. Relation $A_1 = f(G-T)$, Inversions

Fig. 2 shows the scattering of the individual pairs of values for the relation $A_1 = f(G-T)$ in the case of a typical inversion (see Section IV.1.a). The study covers the period of 28 July 1970 through 17 February 1971 with a total of 388 pairs of values. The graph is analogous to Fig. 23 in (1). A comparison of the result depicted in said graph with that in Fig. 2 reveals this: In principle the earlier result agrees quite well with the more recent one except that the relation $A_1 = f(G-T)$ is now much more pronounced, the scattering of the values is narrower.

Fig. 3 shows the analogous result of a statistical analysis with statement of scatter σ_M . This graph clearly shows: There is

a strict, significant relation between an incremental exchange coefficient in the range of $A_1 = 0.1$ through 300, and the temperature gradient in the range of -0.1 through -10 °C/100 m, provided that the inversion in the vertical interval under consideration is a genuine, stable and largely motionless one to which the A_1 and G-T values apply. The relation is independent of season or time of day.

Hence, the function presented in Fig. 3 is already ripe for practical application.

3. Relation $A_1 = f(G-PT)$, Inversions

Fig. 4 shows the scattering of the individual pairs of values for the relation $A_1 = f(G-PT)$. It, again, applies to cases of typical inversions. The principle of selection is defined in Section IV.1. a) and b). The period of the study, again, stretches from 28 July 1970 through 17 February 1971 and covers a total of 289 pairs of values.

With the scattering of values not too wide a surprisingly clearly defined logarithmic relation is found between incremental exchange coefficient A_1 and potential temperature gradient G-PT:

If in a defined vertical interval with an inversion the potential temperature gradient increases, the intensity of exchange in the same layer is decreased. The following rule may be stated: Per 1 °C/100 m increase of G-T, A_1 is decreased by one decimal power.

Fig. 5 shows the analogous result of a statistical analysis with statement of scatter σ_M . True it is greater than in Fig. 3, nevertheless between $A_1 = 100$ and $A_1 = 0.8$ the relation may be considered well established. The volume of data material is to be further increased in order to make the relation come out even more clearly. This particularly applies to the value range of $>A_1 = 150$ and $<A_1 = 1$.

Aside from these statements we feel justified in claiming that this function $A_1 = f(G-PT)$, too, may already be practically applied.

We consider the statement important that now we have two criteria, independent from each other, to quantitatively state the barring effect of an inversion, viz. the function in Fig. 5 and that in Fig. 3. From now on, the layer having the strongest barring effect may immediately be determined from the aerological data: negative temperature gradient and positive potential temperature gradient must be at maximum values.

4. Relation $A_1 = f(G-E)$, Inversions and Other Barring Layers

In studying the relation between the incremental exchange coefficient A_1 and the gradient of water vapor pressure (G-E) we have dropped the demand that a genuine, stable and stationary inversion must exist. With the aid of the gradients of temperature and potential temperature we have merely searched for a barring layer in general (for definition of principle of selection see Section IV.1.c). For the layers thus defined we then studied the above described relation, using data from the period of 28 July 1970 through 17 February 1971. The graph containing the individual pairs of values is shown in Fig. 6; the results of the statistical analysis is presented in Fig. 7. Between $A_1 = 100$ and $A_1 = 1$ we find a significant logarithmic function having a relatively narrow scatter with only 263 pairs of values. It appears desirable to increase the data, particularly in the limiting zones. The rule has been established that A_1 will be increased by one decimal power when the negative value of G-E is increased by 1 mb/100 m - and vice versa. This is comprehensible in principle since the vertical exchange of water vapor, too, is controlled by the amount of the A_1 value. Nevertheless, it is probable that the graphs 6 and 7, for the first time, convey a tangible picture of the function as such as it is actually applicable in the atmosphere.

With this function, now, we have an independent third criterion by which to locate layers having a barring effect, and to determine the degree of their barring effect.

5. Relation $A_1 = f(G-PT)$, General Cases Other than Inversions

Within the scope of the previous studies we had limited ourselves to either typical inversions (Sections 2 and 3) or, at least, layers of such aerologic structure as to permit the inference of a barring effect, including inversions (Section 4). To the extent that criteria of selection were applicable at all, the values of one layer, at maximum two layers, per run were, in this manner, to be taken into consideration. The study now being discussed, however, covers all layers (any number per run) except such containing an inversion. As explained in Section IV.1.d), the following have also been left outside of consideration: cases with a rapid transformation of the weather condition, with precipitations, or runs entirely or partly through clouds.

Preliminary studies have revealed that above all the potential temperature gradient is very well suited for this kind of a general statistics.

In order to determine whether seasonal differences might play a part in this we have subdivided the entire period from which values are available, into 28 July 1970 through 12 November 1970, and 13 November 1970 through 17 February 1971. We shall here dispense with a presentation of the distribution of individual pairs of values, even more so as, in view of the large volume of data, this kind of evaluation would have been too expensive.

The results of the statistical investigations are shown in Figs. 8 and 9, with statements as to the respective scatters. From both figures a clear and strict relation becomes apparent ($A_1 = f(G-PT)$), scattering only insignificantly in the range between $A_1 = 1000$ through $A_1 = 1$. At $A_1 < 1$ the relation becomes

uncertain. As demonstrated by Fig. 10 the results from both periods of time are in good agreement with one another, hence there is no seasonal impact.

The relation depicted in Figs. 8 through 10 applying to those atmospheric layers which are not of the nature of an actual inversion, clearly deviates from that deduced for typical inversions (Fig. 5). The reason for this deviation is yet to be clarified.

The function $A_1 = f(G-PT)$ resulting from Figs. 8 through 10 is suited for deducing, in practical work, from G-PT values the strength of vertical exchange, unless typical inversions are being dealt with in which case the function pursuant to Fig. 5 is to be applied.

V. DETAILED CASE STUDIES

1. Preliminary Remarks

The following case studies are partly based upon the original recordings by cable car telemetry system and pertinent parameter data tables supplied by the computer. The wind profiles are not yet taken into consideration. The graphs and tables are compiled in SECTION X, APPENDIX, 3. They, at the same time, afford a glimpse into our "Analysis Workshop". The numbers of the levels in the recordings correspond to the numbers of levels in the tables. On the other hand, the case studies are based upon a study of changes in time of certain structural elements at the inversion level and pertinent numeric values of functions: $A_1 = f(G-T)$, $A_1 = f(G-E)$ and $A_1 = f(G-PT)$. Separate graphs which are given consecutive figure numbers serve this purpose (SECTION X, APPENDIX, 1). In the graphs on which the values of the aforesaid 3 functions are entered, we have also plotted the "normal progress" of each respective function, in the form of dotted areas, which also, based upon Figs. 3, 5, 7,

indicate the mean scatter. The examples to be discussed hereinafter, are marked with identification letters in alphabetical order. These letters have been transferred to all graphs, recordings and tables (also those in SECTION X, APPENDIX, 3,) so no doubts may arise as to their correlation.

The runs of one and the same day bear encircled numbers in chronological sequence. These, too, are also found on all recordings, graphs and tables, hence the variations of all data in the course of time are readily recognized.

Following are some necessary remarks on the problem of equilibrium. An aerological structure, e.g. the variation of temperature across an inversion layer, and the vertical aerosol profile across the same vertical interval can be in lawful correlation to one another (as shown, e.g. in Figs, 3, 5, 7) only if there had been enough time for the setting of some stationary condition. There are two possibilities:

- a) A temperature inversion is created, very rapidly, say in less than 1 or 2 hours, through some processes which are known in principle. At the moment of its formation, the aerosol profile cannot, as yet, have been accommodated by this inversion to the instantaneous barring effect. The adaptation process requires a few hours; for instance, by evacuation or coagulation of aerosol particles above the inversion and concentration below, until equilibrium has been set based upon the mere eddy diffusion process. Not until then can a lawful correlation exist, and parameters be within the normal ranges of the functions per Figs. 3, 5, 7.
- b) With the barring effect of an already existing inversion being constant in time, the source strength of the aerosol in the layer below the inversion may change, or aerosol may be brought in or drawn off, through advective processes. In this case, again, the equilibrium of diffusion

is temporarily disturbed and time is required for readjustment.

- c) It may happen that aerosol fluxes from some a source into the lower atmospheric layers, where it will spread for the time being. At the upper boundary of the thus developed haze layer, an exradiation of heat begins. This causes continuous cooling and formation of vertical temperature structures which react back upon the vertical aerosol distribution. In this instance, again, the result may be a disturbance of the eddy diffusion equilibrium of long duration.

Other processes are possible which shall not be discussed in this context. The following studies are to show the meteorological conditions under which an equilibrium of diffusion is, as a rule, encountered, or not encountered, and what numeric relations will be found in each respective case.

2. Example A, 5 Oct 1970, Figs 11 through 14

Spontaneous formation in the medium altitude zone of a most sharply defined inversion, which from noon on descends.

Equilibrium of diffusion set from the beginning.

Looking, to begin with, at the original soundings (APPENDIX) we find that at around 1100 CET (run 0) neither T nor TF show any distinctive structure, and this is true of the entire vertical interval. The temperature gradient, however, is weak, and humidity is decreasing with height. The polar conductivities +L and -L are steadily increasing from level 11 to level 5. The causes of their variations above level 5 do not concern us in this report. At 1157 CET (run 1) already, a sharply defined inversion exists immediately below level 6, with humidity decreasing with increasing height, which is clearly reflected in the behavior of +L/-L. In the course of the day the picture of a typical temperature inversion with humidity break

becomes more and more distinct. At the same time the abruptness of the break of $+L$ and $-L$ keeps increasing, and the aerosol concentration below the layer of the break is continually growing (the values of $+L$ and $-L$ are decreasing). Thus the gradient of the aerosol concentration across the layer of the break is getting steeper and steeper. From noon until night the level of the inversion is continuously descending.

Let us now take a look at the detailed analyses in Figs. 11 through 14. Fig. 11 shows changes in regard to altitude and time for:

- a) the layer where the inversion is located ($-G-T_{\max}$);
- b) the level with a minimum value; and
- c) a maximum value of the equivalent potential temperature EPT.

With regard to b) and c) the following must be added: in case of well defined, stable inversions scarcely variable in time or space, or other barring layers, a minimum of EPT is encountered above the barring layer, with a maximum within or below its level. In the absence of barring layers EPT, as a rule, decreases evenly with decreasing height. Thus barring layers normally also identify themselves by an "inversion" in the course of EPT, which, however, proceeds inversely to T . The fact that this EPT inversion (signs: - and = in Fig. 11) is conected together with the temperature inversions proper (sign: $\sqcup$) is clearly apparent from Fig. 11. This figure also reflects the time slope of A_1 . During the descent, to begin with, A_1 increases. In the evening, when the inversion has settled, A_1 drops to a minimum value of only 0.3.

In Fig. 12 we study the variations with time in the range of the function $A_1 = f(G-T)$ within the inversion layer of runs 1 through 4. The variation of the pairs of values occurs exclusively inside the (dotted) normal range. Thus, A during

the descent (run 1 after 2 to 3) increases, it is true, but G-T, too, changes by a corresponding amount as the function shows. Hence, structure of inversion, and control of the vertical aerosol distribution by it are in a true-to-function relation to each other. We arrive at an analogous result if we study Fig. 13 ($A_1 = f(G-PT)$). Except for run 1 the pairs of values are within the normal range of the function. Hence, the behaviour of G-PT and the pattern of the vertical aerosol structure correspond to each other. From these findings we may deduce that equilibrium of diffusion has practically existed from the beginning. No analogous conclusion may be drawn in regard to the relation $A_1 = f(G-E)$, as illustrated by Fig. 14. In this instance all pairs of values are exclusively outside of the normal range. The E-gradients are much greater than would be permissible with regard to the A_1 values. Hence, a change must have occurred in absolute humidity (e.g. caused by advection) above or below the barring layer, which was not related, across the barring layer, to the eddy diffusion process, whereas by contrast the aerosol distribution did obviously correspond to the barring effect of the layer.

Conclusions drawn from this example:

Even in case of spontaneous formation of a very sharply defined inversion, and during a slow and constant descent of the same, the discovered "normal relations" between A_1 and G-T or G-PT, respectively, are being retained. Thus they adequately describe the vertical distribution of aerosol; the state of equilibrium prevailed. A deviation of the values outside of the normal range of $A_1 = (G-E)$ reveals an advective change of humidity. This emphasizes the high value of such detailed analyses. The rule that there has to be a minimum of EPT above the barring layer, and a maximum within its level, is met at all times. It may additionally serve as an indicator of the barring layer.

We are here dealing with a case where rapid formation of an inversion and the transformation of the vertical aerosol structure

went hand in hand together in such a manner, that the condition of equilibrium for the eddy diffusion was met at all times. Conditions for the occurrence of this situation remain yet to be defined.

3. Example B, 18 Nov 1970, Figs. 15 through 19.

Well-defined, spontaneously formed inversion, with height of barring layer and intensity of barring effect varying, Equilibrium of diffusion retained.

The soundings (for a selection pertaining to Example B see X., APPENDIX 3) already show that we are here dealing with a (spontaneously formed) inversion and barring layer which is subject to considerable fluctuations in time as to its structure and height. As demonstrated by Fig. 3 the maxima of -G-T and -G-E are located at the same respective altitude, which level, however, fluctuates considerably. A maximum of EPT above the inversion level and a minimum within the same cannot be found. A_1 , too, fluctuates considerably from run to run, and the rule is confirmed: If the inversion level descends the barring effect of the inversion is temporarily reduced (A_1 is increased), as is demonstrated by runs 3 and 6.

As is recognized from Figs. 16 and 17 the values are fluctuating precisely within the normal range of the function $A_1 = f(G-T)$ and $A_1 = f(G-PT)$. So, with the exception of runs 2 through 4, diffusion equilibrium prevails as in Example A. The change from run 2 via 3 to 4 is indeed striking. Whereas A_1 is changed considerably, G-T and G-PT remain practically constant, the full "normal" scatter of the function is traversed. This means that during the vertical movement of the inversion from 2 via 3 to 4 the aerological structure of T and PT is retained whereas the permeability of the barring layer is reduced temporarily. A local perforation of the layer by turbulences in the early afternoon with relatively extensive insolation, is likely to be the main cause thereof. This example indicates how, by means of such a detailed analysis interfering factors may be isolated.

Considering Fig. 18 we observe that the normal range is considerably exceeded in both directions. The values are not, however, all shifted in one direction as in Example A. This rules out progressive changes of absolute humidity due to advection. Turbulence caused, forced vertical transportation of water vapor is, much rather, to be considered as a cause (transitions from 4 to 5, from 5 to 6). It is, however, a striking observation that the transition 2-3-4 again appears to be out of place: While the water vapor gradient remains constant, A_1 varies considerably. This means again that here a variation of the inversion had temporarily been forced, pointing toward interference with the system from the outside (e.g. heat irradiation). It is interesting that in this case a variation of the vertical aerosol structure was triggered (as a consequence whereof a higher A-value resulted for run 3), not however one of the aerological structure and of the vertical distribution of water vapor.

For the present example we have plotted the vertical profile of the incremental exchange coefficients A_1 and of the water vapor from run to run as a special representation in Fig. 19. This representation does not only show very impressively the variation in the fine structure of the A_1 profiles but also the vertical displacement of the layers having maximum barring effect (A_1 minima). What we want to demonstrate, however, is that just these sometimes very thin layers of highest resistance to eddy diffusion are indicated in the E profile by a peak-like dip of water vapor concentration. This is true even in the case of a splitting into two barring layers with a small difference in height (run 6). At the same time we recognize a problem in the E profile for run 3: Here it is not so much a dip but much rather a sharply confined increase of E with height that is characteristic. Such findings with the aid of detailed analyses are conducive to understanding deviations from normal behavior observed from time to time.

The gathering of such exceptions and the investigation of the conditions under which they occur is important for the following reason: In order to reduce, on a pertinent basis, the scatter in the statistical investigations it is necessary to introduce a more detailed subdivision. Most of all, the complex causal nexus of the various parameters may gradually be cleared up in this manner.

Conclusions drawn from the example:

In the case of spontaneous formation of an inversion and intensive vertical movement of the same, the discovered "normal relations" between A_1 and G-T or G-PT, respectively, are retained and depict the vertical aerosol distribution which corresponds to a stationary setting of equilibrium. A "breaking out" of the values from their normal distribution, insignificant for G-T and G-PT, and quite considerable for G-E, points to a convectively forced variation in the vertical aerosol distribution, which can no longer be lawfully related to the aerological structure. The vertical water vapor gradient, in this case, behaves differently from the vertical aerosol gradient. With the barring layer sharply descending, A_1 is temporarily increased. Barring layers are always indicated by conspicuous deviations of the E profile across very small vertical intervals.

4. Example C, 14 October 1970, Figs. 20 through 23

Sharply defined inversion with predominantly insignificant vertical variations but considerably fluctuating barring effect. Vertical aerosol gradient in stationary equilibrium on basis of aerological structure.

Example C concludes, for the time being, those cases of inversions which are distinguished by the vertical aerosol gradient being in equilibrium on the basis of the given aerological structure, in other words, where the normal relations of our function are fully or largely met.

As is shown by the original recordings in APPENDIX 3 we are, from run 1, dealing with a sharply defined inversion the strong barring effect of which is clearly expressed by the profile of the air conductivities. Only early in the morning (run 0) an inversion is just slightly indicated, with only a very weak barring effect. The vertical movements can be clearly recognized on the original recordings, e.g. by comparison of run 6 with run 8.

To begin with, the variation as a function of time and altitude is shown in Fig. 20. In the forenoon we find that the ascent of the inversion layer from run 1 through run 3 results in a weakening of its barring effect (increase A_1 ; in Examples A and B we had found a weakening with a descent of the inversion), whereas later in the afternoon the opposite relation is found. Hence, vertical movement, no matter in which direction, may lead to either a weakening or a strengthening of the barring effect. We are thus compelled to restrict our statement to the more generalized form: Vertical movement of a barring layer results in a change to its barring effect. The nature of the individual relation is certainly depending upon the causative reasons for the movement of the inversion (advective processes, anticyclonic descent). In order to determine these, further detailed analyses will be needed. In the context of this discussion of the examples the implications of microscale and mesoscale meteorological conditions and processes are not to be discussed in any detail. A deepening of the detailed analyses in this direction is in process.

In the instant case we again find a minimum of EPT above the barring layer, but no marked maximum below.

Figure 21 shows the variations with time of our values relative to function $A_1 = f(G-T)$. We notice that the fluctuations of the values largely remain inside the normal range. An insignificant breaking out (deviation from equilibrium) occurs from 5 through 6 to 7. This means that in Fig. 20, too, we must

consider the temporary descent of the inversion level (run 6) as the expression of an interfering process. At any rate, the fact is again conspicuous that A_1 changes considerably, e.g. from run 5 to 6, although the aerological structure of T has been retained (G-T constant). The attenuation of the aerosol gradient encountered at the same time therefore must be due to an exterior cause, e.g. perforation of inversion by local turbulence. It is striking, however, that, as Fig. 22 indicates, the variation of the values is largely within the normal range of the function $A_1 = f(G-E)$, at least the transition 5-6-7 is within. Again, aerosol and water vapor behave differently in the case on a disturbance of equilibrium. It is surprising, however, that (see Fig. 23) the values for $A_1 = f(G-PT)$ are almost exclusively outside of the normal range of this function. What conclusions have to be drawn therefrom is yet to be investigated, using all available meteorological data.

Conclusions drawn from this example:

Vertical movements of the barring layer result in variations of their barring effect and a disturbance of the diffusion equilibrium. What direction such a variation may take cannot necessarily be predicted, until such time that the causative reasons for the interfering process are known. It appears certain that detailed analyses of the demonstrated kind will lead to recognition if the causes of interference of micro-scale and mesoscale meteorological conditions are included in future investigations. For because three basic function are available, based on G-T, G-E, and G-TE, it is, as a rule, possible from the agreement of the individual values with two of these to isolate the one which is most of all affected by the disturbance of equilibrium. This may later lead to isolation of the interfering cause.

5. Example D, 11 December 1970, Figs. 24 through 26.

Inversion at constant height. Vertical aerosol gradient not in equilibrium on basis of aerological structure.

In Figs. 24 through 26 we are studying the case of cyclic variations (runs 1 through 4) of the value obtained, in the range of functions $A_1 = f(G-T)$, $A_1 = f(G-E)$, and $A_1 = f(G-PT)$. All data obtained from these runs are far outside the normal ranges of the three functions. Vertical movements of the barring layer did not occur. Convective interference, also, must be excluded. To what extent persisted advective processes were determining in this case in the formation of this exception remains to be investigated. For it is striking that this low-lying inversion has been formed through the inflow of near-ground polluted cold air. But although this near-ground cold air layer was soon settled a state of equilibrium did not come about.

To begin with, our example shows that even with inversions at rest and seeming freedom from interferences we must be prepared for extraordinary results. If it were not possible to recognize and isolate these (e.g. on the basis of detailed analysis) such values would result in a considerable scattering within the data material. However, the investigation of the causative relations must not be neglected, particularly in such unclear cases. Exactly these may supply valuable insights.

6. Example E, 17 October 1970, Figs. 27 through 30.

Case of two simultaneously existing inversions one of which is suddenly formed newly and the other disintegrated. State of equilibrium is not attained.

Here we are dealing with a relatively complex case. It is remarkable for the only one initially existing inversion and barring layer descending and disintegrating. Prior to its disappearance, however, a second inversion and barring layer

was forming about 400 m above which also descends and goes on existing to the end - if with a weak barring effect. From the picture of the original recordings (X., APPENDIX, 3), layer I owes its formation to a layer of haze actively conveyed to the spot at little short of 1500 m a.s.l. (minima of +L and -L during run 1). In the following we consider the original recordings along with Fig. 27. This initially striking aerosol structure is reduced to a mere trace during run 2, but during run 3 it is suddenly sharply defined again at 1200 m a.s.l. During runs 3 through 6 layer I is disintegrating or hazily merging with layer II which has meanwhile descended and whose existence was first hinted at during runs 3 and 4. Presumably the transition was as indicated with dots in Fig. 27. As is apparent considerable instability prevails in the present case, impeding the setting of a stationary equilibrium of diffusion, as will also be seen from Figs. 28 through 30. First we cast a glance at the variation with time of A_1 in Fig. 27. We note that with a rapid descent of both I and II the barring effect is reduced (A_1 is steeply increased, runs 2 and 3 for layer I; runs 3 through 5 for layer II). A regeneration of the barring effect for both layers occurs when the descent has come to rest (runs 4 through 6 for layer I; runs 7 and 8 for layer II).

We now consider the position of the data relative to the normal ranges of our three basic functions. In doing so we go from the following sequence of values:

Runs 1 through 6 at level of layer I; runs 7 and 8 at level of II, pursuant to the merger of both layers as indicated (dotted) in Fig. 27.

The poorest agreement is found relative to function $A_1 = f(G-PT)$ in Fig. 28. With regard to the G-PT values the exchange coefficients are too low. i.e. the barring effect appears "too strong". This above all applies to runs 5-6-7. But this is the very vague transition phase from I to II with disintegration of II. It is not surprising, then, if during this time interval

an equilibrium situation does not prevail. The positioning of the individual values relative to the normal distribution is already notably better in Fig. 29 ($A_1 = f(G-E)$). Surprisingly, the above mentioned transition phase of runs 5-6-7, of all things, is not out of line here: water vapor gradient and aerosol gradient are here apparently triggered by the same atmospheric process. By contrast, the values of runs 1 through 3, i.e. during the descent of layer I, are outside of normal distribution: The barring effect of the layer appears to be stronger, during this time interval, for the aerosol than for the water vapor. With regard to function $A_1 = f(G-T)$, however, - see Fig. 30 - all values of our example (excepting run 5) are within its normal range. From this it would be concluded that with regard to vertical aerosol distribution, approximate diffusion equilibrium must have temporarily prevailed, after all.

Conclusions drawn from this example:

Descent movements of barring layers are coupled with a temporary reduction of their barring effect. The latter is regenerated when the vertical movement has come to rest. In case of a double inversion with disintegration of the lower layer and the suggestion of merger with the descending upper one, the diffusion equilibrium is temporarily disturbed. A discrepancy develops between vertical water vapor and aerosol distribution. A minimum of the equivalent potential temperature above the barring layers, and a maximum near the barring layers were not found which also points toward disturbed equilibrium conditions.

7. Example F, 30 September 1970, Figs. 31 through 34

Clearly defined inversion descends slightly, remains at constant level for some time, finally ascends steeply with simultaneous disintegration. Largely set equilibrium of diffusion.

Interest is particularly held by the phases of ascent and disintegration of the inversion or barring layer. For this reason,

we have represented only runs 4 through 7 by their originals and tables in X. APPENDIX,3. As with the previous runs (from which a descending movement of the inversion is recognized), an increase of temperature and decrease of humidity with height within the inversion zone is also very distinct with run 4. Just one hour later (run 5), however, the T-inversion has temporarily disappeared whereas the humidity decrease was retained, although this structural detail has risen somewhat. During run 6, a T inversion with humidity decrease is again recognized, if at a considerably higher level than before. Hardly any structures are left for recognition during run 7, traces are derived from the tables.

In Fig. 31 we consider the variations with altitude and time. In the morning the well defined inversion descends by about 100 m. During this descending movement the A_1 values are relatively high. With the barring layer becoming stationary, however, A_1 reaches minimum values between no more than 0.6 and 1.5. At 1300 hours an ascent commences, immediately triggering a steep increase in the A_1 values which persists until disintegration of the layer. It is remarkable that in instant case, again, a minimum of the equivalent potential temperature is found above the barring layer (100 m - 200 m vertical interval), with a maximum approximately at the level of the barring layer as such. Only in the case of vertical movements of the barring layer, the EPT maximum, too, is above the barring layer (runs 2 and 7).

Fig. 32 shows that the individual values (2 through 4) are in just barely satisfactory relation to the normal range of function $A_1 = f(G-T)$ only during the time that barring layer has nearly or fully settled. But during the vertical movement they are far outside thereof (1 to 2; 4 to 6).

The two pairs of values for a very low-lying near-ground inversion in the morning ("low"), which rapidly disintegrates thereafter, are within the normal range. The values of our example

are quite satisfactory with regard to the normal distribution of $A_1 = f(G-E)$, according to Fig. 33. However, we notice a contrast to Fig. 32: While the inversion has become stationary the values fall insignificantly outside of the normal distribution. The vertical aerosol gradient turns steeper, temporarily, than would correspond to the vapor pressure gradient, whereas agreement prevails during the vertical movements. The values in Fig. 34 are completely outside of the normal distribution of $A_1 = f(G-PT)$. We find a cyclic sequence with extreme deviation during the stationary phase of the barring layer. No explanation can be offered therefor, as yet.

Conclusions drawn from this example:

It was confirmed again: Vertical movement (no matter whether up or down) of the barring layer results in an increase of its permeability, its becoming stationary results in a strengthening of the barring effect. The A_1 values of the example are in satisfactory agreement with functions $A_1 = f(G-T)$ and $A_1 = f(G-E)$, slight deviations from the former are found during the vertical movement, from the latter during the stationary phase. The superpositioned minima or maxima, respectively, of EPT are found to exist throughout, even during the vertical movements.

8. Example G, 28 September 1970, Figs. 35 through 38.

Considerable vertical and structural variations of an inversion. Equilibrium of diffusion is not attained.

This example concludes our discussion of cases with vigorous vertical movements of barring layers. The original runs in APPENDIX, 3 already show that not only does a distinctive vertical unrest prevail but also the pattern of the structures is changing rapidly. We immediately pass on to Fig. 35. The variations with time of A_1 are particularly remarkable. Steep increases of A_1 are invariably found when - in accordance with our earlier experience - the level of the barring layer

is rapidly changed. This is very drastically demonstrated by the transition from run 2 to 3: A_1 is increased by two orders of magnitude and remains persistently high during the subsequent ascent movement (3 to 4). Not until the stationary condition of the barring layer exists during runs 6 and 7, does A_1 arrive at a minimum. Renewed disquiet from runs 7 through 10 causes A_1 to increase again.

The vertical displacements of the various structural elements, according to Fig. 35, present an almost confusing picture: True, minimum and maximum of EPT are still always found and in not too great vertical distance from the inversion level (as a maximum a few 100 m, as an average 200 m), but the maximum is sometimes above, sometimes below the barring layer. Even the respective strongest gradients of T and E are not - as usual - close to each other. Vertical differences up to 150 m are encountered. Besides, maximum G-T is sometimes above, sometimes below maximum G-E.

Due to the sometimes extremely poorly developed temperature gradients Fig. 36 contains only very few individual values. As was to be expected, they are far from the normal distribution of function $A_1 = f(G-T)$.

It is all the more surprising that, in Fig. 37, the individual values fit very well into the normal distribution of $A_1 = f(G-E)$. Although most certainly the setting of an eddy diffusion equilibrium was not possible in our instance the variations with time of the aerosol gradient (expressed in terms of the behavior of A_1) and of the water vapor gradient (G-E) are in good agreement with each other. The interferences, apparent from the continuous structural changes of the aerological pattern and from the vertical variations, thus took effect, in Example G, upon aerosol and water vapor distribution analogous to and on the basis of, the normal relation of $A_1 = f(G-E)$. In Fig. 38, the individual values of our example are far from the normal distribution of $A_1 = f(G-PT)$: We find a cyclic sequence for runs

2 through 7 and another one from 7 through 10. We find it interesting that the first cycle, if defined only from 2 through 6, comprises exactly the first ascent, the second cycle from 7 through 10 the second ascent. It appears to be possible to describe, with the aid of these cyclic "movement graphs", the structural variations connected with processes of ascent or general vertical displacement, and later to improve our understanding of them functionally. At any rate, in retrospect we may say that the cycles in the following earlier figures were connected with vertical movements of the barring layer: Fig. 34 (runs 2 through 6, initial lingering, then rapid ascent with disintegration of the inversion); Fig. 30 (runs 4 through 8, initial descent, then merger with lower disintegrating inversion). An exception relative to all three normal functions is formed by Example D with its cycles. True, there is no vertical movement here, but a clearly cyclic loosening up of the inversion structure (runs 1 through 5) by turbulences in the course of the day. So it will be necessary to pay still more attention to such graphs with a cyclic pattern, in the future, along with further meteorological parameters.

Conclusions drawn from this example:

In the case of violent structural and vertical changes of the barring layer eddy diffusion equilibrium cannot come about. It is, however, possible that variations of the vertical aerosol profiles and water vapor profiles are running in conformity with each other and are jointly controlled by the interferences determining the complex pattern. Again it is shown that vertical movements of the barring layer are connected with a reduction of its barring effect. During vertical movements a cyclic sequence of the values is found relative to the normal distribution of the function $A_1 = f(G-PT)$.

9. Example H, 26 September 1970. Figs. 39 through 43.

Analysis of (a) the upper boundary of the exchange layer, and (b) an ascending inversion located below.

With this example we leave the analysis exclusively of inversions or similar structures analogously belonging thereto. We consider parameters from the zone of the upper boundary of the exchange layer. However, in the present case there is also an inversion which we shall include in our considerations. The inverse temperature gradient is only rather poorly defined, however.

To begin with we go back to the original recordings in X. APPENDIX, 3. The position of the upper boundary of the exchange layer there has been marked with E, that of the inversion with I. In the profile of +L (partly also in that of -L) the increase by increments with height is recognized at the two levels so marked. Until evening (run 6), however, the structures disintegrate to a large extent. What clearly remains is the +L increment in the zone of the loosened-up structure of E.

In Fig. 39 we consider the variations in time and space of the two formations. From the value tables to Example H it is seen that the exchange layer is marked by a minimum and a subjacent maximum of the equivalent potential temperature EPT similar to what we had found with stationary inversions as a rule. The altitudes, of these structural details are entered in Fig. 39. Whereas in the course of the day the upper boundary of the exchange layer perpetually descends, the - weakly defined - inversion ascends in increments, above all from run 5 to 6. The rule found confirmed for stationary inversions, according to which EPT shows a minimum above the inversion, and a maximum below it, applies in the present case, too. These EPT structure details follow suit with the vertical movement of the barring layer proper.

Now a few words on the variation with time of the A_1 values, considering Fig. 39: The barring effect of the inversion is

the greatest in the morning, and steadily decreases in the course of the slow ascent of the layer (A_1 is growing). Another deterioration of the barring effect occurs in the evening as the layer ascends again. By contrast, the diffusion resistance through the upper boundary of the exchange layer increases by increments (A_1 decreases) as the latter descends. In conjunction with the commencing disintegration of both layers in the evening, consistently the A_1 values increase.

Due to the very scantily developed inverse temperature gradient a study with respect to function $A_1 = f(G-T)$ is foregone. Fig. 40 merely reflects the travelling times for the runs, and the symbols for the exchange layer (E) and inversion (I).

On Fig. 41 we note that the individual values for I are within the normal range of function $A_1 = f(G-E)$, even so when $G-E$ is positive (viz. when water vapor pressure is increased with height). The values of E, too, just barely meet these requirements, however, no dependence upon the water vapor gradient is apparent. The values for $A_1 = f(G-PT)$ in Fig. 42 are entirely outside of the normal distribution for this function. This is independent of whether they pertain to the exchange layer or to the inversion.

In Fig. 43 we have plotted the vertical profiles of A_1 for runs 3 and 6. These once more very clearly illustrate the following processes:

- a) descent and increase of barring effect of the upper boundary of the exchange layer;
- b) ascent and decrease of barring effect of the inversion. finally approximate mutual touching of the two layers.

Conclusions drawn from this example:

The upper boundary of the exchange layer, too, is characterized by a minimum of EPT plus a subjacent maximum of EPT. During descent of the upper boundary of the exchange layer its barring effect toward eddy diffusion is increased.

10. Example I. 17 September 1970, Figs. 44 through 48.

Upper boundary of an exchange layer. Slight vertical movement.

In conclusion, we shall discuss the case of an upper boundary of the exchange layer which was sharply defined throughout the day. The original runs in APPENDIX, 3 very convincingly demonstrate the following within the zone of this layer: (a) an increase in temperature; (b) a decrease in humidity; and (c) a sharp increase in conductivity, i.e. a decrease of the aerosol concentration. These vertical variations are very distinct during runs 1,3,4; during run 2 the structure are loosened up, and the level of the layer has descended somewhat. In the following figures we have plotted, for comparison, the values of the main levels separately from the values of the special levels.

In Fig. 44 we first consider the variations with time and with altitude. To begin with it is to be pointed out that this level, too, is distinct for sharply defined minima of EPT above it, and corresponding maxima within its level. These EPT structural details follow suit in the slight vertical movement of the upper boundary of the exchange layer. During the temporary descent of the layer (run 2) A_1 arrives at a maximum, i.e. movement reduces barring effect as with a true inversion.

Fig. 45 shows that all individual values are outside of the normal range of function $A_1 = f(G-T)$. If we concentrate on the main levels only, $G-T$ changes in the course of the day, it is true, but this has little impact on A_1 (maximum during 2, see above). If we consider the special levels we obtain a cyclic progress which is surely coming close to the true conditions and corresponds to our experience gathered so far on the occurrence of cyclic interrelations.

In Fig. 46, using the values from the special levels we find a time slope which, it is true, keeps to the border line of the normal range of our function $A_1 = f(G-E)$ but in agreement with

the latter's direction. If we lean on the values from the main levels only, A_1 remains unaffected by the variations of G-E.

An analogous conclusion is drawn from Fig. 47. If, however, we start from the values of the special levels we are again met with a cyclic pattern, although all the values are far from the normal distribution of function $A_1 = f(G-PT)$.

In Fig. 48 we have plotted the vertical profiles of A_1 for runs 2,3,4. The stepped curves are based upon the values from the main levels. The A_1 values from the special levels within the barring layer level are indicated (solid) and identified by numerals in circles. It is seen how across 1000 m from the barring layer downwards, consistently high A_1 values are encountered which is appropriate for an exchange layer. The loosening-up of the structure of the barring layer is clearly recognizable.

Conclusions drawn from this example:

Values derived from the special levels afford a much deeper insight into the true interrelations. The upper boundary of the exchange layer is again marked by EPT minima and maxima, one above the other with intervals of 150 m up to 300 m (EPT maximum within the levels of the layer). These structural details join in any vertical movements of the barring layer. Vertical movements lead to a weakening of the barring effect of the layer. The vertical movement is expressed in cyclic pattern or part of the values relative to the normal ranges of our main functions.

VI. THEORETICAL SUPPLEMENT FOR COMPUTATION OF EDDY DIFFUSION
COEFFICIENTS

In (1.2) only the asymptote solution to the stationary combined diffusion (D) and coagulation (K) equation

$$D \frac{\partial^2 N(z)}{\partial z^2} - K N(z)^2 = 0 \quad (1)$$

was used for the calculation of the eddy diffusion coefficient D (pages 25 et sequ. of (1)), i.e. that solution which extends to infinity. This presupposition without which the equation (1) is not analytically solvable applies strictly speaking only in that case where D does not vary with height. Where diffusion coefficient D_i differ from layer to layer, the partial solutions $N_i(z)$ to (1) must, for reasons of continuity, be so joined together that the diffusion current $J = -D \partial N / \partial z$ is a steady one. If $N_i(z)$ and D_i designate the nuclei concentration and diffusion coefficient in the vertical interval from z_i through z_{i+1} , the following must then apply:

$$D_i \frac{\partial N_i}{\partial z} (z = z_i) = D_{i-1} \frac{\partial N_{i-1}}{\partial z} (z = z_i) \quad (2)$$

to all i 's.

Under these premises the equation (1) is no longer analytically solvable. The general solution or its inverse function, respectively, has the following form:

$$z - z_0 = \pm \int_{N_0}^N \frac{dx}{\sqrt{(N_0')^2 + \frac{2K}{3D} (x^3 - N_0^3)}} \quad (3)$$

wherein N_0 and N_0' designate the nuclei concentration and its derivative at position z_0 . In the source-free stationary case, N must decrease with height, and the minus sign on the right side of (3) applies.

Solution (3) represents an elliptic integral which is solvable only through methods of approximation. In addition, this is a variation problem. Therein, D in the integrand must be selected so that both members of the equation (3) are in agreement with each other. The derivative $N'(z)$ needed for the joining condition (2) are easily obtained from the inverse function $z(N)$:

$$N' = \sqrt{(N_0')^2 + \frac{2K}{3D} (N - N_0)^3} \quad (4)$$

The equations (2) through (4) must be solved by layers wherein it would be best to use, for the top layer, the asymptote solution, compute therefrom the diffusion current at the lower border, and use the latter to solve the equation (3) for the next lower interval. This operation is to be repeated as often as necessary.

It is true that in principle this problem is solvable with the aid of our Hewlett-Packard desk computer with extended memory. However, this procedure requires such a high expenditure of computation effort that it would hardly appear applicable to the routine analyses. At present, a computer program is being developed for this purpose, with the aid of which it is being studied on a few selected examples what size deviations are permissible from the results of the procedure now applied in a routine manner. It is noted that due to the wide range of variation of D it will not be necessary to strive for a maximum of exactitude.

The procedure can always be used on any cases of particular interest in order to obtain more exact results.

The results of such comparisons will be reported and discussed in the next report.

VII. CONCLUSIONS

So much is certain that it would not have been possible to carry out either the statistical investigations or the detailed analyses on individual examples, without the preceding change-over of the routine evaluations and data processing to a largely electronic basis. It was only this measure that resulted in a sufficiently comprehensive mass of homogenous data from all available soundings for statistical purposes, on the one hand, and in a complete stock of all necessary parameters of the most varied kind per sounding for the detailed analyses, on the other. In practical processing of the electronically obtained data it was revealed that the most recently developed procedure of electronic data processing fully meets requirements.

The statistical analyses in logical continuation of the earlier research work revealed that there is a significant relation between the highest inverse temperature gradient and the incremental exchange coefficient, providing that there are stationary inversions with a set eddy diffusion equilibrium. The scatter of this relation is unimportant. The relation applies to an A_1 range from 0.1 to 300. Under the same conditions (stationary situations, barring layers) a significant relation may be derived, too, between the highest positive gradient of potential temperature and the incremental exchange coefficient. It applies to the A_1 range from 0.8 to 100. The scatter is not great, more data must be added.

The third of the three derived main functions for describing the behavior of A_1 in structural zones (inversions, barring layers, upper exchange boundary), is based upon the vertical gradient of water vapor pressure. Again, a significant relation

was found with relatively insignificant scatter of values. It is valid for the A_1 range from 1 through 100.

Finally it was attempted to discover more general interrelations which would apply quite generally, independently of structural details (inversions, barring layers). In doing so, another relation was found to stand the test: $A_1 = f(G-PT)$ whose scatter is surprisingly small and whose range of applicability is between $A_1 = 1$ and 1000.

Such main functions which are to have more general validity can be derived only if certain criteria of selection are applied. Such criteria, however, can grow only from sufficient experience. With increasing experience, however, the criteria must be improved, supplemented or, if need be, essentially changed, each of which means a reprocessing of the data material. This work, this is understood, can be performed only in a step-by-step manner.

In developing criteria of classification and selection, the detailed analyses proved to be especially helpful. Systematic work according to this method has only just begun. It already resulted in a series of valuable results. Some of these are: A minimum of the equivalent potential temperature is found 100 m or not more than 300 m above an inversion or barring layer resting in eddy diffusion equilibrium, and a maximum is found within the level of the layer proper. Within the zone of the barring layer proper the gradient of potential temperature reaches a maximum, likewise the negative gradient of the vapor pressure. If the temperature gradient is positive, it assumes a minimum, if it is negative, a maximum. As a rule, a minimum of the vapor pressure, narrowly limited in space, is also found within the level of a barring layer. If barring layers get into vertical movement their barring effect is reduced for the duration of such movement. After the barring layer becoming stationary its barring effect may be regenerated. Temporary interferences leading e.g. to a vertical movement of the barring layer or

triggering a loosening-up of its structure due to locally limited convection are usually expressed in cyclic patterns of the time slope of the essential parameters in the level of the barring layer.

The above discussed detailed analyses which, in the form presented, represent only an initial step, are calling, in order to be fecundated, for the use of all available microscale and mesoscale meteorological parameters (radiation, wind profiles, cloudiness) and their variations with time.

Although it is absolutely necessary to further increase the data material, to further develop and improve the scientific procedures of analysis, still some results are available even today which are possible of practical application by the meteorologist. This applies most of all where it is necessary, from the profiles of aerological data as supplied by radiosondes, to draw conclusions, quickly and directly, upon the conditions of vertical propagation of suspended matter.

VIII. OUTLOOK ON ADVANCED STUDIES IN THE NEAR FUTURE

1. General

Since the methodism for recording the exchange conditions up to 3000 m a.s.l. is now completely commanded and fully developed, it would be the obvious next step to expand the investigations at least up to and including the tropopause level. This expansion is further encouraged by another fact:

A program is being carried under contract with US AEC at our institute, for the investigation of stratospheric-tropospheric exchange using the tracers Be7, P32, S35 plus fall-out heavy-metals, the concentration of which is recorded at Zugspitze station for 24 hours at a time (see (3)). The tropopause structures are analyzed with the aid of radio-sonde data from the

Atlantic Ocean and Europe, and attempts are made to calculate the trajectories.

So the following situation exists:

- a) With the aid of cable-car telemetry systems we gain a precise insight into the exchange conditions obtaining between 0.7 and 3 km a.s.l.
- b) With the aid of the above-mentioned tracers we can survey transportation processes from the stratosphere down to the upper troposphere.

If it were possible, now, by means of appropriate additional methods to also cover the exchange conditions between upper troposphere or, even better, between lower stratosphere and 3 km a.s.l., we would have the chance to completely investigate the overall exchange from the stratospheric deposits at 10 or 20 km down to the surface of the earth at 0.7 km a.s.l. Such an expansion of our research program would increase our yield of findings by a multiple in regard to the vital problem of the atmospheric aerosol exchange.

2. Procedures to be applied

a) Lower Troposphere

The procedures per 1) - 5) will be applied without restriction during the continuation period.

b) 3 km a.s.l. up to lower stratosphere

The only one but the most effective and useful method to present itself is a combination of Rawinsonde with Lidar. Under good weather conditions and with a powerful Lidar, aerosol structures as far as the Junge layer at approx. 15 km may be covered and observed in time. The aerological data necessary for the scientific analysis of the aerosol movements are supplied by a special Rawinsonde of our own. The technical requirements for

such an essential and methodic expansion are presently being reviewed. Already now, some preparations are being made, such as the construction of a balloon filling hall adjacent to the new institute building (for radiosonde balloons), the construction of which has just begun, and the erection of a Lidar observation and set-up room on the roof platform of the institute. Negotiations with suppliers have been conducted and technical details agreed on.

IX. REFERENCES

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Contract AT (30-1)-4061, Oct. 1970,
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X. APPENDIX

1. Figures 1 - 48

2. Legends to the Figures 1 - 48

3. Individual Runs and Data Tables
Examples A - J

4. Tables I - XII

RaB concentration measured on the Zugspitze (Z), the
Wank (W) and in Garmisch-Partenkirchen from July 1970
to June 1971

5. The Computer Program

for computation of all necessary values from the
original runs. Tables of results as examples, see 3.

1. Figures 1 - 48

Fig.1

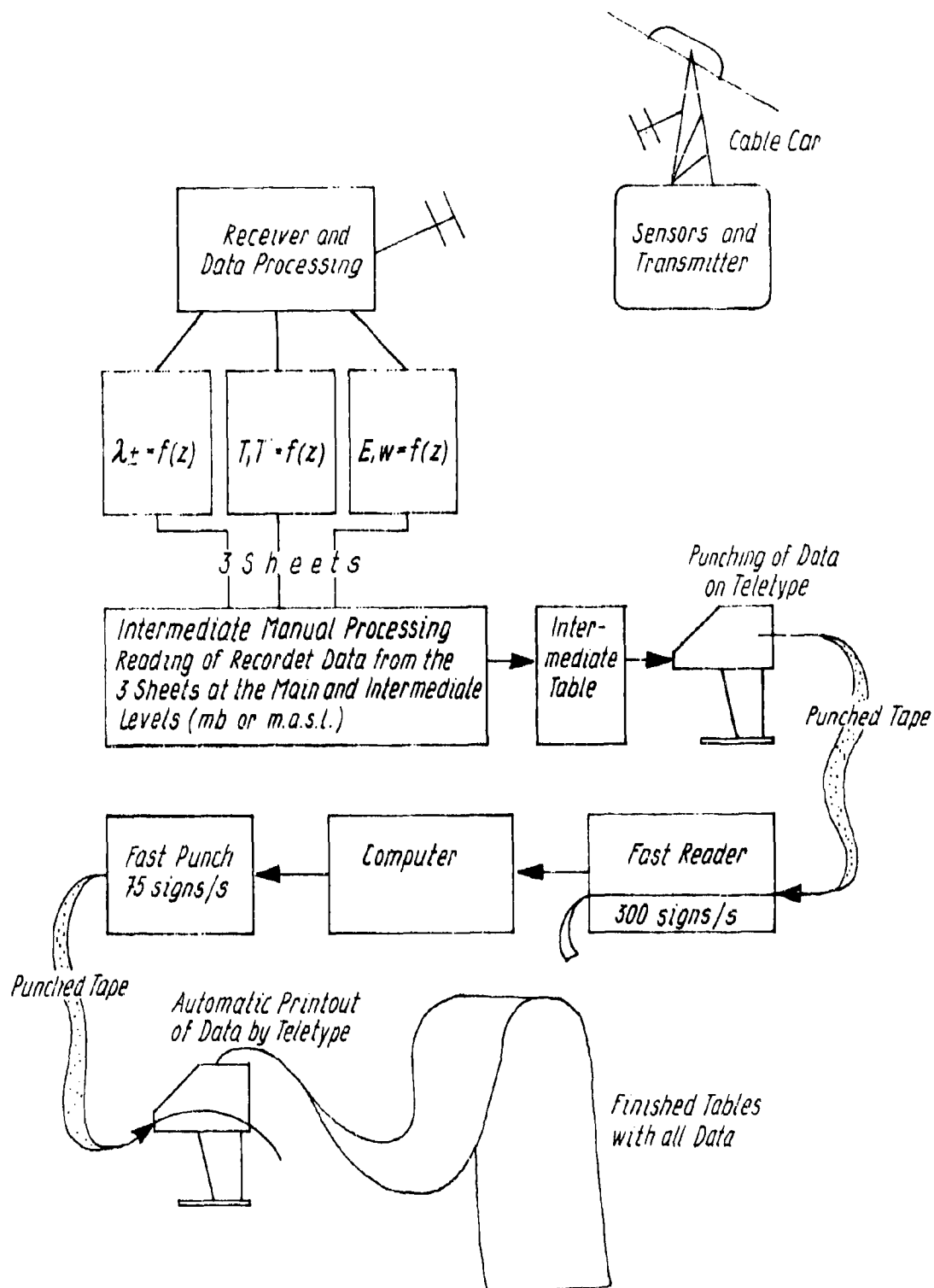


Fig. 2

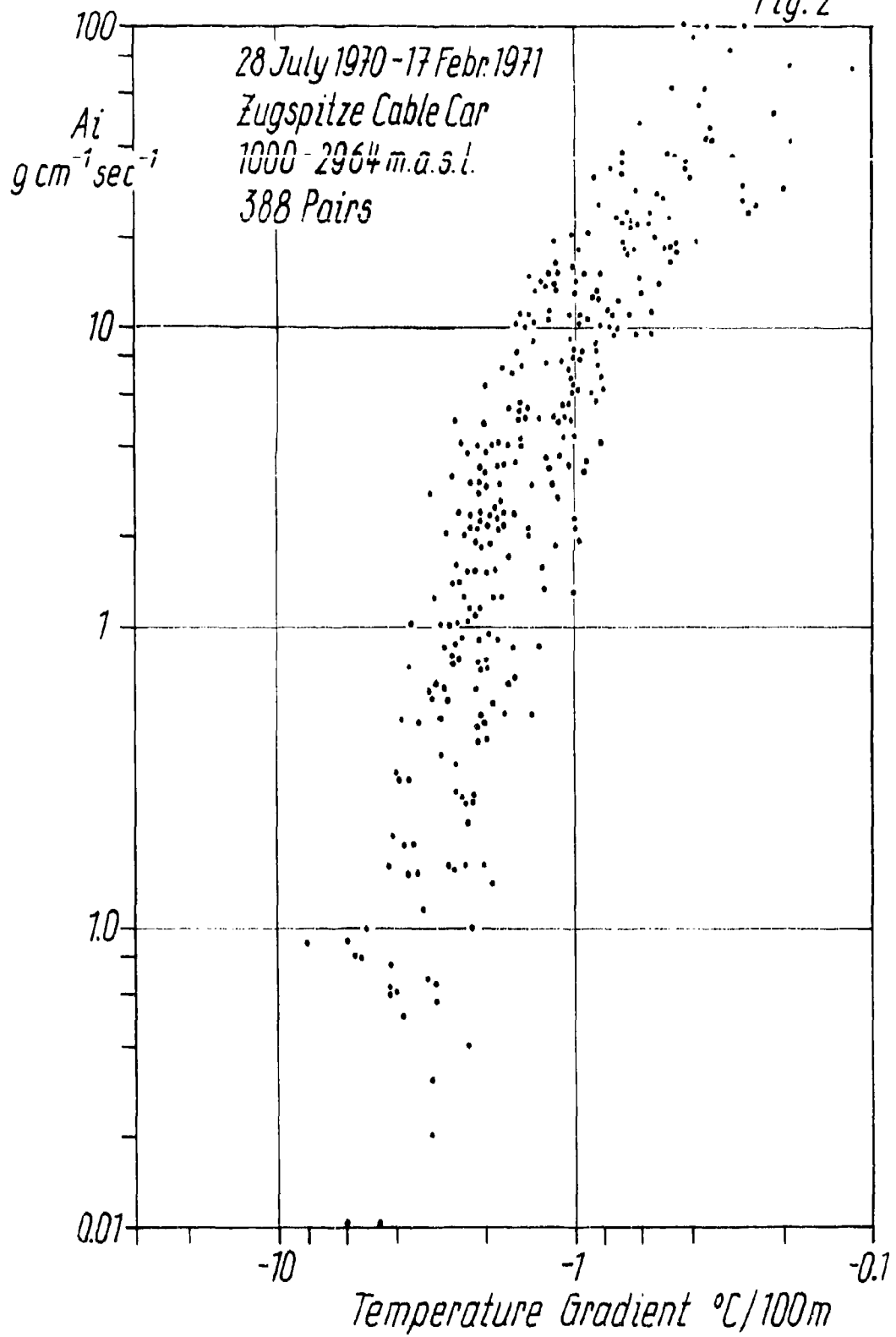


Fig.3

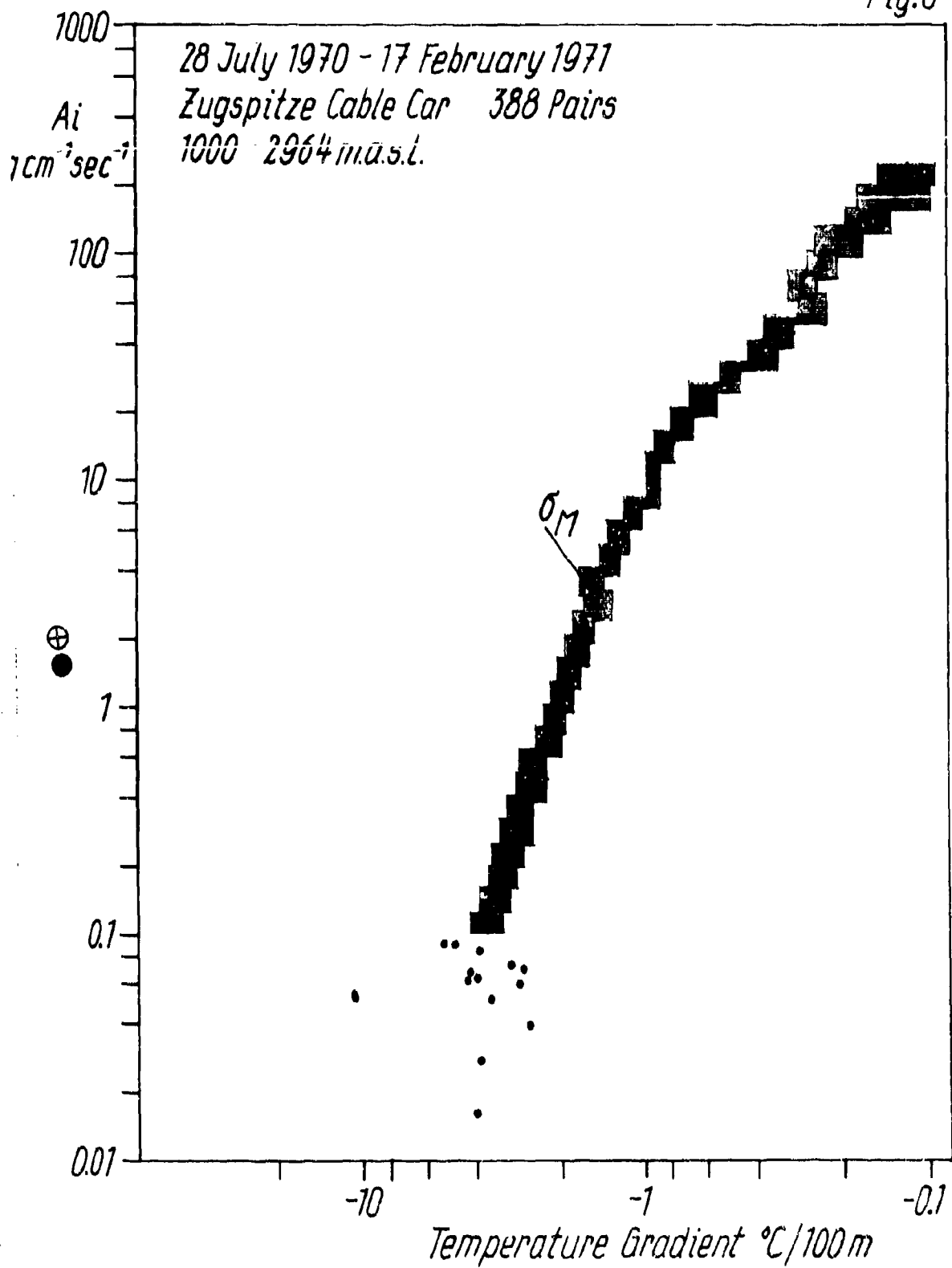


Fig 4

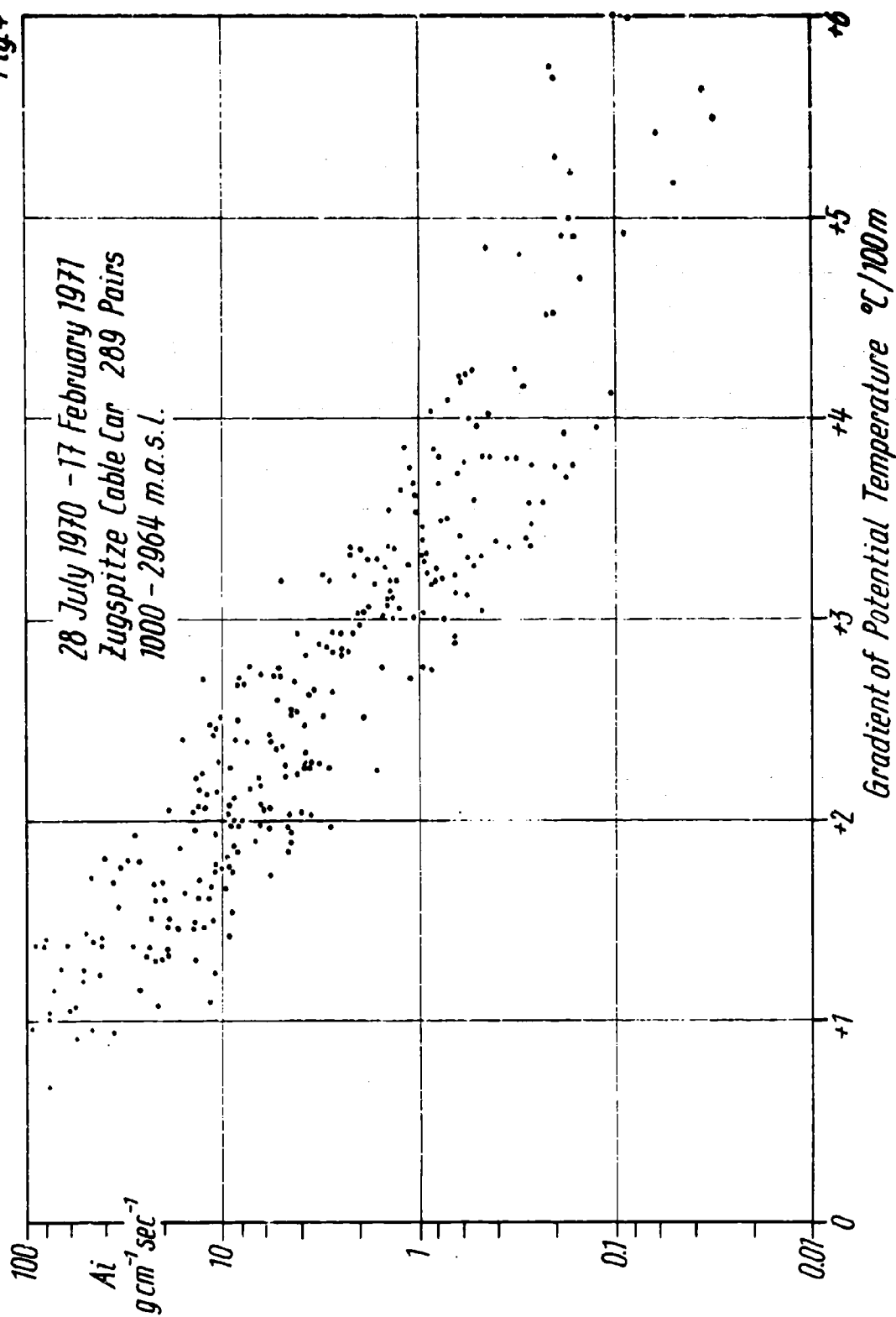
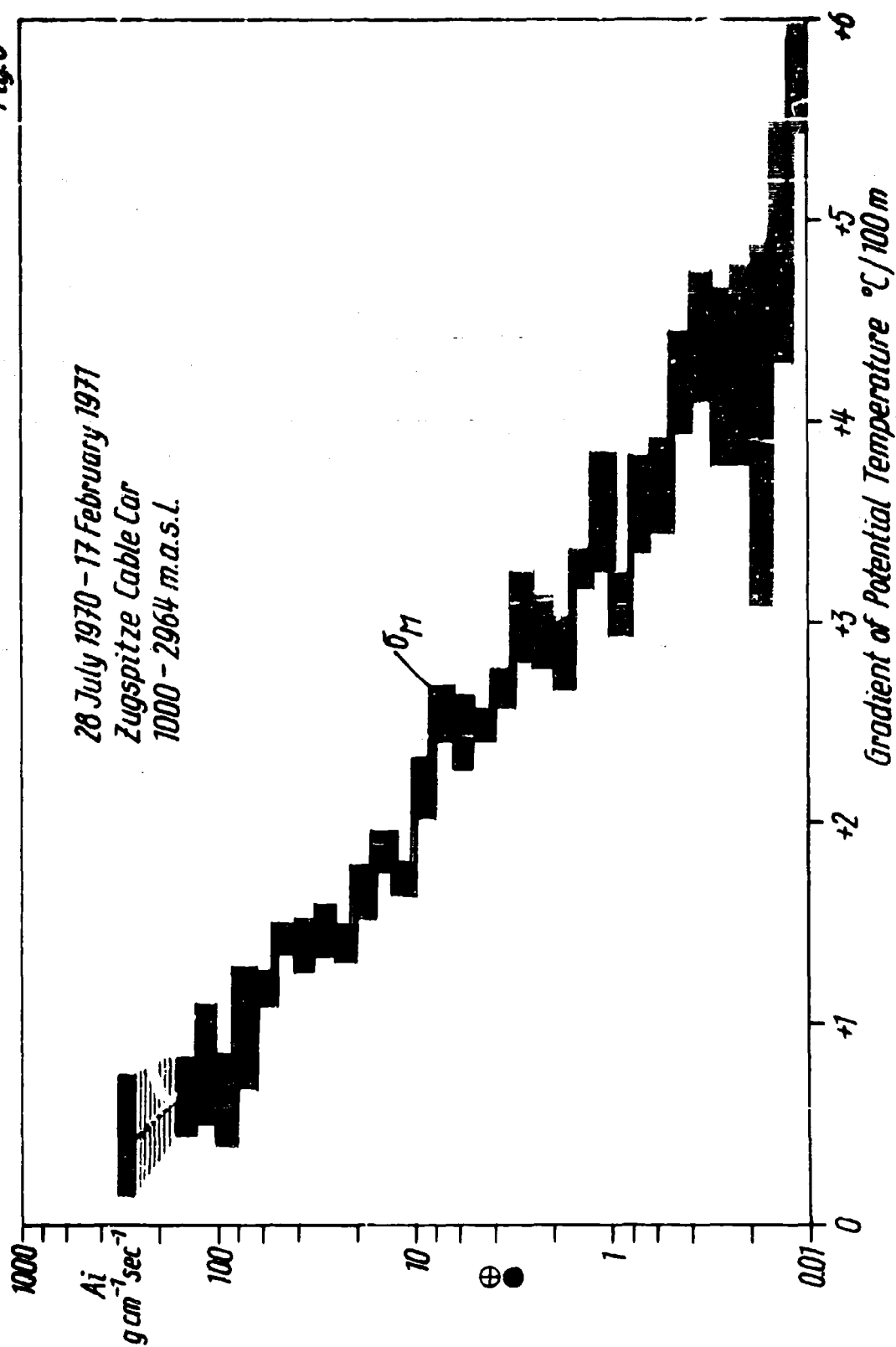


Fig.5



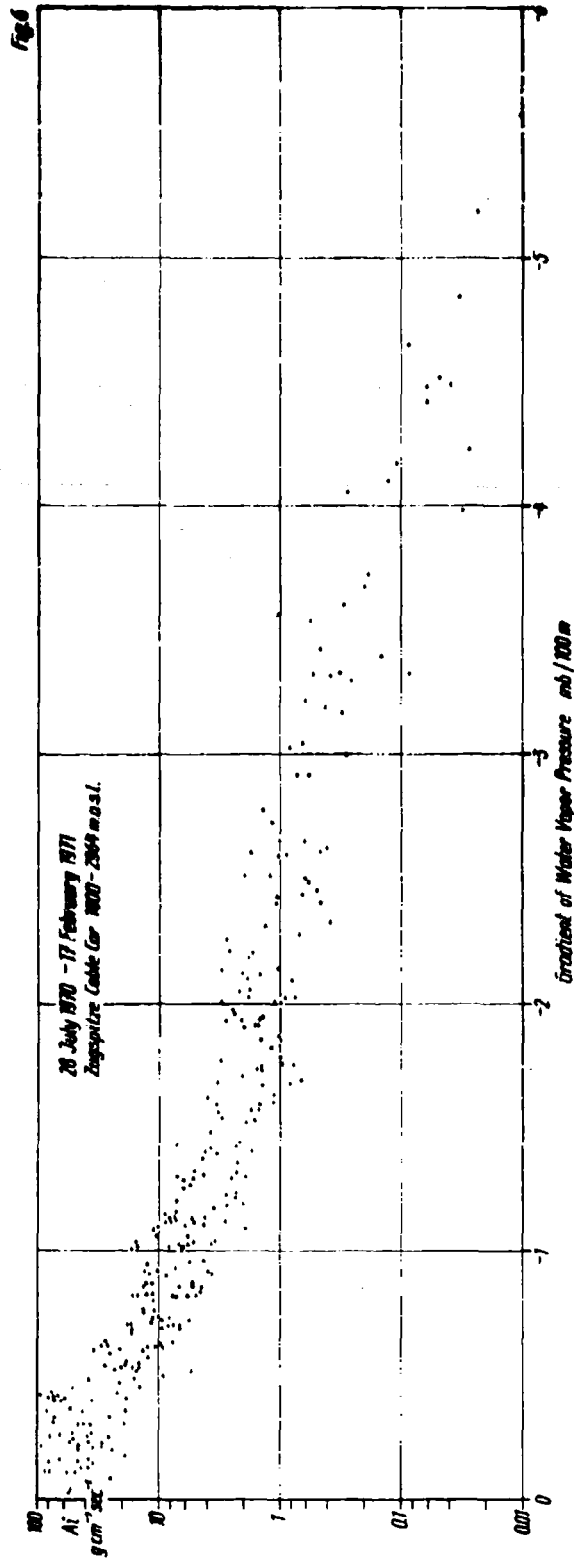


Fig. 2

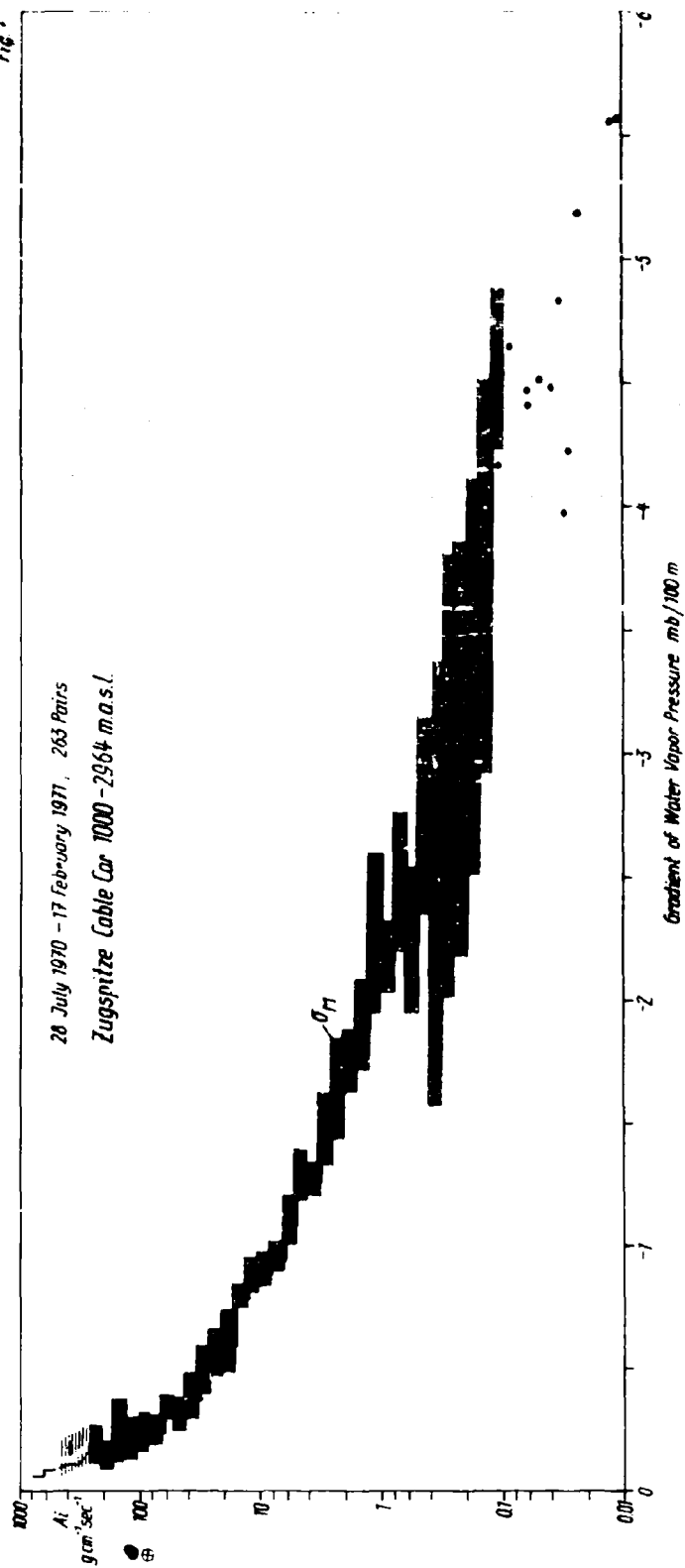
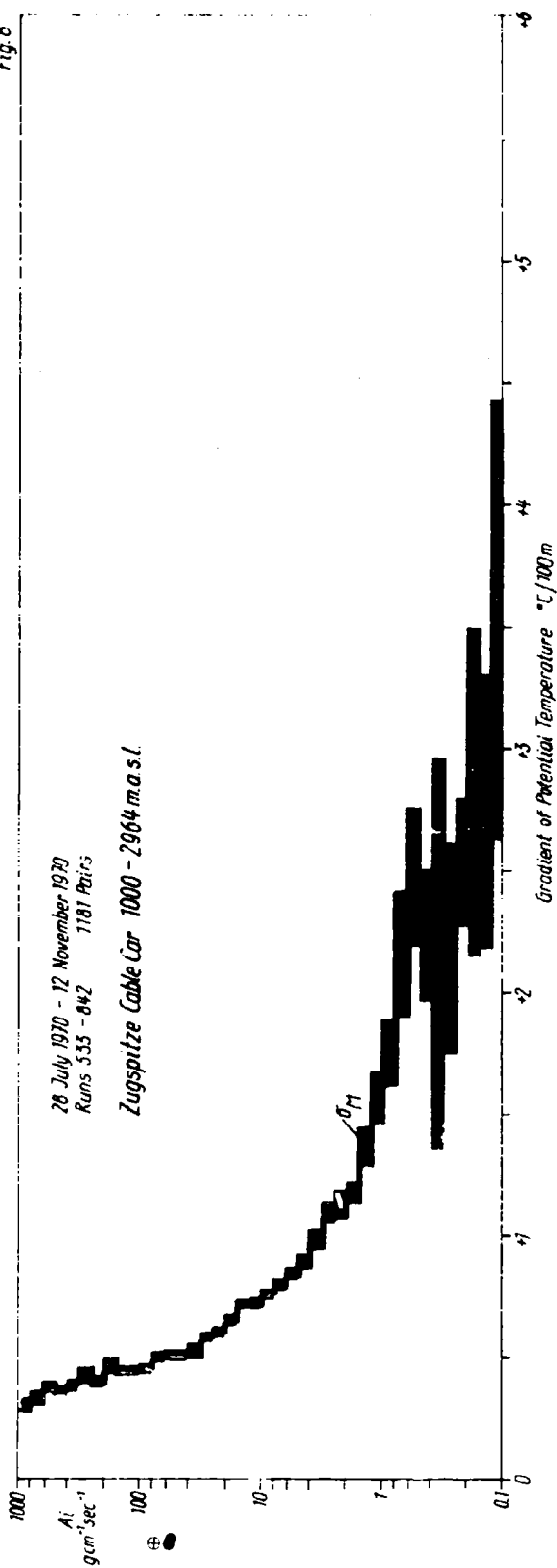


Fig. 8



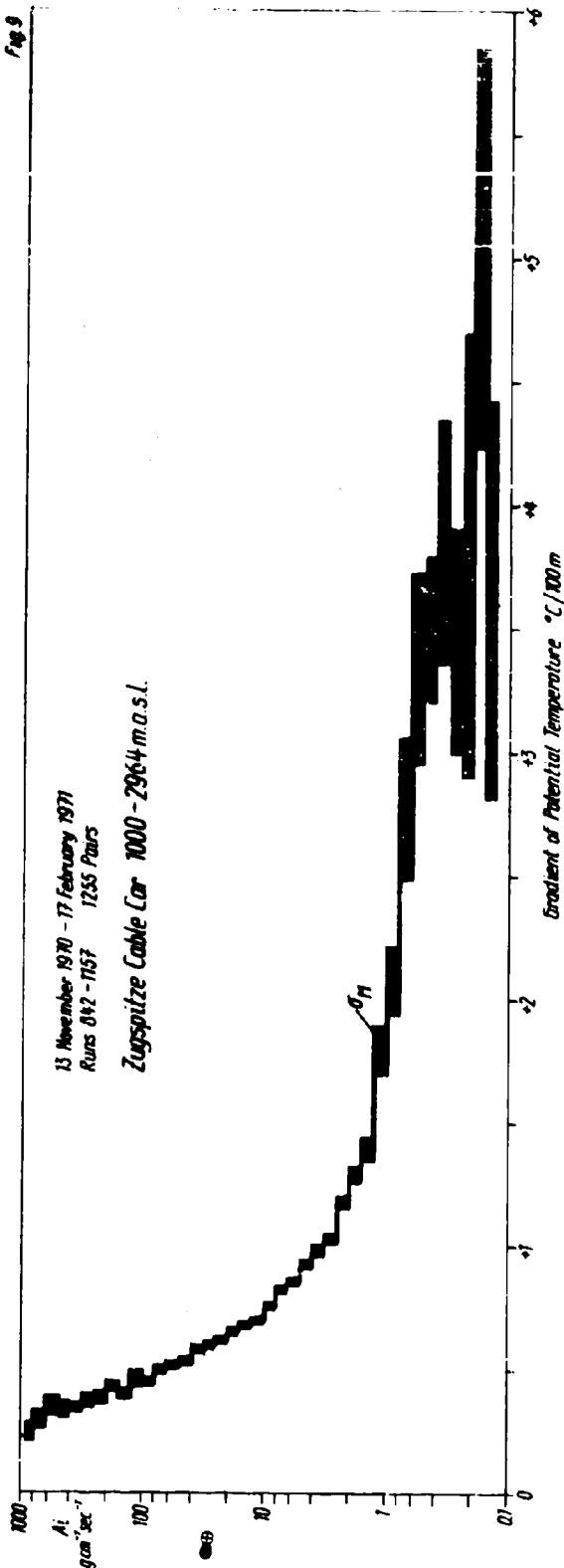
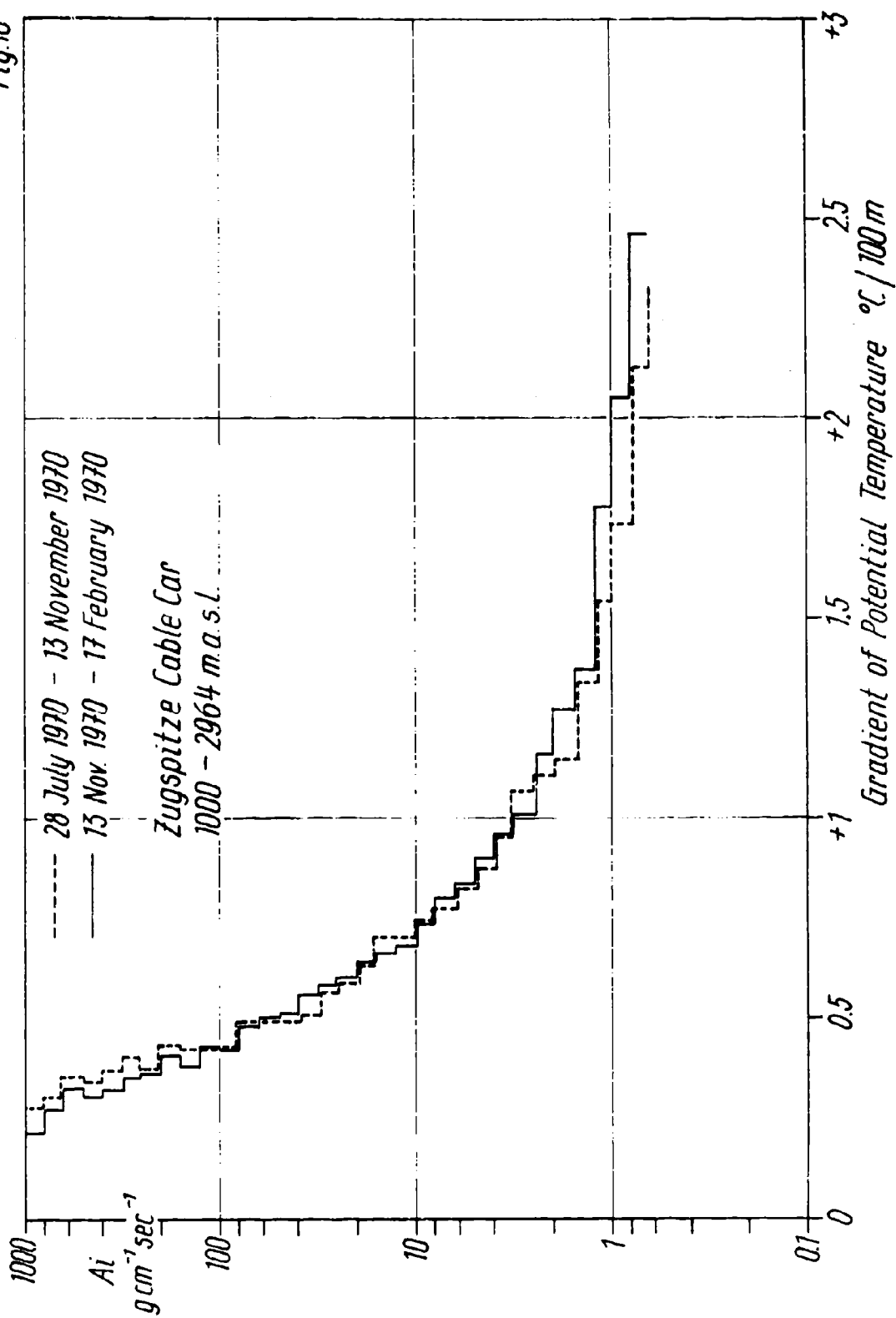
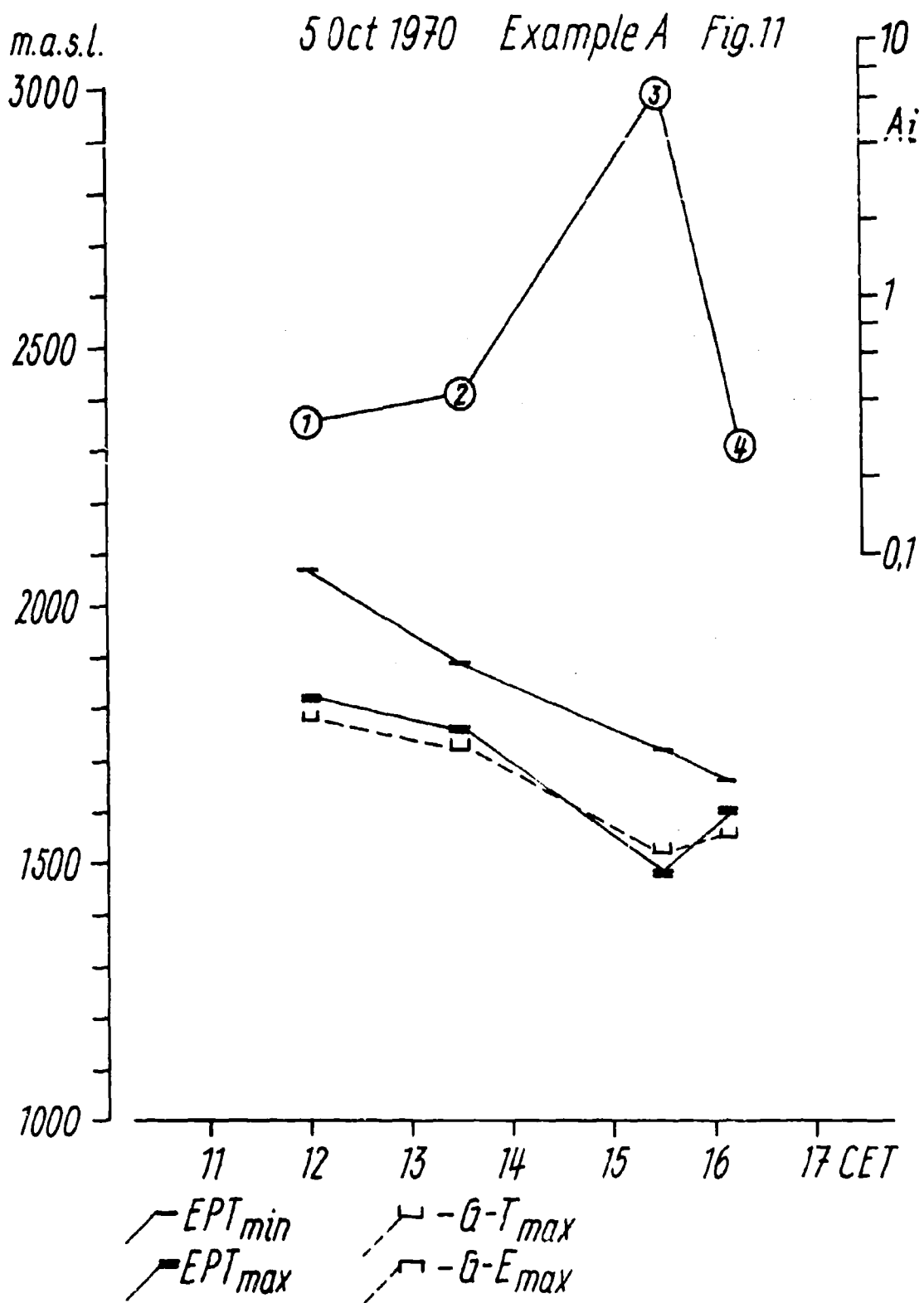
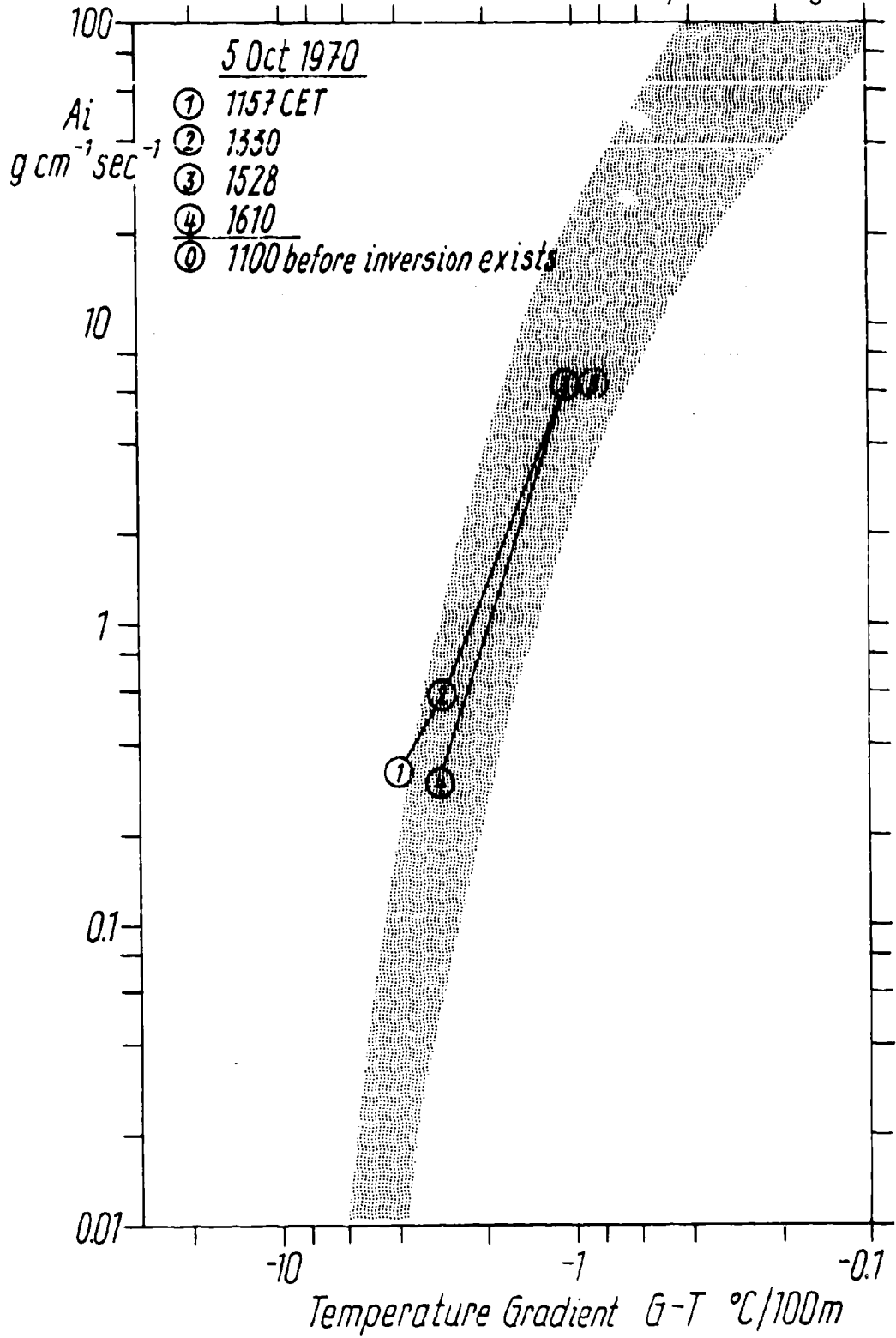


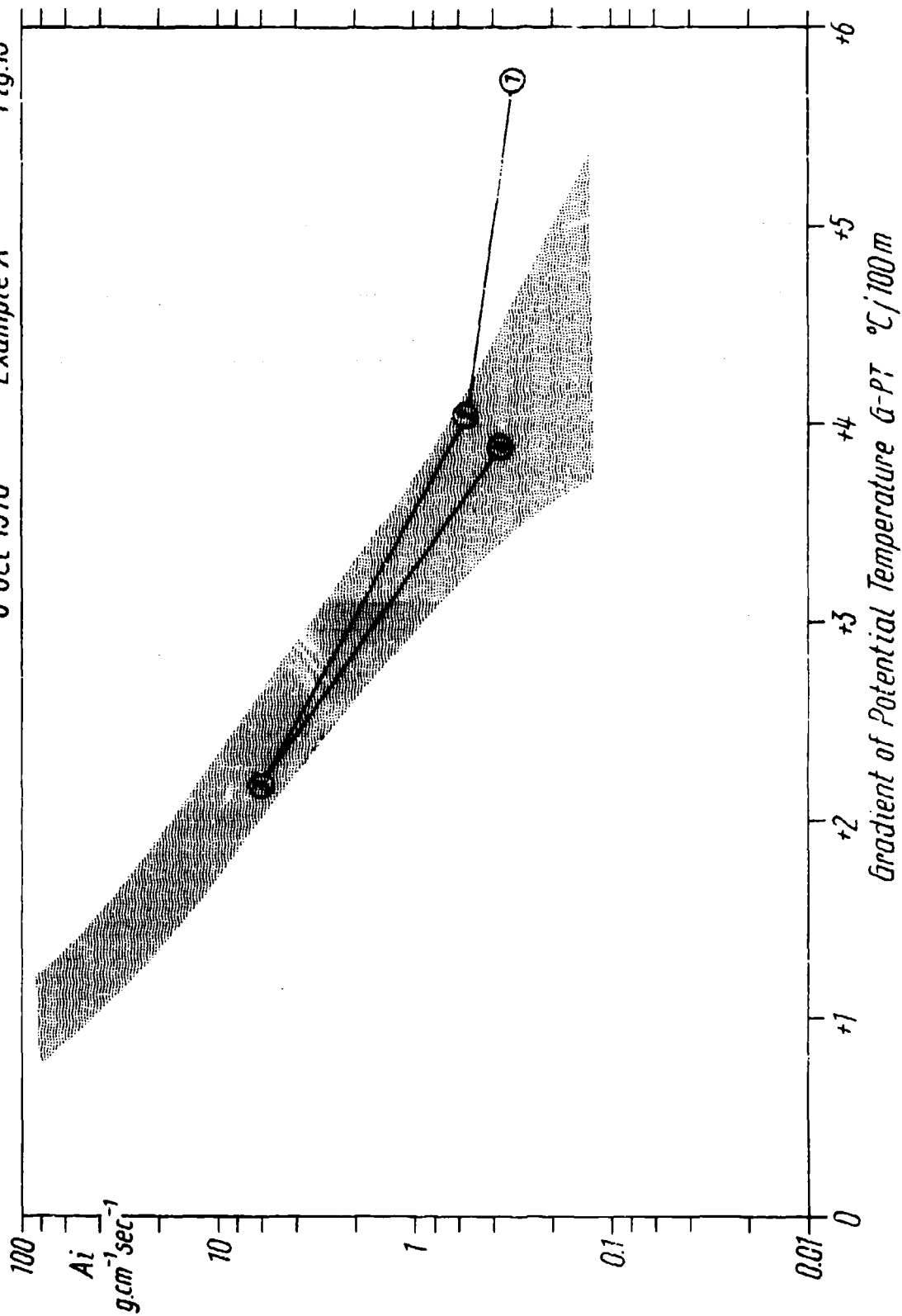
Fig. 10

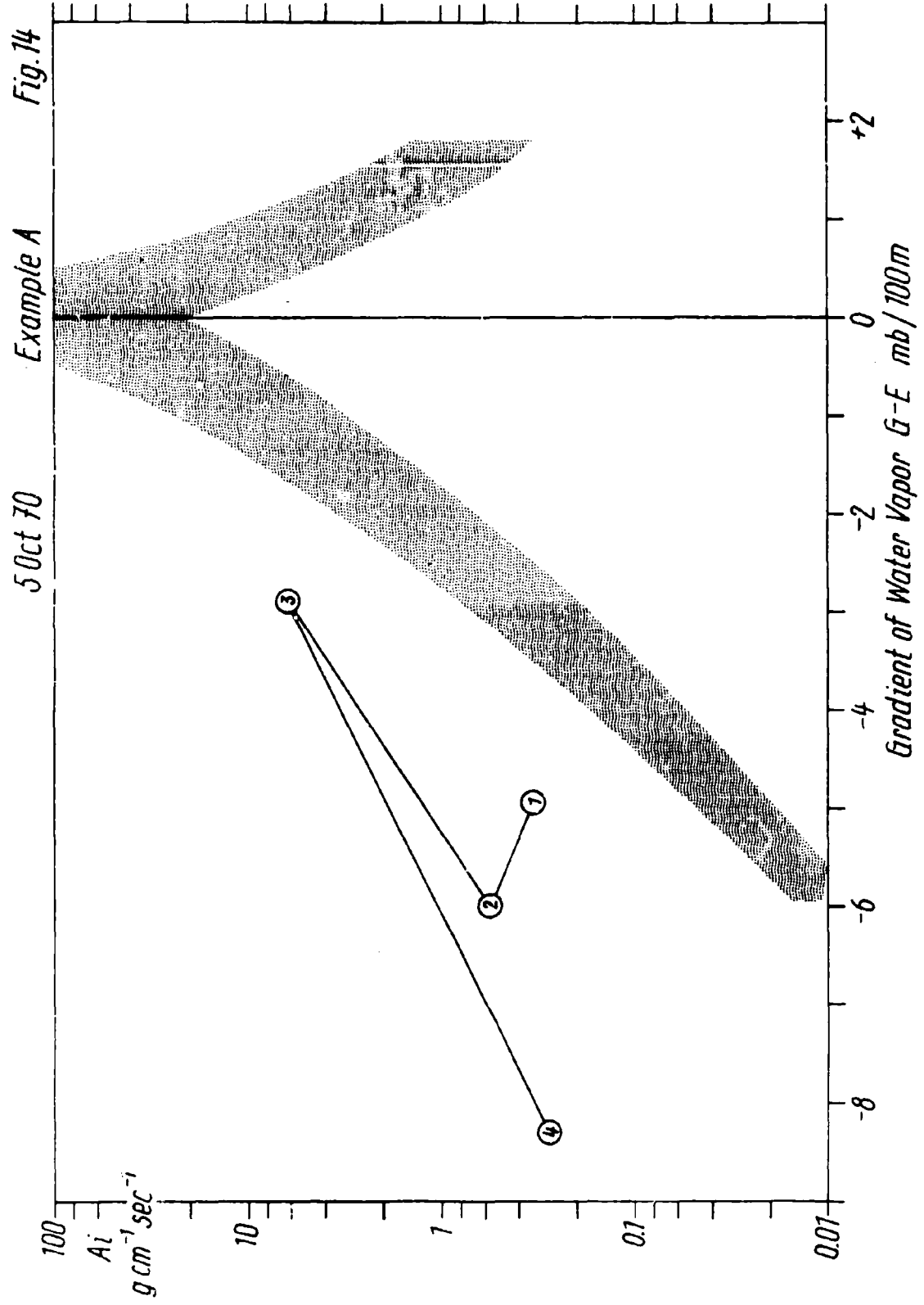




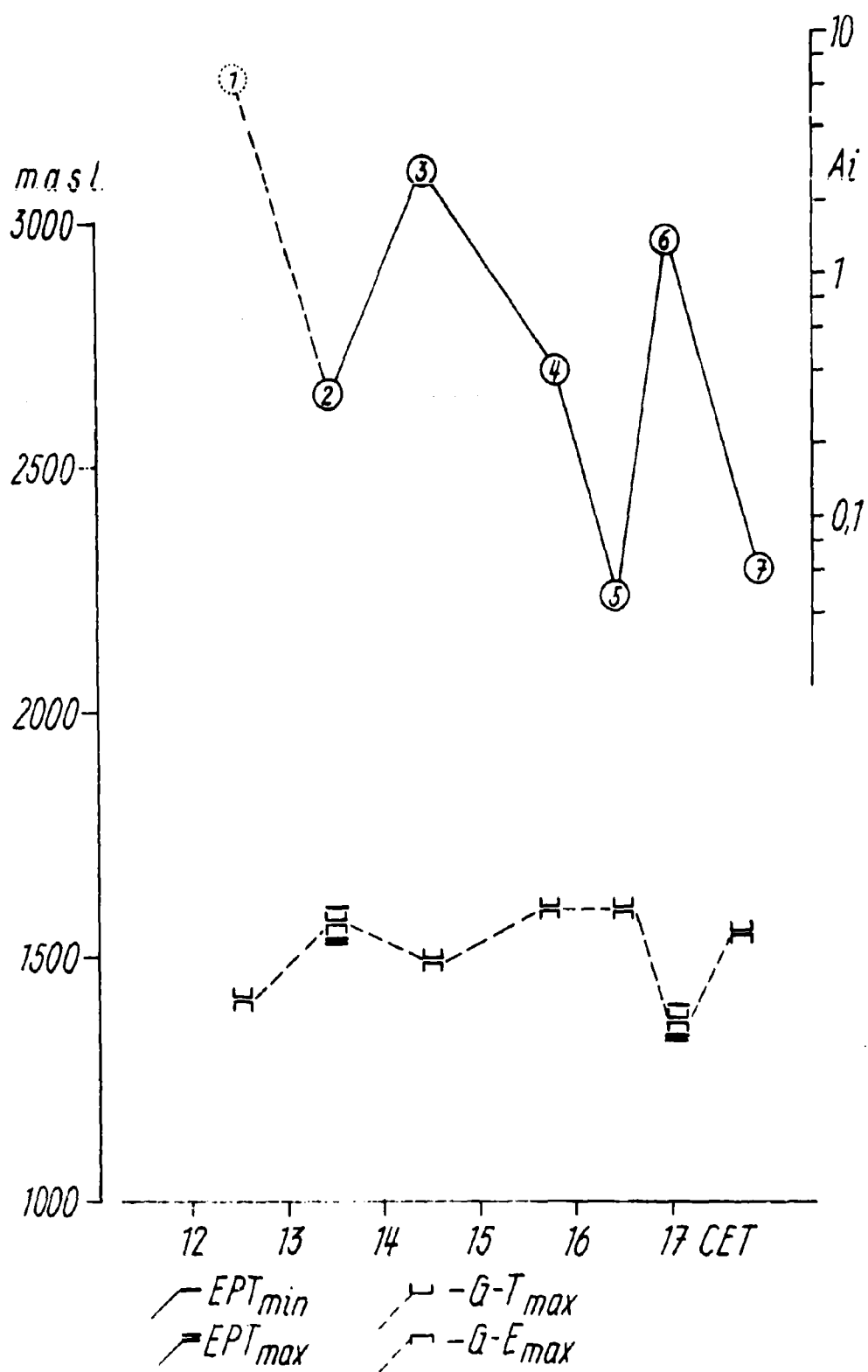
Example A Fig.12



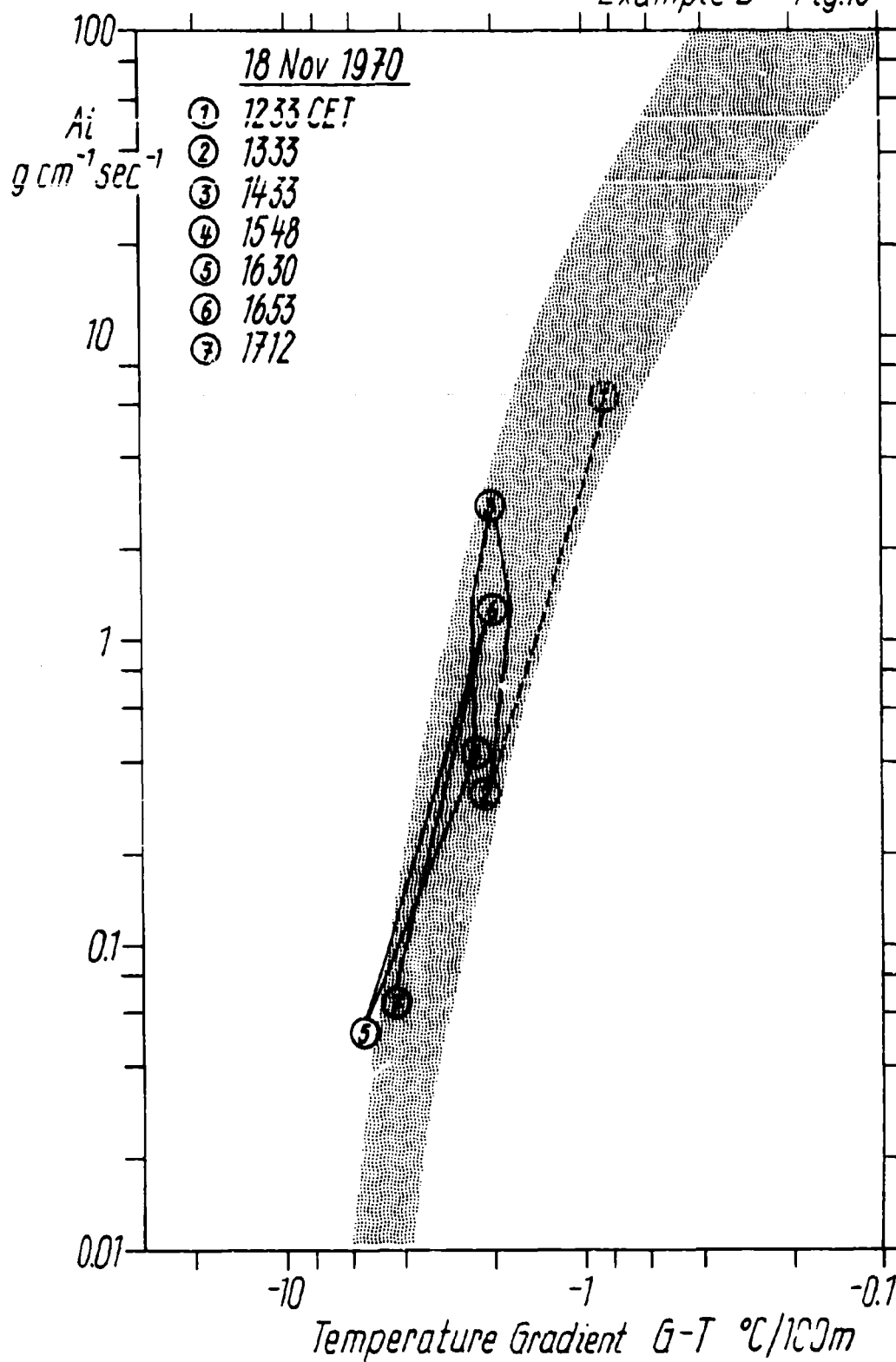




18 Nov 1970 Example B Fig.15



Example B Fig.16



18 Nov 1970

Example B

Fig.17

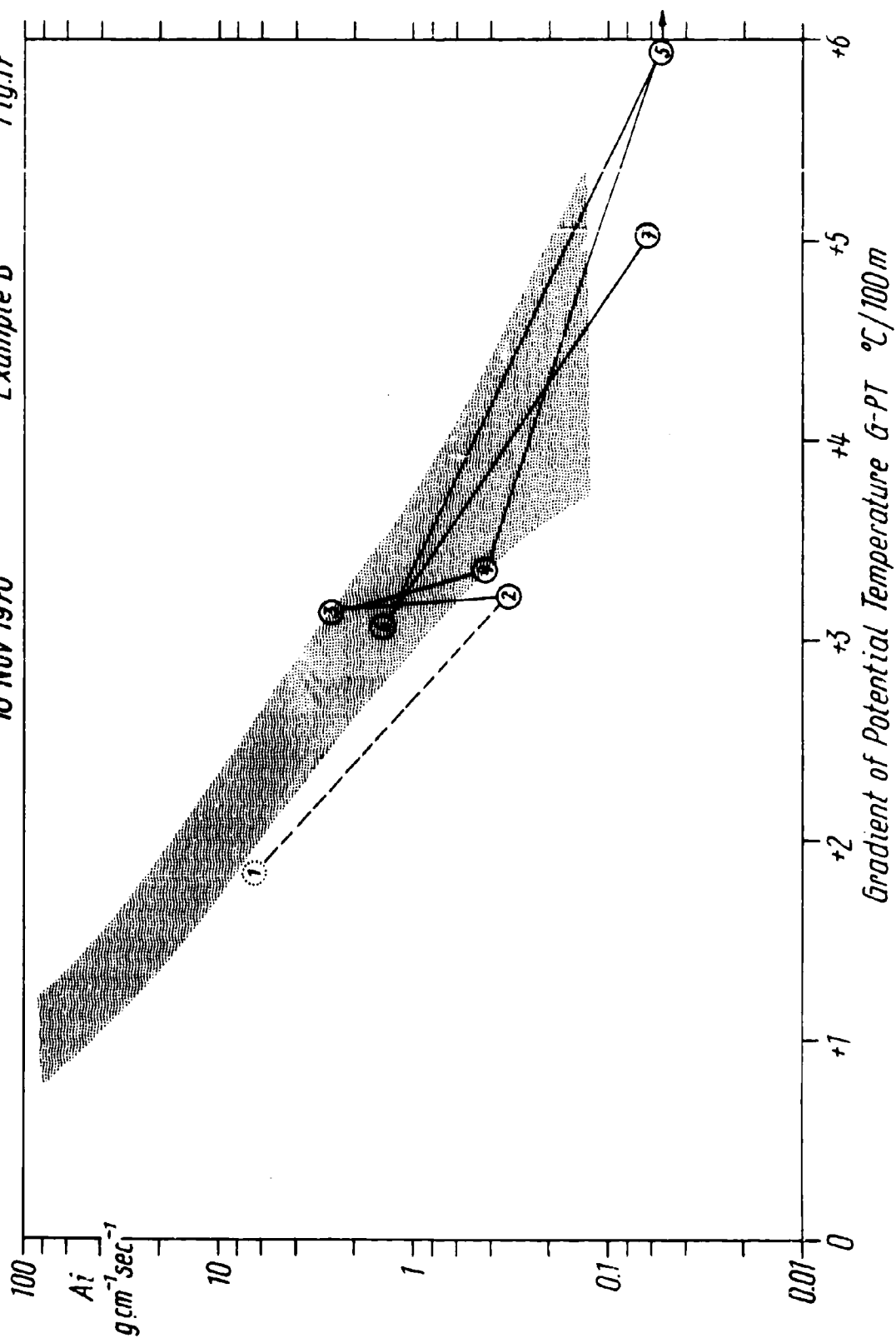
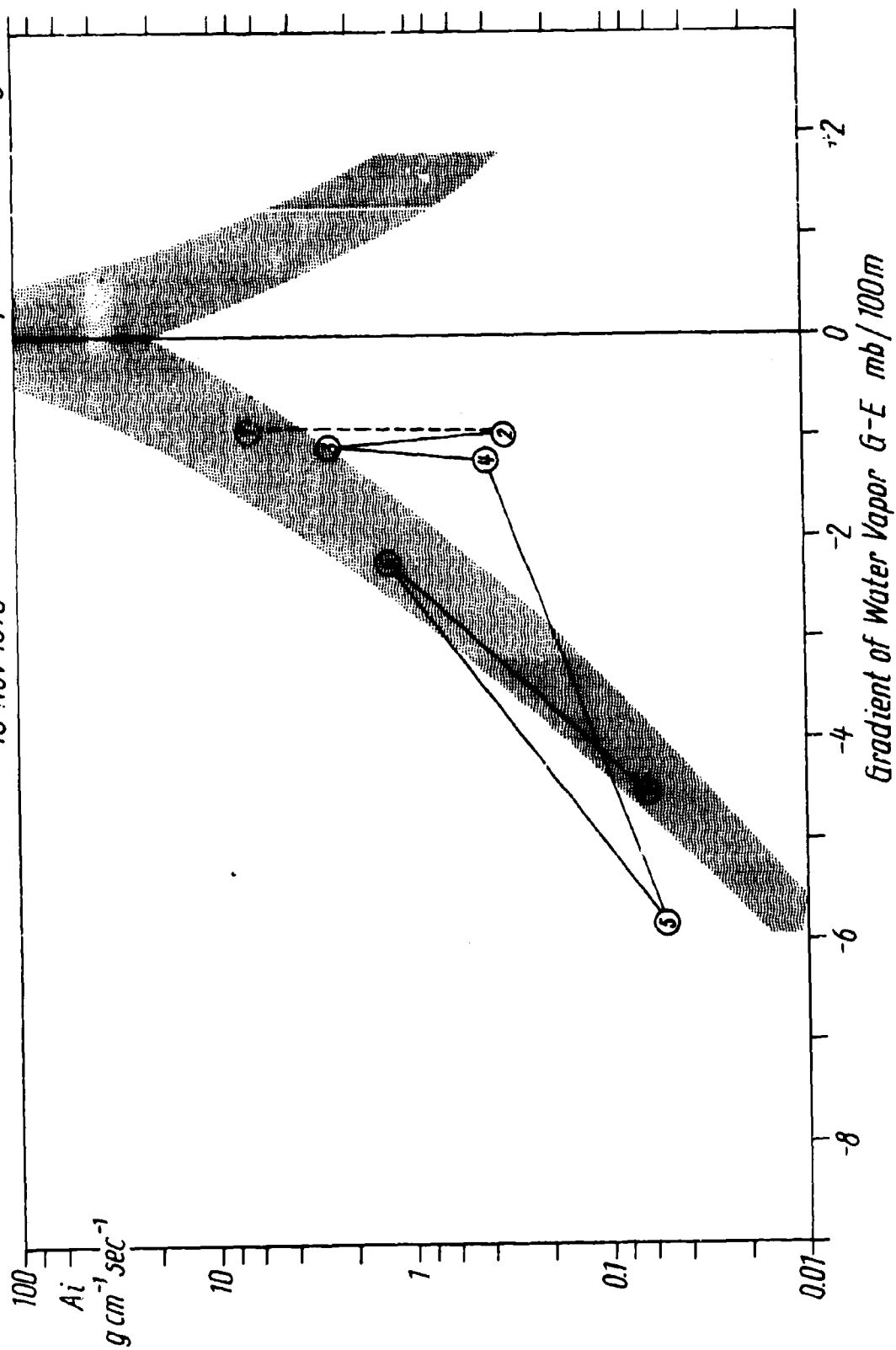


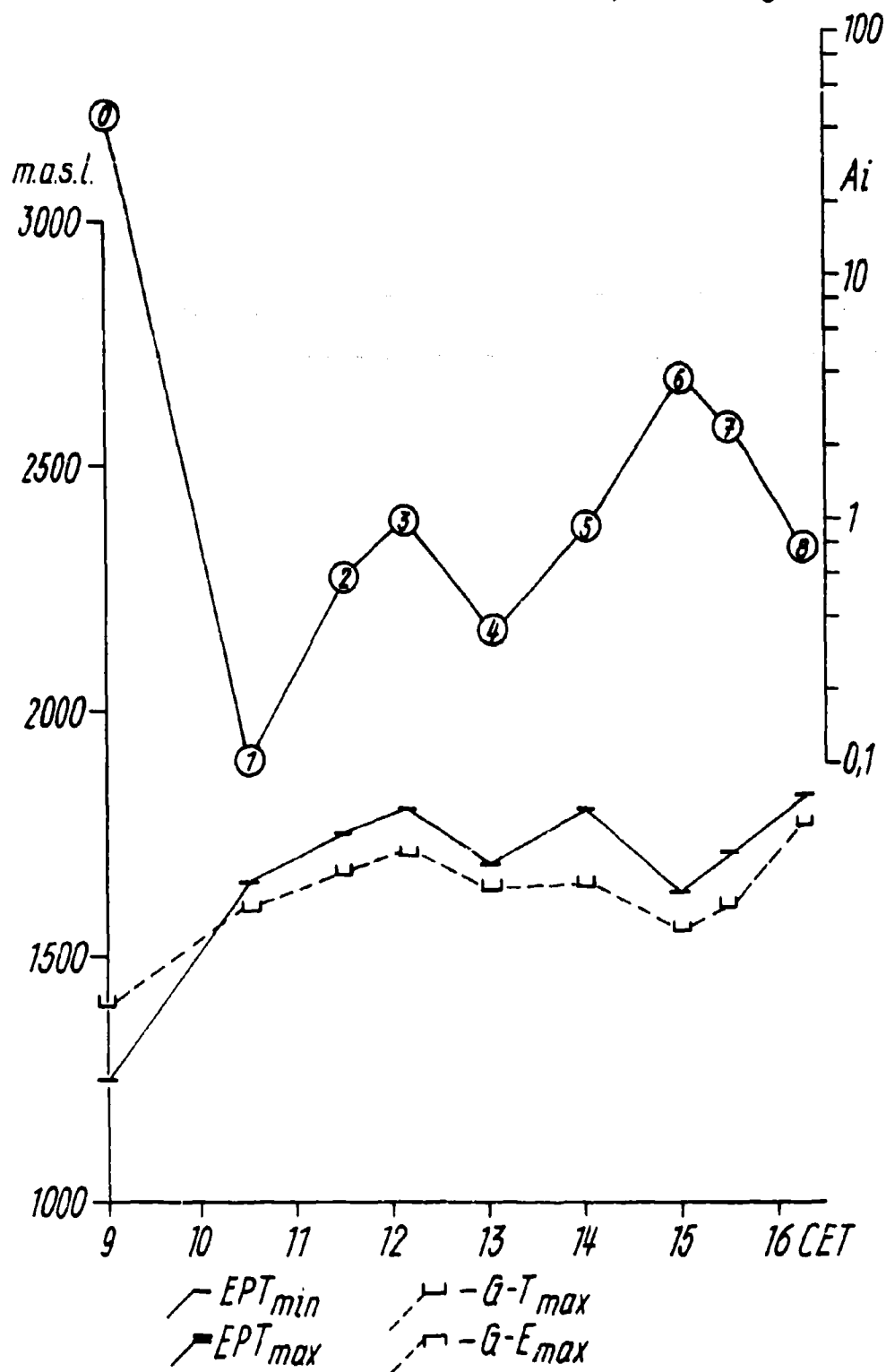
Fig. 18

Example B

18 Nov 1970



14 Oct 1970 Example C Fig. 20



Example C Fig. 21

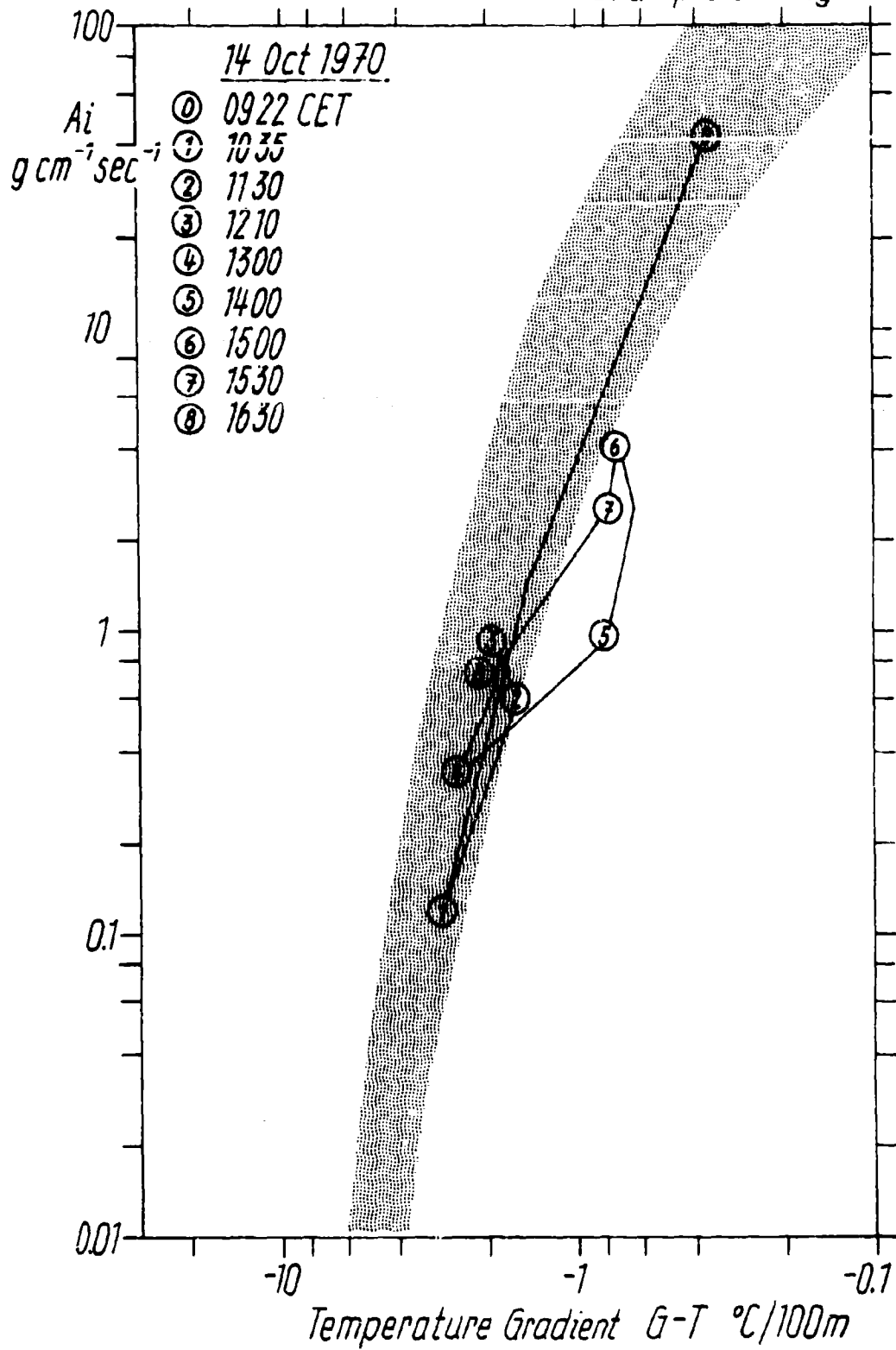
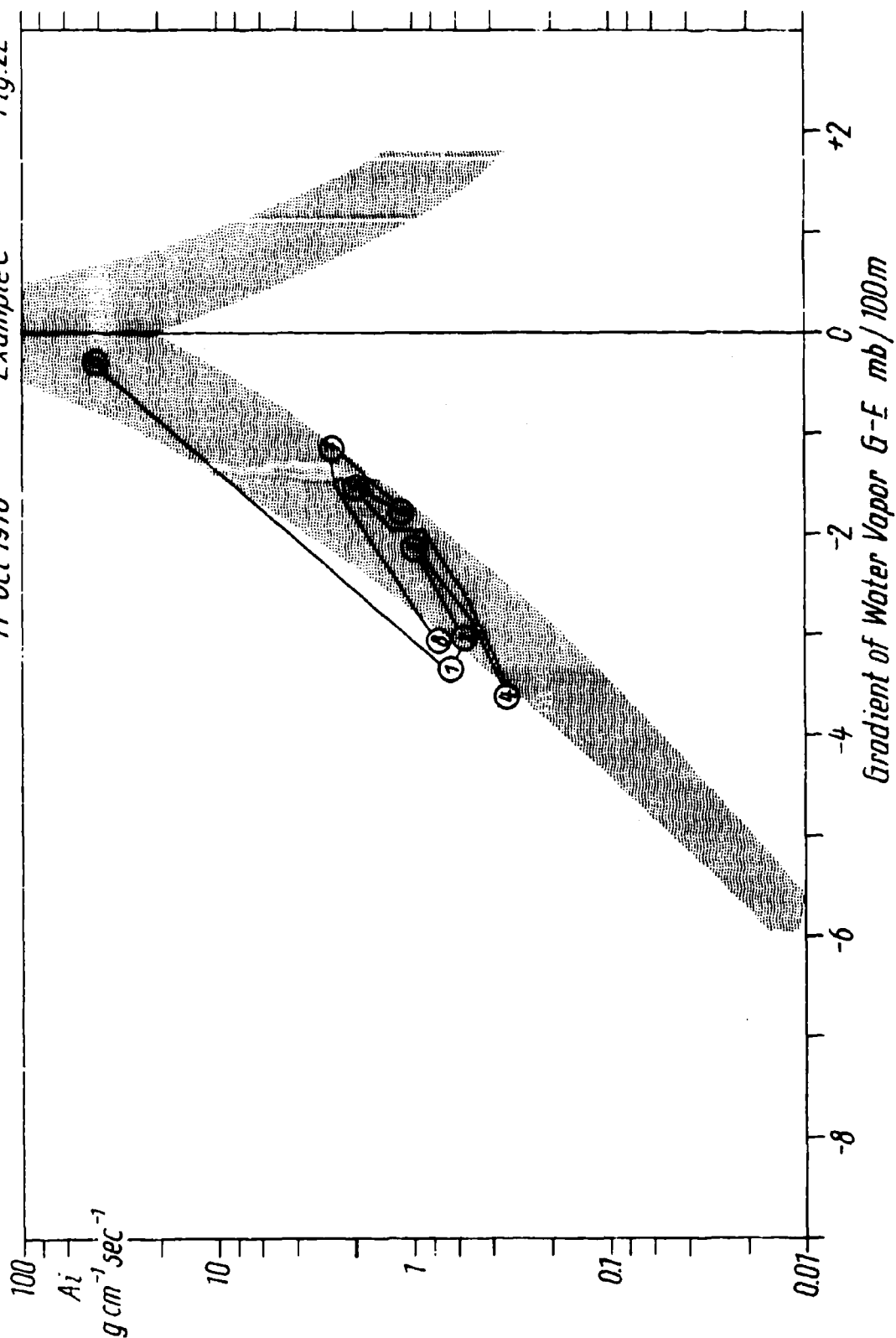


Fig. 22

Example C

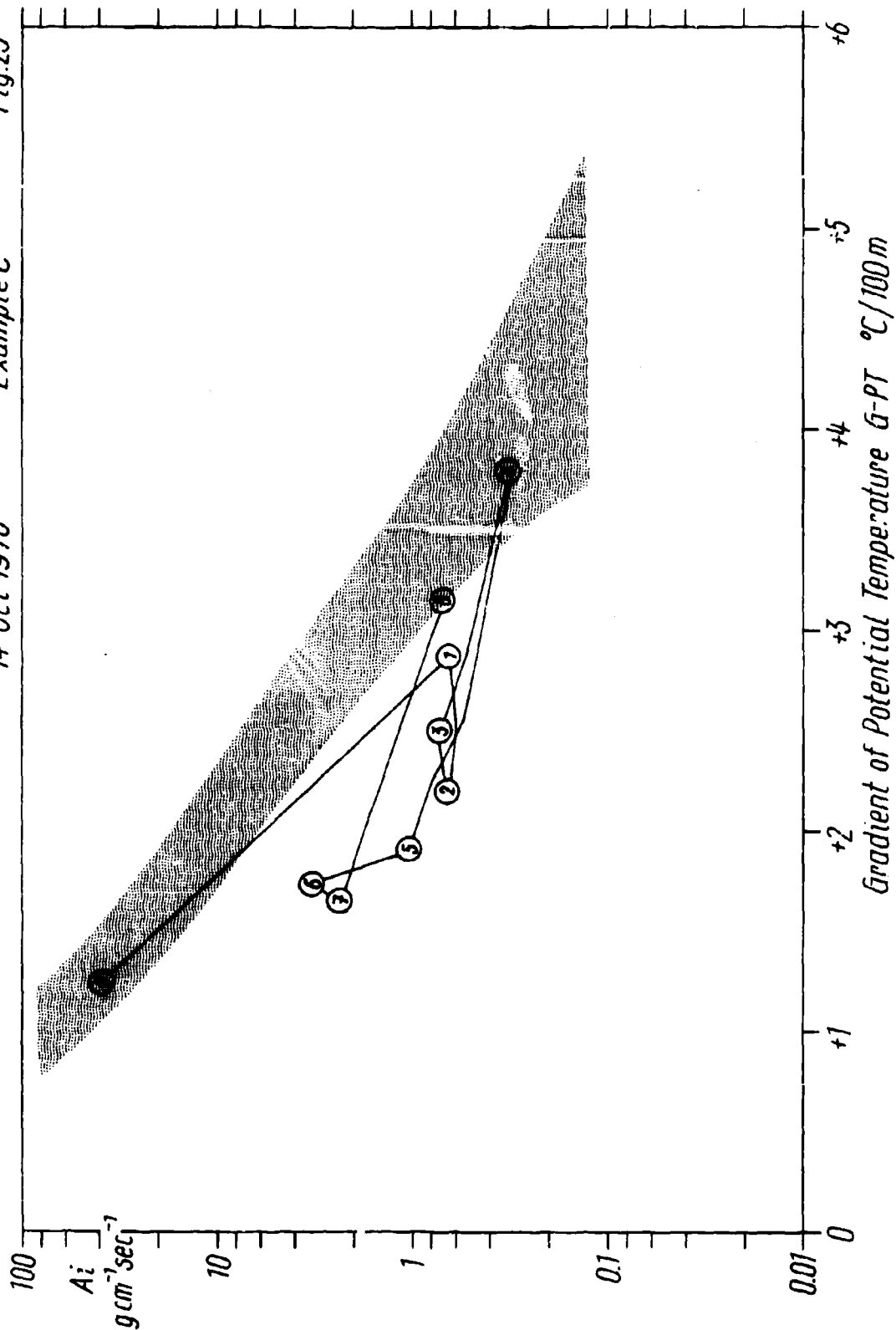
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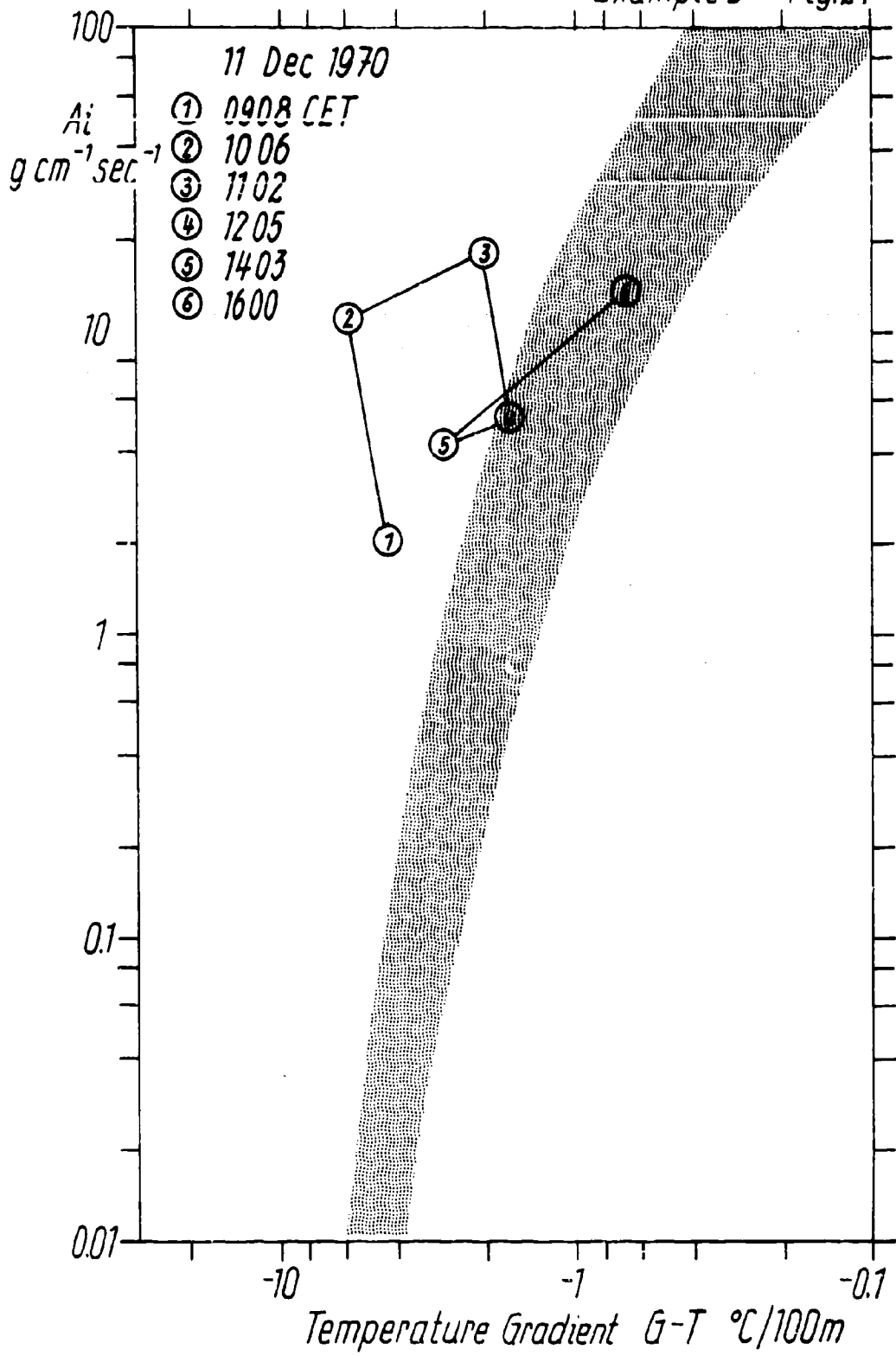
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Example C

Fig. 23



Example D Fig.24



11 Dec 1970

Example D

Fig. 25

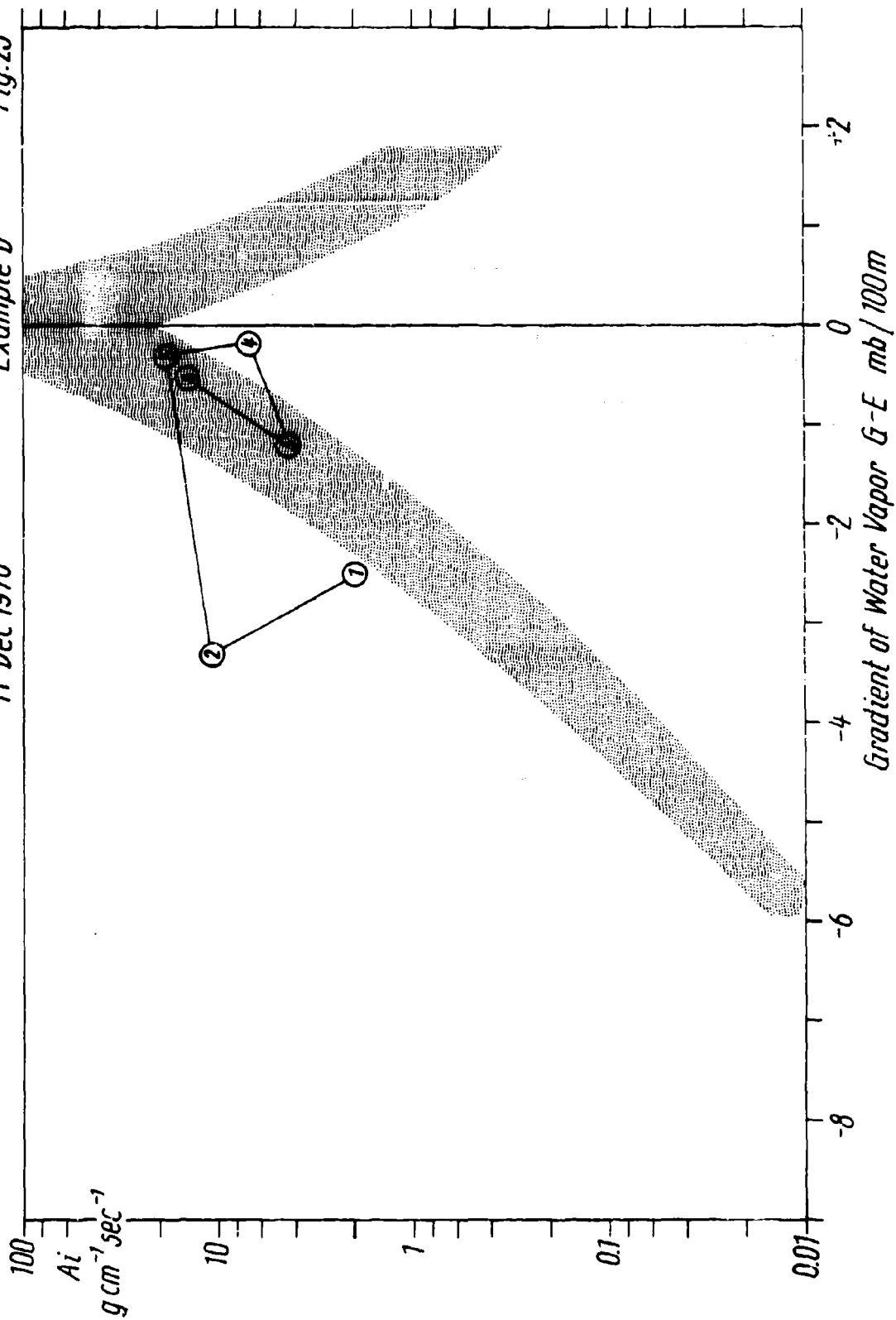
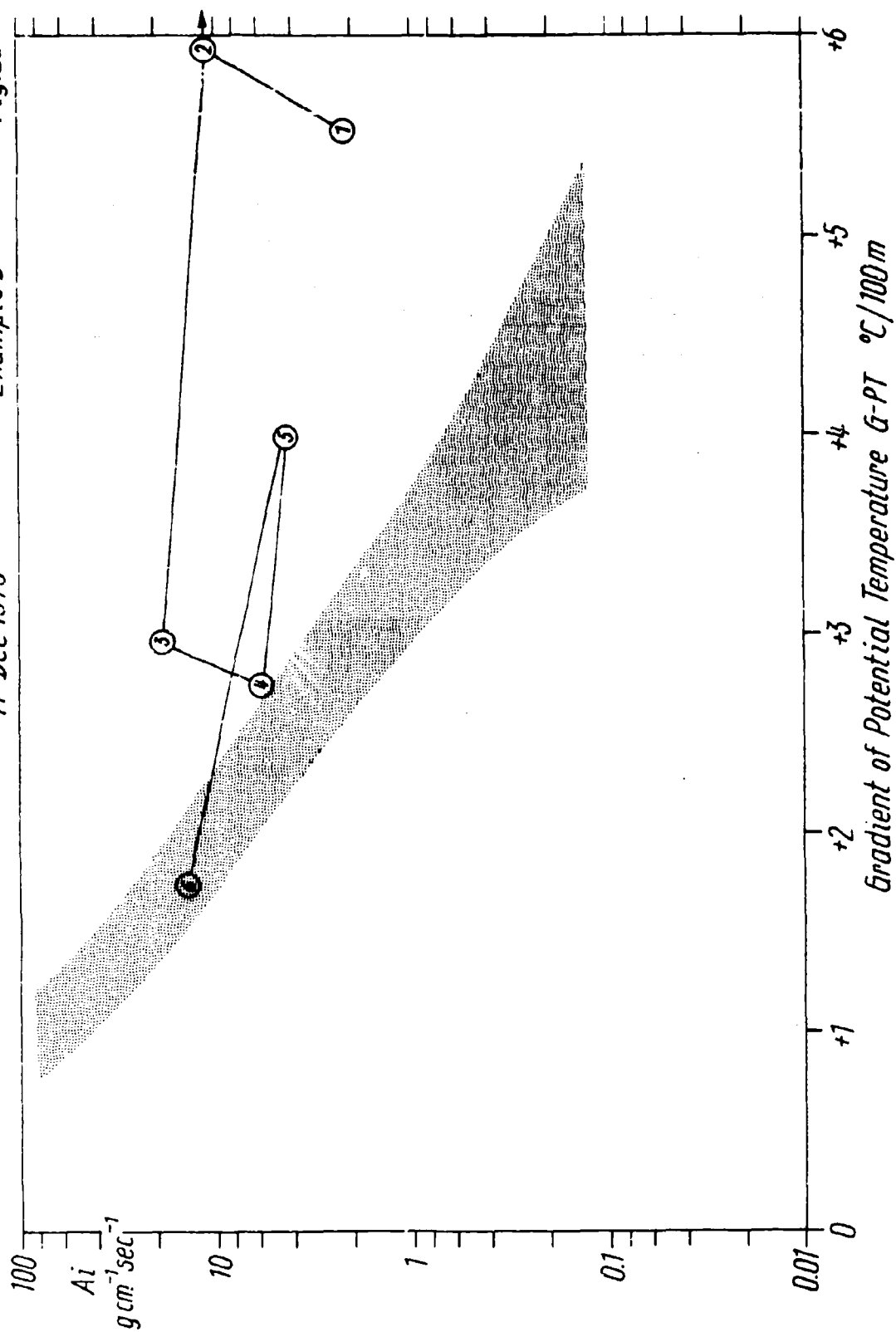


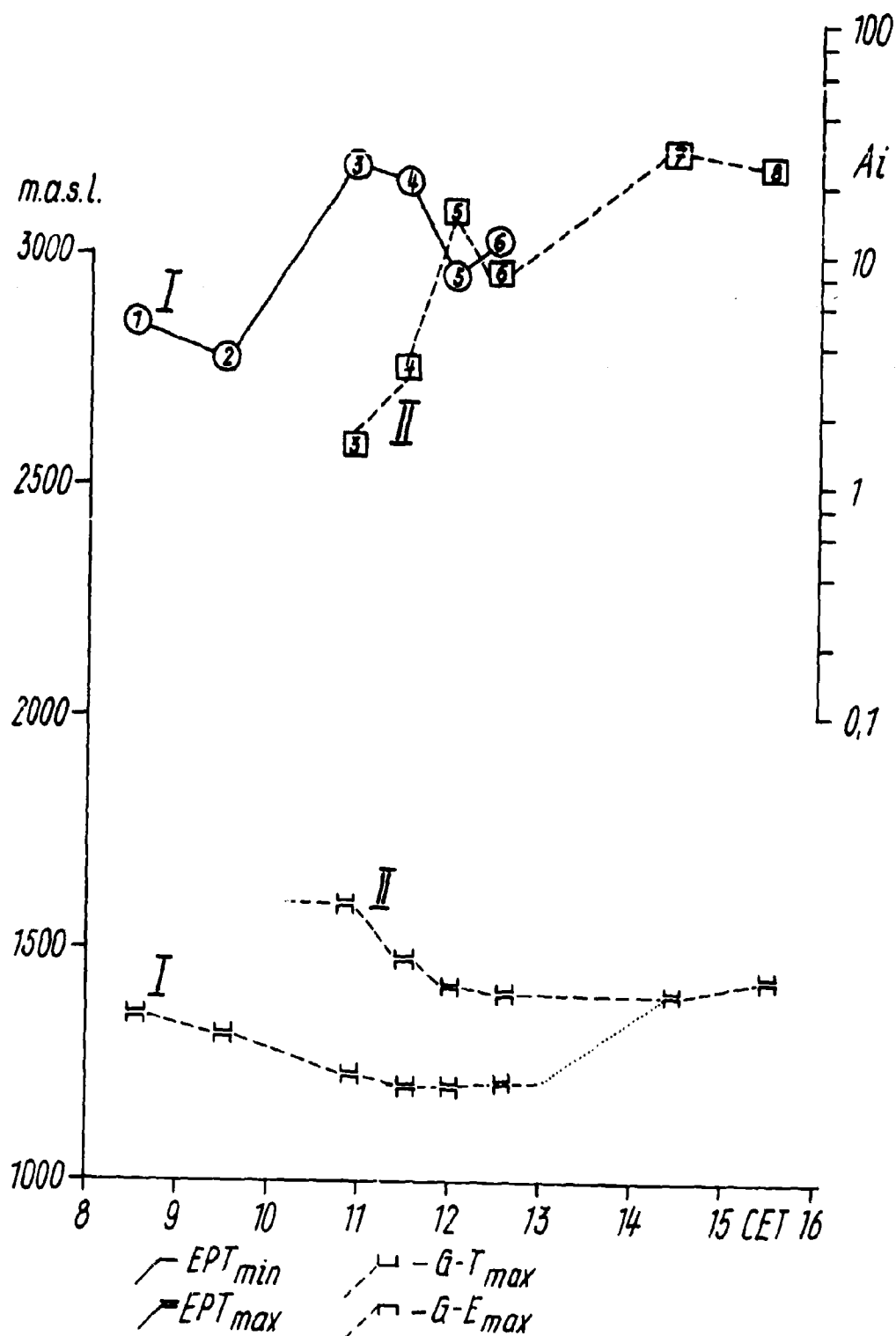
Fig. 26

Example D

11 Dec 1970



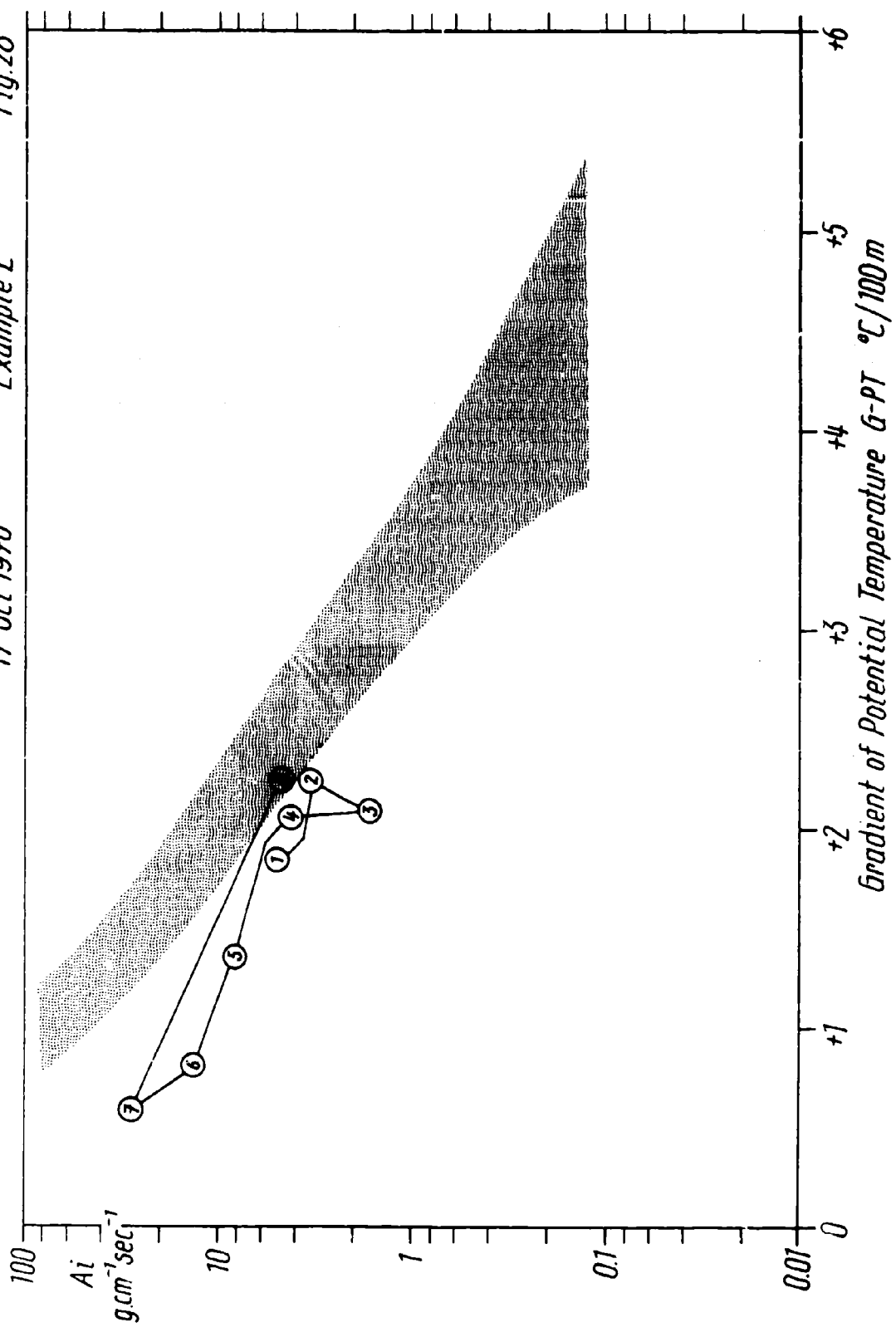
17 Oct 1970 Example E Fig. 27



17 Oct 1970

Example E

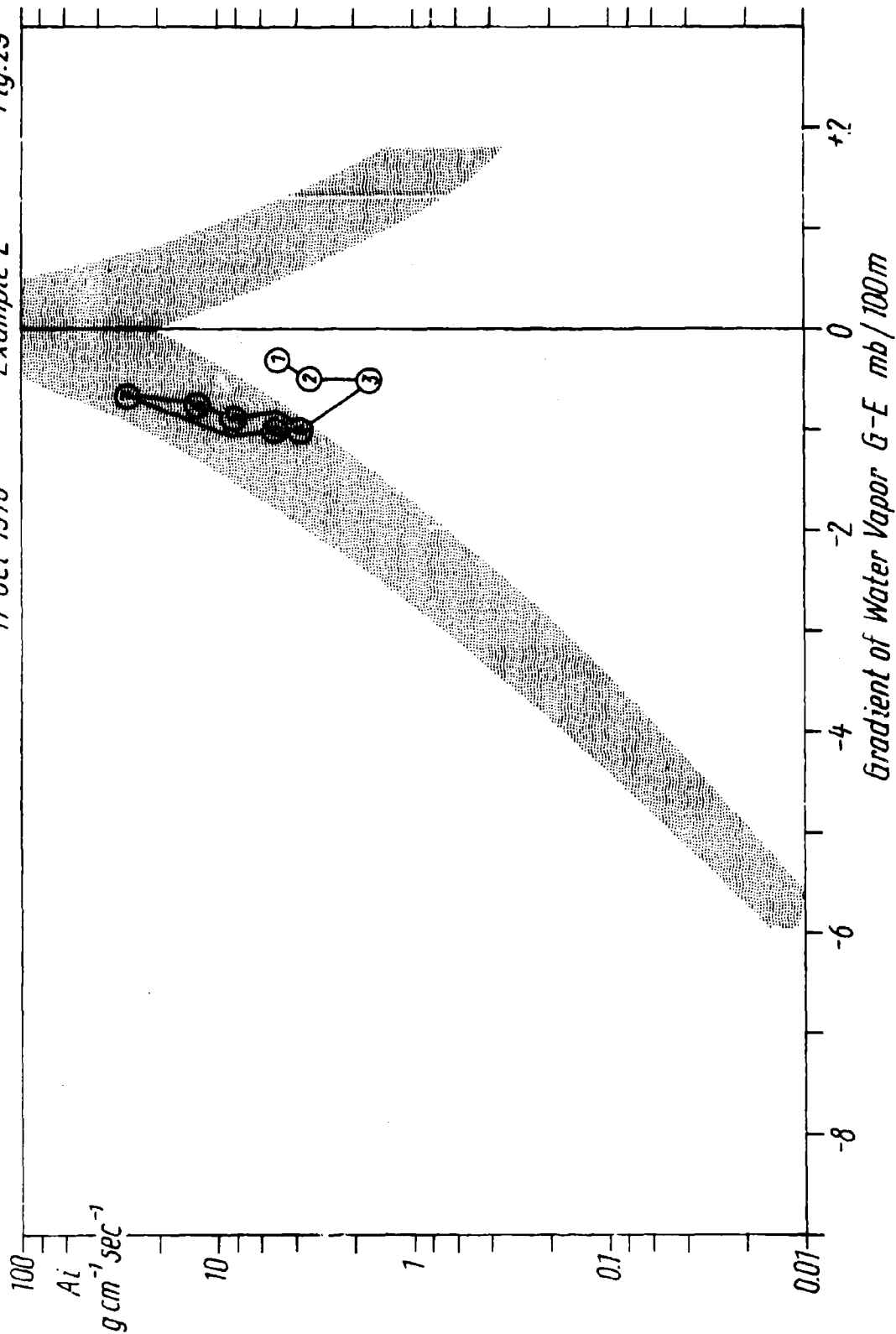
Fig. 28



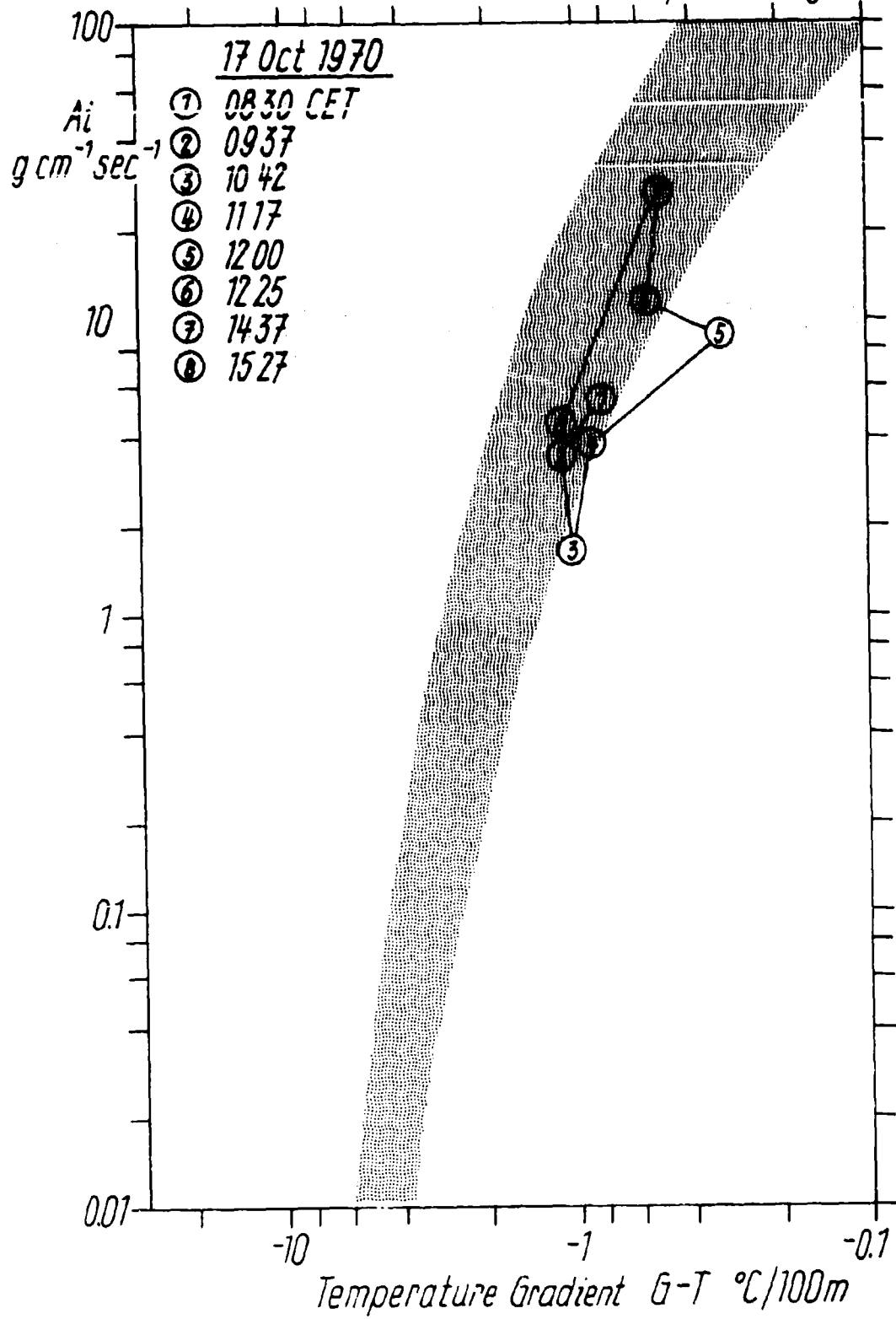
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Example E

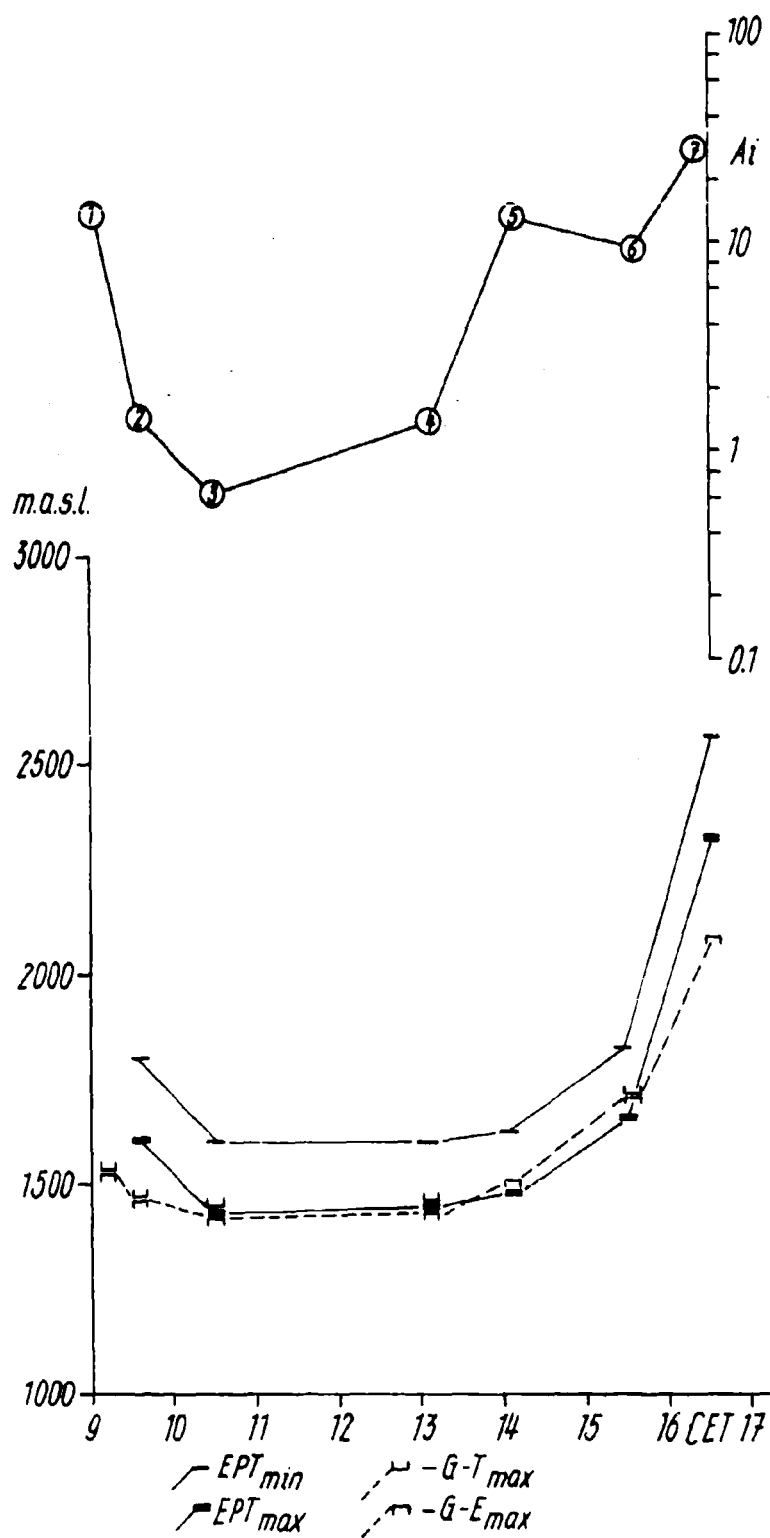
Fig. 29



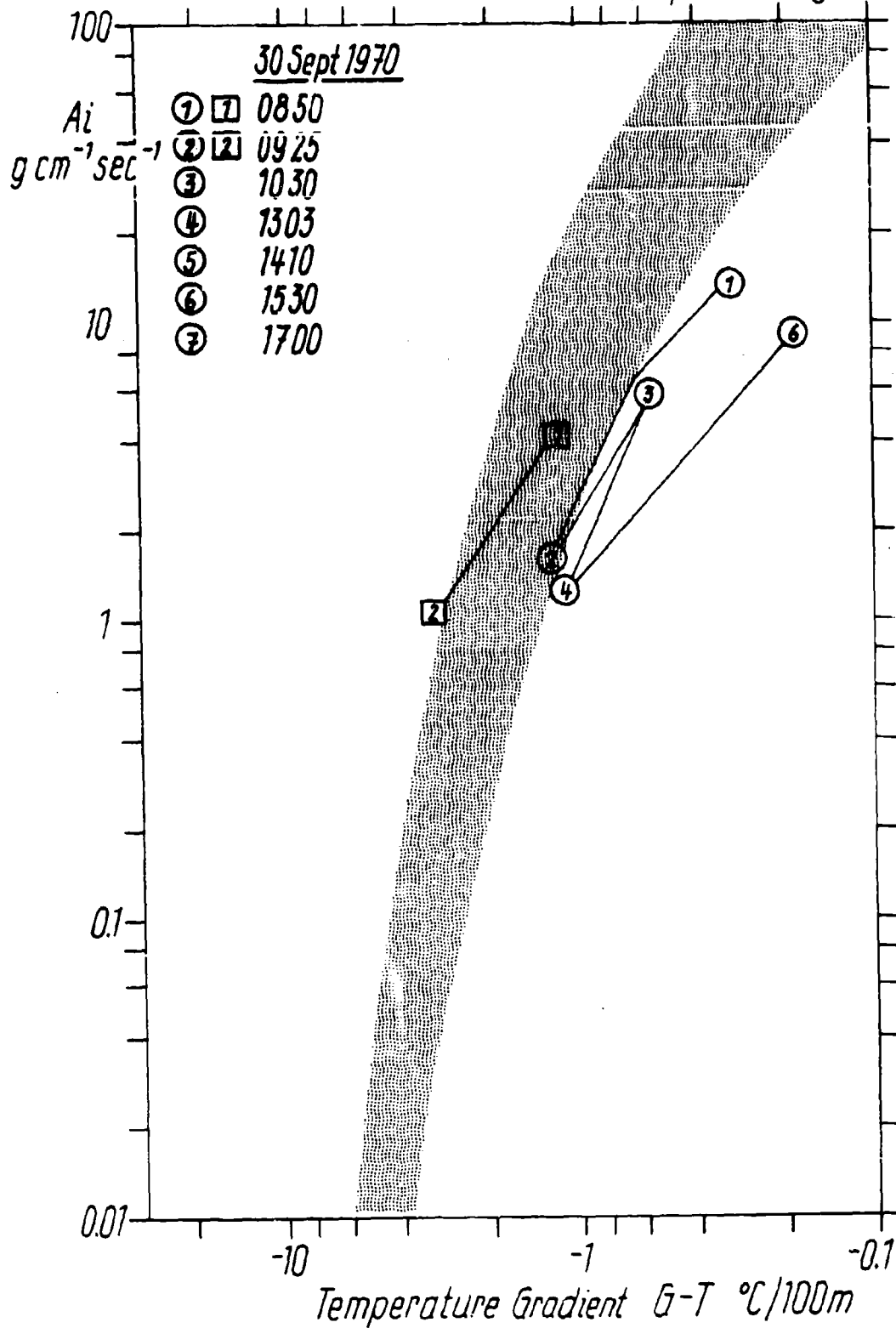
Example E Fig. 30



30 Sept 1970 Example F Fig. 31



Example F Fig.32



30 Sept. 1970

Example F

Fig. 33

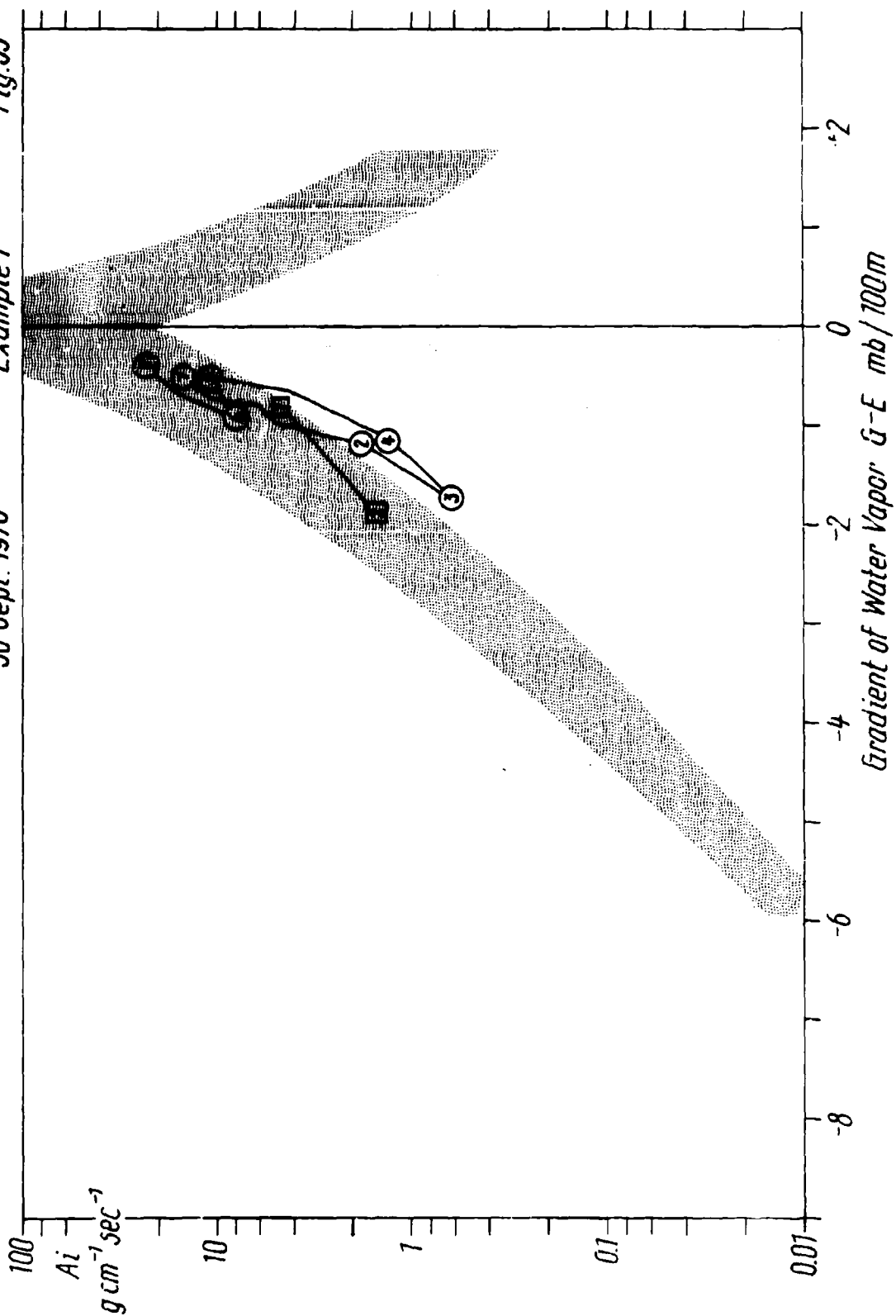
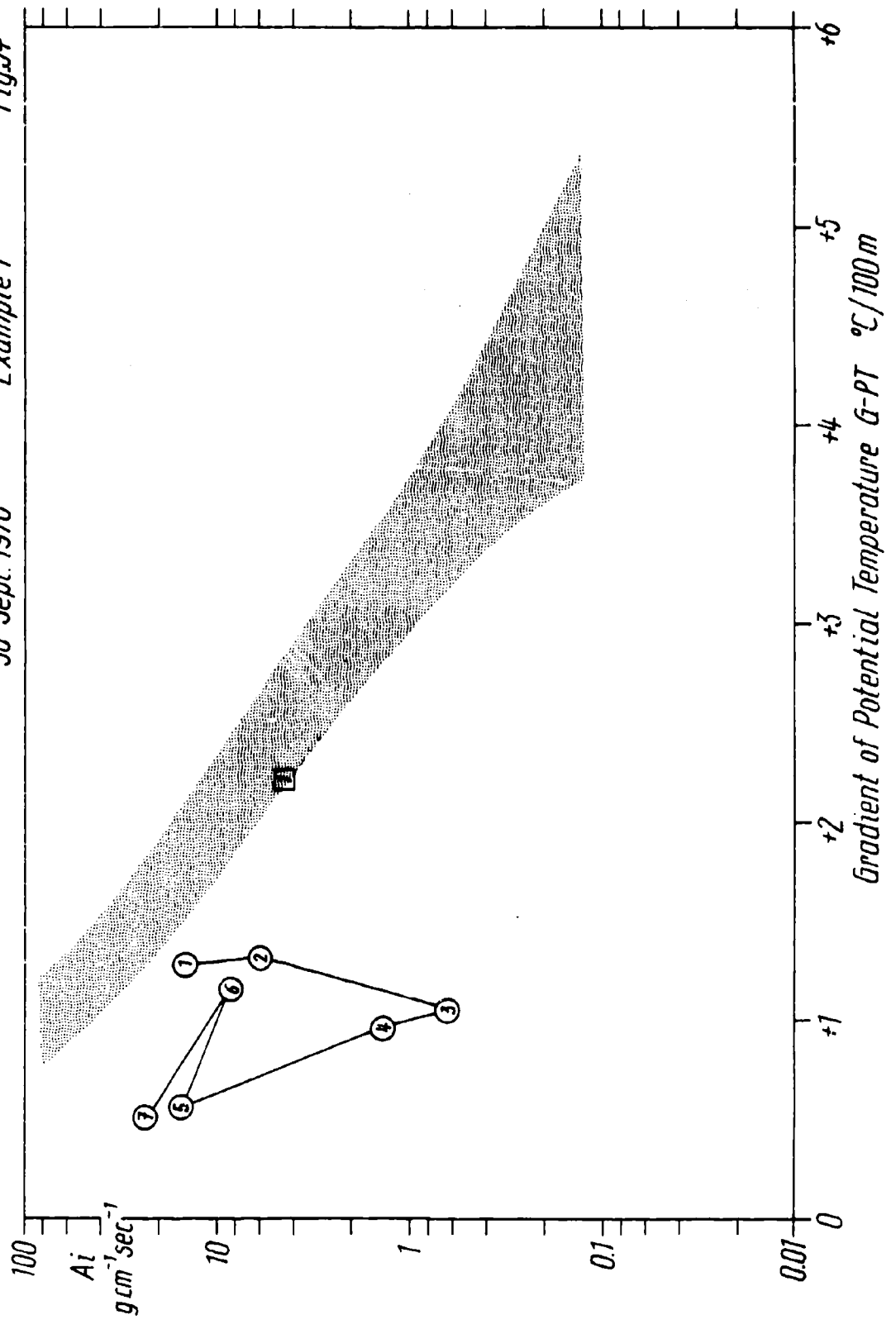


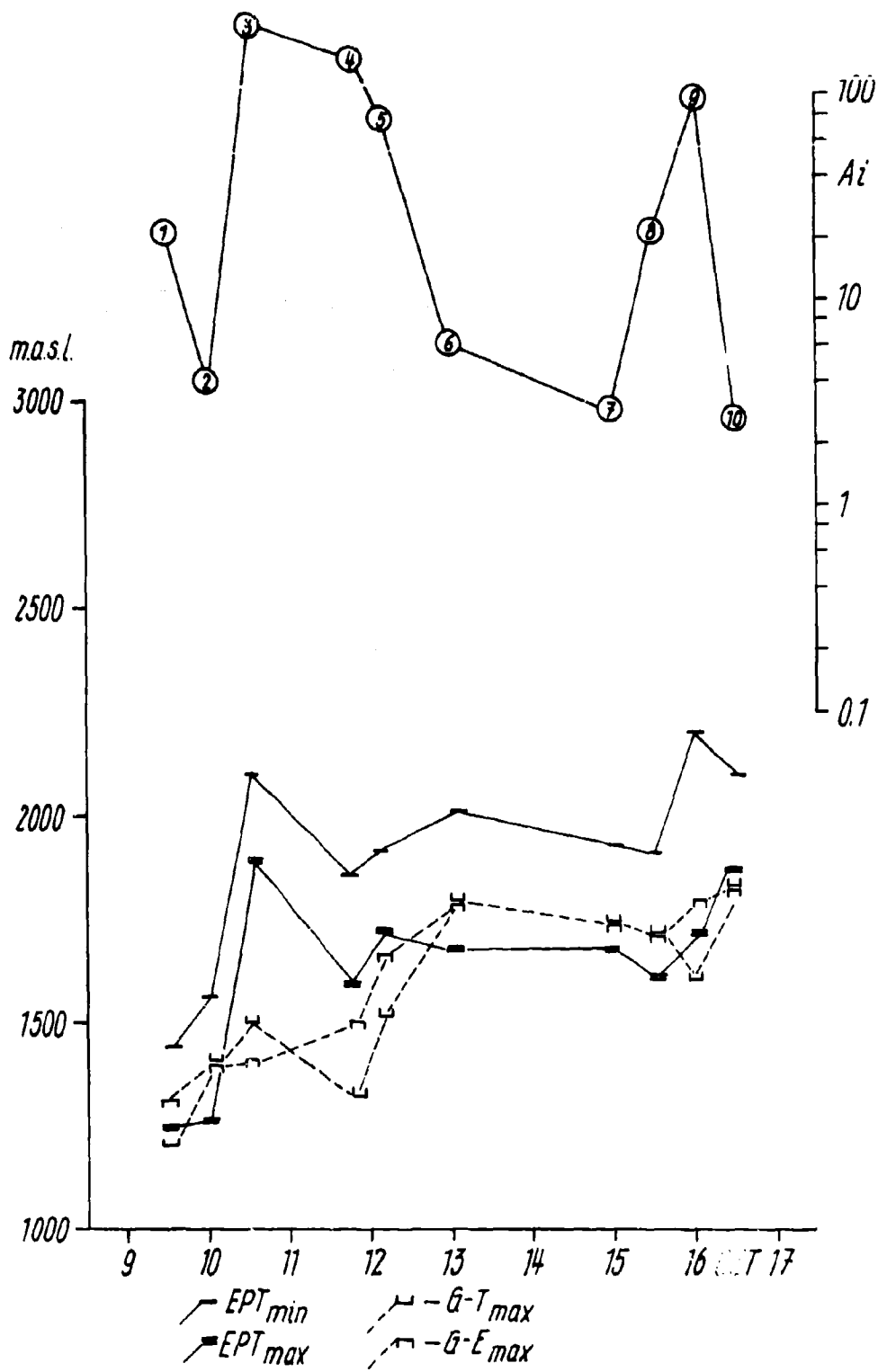
Fig. 34

Example F

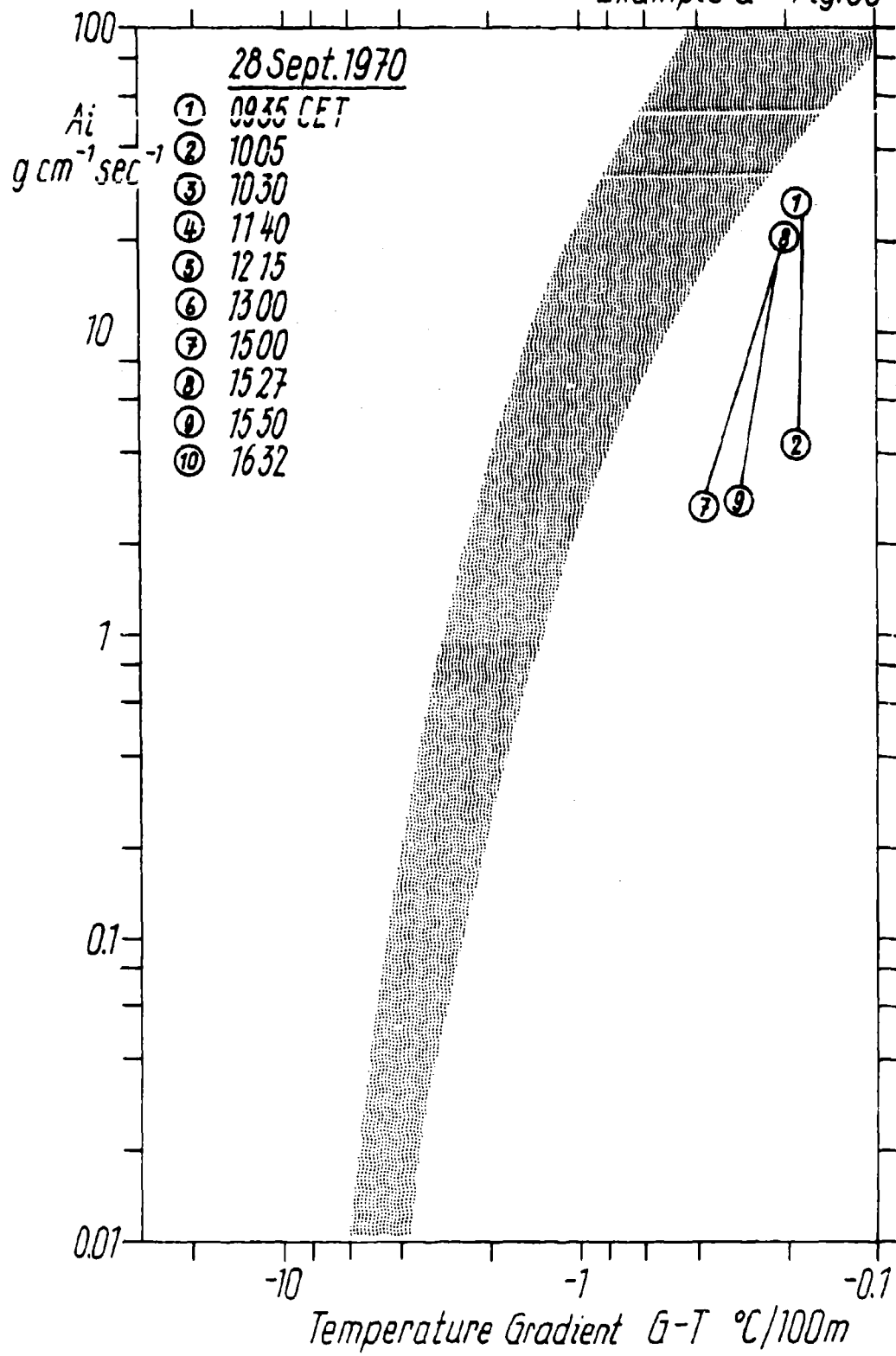
30 Sept. 1970



28 Sept. 1970 Example G Fig. 35



Example G Fig. 36



28 Sept. 1976

Example in Fig. 37

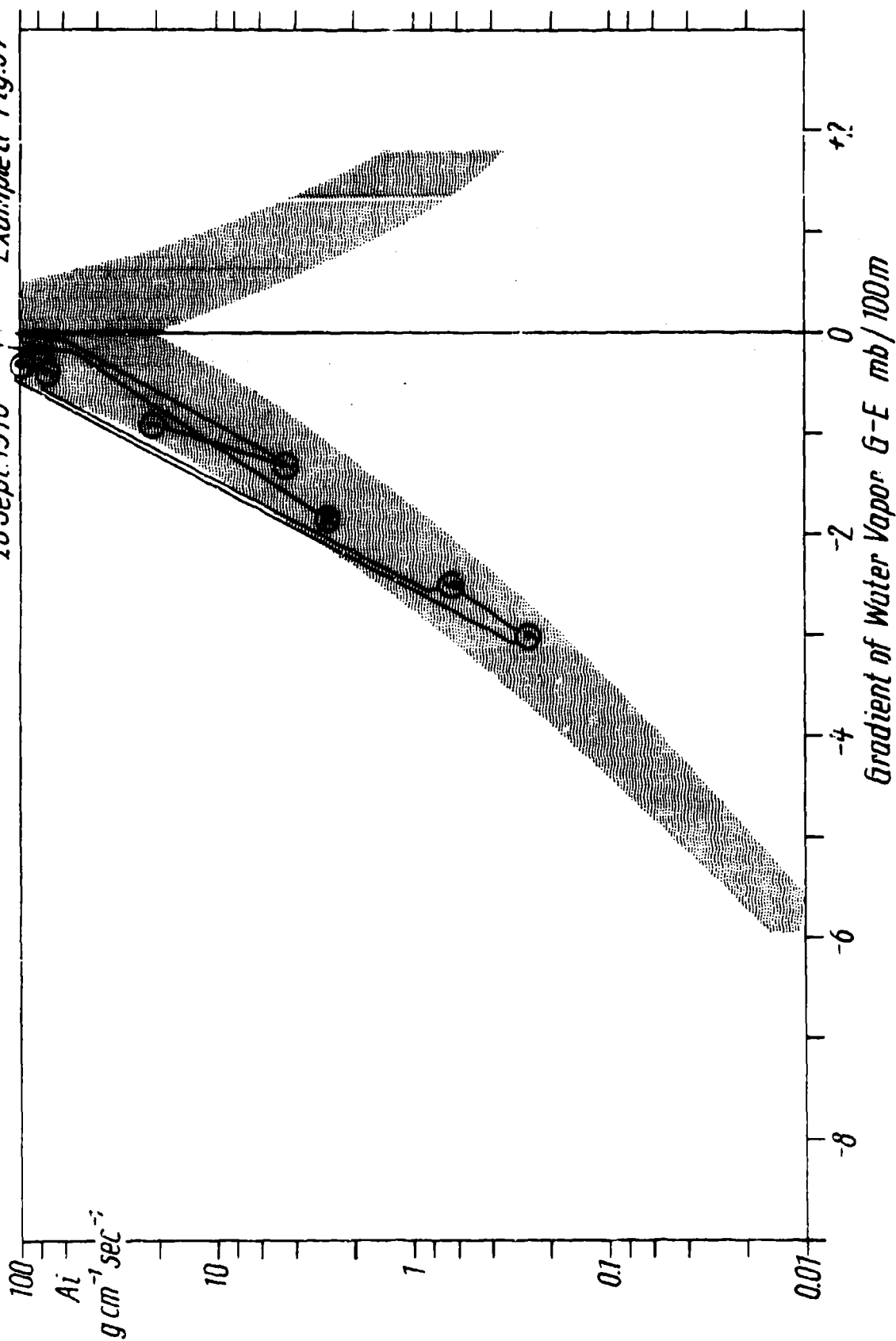
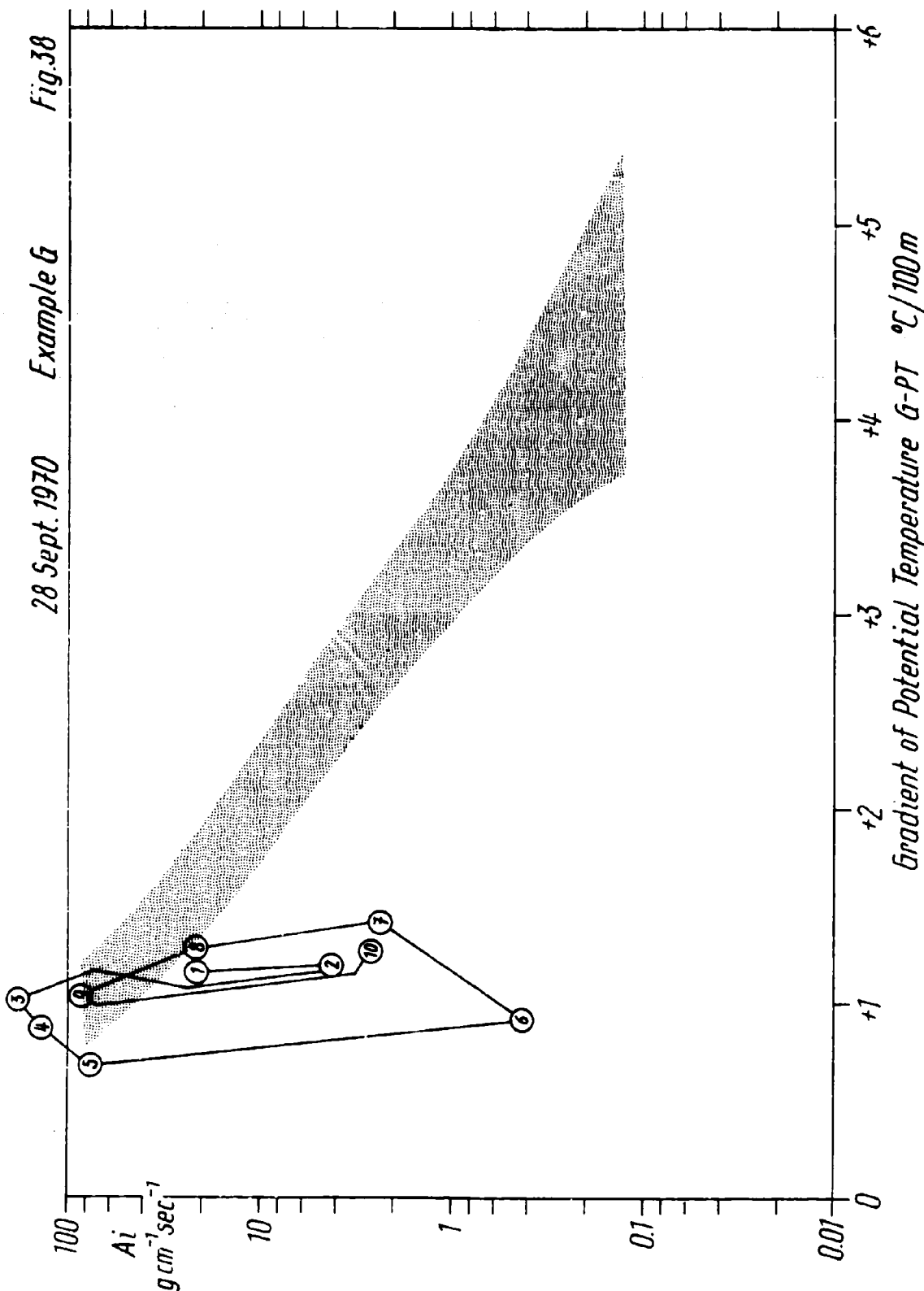


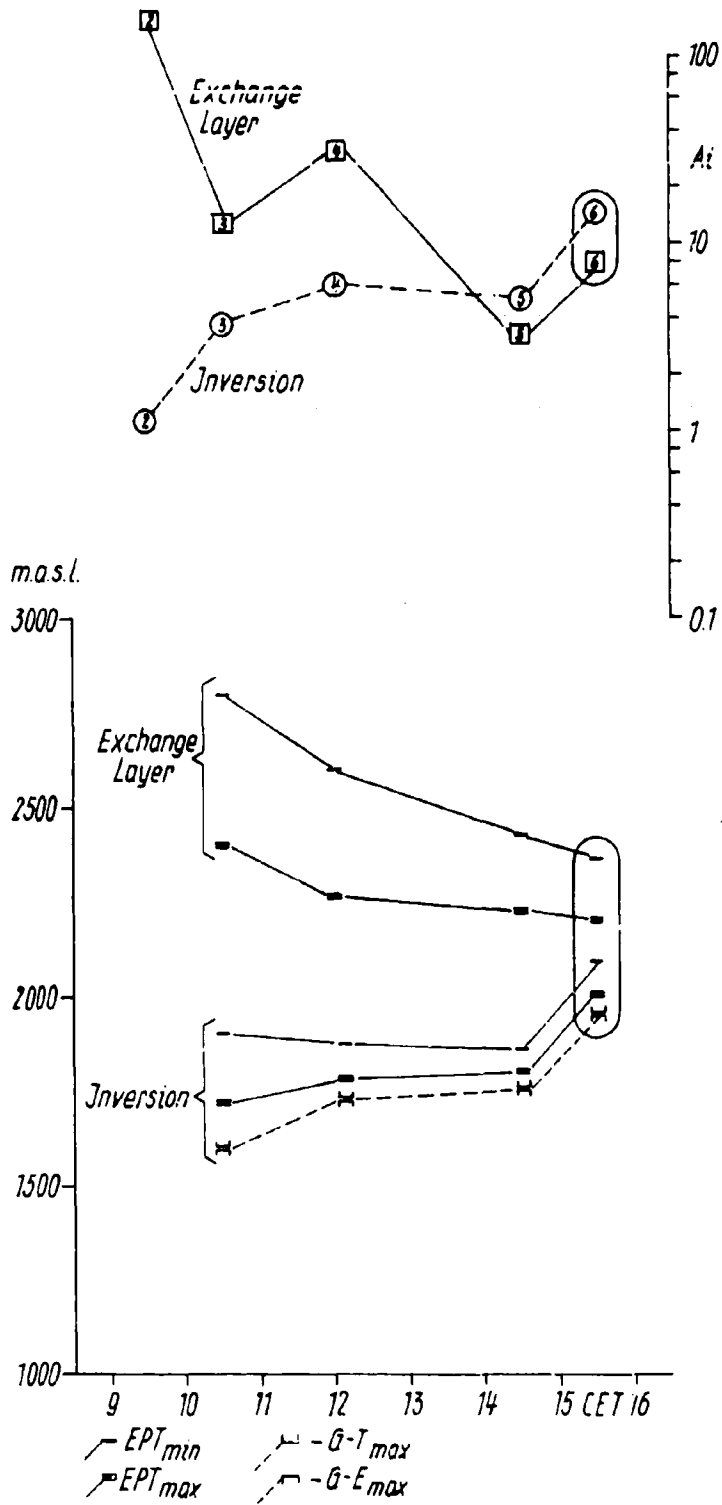
Fig.38

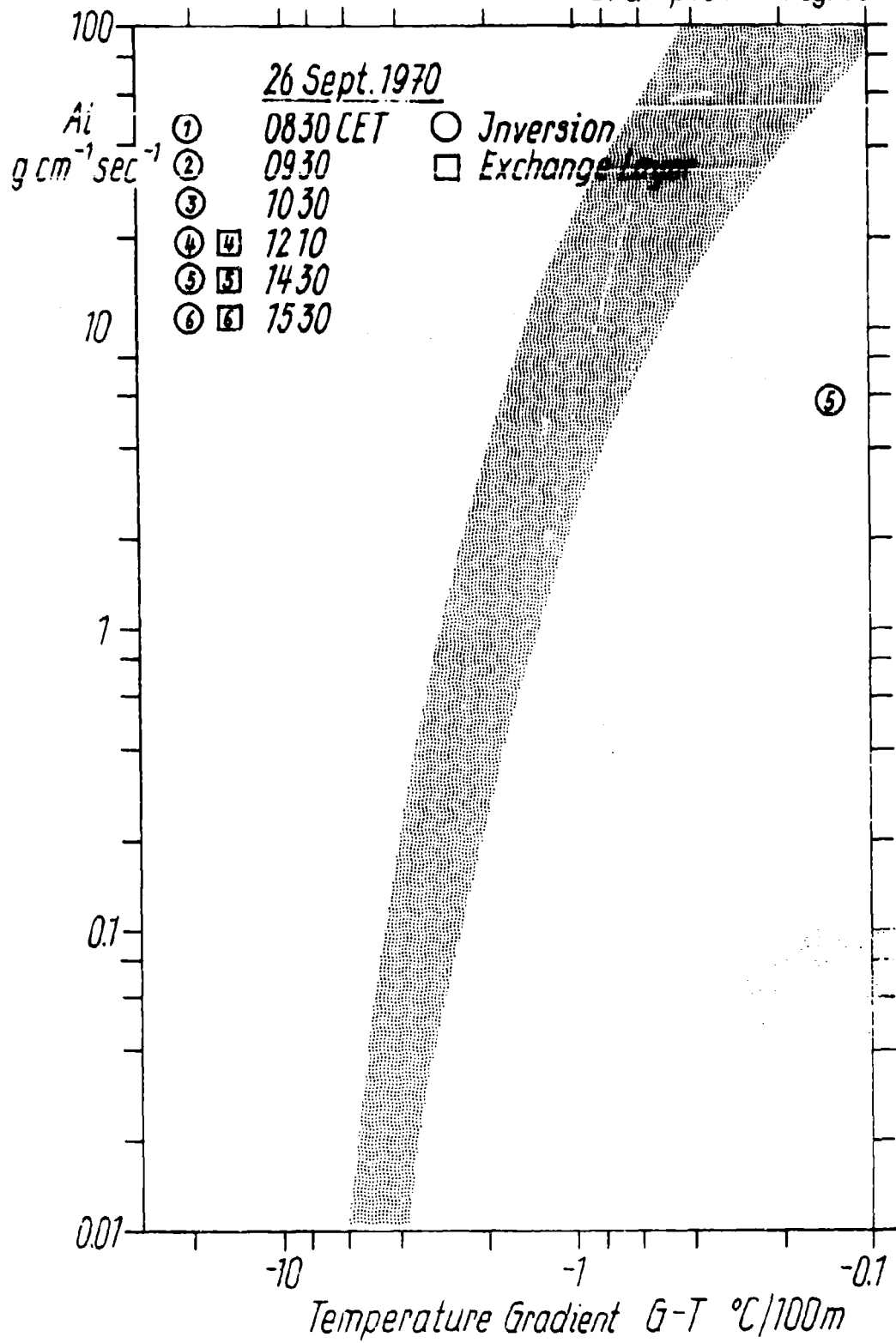
Example G

28 Sept. 1970



26 Sept. 1970 Example H Fig.39





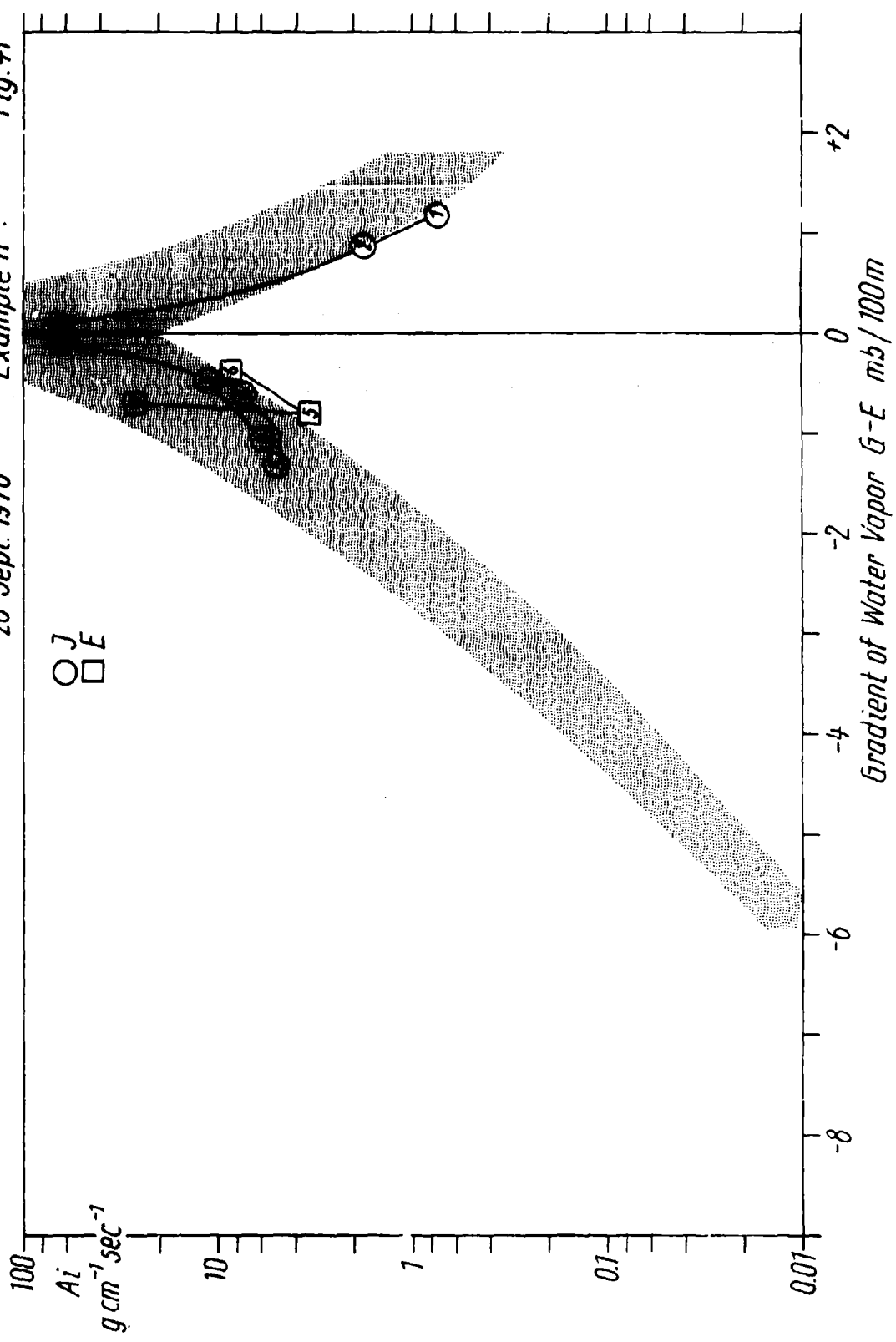
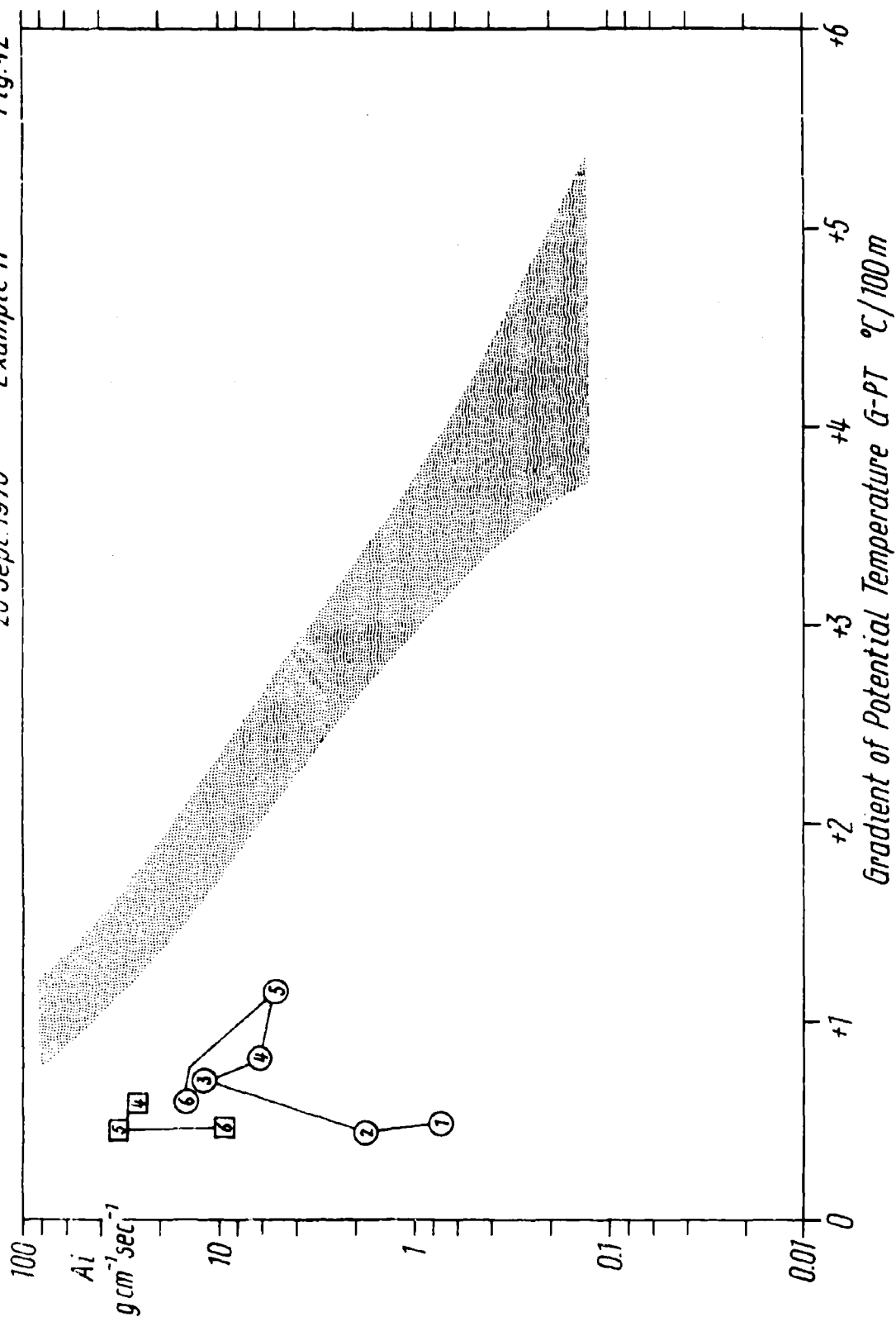


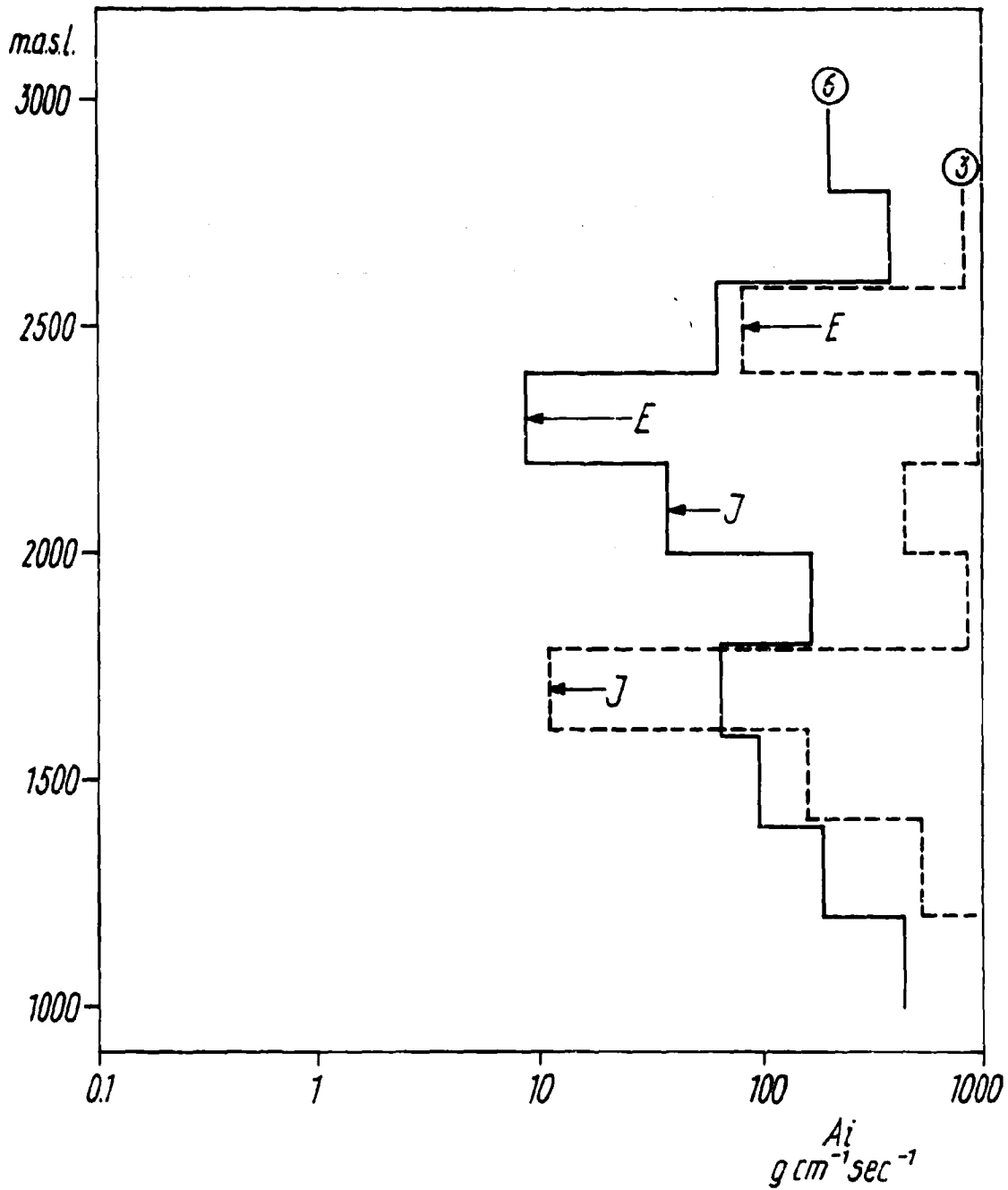
Fig. 42

Example H

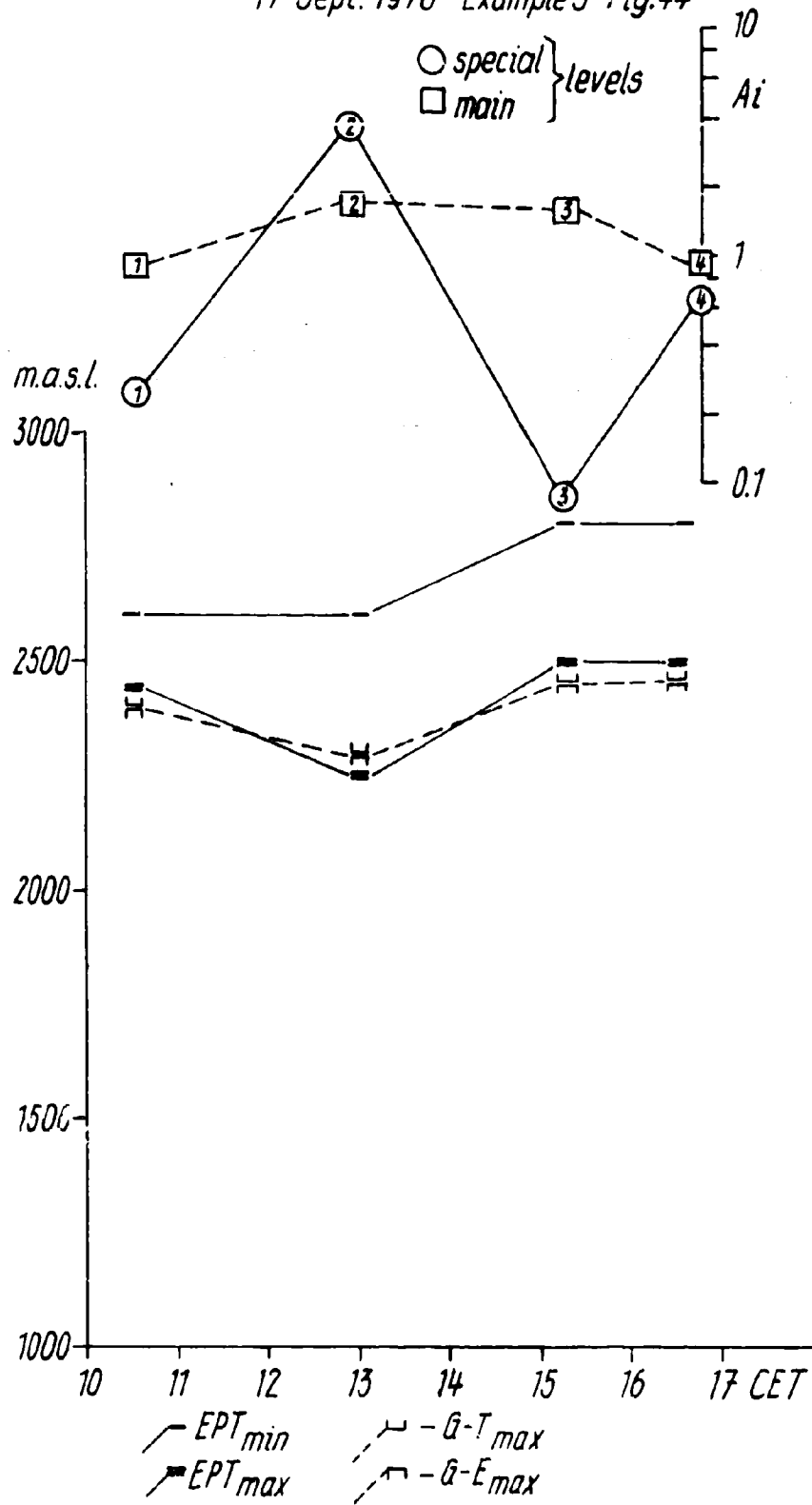
26 Sept. 1970



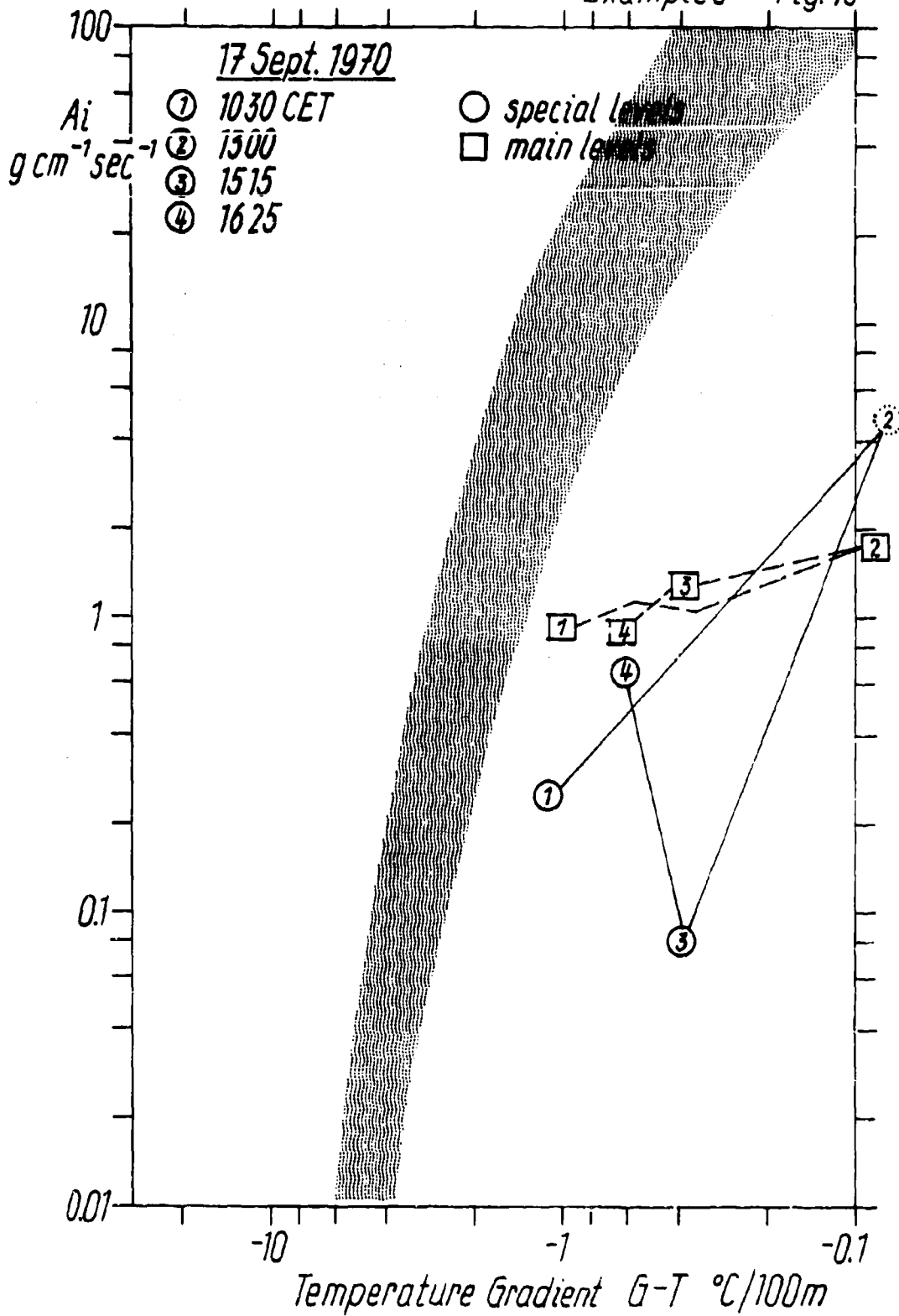
26 Sept. 1970 Example H Fig 43

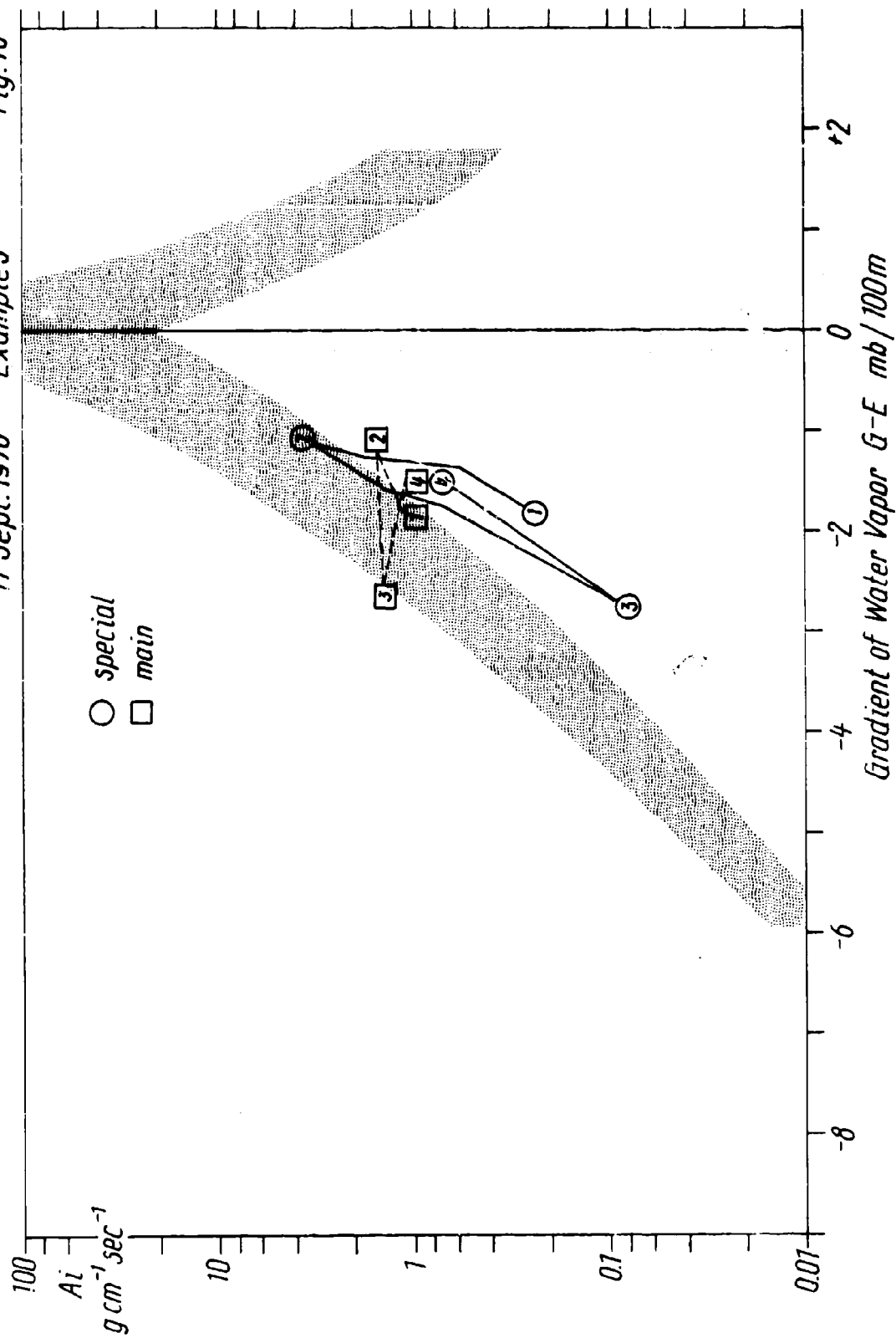


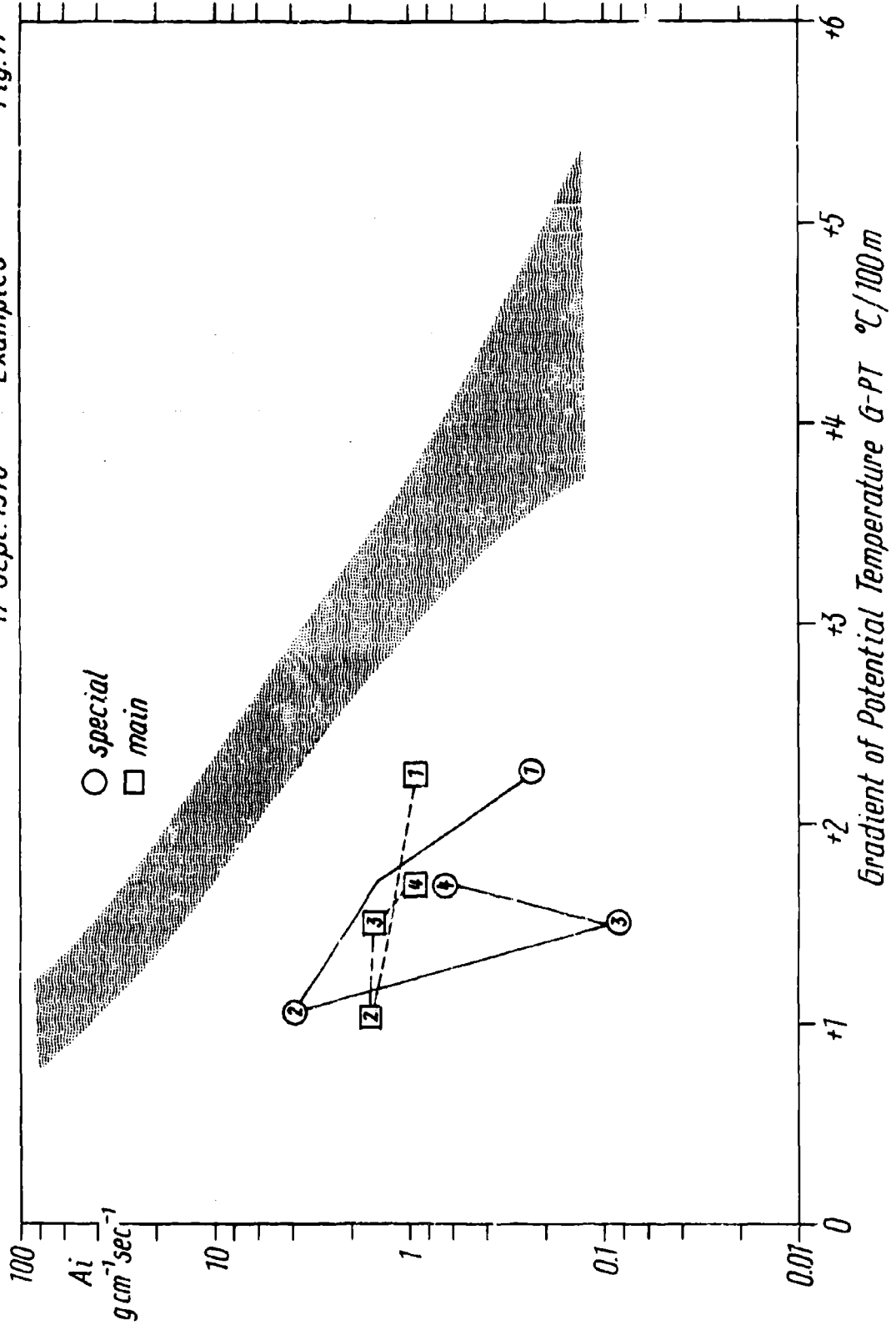
17 Sept. 1970 Example J Fig. 44



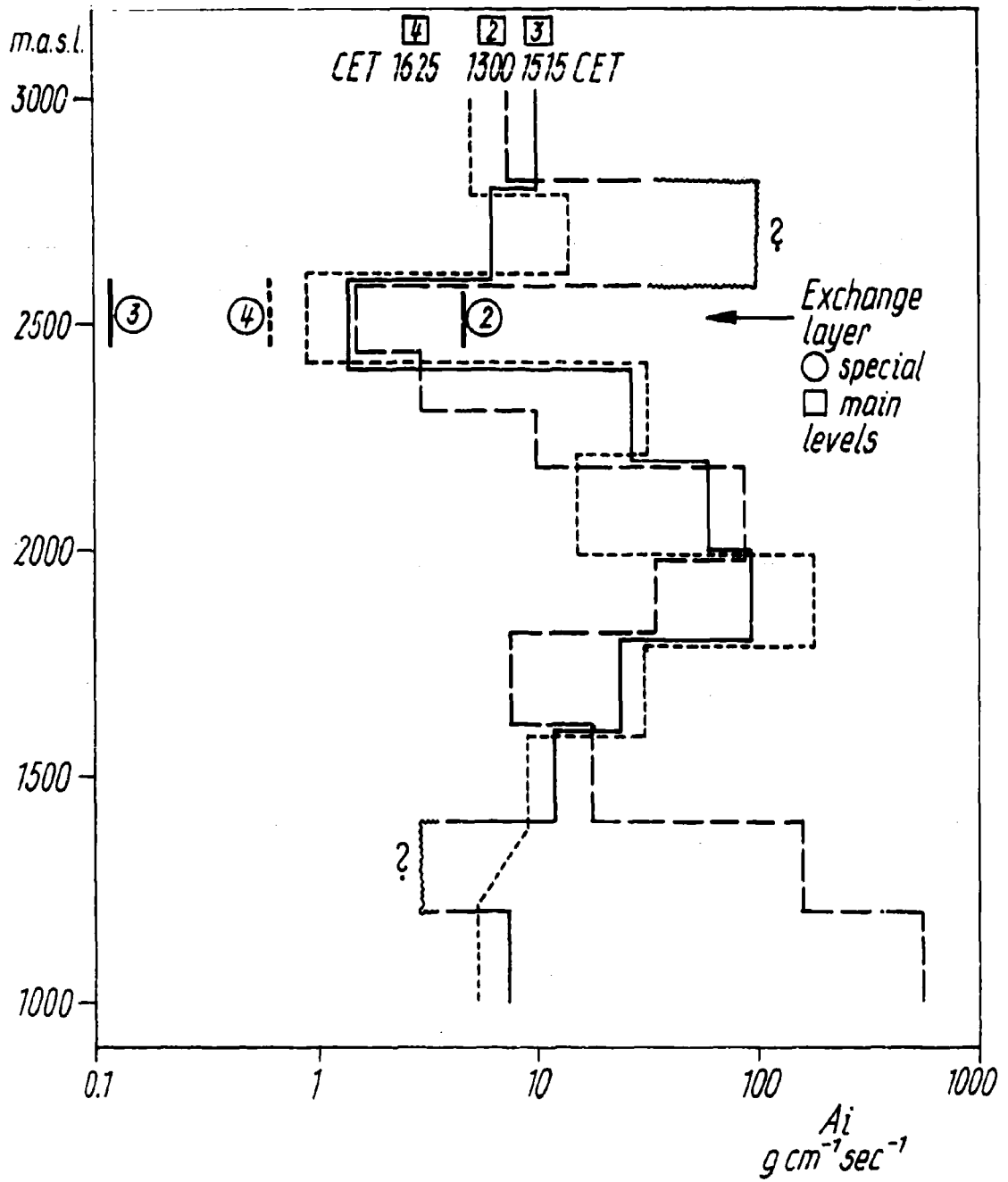
Example J Fig. 45







Example J Fig. 48



2. Legends to the Figures 1 - 48

Legends to Figures

Fig. 1:

Chart on the processing of the recordings transmitted from the moving cable car, up to the final results: the completed tables printed out by teletype.

Fig. 2:

Relation between incremental exchange coefficient A_1 and temperature gradient G-T. Representation of the relation by the individual pairs of values. The relation applies to real inversions exclusively (stationary equilibrium of the eddy diffusion).

Fig. 3:

Result of statistical analyses of Fig. 2 with indication of scatter (red).

Fig. 4:

Relation between incremental exchange coefficient A_1 and gradient of potential temperature G-PT expressed in degrees centigrade per 100 m. Representation of the individual pairs of values. Valid for real inversions (stationary equilibrium of eddy diffusion).

Fig. 5:

Same as Fig. 4, however, result of statistical investigation with indication of scatter (red).

Fig. 6:

Relation between incremental exchange coefficient A_1 and gradient of water vapor pressure G-E in mb/100m. Individual pairs of values. Valid for aerological structures having a G-T minimum with stationary equilibrium.

b

Fig. 7:

Same as Fig. 6, however, result of statistical investigation with indication of scatter (red).

Fig. 8 and 9:

Relation between incremental exchange coefficient A_1 and gradient of potential temperature $G-PT$, divided into two intervals of time. Result of statistical investigation with indication of scatter. Relation of general applicability.

Fig. 10:

Mutual comparison of results from both intervals of time.

Re - Example A

Fig. 11:

Variations in space and time of important structural elements of the inversion newly formed at 1200 CET.

Fig. 12:

Time slope within the range of function $A_1=f(G-T)$. Numerals 1 through 4 identify the runs.

Fig. 13:

Same as Fig. 12, but relation $A_1=f(G-PT)$.

Fig. 14:

Same as Fig. 12, but function $A_1=f(G-E)$.

The following figures to Examples B through H basically correspond to the figures of example A; hence separate legends are not needed.

Fig. 19:

Vertical profiles of the incremental exchange coefficients

A_1 and water vapor pressure E for runs 2 through 7.

Fig. 43:

Vertical profiles of the incremental exchange coefficients

A_1 , recorded during runs 3 and 6.

E : upper boundary of the exchange layer

I : level of inversion.

Fig. 48:

Vertical profile of the incremental exchange coefficients

A_1 for runs 2, 3 and 4.

3. Individual Runs and Data Tables

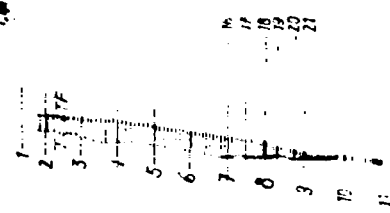
Examples A - J

Example
A
5 Oct. 1970

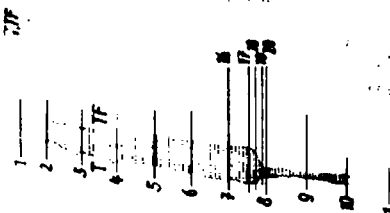
①



②



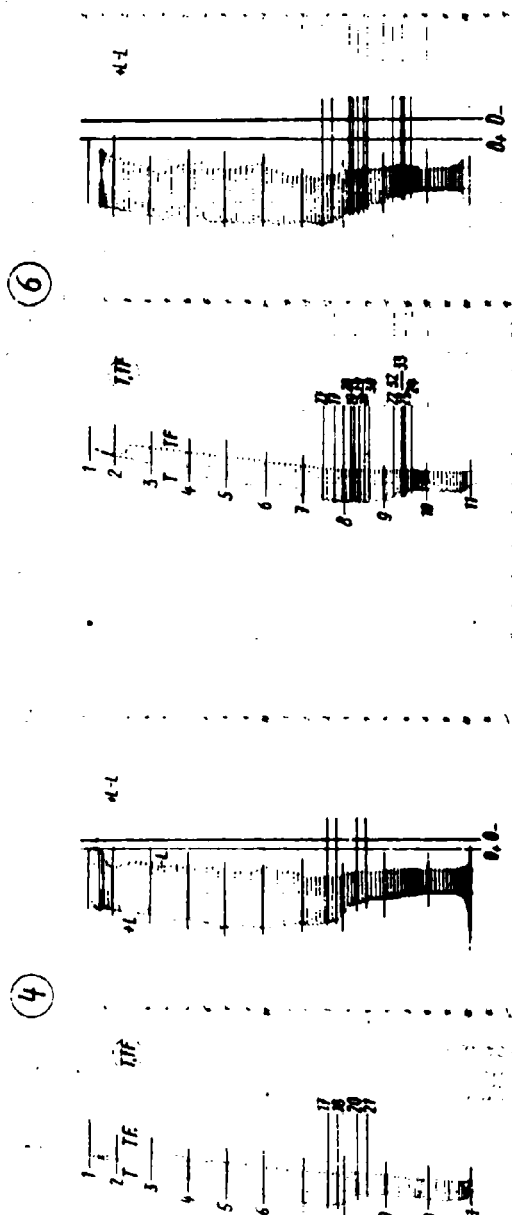
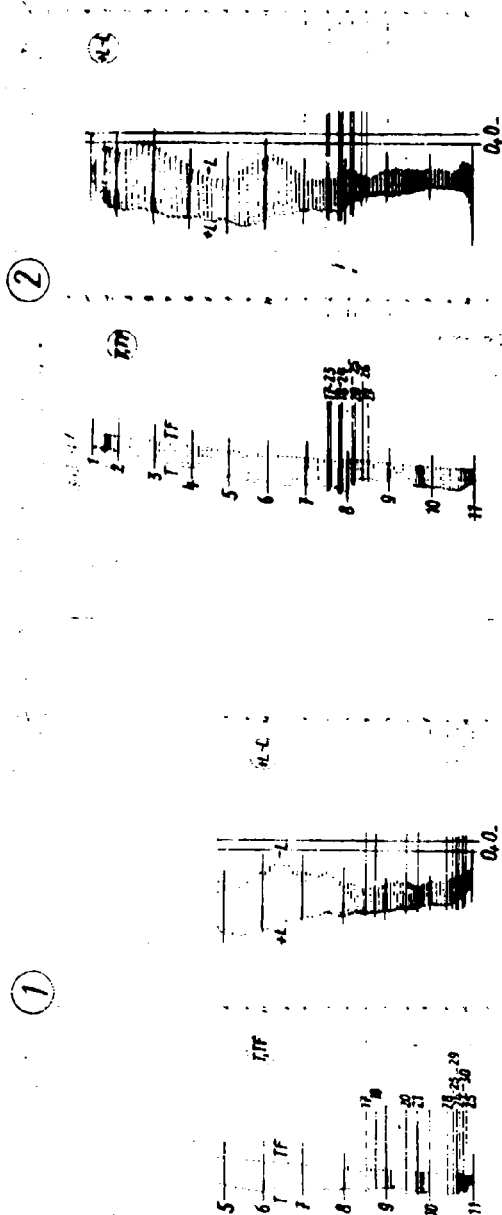
④



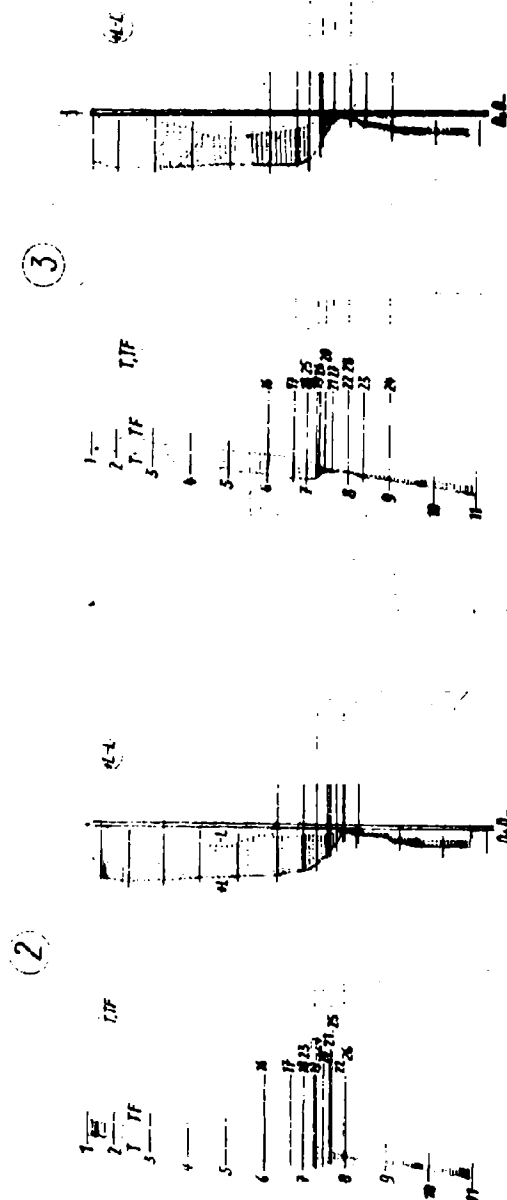
②



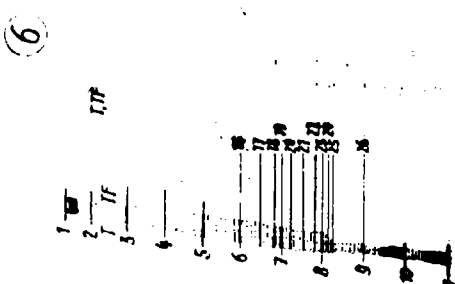
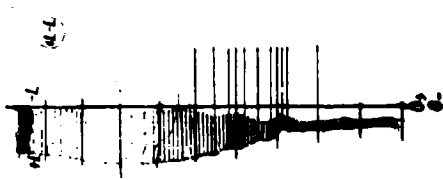
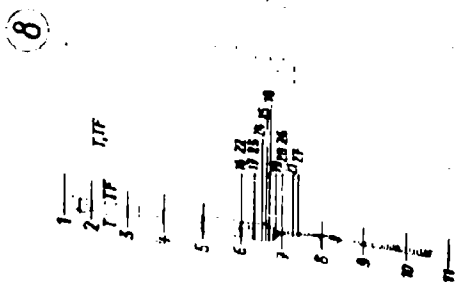
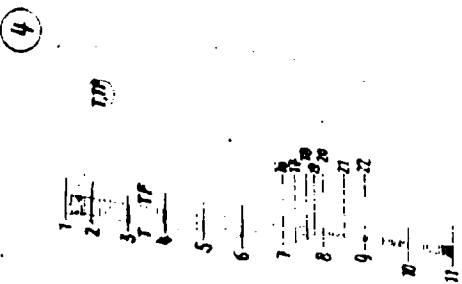
Example
B
13 Nov 1970



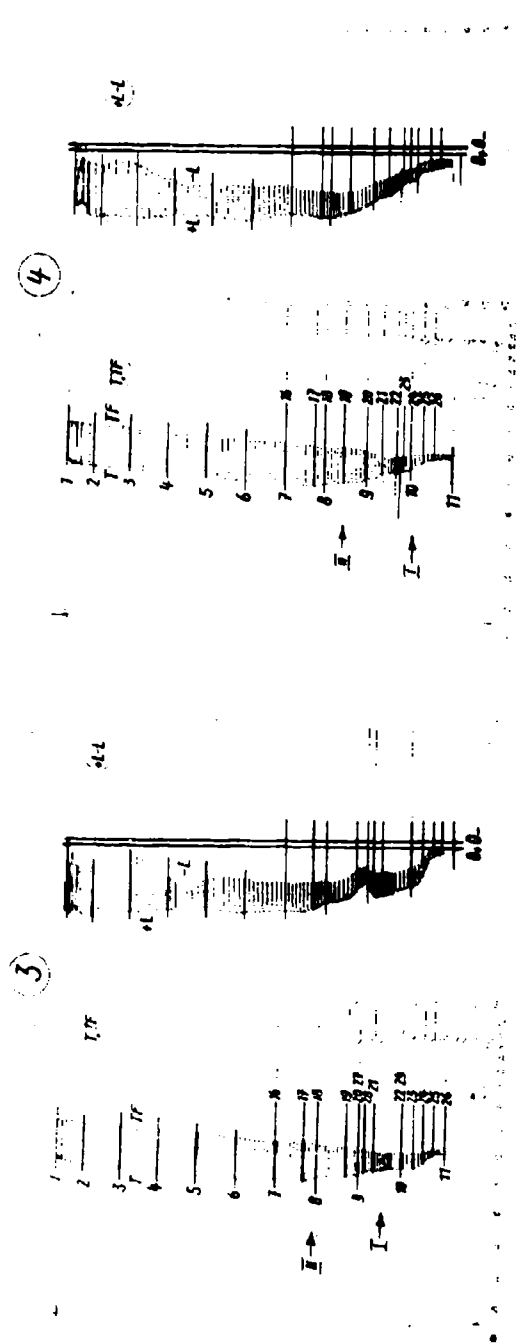
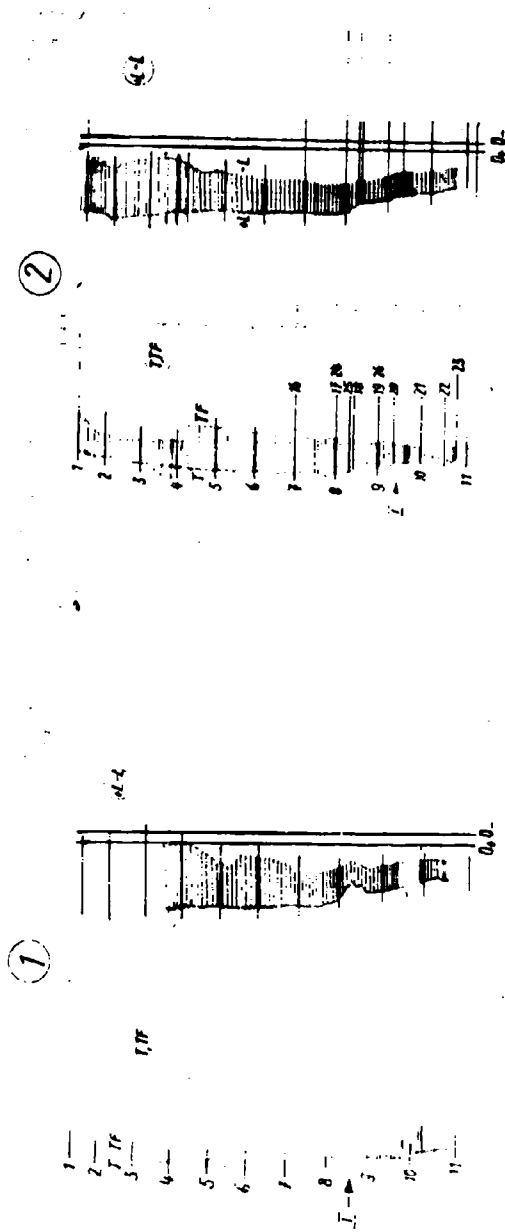
14 Oct 1970



Example
 C_2
 14 Oct 1970



Example
E₁
17 Oct 1970



Example

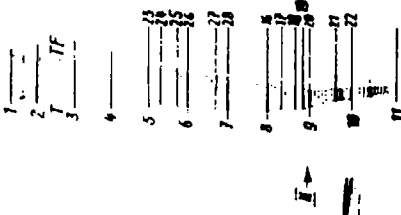
E_2

17 Oct. 1970

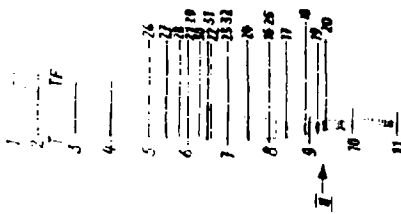
6



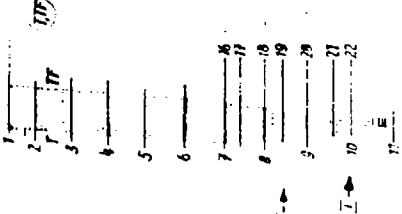
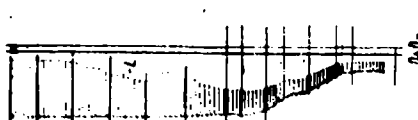
8



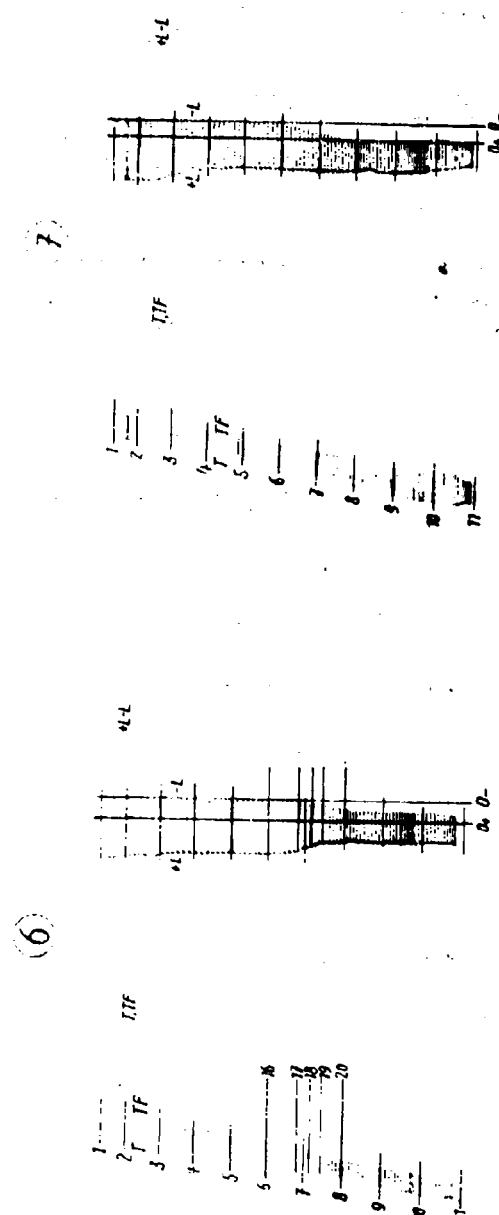
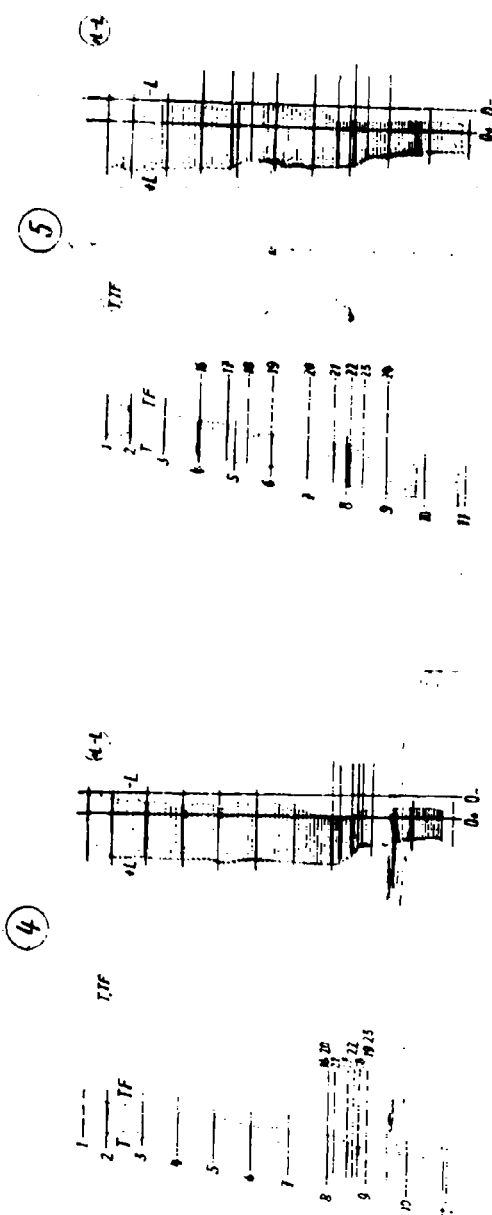
7



5



Example
F
30 Sept 1970



Example

G₁

28 Sept. 1970

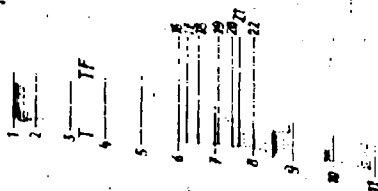
3

100



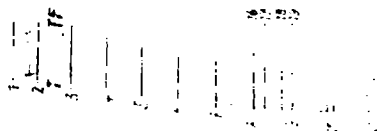
6

100



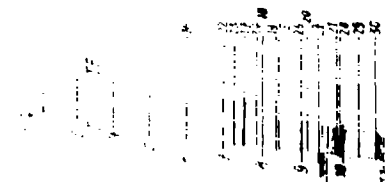
2

100

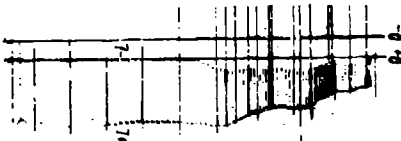


4

100



100



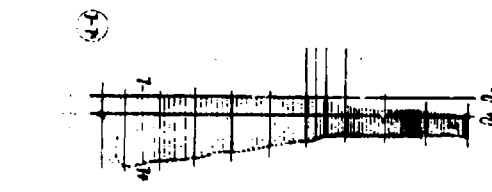
100



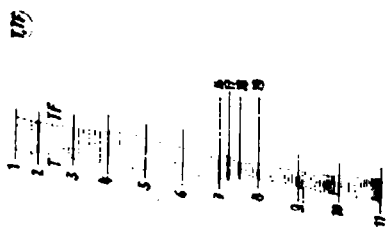
100



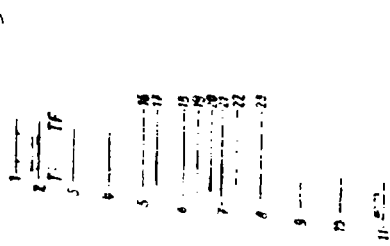
Example
G₂
28 Sept. 1970



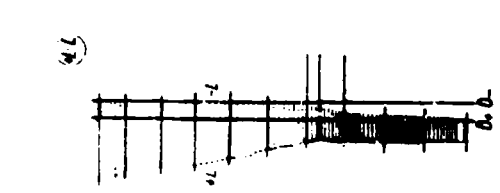
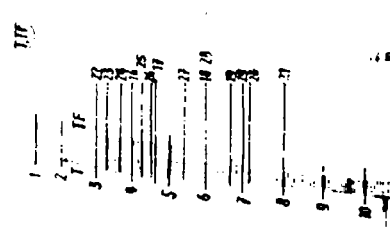
⑦



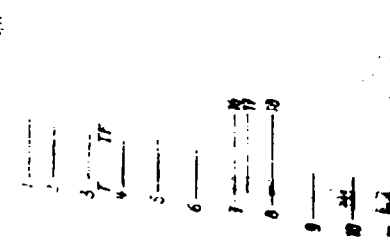
⑧



⑨



⑩



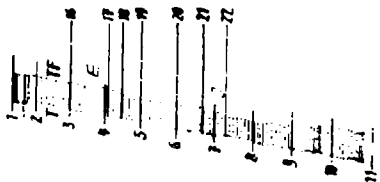
Example

11

26 Sept. 1970

4

43



6

117



3

117



5

117



4, 1



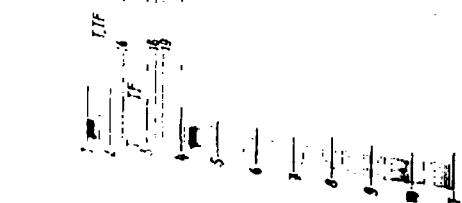
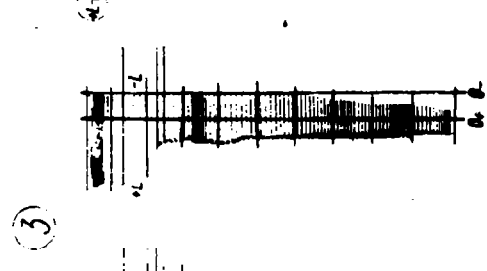
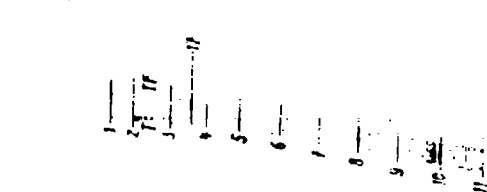
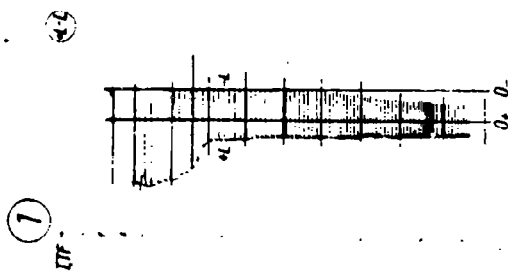
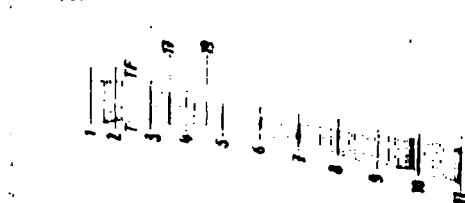
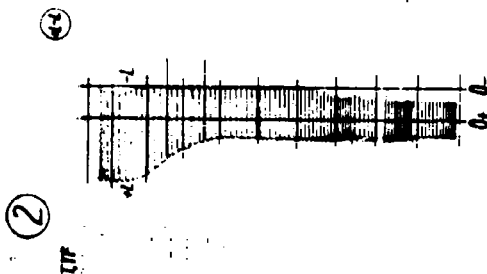
4, 1



Example

7

17 Sept. 1970



Life Cycle Savings

HAJTY-VILAGE

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92-1244-AT(4)

[illegible][illegible]

REF ID: A64166

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48 • J. Biol. Chem. 264:10,100-10,104

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ACKNOWLEDGMENTS

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$$\{f_{\alpha}^{\beta} \mid \alpha \in \mathcal{A}, \beta \in \mathcal{B}\} \subseteq \mathcal{C}$$
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Y	1987	1988	1989	1990	1991	1992	1993	1994	1995	1996	1997	1998	1999	2000	2001	2002	2003	2004	2005	2006	2007	2008	2009	2010	2011	2012	2013	2014	2015	2016	2017	2018	2019	2020	2021	2022	2023	2024	2025	2026	2027	2028	2029	2030	2031	2032	2033	2034	2035	2036	2037	2038	2039	2040	2041	2042	2043	2044	2045	2046	2047	2048	2049	2050	2051	2052	2053	2054	2055	2056	2057	2058	2059	2060	2061	2062	2063	2064	2065	2066	2067	2068	2069	2070	2071	2072	2073	2074	2075	2076	2077	2078	2079	2080	2081	2082	2083	2084	2085	2086	2087	2088	2089	2090	2091	2092	2093	2094	2095	2096	2097	2098	2099	2100	2101	2102	2103	2104	2105	2106	2107	2108	2109	2110	2111	2112	2113	2114	2115	2116	2117	2118	2119	2120	2121	2122	2123	2124	2125	2126	2127	2128	2129	2130	2131	2132	2133	2134	2135	2136	2137	2138	2139	2140	2141	2142	2143	2144	2145	2146	2147	2148	2149	2150	2151	2152	2153	2154	2155	2156	2157	2158	2159	2160	2161	2162	2163	2164	2165	2166	2167	2168	2169	2170	2171	2172	2173	2174	2175	2176	2177	2178	2179	2180	2181	2182	2183	2184	2185	2186	2187	2188	2189	2190	2191	2192	2193	2194	2195	2196	2197	2198	2199	2200	2201	2202	2203	2204	2205	2206	2207	2208	2209	2210	2211	2212	2213	2214	2215	2216	2217	2218	2219	2220	2221	2222	2223	2224	2225	2226	2227	2228	2229	2230	2231	2232	2233	2234	2235	2236	2237	2238	2239	2240	2241	2242	2243	2244	2245	2246	2247	2248	2249	2250	2251	2252	2253	2254	2255	2256	2257	2258	2259	2260	2261	2262	2263	2264	2265	2266	2267	2268	2269	2270	2271	2272	2273	2274	2275	2276	2277	2278	2279	2280	2281	2282	2283	2284	2285	2286	2287	2288	2289	2290	2291	2292	2293	2294	2295	2296	2297	2298	2299	2300	2301	2302	2303	2304	2305	2306	2307	2308	2309	2310	2311	2312	2313	2314	2315	2316	2317	2318	2319	2320	2321	2322	2323	2324	2325	2326	2327	2328	2329	2330	2331	2332	2333	2334	2335	2336	2337	2338	2339	2340	2341	2342	2343	2344	2345	2346	2347	2348	2349	2350	2351	2352	2353	2354	2355	2356	2357	2358	2359	2360	2361	2362	2363	2364	2365	2366	2367	2368	2369	2370	2371	2372	2373	2374	2375	2376	2377	2378	2379	2380	2381	2382	2383	2384	2385	2386	2387	2388	2389	2390	2391	2392	2393	2394	2395	2396	2397	2398	2399	2400	2401	2402	2403	2404	2405	2406	2407	2408	2409	2410	2411	2412	2413	2414	2415	2416	2417	2418	2419	2420	2421	2422	2423	2424	2425	2426	2427	2428	2429	2430	2431	2432	2433	2434	2435	2436	2437	2438	2439
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Year	1971	1972	1973	1974	1975	1976	1977	1978	1979	1980	1981	1982	1983	1984	1985	1986	1987	1988	1989	1990	1991	1992	1993	1994	1995	1996	1997	1998	1999	2000	2001	2002	2003	2004	2005	2006	2007	2008	2009	2010	2011	2012	2013	2014	2015	2016	2017	2018	2019	2020	2021	2022	2023	2024	2025	2026	2027	2028	2029	2030	2031	2032	2033	2034	2035	2036	2037	2038	2039	2040	2041	2042	2043	2044	2045	2046	2047	2048	2049	2050	2051	2052	2053	2054	2055	2056	2057	2058	2059	2060	2061	2062	2063	2064	2065	2066	2067	2068	2069	2070	2071	2072	2073	2074	2075	2076	2077	2078	2079	2080	2081	2082	2083	2084	2085	2086	2087	2088	2089	2090	2091	2092	2093	2094	2095	2096	2097	2098	2099	2100
1971	1971	1972	1973	1974	1975	1976	1977	1978	1979	1980	1981	1982	1983	1984	1985	1986	1987	1988	1989	1990	1991	1992	1993	1994	1995	1996	1997	1998	1999	2000	2001	2002	2003	2004	2005	2006	2007	2008	2009	2010	2011	2012	2013	2014	2015	2016	2017	2018	2019	2020	2021	2022	2023	2024	2025	2026	2027	2028	2029	2030	2031	2032	2033	2034	2035	2036	2037	2038	2039	2040	2041	2042	2043	2044	2045	2046	2047	2048	2049	2050	2051	2052	2053	2054	2055	2056	2057	2058	2059	2060	2061	2062	2063	2064	2065	2066	2067	2068	2069	2070	2071	2072	2073	2074	2075	2076	2077	2078	2079	2080	2081	2082	2083	2084	2085	2086	2087	2088	2089	2090	2091	2092	2093	2094	2095	2096	2097	2098	2099	2100

[illegible]

41 • 21:6-65(34:13)

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40270, 29.1.1970

12410 • J. Neurosci., July 26, 2006 • 26(30):12403–12410

14. *Journal of the American Medical Association*, 1997; 278: 1021-1025.

	1	2	3	4	5	6	7	8	9	10	11	12	13	14	15	16	17	18	19	20	21	22	23	24	25	26	27	28	29	30	31	32	33	34	35	36	37	38	39	40	41	42	43	44	45	46	47	48	49	50	51	52	53	54	55	56	57	58	59	60	61	62	63	64	65	66	67	68	69	70	71	72	73	74	75	76	77	78	79	80	81	82	83	84	85	86	87	88	89	90	91	92	93	94	95	96	97	98	99	100
1	1	2	3	4	5	6	7	8	9	10	11	12	13	14	15	16	17	18	19	20	21	22	23	24	25	26	27	28	29	30	31	32	33	34	35	36	37	38	39	40	41	42	43	44	45	46	47	48	49	50	51	52	53	54	55	56	57	58	59	60	61	62	63	64	65	66	67	68	69	70	71	72	73	74	75	76	77	78	79	80	81	82	83	84	85	86	87	88	89	90	91	92	93	94	95	96	97	98	99	100
2	1	2	3	4	5	6	7	8	9	10	11	12	13	14	15	16	17	18	19	20	21	22	23	24	25	26	27	28	29	30	31	32	33	34	35	36	37	38	39	40	41	42	43	44	45	46	47	48	49	50	51	52	53	54	55	56	57	58	59	60	61	62	63	64	65	66	67	68	69	70	71	72	73	74	75	76	77	78	79	80	81	82	83	84	85	86	87	88	89	90	91	92	93	94	95	96	97	98	99	100
3	1	2	3	4	5	6	7	8	9	10	11	12	13	14	15	16	17	18	19	20	21	22	23	24	25	26	27	28	29	30	31	32	33	34	35	36	37	38	39	40	41	42	43	44	45	46	47	48	49	50	51	52	53	54	55	56	57	58	59	60	61	62	63	64	65	66	67	68	69	70	71	72	73	74	75	76	77	78	79	80	81	82	83	84	85	86	87	88	89	90	91	92	93	94	95	96	97	98	99	100
4	1	2	3	4	5	6	7	8	9	10	11	12	13	14	15	16	17	18	19	20	21	22	23	24	25	26	27	28	29	30	31	32	33	34	35	36	37	38	39	40	41	42	43	44	45	46	47	48	49	50	51	52	53	54	55	56	57	58	59	60	61	62	63	64	65	66	67	68	69	70	71	72	73	74	75	76	77	78	79	80	81	82	83	84	85	86	87	88	89	90	91	92	93	94	95	96	97	98	99	100
5	1	2	3	4	5	6	7	8	9	10	11	12	13	14	15	16	17	18	19	20	21	22	23	24	25	26	27	28	29	30	31	32	33	34	35	36	37	38	39	40	41	42	43	44	45	46	47	48	49	50	51	52	53	54	55	56	57	58	59	60	61	62	63	64	65	66	67	68	69	70	71	72	73	74	75	76	77	78	79	80	81																			

	1	2	3	4	5	6	7	8	9	10	11	12	13	14	15	16	17	18	19	20	21	22	23	24	25	26	27	28	29	30	31	32	33	34	35	36	37	38	39	40	41	42	43	44	45	46	47	48	49	50	51	52	53	54	55	56	57	58	59	60	61	62	63	64	65	66	67	68	69	70	71	72	73	74	75	76	77	78	79	80	81	82	83	84	85	86	87	88	89	90	91	92	93	94	95	96	97	98	99	100
1	1	2	3	4	5	6	7	8	9	10	11	12	13	14	15	16	17	18	19	20	21	22	23	24	25	26	27	28	29	30	31	32	33	34	35	36	37	38	39	40	41	42	43	44	45	46	47	48	49	50	51	52	53	54	55	56	57	58	59	60	61	62	63	64	65	66	67	68	69	70	71	72	73	74	75	76	77	78	79	80	81	82	83	84	85	86	87	88	89	90	91	92	93	94	95	96	97	98	99	100
2	1	2	3	4	5	6	7	8	9	10	11	12	13	14	15	16	17	18	19	20	21	22	23	24	25	26	27	28	29	30	31	32	33	34	35	36	37	38	39	40	41	42	43	44	45	46	47	48	49	50	51	52	53	54	55	56	57	58	59	60	61	62	63	64	65	66	67	68	69	70	71	72	73	74	75	76	77	78	79	80	81	82	83	84	85	86	87	88	89	90	91	92	93	94	95	96	97	98	99	100
3	1	2	3	4	5	6	7	8	9	10	11	12	13	14	15	16	17	18	19	20	21	22	23	24	25	26	27	28	29	30	31	32	33	34	35	36	37	38	39	40	41	42	43	44	45	46	47	48	49	50	51	52	53	54	55	56	57	58	59	60	61	62	63	64	65	66	67	68	69	70	71	72	73	74	75	76	77	78	79	80	81	82	83	84	85	86	87	88	89	90	91	92	93	94	95	96	97	98	99	100
4	1	2	3	4	5	6	7	8	9	10	11	12	13	14	15	16	17	18	19	20	21	22	23	24	25	26	27	28	29	30	31	32	33	34	35	36	37	38	39	40	41	42	43	44	45	46	47	48	49	50	51	52	53	54	55	56	57	58	59	60	61	62	63	64	65	66	67	68	69	70	71	72	73	74	75	76	77	78	79	80	81	82	83	84	85	86	87	88	89	90	91	92	93	94	95	96	97	98	99	100
5	1	2	3	4	5	6																																																																																														

10 51-59

1. The first step is to identify the main components of the system. This includes the hardware, software, and data. The hardware components are the physical devices that make up the system, such as the computer, the network, and the storage devices. The software components are the programs and applications that run on the hardware. The data components are the information that is stored and processed by the system.

[illegible]

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[illegible]

Z	IN	TV	G-I	G-R	G-B	G-SP	I	N	A	L
01	0.04						0.00	0.00		
02	0.04	0.07	0.18	0.04	0.01	0.01	0.00	0.00	0.00	0.00
03	0.04	0.05	0.05	0.05	0.05	0.05	0.05	0.05	0.05	0.05
04	0.00	0.00	0.01	0.01	0.13	0.00	0.00	0.00	0.00	0.00
05	0.00	0.00		0.09	0.10	0.05	0.17	0.00	0.00	0.00
06	0.00	0.05		0.01	0.04	0.00	0.17	0.05	0.00	0.00
07	0.01	0.00	0.22							0.00
08	0.00	0.00		0.33	0.10	0.00	0.24	0.00	0.00	0.00
09	0.01	0.15	0.72	0.40	0.03	0.10	0.21	0.01	0.00	0.00
10	0.00	0.00		0.51	0.20	0.01	0.21	0.05	0.00	0.00
11	0.00	0.15	0.75							0.00

4P-TE'P10ATL4

Z	M	P	T	TF	PT	Exp	b	u	NP	OL	UL	+L
17	1801	8475	+0.61	+0.61	1934	0.019	0.003	0.17	31.4	0.01	0.14	1.25 23.5
18	1801	8475	+0.61	+0.61	1934	0.019	0.003	0.17	31.4	0.01	0.14	1.25 23.5
19	1801	8475	+0.61	+0.61	1934	0.019	0.003	0.17	31.4	0.01	0.14	1.25 23.5
20	1801	8475	+0.61	+0.61	1934	0.019	0.003	0.17	31.4	0.01	0.14	1.25 23.5
21	1801	8475	+0.61	+0.61	1934	0.019	0.003	0.17	31.4	0.01	0.14	1.25 23.5
22	1801	8475	+0.61	+0.61	1934	0.019	0.003	0.17	31.4	0.01	0.14	1.25 23.5
23	1801	8475	+0.61	+0.61	1934	0.019	0.003	0.17	31.4	0.01	0.14	1.25 23.5
24	1801	8475	+0.61	+0.61	1934	0.019	0.003	0.17	31.4	0.01	0.14	1.25 23.5
25	1801	8475	+0.61	+0.61	1934	0.019	0.003	0.17	31.4	0.01	0.14	1.25 23.5
26	1801	8475	+0.61	+0.61	1934	0.019	0.003	0.17	31.4	0.01	0.14	1.25 23.5
27	1801	8475	+0.61	+0.61	1934	0.019	0.003	0.17	31.4	0.01	0.14	1.25 23.5
28	1801	8475	+0.61	+0.61	1934	0.019	0.003	0.17	31.4	0.01	0.14	1.25 23.5

Z	LM	IV	G = I	G = E	G = B	II = B	G = P	I	N	A	Z
16		+3.64						0.24	0.28		
17	0.63	+0.75	+0.67	+0.83	-0.38	-0.37	+0.74	0.61	0.58	10.38	+0.45
18	0.69	+0.73	+0.67	+0.83	-0.38	-0.37	+0.74	0.61	0.58	10.38	+0.45
19	0.66	+0.61	+0.50	+0.13	-0.55	-0.40	+0.66	0.58	0.51	4.34	+0.88
20	0.64	+0.61	+0.50	+0.13	-0.55	-0.40	+0.66	0.58	0.51	4.34	+0.88
21	0.60	+0.68	+0.71	+0.12	+0.37	-0.37	+0.74	+0.66	0.53	2.74	+0.66
22	0.69	+0.87	+0.87	+0.14	+0.66	-0.66	+0.88	+0.87	0.71	0.56	+0.57
23	0.50	+0.68	+0.19	+0.19	+0.66	-0.66	+0.88	+0.87	0.71	0.56	+0.57

Mr. FALCETTI: UNP.

Z	M	P	T	Y	PT	F	g	S	h	OL	UL	OF	L
83	0079	7414	+03.4	+08.6	2676	2201	01.10	00.87	11.6	00.3	///	1.36	5.60
84	0000	7796	+03.2	+01.6	2716	2050	01.30	01.17	14.0	00.8	///	1.26	5.64

85 1694 8094
86 1650 8147

87	1790	5636	+058	+C0.5	5434	1948	DRAB	Clnd	Mar	7	042	7	16	030
88	1861	6409	+069	+C0.1	5743	1490	DR23	Clnd	Mar	7	31.5	Clnd	16	030
89	1861	6409	+069	+C0.1	5743	1490	DR23	Clnd	Mar	7	31.5	Clnd	16	030
90	1861	6409	+069	+C0.1	5743	1490	DR23	Clnd	Mar	7	31.5	Clnd	16	030
91	1861	6409	+069	+C0.1	5743	1490	DR23	Clnd	Mar	7	31.5	Clnd	16	030
92	1861	6409	+069	+C0.1	5743	1490	DR23	Clnd	Mar	7	31.5	Clnd	16	030
93	1861	6409	+069	+C0.1	5743	1490	DR23	Clnd	Mar	7	31.5	Clnd	16	030
94	1861	6409	+069	+C0.1	5743	1490	DR23	Clnd	Mar	7	31.5	Clnd	16	030
95	1861	6409	+069	+C0.1	5743	1490	DR23	Clnd	Mar	7	31.5	Clnd	16	030
96	1861	6409	+069	+C0.1	5743	1490	DR23	Clnd	Mar	7	31.5	Clnd	16	030
97	1861	6409	+069	+C0.1	5743	1490	DR23	Clnd	Mar	7	31.5	Clnd	16	030
98	1861	6409	+069	+C0.1	5743	1490	DR23	Clnd	Mar	7	31.5	Clnd	16	030
99	1861	6409	+069	+C0.1	5743	1490	DR23	Clnd	Mar	7	31.5	Clnd	16	030
100	1861	6409	+069	+C0.1	5743	1490	DR23	Clnd	Mar	7	31.5	Clnd	16	030

4-0011 17-10-1970 16.28 UH 

HAUPT-VERLAUSCH

[illegible]

01	th	74	G-T	G-E	G-b	G-1.3	G-2.1	I	G	A	E	Z
04		+334						G-75	U-1.4			U-1
08		+340	+0.4d	+0.67	+0.57	+1.14	+3.43	G-2.2	U-1.4			U-1

03	+C97	+C
04	+09C	+0

[illegible]

1A	1500	5479	036.9	031.1	WJ29	1010	031.74	03.42	48.4	01.3	02.9	1.5	0.6
1B	1332	5448	036.4	031.4	WJ26	1016	044.91	03.48	51.1	01.3	00.8	1.5	0.6
1C	1447	5363	035.7	031.5	WJ14	1054	034.75	03.44	50.5	01.2	01.4	1.5	0.6
1D	1447	5500	036.6	032.4	WJ11	1011	034.25	03.78	37.4	01.0	01.7	1.5	0.6
MD	1500	8241	036.3	032.3	WJ35	1010	034.55	03.34	61.7	01.0	01.7	1.5	0.6

1	PRC	6751	*02-1	*02-1	2150	8991	05-05	04-24	78-6	Q0-9	Q0-6	1-24	01-5
Md	1800	4632	*02-8	*02-3	8898	R474	0A-6G	C4-14	74-5	C0-7	G0-4	1-24	01-1

[illegible]

22 679 +543 -0.

DATE	TIME	FROM	TO	REMARKS
1964	10:00	10:00	10:00	10:00

74	1917	7478	+05.5	-01.4	24.1	30.12	13.38	01-07	15.66	01-07	15.52	01-04
75	1951	7450	+03.4	-01.4	27.7	17.13	03-06	01-07	15.66	01-07	15.53	01-04
76	1949	5091	+04.3	-00.3	26.9	17.13	01-01	01-04	15.66	01-01	15.53	01-04
77	1956	6143	+05.7	-05.1	22.57	31.17	04-09	09-04	15.66	01-01	15.53	01-04
78	1956	4103	+05.1	+00.2	26.52	31.16	04-09	09-04	15.66	01-01	15.53	01-04

[illegible][illegible]

Z	IN	TG	G-7	G-E	G-3	G-NP	G-27	I	A	D
1	0.517									
2	0.043	-0.26	-0.38	-0.24	-0.50	-0.17	-0.28	0.28	1.30	50
3	0.34	-0.22	-0.01	-0.25	-0.57	-0.78	-0.04	0.62	4.00	01
4	0.52	-0.70	-0.31	-0.62	-0.30	-0.68	-0.10	0.73	1.04	59
5	0.58	-0.23	-0.31	-0.26	-0.22	-0.44	-0.77	0.41	51	15
6	0.15	-0.32	-0.18	-0.18	-0.32	-0.18	0.67	0.60	7.00	03
7	0.08	-0.31	-0.11	-0.31	-0.41	-0.31	-0.31	0.70	1.10	00
8	0.28	-0.13	-0.14	-0.43	-0.45	-0.21	-0.08	0.49	0.90	14
9	0.34	-0.30	-0.31	-0.56	-0.73	-0.30	-0.31	0.41	61	50
10	0.48	-0.70	-0.37	-0.88	-1.36	-0.76	-0.25	3.00	6.80	15

11 1024 = 3

MP-TEMPERATURE														
Z	P	T	II	PI	EPT	B	S	MF	CL	AL	PI	CL		
16	1712	8193	0065	-309	2957	9956	01.08	21.94	16.0	056	///	16.37	0346	
17	1717	9277	0064	-306	2957	9950	01.08	21.94	15.3	063	///	16.37	0346	
18	1808	9369	0063	-300	2923	2931	00.93	21.92	31.6	026	016	16.39	0346	
19	1915	648	0056	-301	2923	9950	00.98	08.98	43.7	017	016	16.38	0346	
20	1989	611	0053	-301	2906	9952	00.98	02.08	50.9	018	016	16.14	0346	
21	1970	8758	0054	-303	2887	9955	00.96	04.14	68.4	008	005	16.48	0116	

[illegible]

82 067 4046 -1.27

2-0416 17.10.1970 10.11 UHh

AUP = 0.7142857142857143													
	Case	Y	X	Y	X	Y	X	Y	X	Y	X	Y	X
01	30020	7189	0.0	0.0	1.056	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
02	00000	7240	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
03	82820	7418	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
04	82820	7418	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
05	14000	7418	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
06	14000	7418	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
07	14000	7418	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
08	14000	7418	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
09	14000	7418	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
10	14000	7418	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
11	14000	7418	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
12	14000	7418	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
13	14000	7418	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
14	14000	7418	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
15	14000	7418	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
16	14000	7418	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
17	14000	7418	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
18	14000	7418	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
19	14000	7418	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
20	14000	7418	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
21	14000	7418	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
22	14000	7418	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
23	14000	7418	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
24	14000	7418	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
25	14000	7418	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
26	14000	7418	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
27	14000	7418	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
28	14000	7418	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
29	14000	7418	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
30	14000	7418	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
31	14000	7418	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
32	14000	7418	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
33	14000	7418	0.0	0.0	0.0</								

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14	15	16	17	18	19	20	21	22	23	24	25	26	27	28	29	30	31	32	33	34	35	36	37	38	39	40	41	42	43	44	45	46	47	48	49	50	51	52	53	54	55	56	57	58	59	60	61	62	63	64	65	66	67	68	69	70	71	72	73	74	75	76	77	78	79	80	81	82	83	84	85	86	87	88	89	90	91	92	93	94	95	96	97	98	99	100
14	15	16	17	18	19	20	21	22	23	24	25	26	27	28	29	30	31	32	33	34	35	36	37	38	39	40	41	42	43	44	45	46	47	48	49	50	51	52	53	54	55	56	57	58	59	60	61	62	63	64	65	66	67	68	69	70	71	72	73	74	75	76	77	78	79	80	81	82	83	84	85	86	87	88	89	90	91	92	93	94	95	96	97	98	99	100

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1	2	3	4	5	6	7	8	9	10	11	12	13	14	15	16	17	18	19	20	21	22	23	24	25	26	27	28	29	30	31	32	33	34	35	36	37	38	39	40	41	42	43	44	45	46	47	48	49	50	51	52	53	54	55	56	57	58	59	60	61	62	63	64	65	66	67	68	69	70	71	72	73	74	75	76	77	78	79	80	81	82	83	84	85	86	87	88	89	90	91	92	93	94	95	96	97	98	99	100
1	2	3	4	5	6	7	8	9	10	11	12	13	14	15	16	17	18	19	20	21	22	23	24	25	26	27	28	29	30	31	32	33	34	35	36	37	38	39	40	41	42	43	44	45	46	47	48	49	50	51	52	53	54	55	56	57	58	59	60	61	62	63	64	65	66	67	68	69	70	71	72	73	74	75	76	77	78	79	80	81	82	83	84	85	86	87	88	89	90	91	92	93	94	95	96	97	98	99	100

22	049	+099	+0110	-0105	-1174	+0264	+1117	0024	0033	17818	14238	0
23	150	+060	+0120	-0127	+0130	+0150	0018	0021	0040	15199	15635	0
25	099	+050	+0110	-0125	+0122	+0128	+0116	0024	0028	10210	17264	0

MS-DOS 5.0 FILES

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DATE	TIME	TYPE	FROM	TO	REMARKS	INITIALS	REMARKS	INITIALS
17	05:30	1	100	100	100	100	100	100
18	05:30	1	100	100	100	100	100	100
19	05:30	1	100	100	100	100	100	100
20	05:30	1	100	100	100	100	100	100
21	05:30	1	100	100	100	100	100	100
22	05:30	1	100	100	100	100	100	100
23	05:30	1	100	100	100	100	100	100
24	05:30	1	100	100	100	100	100	100
25	05:30	1	100	100	100	100	100	100
26	05:30	1	100	100	100	100	100	100
27	05:30	1	100	100	100	100	100	100
28	05:30	1	100	100	100	100	100	100
29	05:30	1	100	100	100	100	100	100
30	05:30	1	100	100	100	100	100	100
31	05:30	1	100	100	100	100	100	100

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MP-LFI TRAINING-FIT

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10-2-2 UN

КАЛУЖА - 41 СЪСТАВИ

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• **100% Satisfaction Guarantee**

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Test Methodology

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15528 86

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4.2.1. *Model 1*

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19 507 +104 +0.16

$M = \{a, b, c, d, e, f, g, h, i, j, k, l, m, n, o, p, q, r, s, t, u, v, w, x, y, z\}$
 $....$
 $L = \{0, 1, 2, 3, 4, 5, 6, 7, 8, 9\}$ $\{a, b, c, d, e, f, g, h, i, j, k, l, m, n, o, p, q, r, s, t, u, v, w, x, y, z\}$
 $q_1(a)(b) = \{a, b, c, d, e, f, g, h, i, j, k, l, m, n, o, p, q, r, s, t, u, v, w, x, y, z\}$

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93 = 16, 17, 18, 19, ...

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Year	1971	1972	1973	1974	1975	1976	1977	1978	1979	1980	1981	1982	1983	1984	1985	1986	1987	1988	1989	1990	1991	1992	1993	1994	1995	1996	1997	1998	1999	2000	2001	2002	2003	2004	2005	2006	2007	2008	2009	2010	2011	2012	2013	2014	2015	2016	2017	2018	2019	2020	2021	2022	2023	2024	2025	2026	2027	2028	2029	2030	2031	2032	2033	2034	2035	2036	2037	2038	2039	2040	2041	2042	2043	2044	2045	2046	2047	2048	2049	2050	2051	2052	2053	2054	2055	2056	2057	2058	2059	2060	2061	2062	2063	2064	2065	2066	2067	2068	2069	2070	2071	2072	2073	2074	2075	2076	2077	2078	2079	2080	2081	2082	2083	2084	2085	2086	2087	2088	2089	2090	2091	2092	2093	2094	2095	2096	2097	2098	2099	2100
1971	1971	1972	1973	1974	1975	1976	1977	1978	1979	1980	1981	1982	1983	1984	1985	1986	1987	1988	1989	1990	1991	1992	1993	1994	1995	1996	1997	1998	1999	2000	2001	2002	2003	2004	2005	2006	2007	2008	2009	2010	2011	2012	2013	2014	2015	2016	2017	2018	2019	2020	2021	2022	2023	2024	2025	2026	2027	2028	2029	2030	2031	2032	2033	2034	2035	2036	2037	2038	2039	2040	2041	2042	2043	2044	2045	2046	2047	2048	2049	2050	2051	2052	2053	2054	2055	2056	2057	2058	2059	2060	2061	2062	2063	2064	2065	2066	2067	2068	2069	2070	2071	2072	2073	2074	2075	2076	2077	2078	2079	2080	2081	2082	2083	2084	2085	2086	2087	2088	2089	2090	2091	2092	2093	2094	2095	2096	2097	2098	2099	2100

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14,150 Adult education students are enrolled in 100,000 credit hours with 100,000 credit hours

4. Tables I - XII

TABLE I

RaB in air; units: $10^{-12} \mu\text{c}/\text{cm}^3$

Station designations: Z Zugspitze 2964 m a.s.l.
 W Wank 1780 m a.s.l.
 G Garmisch-Partenkirchen 730 m a.s.l.

Measurement periods : a = 2150 - 0920, b = 0950 - 1220,
 c = 1250 - 1620. (Central European Time)

Month: July 1970 ("No measurement" is denoted by: -)

	a			b			c		
	Z	W	G	Z	W	G	Z	W	G
1.	8,5	7	130	14,2	16	82	7,9	61	85
2.	-	57	233	12,2	24	64	14,8	-	58
3.	8,1	28	108	8,8	16	43	7,1	10	37
4.	3,3	18	130	1,9	11	46	3,9	17	57
5.	-	-	171	-	-	54	-	-	-
6.	4,2	-	263	4,9	57	145	17,5	64	92
7.	4,1	49	299	20,4	109	160	60,6	108	150
8.	5,8	116	256	71,7	145	237	90,3	154	229
9.	-	75	194	55,9	76	128	31,8	11	57
10.	32,7	90	304	58,7	81	116	62,9	81	86
11.	100,3	135	333	86,5	146	151	92,8	90	129
12.	-	-	268	-	-	115	-	-	-
13.	49,7	150	-	80,9	171	228	96,2	141	189
14.	65,5	87	313	68,2	113	180	66,9	103	131
15.	3,7	4	35	2,5	7	39	1,5	10	53
16.	3,2	7	80	1,3	4	58	0,8	-	38
17.	1,7	2	54	0,5	4	50	0,6	-	46
18.	2,6	15	180	7,7	35	93	5,4	21	53
19.	-	-	172	-	-	88	-	-	-
20.	75,1	-	306	99,9	157	208	93,2	168	153
21.	63,7	63	148	79,3	77	135	55,9	86	89
22.	69,7	156	336	64,0	150	200	85,1	113	138
23.	68,3	107	320	78,3	157	291	114,2	105	147
24.	75,8	75	200	84,4	126	-	88,2	104	183
25.	1,6	14	-	1,6	9	-	1,2	14	38
26.	-	-	-	-	-	38	-	-	-
27.	1,4	15	173	2,1	15	128	3,4	23	113
28.	5,2	74	316	13,3	46	190	69,3	108	186
29.	77,4	68	307	111,4	110	131	17,6	32	96
30.	49,3	97	304	58,1	150	213	73,1	118	152
31.	44,2	89	240	55,2	126	184	65,4	46	-

TABLE II

RAB in air; units: $10^{-12} \mu\text{C}/\text{cm}^3$

Station Designations: Z Zugspitze 2964 m a.s.l.
 W Wank 1780 m a.s.l.
 G Garmisch-Partenkirchen 730 m a.s.l.

Measurement periods : a = 2150 - 0920, b = 0950 - 1220,
 c = 1250 - 1620. (Central European Time)

Month: August 1970 ("No measurement" is denoted by: -)

	a			b			c		
	Z	W	G	Z	W	G	Z	W	G
1.	22,1	49	297	40,5	112	246	51,2	44	65
2.	-	-	-	-	-	174	-	-	-
3.	31,8	101	-	63,2	111	-	67,2	77	-
4.	38,8	69	286	41,0	36	148	70,4	108	145
5.	110,3	147	359	99,8	125	143	89,7	82	-
6.	71,3	160	333	76,2	166	268	97,0	137	191
7.	142,6	107	349	87,5	95	188	73,0	94	-
8.	50,4	49	242	45,6	87	120	25,2	48	96
9.	-	-	161	-	-	130	-	-	-
10.	1,5	-	569	1,6	4	338	-	7	431
11.	0,3	32	466	-	58	174	-	26	112
12.	2,7	18	165	10,7	50	107	46,6	56	-
13.	28,4	98	-	39,8	105	-	78,5	122	132
14.	83,7	107	388	73,5	136	200	53,4	77	100
15.	-	-	-	-	-	-	-	-	-
16.	-	-	-	-	-	-	-	-	-
17.	28,3	98	268	0,8	17	153	3,6	18	109
18.	15,6	67	215	13,3	103	169	40,0	99	133
19.	70,5	172	426	85,6	232	295	105,0	154	134
20.	77,8	110	449	101,7	132	150	53,7	23	110
21.	3,8	10	90	4,4	9	66	-	9	61
22.	5,1	25	151	7,8	35	93	28,6	41	68
23.	-	-	258	-	-	109	-	-	-
24.	-	-	194	1,7	20	243	1,2	19	102
25.	21,9	62	304	19,9	91	158	35,3	76	93
26.	15,5	101	306	24,1	128	207	82,2	86	117
27.	72,0	100	382	78,3	104	203	86,3	125	189
28.	28,3	48	217	50,7	58	130	82,6	70	153
29.	75,2	115	318	94,5	140	240	87,4	156	210
30.	-	-	503	-	-	174	-	-	-
31.	33,1	-	306	36,0	121	198	39,5	101	133

TABLE III

RaB in air; units: $10^{-12} \mu\text{C}/\text{cm}^3$

Station Designations: Z Zugspitze 2964 m a.s.l.
 W Wank 1780 m a.s.l.
 G Garmisch-Partenkirchen 730 m a.s.l.

Measurement periods : a = 2150 - 0920, b = 0950 - 1220,
 c = 1250 - 1620. (Central European Time)

Month: September 1970 ("No measurement" is denoted by: -)

	a			b			c		
	Z	W	G	Z	W	G	Z	W	G
1.	8,3	56	535	3,7	33	117	4,8	85	198
2.	38,4	80	167	45,5	114	154	88,5	129	120
3.	76,4	124	448	104,4	127	250	60,8	123	247
4.	2,4	41	279	3,7	22	67	13,2	30	59
5.	2,7	22	342	-	4	114	1,4	16	61
6.	-	-	92	-	-	57	-	-	-
7.	3,1	40	136	10,4	54	115	26,4	82	81
8.	65,6	120	276	44,5	131	230	84,7	153	160
9.	67,9	136	279	69,7	181	255	121,5	163	158
10.	54,0	112	301	59,7	101	157	50,9	91	181
11.	63,2	91	326	62,8	65	105	15,4	12	75
12.	37,8	20	88	51,8	35	104	65,3	47	99
13.	-	-	107	-	-	44	-	-	-
14.	29,4	58	258	26,3	153	160	63,4	101	103
15.	83,7	-	394	124,8	131	208	138,7	156	159
16.	4,3	16	74	2,3	11	84	1,8	12	37
17.	6,3	38	235	3,5	28	53	1,7	33	59
18.	5,5	26	275	11,1	53	184	24,0	83	124
19.	27,3	108	380	26,9	177	376	43,9	193	318
20.	-	-	504	-	-	237	-	-	-
21.	8,0	46	527	1,6	92	150	-	91	134
22.	1,6	48	697	12,4	143	201	36,8	94	119
23.	4,8	89	273	5,9	117	153	56,3	84	122
24.	8,4	84	410	0,9	109	201	24,3	132	167
25.	2,6	10	331	5,6	70	175	14,3	61	133
26.	5,4	41	401	3,2	138	230	31,2	170	164
27.	-	-	419	-	-	257	-	-	-
28.	5,8	19	311	6,6	74	265	22,3	144	211
29.	13,6	64	384	36,2	109	315	64,3	200	333
30.	5,9	44	329	12,8	50	200	41,6	146	226

TABLE IV

RaB in air; units $10^{-12} \mu\text{c}/\text{cm}^3$

Station Designations: Z Zugspitze 2964 m a.s.l.
 W Wank 1780 m a.s.l.
 G Garmisch-Partenkirchen 730 m a.s.l.

Measurement periods : a = 2150 - 0920, b = 0950 - 1220,
 c = 1250 - 1620. (Central European Time)

Month: October 1970 ("No measurement" is denoted by: -)

	a			b			c		
	Z	W	G	Z	W	G	Z	W	G
1.	-	4	130	-	6	47	-	2	53
2.	1,7	16	99	0,2	9	96	0,1	8	34
3.	2,7	3	45	0,9	-	35	2,4	2	15
4.	-	-	59	-	-	56	-	-	-
5.	1,1	11	175	0,6	11	145	0,9	27	126
6.	15,5	22	190	20,6	74	143	50,3	103	174
7.	24,8	86	210	29,3	77	107	65,6	80	109
8.	40,3	43	251	26,3	61	245	34,8	50	294
9.	40,7	49	631	40,5	40	410	45,9	74	401
10.	27,7	71	276	28,6	74	123	52,8	103	392
11.	-	-	359	-	-	653	-	-	-
12.	12,4	51	238	9,9	124	431	14,0	155	399
13.	8,8	84	407	7,0	112	382	15,9	168	270
14.	16,4	65	727	20,2	232	373	46,4	192	392
15.	4,7	78	268	5,4	150	200	5,7	34	141
16.	7,4	61	412	9,9	82	322	16,7	78	141
17.	6,8	61	700	4,4	79	361	16,6	145	354
18.	-	-	395	-	-	464	-	-	-
19.	11,8	18	297	13,6	31	282	19,6	69	203
20.	13,1	11	71	6,0	4	136	8,7	3	171
21.	18,9	25	159	5,2	21	74	4,1	20	74
22.	11,1	32	226	14,7	19	63	13,4	40	90
23.	-	7	129	3,5	3	57	3,2	4	18
24.	14,2	23	175	9,6	61	112	12,9	69	102
25.	-	-	221	-	-	117	-	-	-
26.	-	9	-	2,3	10	345	6,6	36	200
27.	13,3	44	493	13,4	33	345	16,4	105	193
28.	7,2	38	487	6,0	66	264	16,1	60	169
29.	3,4	17	302	0,8	9	138	2,0	7	149
30.	6,5	15	286	2,5	35	208	8,4	25	205
31.	7,6	13	208	7,3	18	305	14,6	21	252

TABLE V

Station Designations: Z Zugspitze 2964 m a.s.l.
W Wank 1780 m a.s.l.
G Garmisch-Partenkirchen 730 m a.s.l.

Measurement periods : a = 2150 - 0920, b = 0950 - 1220,
c = 1250 - 1620. (Central European Time)

Month: November 1970 ("No measurement" is denoted by: -)

	a			b			c		
	Z	W	G	Z	W	G	Z	W	G
1.	-	-	221	-	-	145	-	-	-
2.	4,8	10	271	9,8	7	82	12,0	10	56
3.	14,4	22	276	11,6	9	35	12,2	11	34
4.	26,9	51	319	19,4	65	129	24,8	77	134
5.	7,0	15	416	8,3	10	196	4,2	5	36
6.	14,2	9	54	4,5	8	41	9,3	16	58
7.	20,1	65	384	19,7	125	320	49,5	193	289
8.	-	-	446	-	-	383	-	-	-
9.	35,2	48	484	15,2	47	373	13,5	29	91
10.	10,8	17	64	8,6	14	45	7,8	13	46
11.	4,7	28	334	2,6	25	196	2,3	6	270
12.	11,0	30	366	11,4	41	222	17,7	83	177
13.	82,9	83	415	80,1	136	241	86,8	116	183
14.	65,3	88	521	61,1	136	310	67,8	26	263
15.	-	-	211	-	-	91	-	-	-
16.	7,9	5	109	10,3	4	78	17,0	7	79
17.	3,3	21	315	1,3	10	228	0,7	2	227
18.	16,6	36	327	18,5	70	291	28,1	78	245
19.	32,2	32	302	39,4	29	209	48,1	70	146
20.	11,7	10	193	10,6	7	189	9,3	28	99
21.	45,8	-	384	57,9	-	174	44,3	-	61
22.	-	-	194	-	-	86	-	-	-
23.	3,4	19	306	2,9	14	145	-	23	146
24.	2,8	7	236	1,9	5	188	1,3	5	209
25.	4,7	22	245	3,6	16	189	1,3	83	294
26.	5,9	11	285	0,3	25	271	-	47	247
27.	3,4	24	551	0,7	40	308	0,4	61	403
28.	2,9	-	202	2,1	-	187	23,5	-	158
29.	-	-	197	-	-	340	-	-	-
30.	57,7	33	546	7,6	5	620	7,2	14	261

TABLE VI

Station Designation: Z Zugspitze 2964 m a.s.l.
W Wank 1780 m a.s.l.
G Garmisch-Partenkirchen 730 m a.s.l.

Measurement periods: a = 2150 - 0920, b = 0950 - 1220.
b = 1250 - 1620. (Central European Time)

Month: December 1970 ("No measurement" is denoted by: -)

	a			b			c		
	Z	W	G	Z	W	G	Z	W	G
1.	22,9	70	350	24,9	85	292	55,4	92	315
2.	6,7	9	285	7,5	18	175	7,5	22	146
3.	10,9	27	279	-	21	224	-	4	135
4.	-	11	233	4,1	15	250	-	-	181
5.	15,0	-	317	4,7	-	371	2,1	-	59
6.	-	-	343	-	-	183	-	-	-
7.	57,0	-	505	73,9	85	648	74,6	226	333
8.	42,0	23	187	38,7	23	205	9,4	72	233
9.	3,2	17	513	1,3	10	397	1,5	10	401
10.	2,9	6	746	0,2	12	672	-	38	551
11.	3,1	16	1045	1,6	13	557	1,4	36	509
12.	6,8	49	239	6,1	23	228	9,7	33	344
13.	-	-	317	-	-	570	-	-	-
14.	17,6	142	824	14,2	148	825	30,7	140	750
15.	32,3	10	173	5,4	12	108	16,3	15	102
16.	-	11	257	2,1	16	332	3,2	27	234
17.	2,4	7	804	0,9	9	830	-	50	630
18.	7,8	16	686	5,6	17	670	7,4	33	430
19.	5,8	54	429	5,3	74	378	3,6	105	280
20.	-	-	554	-	-	715	-	-	-
21.	2,7	15	555	1,4	26	685	6,5	16	485
22.	11,3	6	88	8,9	14	84	18,7	16	102
23.	20,4	44	436	21,7	21	260	15,9	28	152
24.	11,2	28	-	13,8	43	598	-	-	272
25.	-	-	490	-	-	495	-	-	-
26.	-	-	-	-	-	750	-	-	-
27.	-	-	328	-	-	850	-	-	-
28.	52,0	52	559	47,4	57	510	62,2	79	546
29.	34,8	88	504	31,9	93	466	25,3	38	434
30.	39,3	113	430	63,9	168	488	71,1	68	494
31.	78,5	61	713	54,2	103	292	-	-	278

TABLE VII

RaB in air; units $10^{-12} \mu\text{c}/\text{cm}^3$

Station Designations: Z Zugspitze 2964 m a.s.l.
 W Wank 1780 m a.s.l.
 G Garmisch-Partenkirchen 730 m a.s.l.

Measurement periods : a = 2150 - 0920, b = 0950 - 1220,
 c = 1250 - 1620, (Central European Time)

Month: January 1971 ("No measurement" is denoted by: -)

	a			b			c		
	Z	W	G	Z	W	G	Z	W	G
1.	-	-	182	-	-	120	-	-	-
2.	3,2	12	444	2,3	8	415	3,1	10	459
3.	-	-	1150	-	-	721	-	-	-
4.	17,6	6	514	31,5	54	368	34,6	133	330
5.	25,1	74	650	12,5	79	1006	17,0	120	844
6.	-	-	402	-	-	481	-	-	-
7.	8,6	51	347	6,7	19	319	7,4	45	331
8.	4,4	8	247	3,0	13	291	0,7	13	399
9.	1,9	5	256	0,1	4	350	5,2	22	345
10.	-	-	279	-	-	413	-	-	-
11.	30,8	30	290	32,5	30	292	25,5	30	287
12.	34,1	85	508	25,6	111	445	31,4	186	344
13.	21,4	124	661	14,3	145	880	17,1	160	509
14.	26,4	36	334	30,6	26	371	32,5	30	558
15.	84,1	198	516	63,4	195	536	109,2	128	150
16.	30,2	127	588	25,3	151	599	25,5	147	304
17.	-	-	560	-	-	420	-	-	-
18.	12,8	107	434	10,2	79	445	25,9	217	443
19.	57,5	81	441	52,6	73	497	48,3	95	327
20.	52,9	91	480	35,1	77	511	34,9	107	431
21.	31,2	53	388	27,2	25	498	36,4	22	369
22.	11,7	31	-	11,7	19	403	19,4	24	416
23.	30,9	91	400	23,8	119	342	21,1	124	213
24.	-	-	314	-	-	139	-	-	-
25.	17,3	-	321	32,9	62	348	36,3	37	198
26.	11,3	27	350	9,9	12	338	9,0	22	291
27.	32,7	59	330	27,5	77	297	35,1	71	160
28.	11,7	11	176	11,7	12	207	12,3	17	109
29.	10,3	119	467	7,4	122	278	12,6	145	184
30.	26,5	99	336	29,4	86	294	24,2	61	198
31.	-	-	144	-	-	78	-	-	-

TABLE VIII

RaB in air; units $10^{-12} \mu\text{c}/\text{cm}^3$

Station Designations: Z Zugspitze 2964 m a.s.l.
 W Wank 1780 m a.s.l.
 G Garmisch-Partenkirchen 730 m a.s.l.

Measurement periods : a = 2150 - 0920, b = 0950 - 1220,
 c = 1250 - 1620. (Central European Time)

Month: February 1971 ("No measurement" is denoted by: -)

	a			b			c		
	Z	W	G	Z	W	G	Z	W	G
1.	20,5	26	498	41,8	102	390	45,5	72	235
2.	4,1	27	57	2,0	26	60	5,7	86	74
3.	3,0	7	430	2,1	10	333	0,2	12	352
4.	11,9	3	65	16,1	4	64	11,1	15	85
5.	5,2	5	388	1,0	10	308	-	26	305
6.	8,6	2	224	-	6	445	0,8	26	392
7.	-	-	772	-	-	283	-	-	-
8.	2,8	5	414	0,3	22	212	1,0	34	134
9.	3,1	31	457	0,5	66	415	-	121	298
10.	2,0	91	635	-	117	644	-	192	478
11.	2,4	156	677	0,8	158	910	0,1	170	592
12.	5,1	76	487	-	31	364	24,3	26	135
13.	8,1	64	388	9,1	67	425	33,6	153	184
14.	-	-	615	-	-	277	-	-	-
15.	103,7	85	362	75,4	149	-	104,3	185	362
16.	54,7	56	281	35,2	34	184	53,5	16	33
17.	7,1	20	264	5,3	12	-	6,6	22	-
18.	7,3	17	-	4,1	8	280	9,9	21	275
19.	12,3	20	339	11,6	21	367	24,5	22	141
20.	12,1	7	47	12,9	20	79	12,1	16	66
21.	-	-	348	-	-	210	-	-	-
22.	8,6	23	388	3,0	15	96	6,0	7	38
23.	5,1	7	47	4,6	7	25	5,5	5	36
24.	5,7	15	236	1,5	29	117	1,1	78	93
25.	1,8	19	367	6,1	14	203	6,2	7	63
26.	-	-	46	8,6	7	34	12,0	6	40
27.	-	14	33	-	17	48	22,5	21	63
28.	-	-	338	-	-	109	-	-	-

TABLE IX

RnB in air; units $10^{-12} \mu\text{C}/\text{cm}^3$

Station Designations: Z Zugspitze 2964 m a.s.l.
 W Wank 1780 m a.s.l.
 G Garmisch-Partenkirchen 730 m a.s.l.

Measurement periods : a = 2150 - 0920, b = 0950 - 1220,
 c = 1250 - 1620. (Central European Time)

Month: March 1971 ("No measurement" is denoted by: -)

	a			b			c		
	Z	W	G	Z	W	G	Z	W	G
1.	58,9	72	85	57,4	73	117	65,2	70	128
2.	14,6	90	482	22,9	128	255	6,0	133	210
3.	41,1	71	470	-	113	349	98,4	112	215
4.	62,9	133	147	45,6	117	170	13,9	107	130
5.	49,2	75	553	51,2	109	209	69,6	-	129
6.	45,4	95	461	52,9	169	420	55,8	176	230
7.	-	-	613	-	-	198	-	-	-
8.	28,9	136	470	16,7	136	312	16,4	182	320
9.	11,3	130	756	19,9	192	544	14,7	101	266
10.	5,8	27	272	34,3	13	137	18,7	18	54
11.	-	18	101	9,2	15	87	-	13	64
12.	5,5	19	124	1,5	12	79	7,1	16	101
13.	4,4	16	266	1,7	35	207	5,0	65	189
14.	-	-	336	-	-	242	-	-	-
15.	23,1	180	407	17,1	169	334	62,9	157	196
16.	64,2	12	128	23,0	6	69	24,1	8	58
17.	10,2	18	232	0,8	38	125	22,4	24	138
18.	27,1	90	270	29,8	79	221	100,3	69	162
19.	24,6	30	292	22,2	37	176	41,2	47	158
20.	32,8	37	201	42,7	35	105	31,3	23	63
21.	-	-	93	-	-	41	-	-	-
22.	16,4	28	195	13,4	39	165	85,8	75	162
23.	24,9	8	60	26,3	21	61	-	5	80
24.	8,7	92	274	0,6	91	212	2,9	63	124
25.	5,7	42	381	16,4	80	201	29,4	51	78
26.	10,5	18	91	15,8	19	57	18,3	13	42
27.	16,8	3	57	17,2	12	38	13,1	13	35
28.	-	-	138	-	-	61	-	-	-
29.	6,2	-	259	1,4	58	108	37,1	36	71
30.	4,4	86	350	10,7	122	201	1,2	122	108
31.	87,8	-	317	94,7	116	352	147,6	106	266

TABLE X

RaB in air; units $10^{-12} \mu\text{C/cm}^3$

Station Designations: Z Zugspitze 2964 m a.s.l.
 W Wank 1780 m a.s.l.
 G Garmisch-Partenkirchen 730 m a.s.l.

Measurement periods : a = 2150 - 0920, b = 0950 - 1220,
 c = 1250 - 1620. (Central European Time)

Month: April 1971 ("No measurement" is denoted by: -)

	a			b			c		
	Z	W	G	Z	W	G	Z	W	G
1.	13,4	75	324	3,2	87	210	64,6	85	189
2.	39,0	132	388	44,2	134	259	73,5	152	242
3.	82,1	122	382	108,0	141	229	115,3	98	179
4.	-	-	-	-	-	61	-	-	-
5.	35,5	74	316	38,7	98	139	73,0	67	104
6.	26,2	67	330	29,5	68	103	51,3	47	92
7.	56,2	109	276	75,1	99	147	105,4	111	157
8.	37,3	109	265	15,4	42	149	62,2	95	129
9.	-	-	311	-	-	156	-	-	-
10.	61,4	97	507	35,7	153	263	70,6	174	255
11.	-	-	203	-	-	132	-	-	-
12.	-	-	228	-	-	90	-	-	-
13.	8,5	44	261	1,6	47	94	30,3	56	75
14.	5,8	56	304	1,0	63	130	23,5	22	53
15.	8,5	63	242	1,8	76	162	34,7	105	130
16.	53,9	76	332	54,6	81	101	47,6	76	102
17.	-	145	381	52,6	24	88	16,0	19	91
18.	-	-	178	-	-	102	-	-	-
19.	8,1	122	256	32,4	145	203	92,2	139	166
20.	77,1	129	372	109,3	140	145	75,0	85	101
21.	-	122	289	48,4	117	187	61,9	110	138
22.	67,2	142	286	70,1	118	152	77,0	123	113
23.	54,4	92	254	62,2	72	104	45,4	63	90
24.	-	-	176	22,8	-	60	11,5	-	80
25.	-	-	49	-	-	58	-	-	-
26.	32,1	-	102	16,7	15	51	18,5	19	42
27.	52,5	87	295	50,4	79	169	39,8	54	73
28.	32,8	18	52	35,0	13	56	28,1	16	89
29.	28,8	71	245	41,0	76	185	49,7	88	114
30.	31,9	124	402	24,6	109	250	18,5	120	138

TABLE XI

RaB in air; units $10^{-12} \mu\text{C}/\text{cm}^3$

Station Designations: Z Zugspitze 2964 m a.s.l.
 W Wank 1780 m a.s.l.
 G Garmisch-Partenkirchen 730 m a.s.l.

Measurement periods: a = 2150 - 0920, b = 0950 - 1220,
 c = 1250 - 1620. (Central European Time)

Month: May 1971 ("No measurement" is denoted by: -)

	a			b			c		
	Z	W	G	Z	W	G	Z	W	G
1.	-	-	282	-	-	193	-	-	-
2.	-	-	203	-	-	190	-	-	-
3.	51,5	-	220	49,6	118	207	53,9	100	155
4.	67,6	91	220	70,0	98	195	92,6	95	194
5.	75,8	90	329	64,8	68	207	45,7	76	198
6.	39,3	64	299	53,1	49	140	61,4	79	110
7.	36,6	52	306	31,8	68	126	62,3	65	100
8.	30,3	-	429	37,3	-	317	26,4	-	223
9.	-	-	336	-	-	100	-	-	-
10.	49,2	-	395	65,6	130	247	32,7	47	144
11.	39,4	77	204	41,1	81	179	47,4	49	95
12.	28,5	56	185	49,0	53	68	53,0	80	103
13.	131,3	134	340	90,4	125	128	42,7	66	91
14.	35,9	29	260	13,6	8	128	28,0	61	139
15.	33,9	43	398	39,7	87	154	52,3	73	81
16.	-	-	228	-	-	75	-	-	-
17.	19,5	111	172	34,9	45	113	52,3	86	117
18.	29,0	129	277	40,2	39	103	45,4	45	111
19.	23,0	50	150	44,8	52	79	67,5	54	97
20.	-	-	251	-	-	158	-	-	-
21.	58,0	128	203	62,8	94	94	44,1	27	73
22.	19,5	33	174	35,5	46	69	45,8	39	78
23.	-	-	213	-	-	159	-	-	-
24.	85,2	110	289	84,7	97	118	72,2	53	114
25.	71,5	70	168	57,6	74	116	45,9	69	112
26.	31,1	42	128	48,8	41	102	35,8	25	58
27.	25,7	29	190	40,5	15	153	55,2	39	144
28.	24,6	44	147	25,0	56	89	14,3	21	69
29.	19,7	24	111	33,2	23	78	30,7	35	80
30.	-	-	148	-	-	121	-	-	-
31.	-	-	308	-	-	145	-	-	-

TABLE III

RaB in air; units $10^{-12} \mu\text{C}/\text{cm}^3$

Station Designations: Z Zugspitze 2964 m a.s.l.
 W Wank 1780 m a.s.l.
 G Garmisch-Partenkirchen 730 m a.s.l.

Measurement periods : a = 2150 - 0920, b = 0950 - 1220,
 c = 1250 - 1620. (Central European Time)

Month: June 1971 ("No measurement" is denoted by: -)

	a			b			c		
	Z	W	G	Z	W	G	Z	W	G
1.	25,7	101	450	25,7	89	165	21,0	75	104
2.	31,6	54	223	64,7	130	198	93,3	65	123
3.	54,8	102	368	87,3	98	158	97,6	78	142
4.	72,8	107	274	90,4	146	203	94,8	115	159
5.	42,4	83	201	47,4	68	96	72,2	49	97
6.	-	-	346	-	-	147	-	-	-
7.	26,9	62	251	22,7	40	150	21,6	70	113
8.	16,5	17	300	20,1	21	228	20,9	39	102
9.	30,1	47	299	32,4	47	193	47,6	65	96
10.	-	-	209	-	-	37	-	-	-
11.	34,4	32	121	10,0	12	54	30,2	10	55
12.	8,3	13	73	17,3	17	53	21,5	21	37
13.	-	-	169	-	-	39	-	-	-
14.	28,7	69	171	44,0	46	110	75,8	68	128
15.	60,6	80	99	12,4	4	29	2,0	5	27
16.	7,3	17	73	6,6	21	34	19,6	16	30
17.	-	-	150	-	-	67	-	-	-
18.	12,7	71	329	11,6	12	74	9,1	8	39
19.	-	26	206	5,9	7	91	1,3	13	89
20.	-	-	51	-	-	73	-	-	-
21.	-	39	110	-	14	43	-	12	48
22.	28,9	59	218	55,6	101	116	26,1	17	53
23.	44,4	58	111	42,5	37	81	43,5	32	76
24.	71,5	96	379	87,3	112	261	84,0	53	252
25.	40,3	40	391	48,1	93	187	71,8	71	96
26.	63,5	87	297	84,7	81	135	107,1	81	101
27.	-	-	61	-	-	55	-	-	-
28.	52,9	53	134	53,6	61	116	54,2	44	74
29.	5,1	9	26	7,2	11	25	6,9	8	33
30.	6,3	14	101	8,0	16	77	4,3	11	57

5. The Computer Program

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PMT 1-TEEL TAB IN MM

KOPF 2
PRINT 2: TAIL TAB IN MM

MP-1

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TAB • SPE •

PAGE 010

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050B ZSAA00      LDV/ TAB,X
050C ZSFFA0      LDW/ X'FFA0'
050F 1708        JAX  SH6
0511 61F0        JNP  SH6

      4
      4
      4 MP 1 PRINT PORTSETING 1 NR. 0 6
      0

0513 051170      SRB  RTJ/  WRE
0516 451049      RTJ/  WRE
0519 451049      RTJ/  P13
051C 450506      RTJ/  MET
051E ZSAA00      LDV/ TAB,X
0520 ZSFFA0      LDW/ X'FFA0'
0528 1708        JAX  SH6
052F 61F0        JNP  SH6+6
0536 051170      RTJ/  WRE
053C 87000F      LDW/ 13
053F 450506      PAO  LDW/ ANA,X
053E 45250F      RTJ/  OUT
0538 45      DEC
0536 150F      RXH  PAO
053E 051170      RTJ/  WRE
0538 451049      JNP/  WRE
053E 451049      RTJ/  WRE
0541 05100B      SRB  RTJ/  M01
0544 050506      RTJ/  MET
0547 7F0A00      LDV/ TAB,X
0546 ZSFFA0      LDW/ X'FFA0'
054B 1706        JAX  SH6
054F 61F0        JNP  SH6
0551 451170      SRB  RTJ/  WRE
0556 870049      RTJ/  WRE
0557 451199      SRA  RTJ/  P13
055A 850506      RTJ/  MET
0560 ZSAA00      LDV/ TAB,X
0560 ZSFFA0      LDW/ X'FFA0'
0563 1708        JAX  SH6
0565 61F0        JNP  SH6
0567 051170      RTJ/  WRE
0568 870049      LDW/ 17
0560 ZSAA7F      PAO  LDV/ ANA,X
0567 05250F      RTJ/  OUT
0572 45      DEC
0574 150F      RXH  WRE
0576 051170      RTJ/  WRE
0578 051056      JNP/  WRE
057C 451049      SNU  RTJ/  M01
057F 451056      SRB  RTJ/  M01
0586 850506      RTJ/  MET
058E ZSAA00      LDV/ TAB,X
0590 ZSFFA0      LDW/ X'FFA0'
0592 1708        JAX  SH6
0590 61F0        JNP  SH6
059F 051170      SRB  RTJ/  WRE
05A2 051049      RTJ/  M03
05A5 051199      SNE  RTJ/  P13
05A8 450506      RTJ/  MET
0590 ZSAA00      LDV/ TAB,X

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05P7	620003	RTJ	N13
05P8	620003	STV	TAB+9.X
05P9	620003	RTJ	N14
0600	63P7	JRPO	N1P
0608	0000		
060A	6201E3	DC	003
0607	620009	RTJ	N13
060A	6201E3	STV	TAB+9.X
060D	63P7	RTJ	N14
		JRPO	N1P
0609	0000		
0611	6201E3	DC	007
061A	6201E3	RTJ	N13
0614	6201E3	STV	TAB+13.X
0617	6201E3	RTJ	N14
061A	63P7	JRPO	N1P
061C	0000		
061E	6201E3	DC	013
0621	620009	RTJ	N13
062A	6201E3	STV	TAB+37.X
062B	63P7	RTJ	N14
		JRPO	N1P
0629	0000		
062E	6201E3	DC	007
062E	620009	RTJ	N13
0631	6201E3	STV	TAB+40.X
063A	63P7	RTJ	N14
		JRPO	N1P
		UP-4	N1LP5PPO. 0
063E	0000		
063E	0E	N16	00
0639	620000		
063C	620003	LDW	SPX.X
0639	620000	RTJ	N1P
063E	62000A	LDW	SPX+X.X
064B	620108	RTJ	N1T
064E	62000E	LDW	SPX+0.X
064E	62000F	RTJ	N1P
064E	0A	RD3	
064C	62000E	LDW	SPX+18.X
064F	62010C	RTJ	N1P
064E	62000F	LDW	SPX+18.X
065E	620009	RTJ	N1P
065E	62000E	LDW	SPX+0.X
065E	007E	LDX	2TA
065D	6200F3	RD1	ENR
0660	0E	LDX	23P
066A	62010B	LDW	SPX+18.X
066A	63E2	JRPO	N1P
066B	0000		
066B	62010B	RTJ	N1P
066D	6201E3	RTJ	N1T
0670	63P6	JRPO	N1T

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0448 0000 N17 DC 00
0449 000000 22 14 1524

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PF3 : VELONE 21.28 VERDEN BENUTZT : NR. 17 *

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DATA 21.28 RAD-1. RAD-2.2 IN FL : NR. 18 *

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PF3 : VELONE 21.28 VERDEN BENUTZT : NR. 17 *

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DATA 21.28 RAD-1. RAD-2.2 IN FL : NR. 18 *

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PF3 : VELONE 21.28 VERDEN BENUTZT : NR. 17 *

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DATA 21.28 RAD-1. RAD-2.2 IN FL : NR. 18 *

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DATA 21.28 RAD-1. RAD-2.2 IN FL : NR. 18 *

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DATA 21.28 RAD-1. RAD-2.2 IN FL : NR. 18 *

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13. ABSTRACT

The investigations continued during the reporting period essentially as
described in preceding reports. The objectives were:
a. To shield the cable car telemetry systems against all effects of the
weather and to obtain by them homogeneous series of recordings;
b. to develop and apply the physical and mathematical bases for the
complete numeric evaluation of the recording runs, up to and including
computation of the incremental exchange coefficients;
c. to continually record at the three stations the Aitken nuclei con-
centrations, natural radioactivity (RaB, RaC), and polar conductivities,
and to utilize them in the evaluations; and
d. to derive initial deductions from the results.
The entire evaluation technique was completely changed. A new computer
(Intertechnique Multi 8) with 12 k words (1 word = 9 bit) affords us
the possibility to perform the entirety of all computation processes in
one single operation, and to print them out in tables. ()

Key words:

Aerosol; gradient; troposphere; boundary layer; exchange coefficient.

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1 NOV 66

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