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AN ANALYTICAL AND COMPUTATIONAL MODEL
OF THE GROWTH OF WATER AND ICE
PARTICLES IN NUCLEAR-EXPLOSION CLOUDS

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SUMMARY

This report describes development of the AEROSOL program for modeling growth of water and ice particles in nuclear-explosion clouds. AEROSOL has been developed for incorporation into an expanded version of the SHELL hydrocode which describes the evolution of the nuclear cloud environment. The report summarizes the physics, mathematical formulation, and particle-size classification scheme used in the program. It includes a listing of the AEROSOL program and recommendations for further program development.

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SECTION 1
INTRODUCTION, BACKGROUND, AND OBJECTIVE

1.1 NUCLEAR CLOUD PHENOMENA AND HAZARDS

The fireball produced by an atmospheric nuclear explosion rises as a buoyant nuclear cloud, which stabilizes near its maximum altitude, continues to expand and diffuse, and eventually dissipates. Associated with the cloud development are physical phenomena including ionization of the air, intense turbulence, and formation and dispersion of cloud debris, including radioactive fallout particles. The cloud contains water vapor, whose potential sources are entrained atmospheric moisture, and also, if the explosion occurs near the sea or land surface, sea water or ground water. As the cloud rises, entrains air, expands and cools, much of the water vapor condenses into water and ice. The water-substance particles continue to grow by condensation and agglomeration. Besides water, the particles may contain debris from the nuclear device, and dust (i.e., soil and rock materials) or sea salt, in amounts depending on the burst location.

The nuclear cloud and its particulate contents present potential hazards, including:

1. Mechanical damage to missiles and aircraft passing through the cloud.
2. Degradation or blackout of communications and reconnaissance systems.
3. Radiation from fission products and other radionuclides.

The extent of each of these hazards depends on the mass-concentration and particle-size distribution of the nuclear cloud particles.

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1.2 COMPUTATIONAL CLOUD MODELS

Analytical models of nuclear-cloud development have been developed to predict and estimate the magnitude of the hazards mentioned above. As with all theoretical models, it is desirable to compare the predictions with experimental data. Actual data available from nuclear-explosion tests are highly scattered, incomplete and uncertain. Useful types of data include photographic and radar observations of nuclear clouds, and samples of particle mass and radioactivity from airborne and ground collectors. Radar observations are especially important for checking cloud models that consider the cloud water content, because the intensity of the radar return from a nuclear or natural cloud depends on the number and size of cloud particles, and also on the dielectric constants of the particle materials. The radar cross sections, and thus the radar return, of particles of a given size differ according as the particles consist of water, ice, or ice with a coating of water, i.e., wet ice. These differences, in turn, vary with both particle size and radar wave length.

Existing cloud models differ considerably in their purpose and in their physical and computational level of complexity. The SHELL model and computer program is a two-dimensional hydrodynamic model. The region of space in which the cloud develops is covered by a fixed (Eulerian) axisymmetric grid. To compute the cloud history, the mass, momentum, and energy conservation equations are solved for each zone or cell in the grid, in a sequence of discrete time steps covering the period of the cloud history. Versions of SHELL are available which compute the concentrations of dust and water in the cloud. In particular, the SHELL-WV program computes the concentrations of dry air, water vapor, liquid water, and ice. The condensed water is assigned to a single arbitrary particle size. The possible change in particle size and the evolution of a particle-size distribution are not considered.

Associated with SHELL is the DUSTY program which computes the trajectories of particles of specified mass in the flow field computed by SHELL. Again, the possible changes in particle size during the trajectories are not considered.

1.3 OBJECTIVE AND GENERAL APPROACH

The objectives of the present study were to review the physical processes affecting growth of water and ice particles in nuclear clouds, to identify significant growth mechanisms, and to develop a computer program modeling particle growth, suitable for eventual incorporation into the SHELL program.

The general approach was to make both the review of particle physics and program development open-ended. The AEROSOL program described in this report is intended to be combined with the SHELL program and the DUSTY program, to form a SUPER-SHELL program. SHELL computes cloud hydro- and thermodynamics. DUSTY computes particle trajectories. AEROSOL will be interactive with these programs, and will communicate with them through COMMON blocks. At each time step, for each space zone, AEROSOL will use data generated by the other programs, and return new data to them. For instance, AEROSOL will use the gas temperature and pressure computed by SHELL, and the particle velocities and falling rates computed by DUSTY. As a future development, additional particle growth mechanisms may be incorporated in AEROSOL, for instance, turbulence effects. The particle allocation (bookkeeping) procedures in AEROSOL are designed to accept any such mechanisms.

1.4 PLAN OF THIS REPORT

Particle growth processes are reviewed and processes to be programmed are given mathematical formulation in Section 2. The algorithms for representing particle growth by a specified number of discrete particle-size classes are also developed in this section.

The AEROSOL computer program is described in Section 3. This section also contains instructions for the intended incorporation of AEROSOL into SHELL.

Section 4 describes test computations made with AEROSOL. These computations were made to ensure that the program was operating correctly.

Section 5 gives recommendations and procedures for further development of the AEROSOL program, including extension to growth of multi-material particles, and consideration of additional growth mechanisms.

The Appendix contains a listing of the AEROSOL program. The Bibliography lists several references on cloud physics which give the background to the particle-growth equations used in AEROSOL.

SECTION 2

PHYSICAL PROCESSES AND MATHEMATICAL FORMULATION OF PARTICLE GROWTH

2.1 PHYSICAL PROCESSES

The size of water-substance particles can change by condensation/evaporation of water vapor, and by agglomeration (also called coagulation) and breaking of the particles. Both these types of mechanism change the particle-size distribution. The distribution within a fixed volume of space, such as a cell in an Eulerian computational grid, also changes as existing particles enter and leave the volume. Thus the size distributions are affected by condensation, coagulation, and transport processes. The three types of process are interactive, and in addition are affected by phase changes, e.g., freezing and solution processes, which do not directly change particle size.

Attention is centered on air bursts, for which entrained atmospheric moisture contributes most of the eventual particulate mass. For land surface bursts, the cloud particles consist mainly of so-called dust, i.e., soil and rock material, and for sea surface bursts, salt and the ocean water as well as atmospheric moisture, may play important parts in particle growth. In the present treatment, particles are taken as consisting principally or entirely of water.

In a review of water-particle growth processes, the question arises whether any such processes are important in nuclear clouds which are not important, and possibly non-existent, in natural clouds. Nuclear clouds, even after they have cooled to the water-condensation point, differ from natural clouds in their extreme turbulence, high updraft velocities, and rapid temperature changes. The intensities of these phenomena in high-yield nuclear clouds far exceed the intensities in even the largest

natural convective clouds. Thus coagulation by turbulent acceleration may contribute significantly to particle growth in nuclear clouds, while in natural clouds this mechanism is generally much less important than coagulation due to gravitational acceleration, i.e., differential falling rates of large and small particles.

Would it be possible for droplets or crystals of diameter ten to a hundred microns to grow to precipitation size, i.e., diameter of order at least a thousand microns, by condensation alone, because of the intense updrafts and rapid cooling of the nuclear cloud? Such purely condensational growth does not occur in natural clouds. The answer is probably negative. The rapid cooling and resulting supersaturation of the air tend to activate additional condensation nuclei. There is not time for more condensing vapor to diffuse to the droplet surface. Thus growth is diffusion-limited. Also, the increased droplet or crystal population results in more frequent collisions; and thus more particle growth by agglomeration. However, the intense updrafts and turbulent mixing do prolong the residence times of particles in the nuclear cloud, and thus permit an increase in growth by condensation. This effect on growth is perhaps analogous to the growth of hailstones in updrafts in natural convective clouds.

2.2 CONDENSATION/EVAPORATION PROCESSES

In modeling water condensation, it is assumed that condensation nuclei already exist. Even in the natural atmosphere the concentration of nuclei is much greater than the concentration of droplets. Thus only a very small fraction of the nuclei are involved in droplet formation. This oversupply of nuclei will be at least equally true in the nuclear cloud even in the case of an air burst for which these nuclei will be provided primarily by the recondensed device materials. Surface bursts, of course, provide a much larger supply of nuclei.

For a single, spherical water or ice particle of radius r , the rate of change of particle mass, m , due to evaporation/condensation is given by

$$\frac{dm}{dt} = 4\pi r D(\rho_a - \rho_s)(VM)$$

where

D = diffusivity of water vapor in air

ρ_a = water vapor density in the ambient gas

ρ_s = water vapor density at particle surface

VM = ventilation factor for mass transfer

$$VM = 1 + CVM Re^{1/2} Sc^{1/3}$$

Re is the Reynolds number $2rv/\nu$, where v is relative particle velocity and ν is gas viscosity. Sc is the Schmidt number ν/D . CVM is a non-dimensional factor of order 0.3. The value of ρ_s varies with the temperature at the particle surface. Therefore, unless certain simplifying assumptions are made, as described below, it is necessary to solve the heat transfer and mass transfer equations simultaneously. The rate of change of particle heat content H due to evaporation/condensation is given by

$$\frac{dH}{dt} = 4\pi r K(T_s - T_a)(VH)$$

K = heat conductivity of air

T_a = temperature in the ambient gas

T_s = temperature at particle surface

VH = ventilation factor for heat transfer,

$$VH = 1 + CVH Re^{1/2} Pr^{1/3}$$

Pr is the Prandtl number $C_p \nu / K \rho_a$, where C_p is the specific heat of the gas. CVH is of order 0.3.

The complexity of solving the mass and heat transfer equations lies largely in the interaction of surface temperature and surface vapor density. If the particles are micron-size or smaller, vapor pressure at the surface is increased significantly by surface tension. Thus, for instance, in an atmosphere at 100% relative humidity larger drops may grow at the expense of smaller ones.

If droplets contain dissolved salts or other contaminants, then vapor pressure varies with the concentration and properties of these contaminants. Thus the phase transition from dry to wet particles can occur at less than 100% relative humidity due to lowering of vapor pressure over a solution. Vapor-pressure data is available for only a limited number of solutions of pure compounds, while nuclear air-burst and land-burst particles may contain a variable mixture of metal oxides and silicates and other ground materials.

It has been assumed in the development of the AEROSOL program that that particles consist of pure water or ice. Surface tension effects are neglected. Under these circumstances, the rates of mass transfer and heat transfer to a particle are directly related by

$$\frac{dH}{dt} = -L \frac{dm}{dt}$$

Here L is the latent heat of evaporation or sublimation, according as the drop is liquid or frozen. It is further assumed that water-vapor pressure at the drop surface is the saturation pressure of water vapor at the drop surface temperature. The quantities $\rho_a - \rho_s$ and $T_a - T_s$ can then be related to each other, and the mass growth rate can be defined entirely in

terms of known ambient quantities. The mathematical procedure involves manipulation of the ideal gas law for the water vapor densities, and use of the Clausius-Clapeyron equation to eliminate the drop-surface parameters. For details see the references cited in the Bibliography.

The resulting particle-growth equation, after making certain approximations, can be written

$$\frac{dm}{dt} = \frac{4\pi r D(VM)\rho_e(S-1)}{1 + DL^2 \rho_e M_w (VM) / RT_a^2 K(VH)}$$

In addition to the symbols previously defined, R is the ideal gas constant, M_w is the molecular weight of water, and S is the saturation ratio, or ratio of ambient water-vapor pressure ρ_a to saturation water-vapor pressure ρ_e at the cloud gas temperature, T_a . Thus condensation and evaporation occur according as S exceeds or is less than unity, respectively.

Particle density is taken as that of liquid water. Since ice is about 9% less dense, an ice particle has about 3% greater radius than a water drop of the same mass. Thus use of constant water density tends to underestimate the rate of evaporation/condensation of ice particles by about 3%. Since the program uses the correct values of vapor pressure and latent heat for water and ice, the net error should be considerably less than 3%.

2.3 AGGLOMERATION PROCESSES

Very small, sub-micron processes collide and adhere primarily because of Brownian motion. The largest water and ice particles, of diameters in the millimeter range, collide with smaller particles primarily due to gravity, that is, because of differential falling rates of larger and smaller particles. In the size range between microns and millimeters several processes may operate to cause agglomeration. Electrical attraction may cause agglomeration between particles bearing different charges.

Such charge effects may be enhanced by ionizing radiation from the fission products and other radionuclides carried by particles in the nuclear cloud.

Two types of turbulence mechanism cause collision of particles: turbulent shear and turbulent acceleration. Because different nearby elements of a turbulent fluid move with different velocities, particles moving with the fluid may collide due to turbulent shear. Because particles of different sizes have different inertia, two such neighboring particles may collide due to turbulent acceleration of a fluid element. Thus turbulent acceleration acts like a randomly directed gravity to produce collisions between particles of different sizes.

Agglomeration may produce nuclei, for further condensation, consisting of a mixture of different materials. The surface of an agglomerated nucleus will contain cracks and cavities which are favored condensation sites. If the mixture includes soluble and insoluble materials, then above the phase transition for the soluble material, the particle behaves as a solution droplet. The presence of the insoluble components enhances the effectiveness of the soluble salts by increasing the particle size and therefore the mass rate of condensation. Thus agglomeration and condensation interact in particle growth.

Not all particles on a collision course actually collide, and not all colliding particles actually adhere to each other. In the first case, aerodynamic forces may deflect one particle around another. On the other hand, the deflecting forces may be balanced or outweighed by electrical attraction between two particles. As to adhesion efficiencies, some particle-growth model calculations assume that colliding ice particles do not adhere, while collisions between an ice particle and a water particle do result in adhesion.

Large drops, of diameter several millimeters, are likely to break up into two or more smaller drops. Small ice particles may split up due to stresses caused by freezing and temperature changes. The resulting multiplication of ice particles may contribute to condensational growth, because two small ice particles will accrete more water vapor by condensation than one large particle of the same total mass.

In the formulation of agglomeration in the AEROSOL program, particle breakup is neglected. Two agglomerating, spherical particles are treated as forming a single, larger spherical particle. Collision-efficiency and adhesion effects are represented by a collision-efficiency factor which in test computations was taken as unity. For developmental purposes the collision mechanisms incorporated in the program were Brownian movement and gravity. The mathematical formulation of agglomeration is described in Section 2.5, "Particle-Size Classification Algorithms."

2.4 TRANSPORT PROCESSES

Particles can move into and out of a fixed volume of space such as a cell of the SHELL Eulerian hydrodynamic computation, due to two types of mechanisms. The first mechanism is that particles tend to move with the fluid. This particle motion occurs in steady fluid flow, and, to the extent that particle inertia can be neglected, also in accelerated flow. Secondly, particles move relative to the fluid, if the particles, the fluid, or both are subject to acceleration. Since large particles fall faster than small ones, they will enter and leave the fixed volume at different rates. Also, non-gravitational acceleration will move large and small particles into and out of the volume at different rates.

2.5 PARTICLE-SIZE CLASSIFICATION ALGORITHMS

For computational purposes, the particle-size distribution is represented by a set of discrete size classes. The smallest particles are assigned to the first class containing particles of mass $m(1)$. Particle

mass doubles from one class to the next so that $m(i) = 2m(i-1)$. The number of particles in class i , per unit volume, is $n(i)$. There are N classes, and N is chosen as large as required. For instance, $N = 30$ spans a mass range of 2^{30} , or a particle-diameter range of about one thousand, as from ten microns to one centimeter.

2.5.1 Condensation/Evaporation. Particles grow or shrink by condensation or evaporation as described by the final equations for $\frac{dm}{dt}$ given in Section 2.2. Thus, in a time interval Δt the mass of particles of class i changes by an amount $\Delta m(i)$, where

$$\Delta m(i) = \frac{dm}{dt} \Delta t$$

and the values of m , radius r , etc., appropriate to class i are used. However, the mass per particle in class i retains its fixed value $m(i)$, so that the particle of mass $m(i) + \Delta m(i)$ must now be allocated partly to the next larger or smaller class, i.e., to class $i+1$ or class $i-1$. The following allocation method was developed. In a time step Δt , the number of particles $n(i)$ in class i changes by an amount

$$\begin{aligned} \Delta n(i) = & \mp \frac{n(i)\Delta m(i)}{m(i+1) - m(i)} + \max \left(0, \frac{n(i-1)\Delta m(i-1)}{m(i) - m(i-1)} \right) \\ & + \max \left(0, - \frac{n(i+1)\Delta m(i+1)}{m(i+1) - m(i)} \right) \end{aligned}$$

In the first term on the right side of the equation, the minus and plus signs are used according as $\Delta m(i)$ is positive or negative. Thus the term is always negative, since any change in size of an i class particle must result in some of its mass being allocated to the next smaller or larger class. If $\Delta m(i-1)$ is positive, $n(i)$ thereby increases, as shown in the second term. If $\Delta m(i+1)$ is negative, $n(i)$ increases as shown in the third

term. If these two Δ -amounts have the opposite signs they do not affect $n(i)$, but instead affect classes $i-2$ and $i+2$, respectively.

Class number one requires special treatment because there is no class $i-1$. To take care of this case, $n(0)$ is defined as zero. Furthermore, particles are not allowed to shrink below mass $m(1)$, which is treated as a dry condensation nucleus. Therefore, if $\Delta m(1)$ is negative, only the third term above is non-zero:

$$\Delta m(1) = \max \left(0, \frac{n(2) \Delta m(2)}{m(2) - m(1)} \right)$$

In the foregoing algorithm, the total gain in particle mass in time Δt is exactly equal to the mass of water vapor condensing, summed over all size classes, and the total number of particles is unchanged.

2.5.2 Agglomeration. If two particles of class i agglomerate, evidently $n(i)$ is reduced by 2 and some other class $n(j)$ is increased. If an i and a j particle agglomerate, adjustments must be made in $n(i)$ and $n(j)$ and in the classes to which the new particle is assigned. The frequency of collision of i and j particles, per unit concentration of each class, is called $F(i,j)$. Then the number of such collisions per unit time, per unit volume, is $F(i,j)n(i)n(j)$. The following new algorithm was developed which conserves total particle mass and reduces particle number by one for each collision. In a time Δt , $n(i)$ changes by agglomeration by an amount $\Delta n(i)$:

$$\begin{aligned} \Delta n(i) = & - \sum_{j=1}^{i-1} F(i,j)n(i)n(j) 2^{j-i} - \sum_{j=i}^N F(i,j)n(i)n(j) \\ & + \sum_{j=1}^{i-1} F(i-1,j)2^{j-i} n(i-1)n(j) \end{aligned}$$

Thus, a j,i collision, for $j \leq i$, reduces $n(j)$ by 1, reduces $n(i)$ by 2^{j-i} , and increases $n(i+1)$ by 2^{j-i} .

2.5.3 Total Change in Particle-Size Distribution. The total change in size distribution in a time Δt , due to both condensation and agglomeration, is given by the sum of the $\Delta n(i)$ for the two mechanisms, as specified in Sections 2.5.1 and 2.5.2, respectively. In addition, the $n(i)$ may change by transport processes as described in Section 2.4. It is intended that transport effects be computed by the DUSTY portion of the eventual SUPER-SHELL program mentioned in Section 1.3.

SECTION 3

DESCRIPTION OF THE AEROSOL PROGRAM

3.1 GENERAL APPROACH AND PROGRAM STRUCTURE

The AEROSOL program is intended primarily for incorporation into the SHELL/DUSTY system. The program can be run independently for development and checkout purposes, if initial values of variables, values of cloud parameters and physical constants are specified. All these variables are in COMMON so that they may be updated either within AEROSOL or through the appropriate subroutine of SHELL or DUSTY which contains the same COMMON statements.

The program is designed to permit incorporation of additional particle-growth mechanisms besides those now computed. Thus, it includes flags (true-false variables) which may be set to compute or omit ventilation factors, and surface-tension effects on evaporation/condensation. The logic for these computations may subsequently be developed and inserted when feasible. Similarly, the collision frequencies contributing to agglomeration are computed in a separate subroutine. The logic for collision frequencies due to Brownian motion, electrostatic attraction, turbulent acceleration, gravity, for collision efficiency and for particle breakup may be inserted step-by-step.

AEROSOL consists of the following elements. The main program, TEST, includes most of the variable data and "flag" options. TEST calls subroutines SETUP and AEROSOL. SETUP includes values of physical constants. AEROSOL contains the evaporation/condensation equations and the tests for most of the flags and options. AEROSOL calls subroutine NCONCN or alternatively, CONCEN, which reclassifies particles at each time step, and updates the size distribution to allow for the effects of condensation and agglomeration. CONCEN calls subroutine COLLIS to compute the collision frequencies used in the agglomeration equations. COLLIS calls function FGT to compute

the collision frequencies due to different velocities or falling rates of particles of different size classes. Brief descriptions of these sub-routines are given below.

3.2 TEST

TEST is the main program for AEROSOL. It serves primarily as an executive program. It is used to (1) define input data, (2) define values of flags that are set as options, (3) define certain auxiliary variables as functions of input data, (4) initialize the computation by calling SETUP, (5) call AEROSOL and update problem time and terminate the computation when the specified problem time or number of time steps is reached, (6) print values of some of the variables.

When AEROSOL is incorporated in SHELL, purposes (1) and (6) will be performed by SHELL. The auxiliary variables (3) can be computed by SETUP. Problem-time control (5) must be compatible with the cloud computation and therefore will be replaced by SHELL time control. If the data supplied by SETUP is available in the main program, (4) is unnecessary. The values of flags (2) can either be given as part of the SHELL data, or be incorporated semi-permanently as data statements in AEROSOL. Thus, it will be feasible for SHELL to call subroutine AEROSOL directly and TEST can be discarded.

3.3 SETUP

SETUP is called once by TEST at the start of computation. It contains the values of physical constants as data statements. These are placed in SETUP so that the rest of the operating program, subroutine AEROSOL, etc., can be used with any consistent system of units. The data now given in SETUP are in cgs units. For SHELL operation all these data should be available from the main program through COMMON. SETUP also contains the definition of the class-to-class difference in particle mass, $DIFCM(I-1) = CM(I) - CM(I-1)$. For SHELL operation the definition of particle class radius $RAD(I)$ should be transferred to SETUP from TEST.

3.4 AEROSOL

AEROSOL is the principal operating subroutine. It is called at each time step by TEST. For SHELL operation, AEROSOL will be called directly by the appropriate subroutine of SHELL.

AEROSOL computes, for each size class, the rate of mass change due to evaporation/condensation. It also computes the corresponding rate of heat transfer. Before making these computations, it checks the status of the flags IMPURE and PSMALL. A TRUE value for IMPURE indicates that solute effects are to be allowed for. A TRUE value for PSMALL indicates that small particle effects, i.e., curvature and surface tension are to be allowed for. The TRUE logic has not yet been developed; hence a TRUE answer stops the program. If both these flags are set to FALSE, then heat transfer is taken as proportional to mass transfer (Section 2.2). AEROSOL also tests whether particles are liquid or frozen and resets the corresponding flags, FROZEN and TFFLAG.

AEROSOL calls the particle-classification subroutine NCONCN, or alternatively, CONCEN, which adjusts the size distribution for changes due to both evaporation/condensation and agglomeration.

3.5 NCONCN

Subroutine NCONCN is called by AEROSOL. NCONCN calls COLLIS to determine the collision frequencies, $CF(I,J)$, between each pair of particle size classes. The term collision frequency here is taken to allow for collision efficiency, sticking coefficients, etc., so that $CF(I,J)$ are the effective collision efficiencies. NCONCN reallocates mass between size classes to take into account evaporation/condensation of particles of each class as computed in AEROSOL, and destruction and creation of particles by agglomeration, at the frequencies computed in COLLIS. The allocation scheme

ensures that the number of particles is conserved in evaporation/condensation and that the mass of particles is conserved in agglomeration and that each collision reduces particle number by exactly one unit.

CONCEN is a similar, alternative subroutine. It uses an earlier allocation scheme which lacks the exact treatment of particle number in agglomeration. Furthermore, it may be valid only for condensing, not evaporating particles, until revisions are made to bring it into line with NCONCN. Use of NCONCN is therefore recommended.

3.6 COLLIS

Subroutine COLLIS is called by NCONCN or CONCEN. It computes the collision frequencies $CF(I,J)$ of all pairs of classes I,J. At present, the collision frequencies specified in COLLIS are those due to Brownian motion, FB, and due to gravity, FGT. COLLIS calls function FGT(I,J) to compute gravitational collision frequencies. Collision frequencies due to other mechanisms, such as turbulence effects and zone-to-zone variations in velocities and accelerations, as computed by SHELL, can be inserted in the $CF(I,J)$ computation. If collision efficiencies and particle breakup frequencies are considered, they can also be inserted in COLLIS.

3.7 FGT

Function FGT(I,J) is called by COLLIS to compute collision frequencies between particles of classes I and J, due to differences in either gravitational fall rates $VG(I)$ or velocities $V(I)$ between I and J particles. VG and V are in COMMON and it is intended that their values will be initialized and updated by sub-program DUSTY. For stand-alone use of AEROSOL, equations to compute falling rates should be inserted. At present, FGT contains the dummy value $FGT = 0$.

SECTION 4 TEST COMPUTATIONS

4.1 TYPES OF TEST MADE

Test computations were performed to verify that the programming correctly reproduced the algebraic form of the particle-growth equations, and that certain of the internal program controls operated correctly. The controls tested included those that specify whether particles were frozen or liquid, and those that determine the time step DELTIM used in integrating the particle-growth equations.

For most of the tests, the numerical values used for input data were as shown in the listing of TEST, except that for ambient air temperature, TAMB, the two values 273.5 and 273. were both used to test the freeze/melt logic, since the freezing point used was $TF = 273.16$ as shown in the listing of SETUP. Other physical constants were also as shown in SETUP.

Hand computations were made at the first time step, of the rate of condensation of class 1 particles, and of the Brownian collision frequency. The printouts were checked against the hand computation.

4.2 FLAGS TESTED

Test runs were made with both TRUE and FALSE values of NAGLOM, TFIX, CWF, FROZEN, and SKIP. NAGLOM omits the agglomeration computation, i.e., the collision frequencies are set to zero. If SKIP = FALSE the condensation/evaporation computation is in effect omitted (skipped). SKIP is not an input option, but can be set to FALSE by proper choice of the input CWF and of SPEMIN, the threshold value of saturation ratio S.

When TFIX is FALSE, the value of time step DELTIM is halved or doubled from its initial value until the fractional change in mass, of the smallest particle, of mass CM(1), in a time DELTIM, is between C/2 and C. The value C = .1 was used, i.e., DELTIM was set to permit between 5% and 10% mass-change in the smallest particles. With the input data shown in the listing, DELTIM is halved six times from its initial value of 0.1 seconds, before the test is passed. The data includes a supersaturation value of 10% which was chosen for test purposes and is not to be considered physically realistic.

If CWF = FALSE, the program assumes that no condensed water is present. This option is intended for use when only particles of the smallest size, i.e., class 1, are present. The convention is that class 1 particles are

dry nuclei which cannot evaporate. If they were permitted to shrink by evaporation, smaller-size classes would be needed in the allocation scheme. If they were permitted to disappear completely by evaporation, an additional algorithm would be needed to assign condensing vapor if water started to condense at a later time step. CWF is automatically reset to TRUE after any time step in which the mass of a particle of at least one class increases. CWF is automatically reset to FALSE after a time step such that all size classes except class 1 are empty of particles and also saturation ratio S was less than the specified threshold value.

FROZEN was given initial TRUE and FALSE values with TAMB set to values 273.5 and 273, i.e., just above and below the freezing point $T_F = 273.16$, thus giving FALSE and TRUE values of TFFLAG. If, for instance, FROZEN = FALSE and TFFLAG is found to be TRUE, then the vapor-pressure constants and latent heat value are reset from liquid-water to ice values, and the release of heat of fusion is computed. The other logical combinations are treated similarly.

4.3 MASS AND NUMBER CONSERVATION

It was verified that total particle mass remains constant in agglomeration, that total particle number remains constant in evaporation condensation (because each particle is assumed to contain a non-volatile condensation nucleus which remains after all water has evaporated), and that each collision (at 100% efficiency) reduces the total number of particles by one. Total mass of particles is defined as SUMA in TEST and is printed by TEST. Total number of particles is the sum of all CN(I). For checkout purposes, this sum was computed and printed by TEST.

4.4 COMPUTATIONAL DIFFUSION OF THE SIZE DISTRIBUTION

The particle-size classes are Eulerian, that is, each size class represents particles of a fixed size. If the size distribution initially consists of class 1 particles only, then if any condensation takes place in the first time step, the effect of condensation is represented by the creation of a few class 2 particles. An alternative Lagrangian classification scheme would be to increase slightly the size represented by class 1, after the first time step. The terms Eulerian and Lagrangian are used in analogy to fixed-grid and fixed-mass cells in hydrodynamic computations. In the present, Eulerian method, if at time step 2 class 2 particles grow by condensation, then some class 3 particles are created. After 10 time steps there will be particles of classes 1 through 10, even though a single particle of class 1 size might not even have grown to class 2 size. Thus, the particle size distribution tends to diffuse through the size classes. The effect appears in computing agglomeration as well as condensation. This effect is similar to mass diffusion in Eulerian hydrodynamics.

As an example of the diffusion effect, with the test data shown in the listing, a class 1 particle has mass $CM(1) = 5.236 \text{ E-13}$ while $CM(2) = 1.047 \text{ E-12}$. The mass condensation rate for class 1 is 2.930 E-11 , so that it would take about 1.8 E-2 seconds for the class 1 particle to reach class 2 size. This

time interval is nearly 12 of the time steps DELTIM used in the computation, with DELTIM controlled to permit a 10% mass growth per time step.

The rate of advance of the diffusion front is just one class per time step, no matter how small the time step. However, the possible actual error produced in the size distribution by the Eulerian method is very small. In the present system of sizes, class I +1 has twice the mass per particle that class I does. The resulting diffusive front is very steep. After 10 time steps the concentration, CN, at the front may decrease by a factor of over 1000 per size class. For instance, in one run for condensation alone without agglomeration, the values were of order CN(2) = 1.E2, CN(3) = 4.E0, CN(4) = 6.E-2, down to CN(9) = 8.E-17, CN(10) = 6.E-21, CN(11) = 1.E-25, CN(12) = 0.

The point is that the Eulerian method gives both the correct mean mass of a particle, and the correct number of particles, and that the smearing of the distribution amounts to a negligible fraction of the total mass.

Furthermore, the alternative Lagrangian method is not feasible if agglomeration is to be computed, because it results in an increase in the number of size classes at each time step. Thus, an Eulerian allocation scheme is a practical necessity for computing development of a particle size distribution.

SECTION 5
RECOMMENDATIONS FOR FURTHER DEVELOPMENT

AEROSOL should be extended to handle the simultaneous presence of liquid-water and ice particles in the same computational cell. The appropriate treatment may be to carry two different species of particles, or, in each cell, to allocate all water particles to either the liquid or the frozen state, with the temperature adjusted for latent-heat changes so as to conserve total energy.

Particles from land-surface bursts consist initially mainly of soil or rock materials, on which moisture gradually condenses. AEROSOL should be extended to such multi-material particles. Interactions between dust and water should be considered. If the dust fraction of total particle mass is significant, an averaging method must be developed so that all particles of one size class have the same density. A statistical algorithm may be required, to decide what fraction of the dry dust nuclei are activated at each time step in the condensation of water vapor.

Turbulence effects on agglomeration should be incorporated in the program. To generate turbulence variables for each computational cell, AEROSOL will have to look at the fluid velocities in the adjoining cells of the SHELL grid. Alternatively, the turbulence estimation routine may be incorporated directly in SHELL, and the turbulence parameters made available through COMMON. Collision-efficiency factors and drop-breakup frequencies should be included in the program on a simplified basis.

In addition to extensions of AEROSOL, test runs using a range of realistic burst and atmospheric conditions should be performed to detect possible program errors and to verify the realism and reliability of the SUPER-SHELL program.

BIBLIOGRAPHY

The following standard works on cloud physics were used as references on particle growth.

1. Fletcher, N. H., The Physics of Rainclouds, Cambridge University Press, 1962.
2. Mason, B. J., The Physics of Clouds, Oxford University Press, 1957.

The form given for condensational particle growth is based on that used in the following reference.

3. Bird, R. G., Stewart, W. E., and Lightfoot, E. N., Transport Phenomena, Wiley, 1960.

The following reports and articles were also used in formulating the particle-growth equations.

4. Danielsen, E. F., et al., "Hail Growth by Stochastic Collection in a Cumulus Model," J. Atmos. Sci., Vol. 29, p. 135, January 1972.
5. Huebsch, I. O., "Relative Motion and Coagulation of Particles in a Turbulent Gas," USNRDL-TR-67-49, May 1967.
6. Koenig, L. R., "Numerical Modeling of Ice Deposition," RAND Corporation RM-5715-NSF, August 1968.
7. Ogura, Y., and Takahashi, T., "Numerical Simulation of the Life of a Thunderstorm Cell," Monthly Weather Review, Vol. 99, No. 12, p. 895, December 1971.
8. Slinn, W. G. N., and Gibbs, A. G., "The Stochastic Growth of a Rain Droplet," J. Atmos. Sci., Vol. 28, p. 973, September 1971.

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APPENDIX

LISTING OF THE AEROSOL PROGRAM

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PROGRAM TEST(INPUT,OUTPUT,TAPE8=INPUT,TAPE1=OUTPUT)          TES00000
COMMON/FLAGS/FIRST,IMPURE,INTRCT,INSTCN,FROZEN,PSMALL,NUM1,NUM2, TES00100
1TFIX,VFSIG,CDHC,CWF,NAGLOM,SKIP,INSTFM                      TES00200
LOGICAL FIRST,IMPURE,INTRCT,INSTCN,FROZEN,PSMALL,NUM1,NUM2, TES00300
1TFIX,VFSIG,CDHC,CWF,NAGLOM,SKIP                            TES00400
COMMON/CONSTN/C,CK,CLF,CLL,CM(50),CMW,CSW,D,DELTIN,FL,PI,    TES00500
1RG,RHOW,ONETRD,CVM,CVH,DIFCM(50),NUMPAR,RHOP,CP,           TES00600
2SPEMIN,RK,TNOT,PNOT,DNOT,AF,BF,TF,AL,BL,CSI,ESO           TES00700
COMMON/VARIAB/DELTIM,RHOVA,TAMB,TIM,CNU,S,A,B,              TES00800
1RHUES,P                                                      TES00900
COMMON/ARRAYS/CN(50),CSD(50),DELHA(50),DELWV(50),HD(50),    TES01000
1RHUVS(50),RAD(50),TSUR(50),V(50),DELCM(50),              TES01100
2CF(50,50),DMDT(50),STATE(50),VG(50)                       TES01200
C*****TES01300
C                                                                TES01400
C  VARIABLES NORMALLY SET BY 'SHELL'                          TES01500
C                                                                TES01600
C*****TES01700
C                                                                TES01800
C  FLAGS SET AS OPTIONS --                                     TES01900
C                                                                TES02000
C*****TES02100
C                                                                TES02200
C-----TES02300
C  CHANGE IN DROP HEAT CONTENT CONSIDERED                     TES02400
  DATA CDHC / .FALSE. /                                       TES02500
C-----TES02600
C  NO AGGLOMERATION CONSIDERED                                TES02700
  DATA NAGLOM / .FALSE. /                                       TES02800
C-----TES02900
C  PARTICLES CONTAIN IMPURITIES                                TES03000
  DATA IMPURE / .FALSE. /                                       TES03100
C-----TES03200
C  CONDENSATION/EVAPORATION IS INSTANTANEOUS                 TES03300
  DATA INSTCN / .FALSE. /                                       TES03400
C-----TES03500
C  FREEZING/MELTING IS INSTANTANEOUS                          TES03600
  DATA INSTFM / .TRUE. /                                         TES03700
C-----TES03800
C  CALCULATION IS INTERACTIVE WITH ATMOSPHERE                 TES03900
  DATA INTRCT / .FALSE. /                                       TES04000
C-----TES04100
C  PARTICLE IS SMALL ENOUGH FOR SURFACE TENSION OR CURVATURE TO AFFECT TES04200
C  WATER VAPOR DENSITY OVER SURFACE                           TES04300
  DATA PSMALL / .FALSE. /                                       TES04400
C-----TES04500
C  TIME STEP IS FIXED                                         TES04600
  DATA TFIX / .FALSE. /                                           TES04700
C-----TES04800
C  VENTILATION FACTORS ARE SIGNIFICANT                        TES04900
  DATA VFSIG / .FALSE. /                                           TES05000
C-----TES05100
C                                                                TES05200
C*****TES05300
C                                                                TES05400
C  FLAGS THAT MAY BE RESET BY PROGRAM                          TES05500

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C
C*****TES05600
C*****TES05700
C*****TES05800
C-----TES05900
C CONDENSED WATER PRESENT TES06000
  DATA CWF / .TRUE. / TES06100
C-----TES06200
C PARTICLES ARE FROZEN TES06300
  DATA FROZEN / .TRUE. / TES06400
C-----TES06500
C-----TES06600
C*****TES06700
C-----TES06800
C VARIABLES THAT MAY BE SET, INITIALIZED OR UPDATED BY SHELL TES06900
C BUT MUST BE DEFINED HERE FOR STAND ALONE RUN TES07000
C-----TES07100
C*****TES07200
C-----TES07300
C-----TES07400
C MAXIMUM PERMISSIBLE CHANGE PER TIME STEP TES07500
  DATA C / .1 / TES07600
C-----TES07700
C PARTICLE CONCENTRATIONS TES07800
  DATA CN / 1.0E6,49*0.0 / TES07900
C-----TES08000
C KINEMATIC VISCOSITY OF AIR AT CONSTANT PRESSURE TES08100
C CNU VALUE MATCHES P AND TAMB TES08200
  DATA CNL / .14 / TES08300
C-----TES08400
C TIME STEP VALUE TES08500
  DATA DELTIM / .1 / TES08600
C SECONDS TES08700
C-----TES08800
C NUMBER OF TIME STEPS TES08900
  DATA NSTEP / 20 / TES09000
C-----TES09100
C NUMBER OF PARTICLE CLASSES TES09200
  DATA NUMPAR / 20 / TES09300
C-----TES09400
C PRESSURE TES09500
  DATA P / 1.013E6 / TES09600
C DYNES / CENTIMETER 2 TES09700
C-----TES09800
C RATIO BETWEEN MASS OF PARTICLES IN CONTIGUOUS CLASSES TES09900
C AGGLOMERATION LOGIC VALID ONLY FOR RATIO = 2. TES10000
  DATA RATIO / 2. / TES10100
C-----TES10200
C PARTICLE DENSITY TES10300
  DATA RHOP / 1.0 / TES10400
C GRAMS / CENTIMETER 3 TES10500
C-----TES10600
C WATER VAPOR DENSITY TES10700
C 10 PERCENT SUPERSATURATION AT TAMB = 273. TES10800
  DATA RHOVA / 5.34E-6 / TES10900
C GRAMS / CENTIMETER 3 TES11000
C-----TES11100

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C WATER VAPOR DENSITY OVER DROP SURFACE TES11200
  DATA RHOVS / 50*0.0 / TES11300
C GRAMS / CENTIMETER 3 TES11400
C-----TES11500
C MINIMUM VALUE FOR TEST OF SATURATION RATIO TES11600
  DATA SPEMIN / 1.0 / TES11700
C-----TES11800
C AMBIENT AIR TEMPERATURE TES11900
  DATA TAMB / 273.5 / TES12000
C DEGREES KELVIN TES12100
C-----TES12200
C INITIAL TIME TES12300
  DATA TIM / 0.0 / TES12400
C SECONDS TES12500
C-----TES12600
C MAXIMUM TIME CONSIDERED TES12700
  DATA TIMMAX / 1.0 / TES12800
C SECONDS TES12900
C-----TES13000
C PARTICLE VELOCITIES TES13100
  DATA V / 50*0.0 / TES13200
C CENTIMETERS / SECOND TES13300
C-----TES13400
C ONETRD=1./3. TES13500
  PI=4.0*ATAN(1.0) TES13600
C-----TES13800
C DATA -- SMALLEST PARTICLE CLASS TES13900
C VALUE MEANS RAD(1) = .5E-4 TES14000
C-----TES14100
  CM(1)=4.*PI*(.125E-12)/3. TES14200
  PRAD=.75/(PI*RHOP) TES14300
  RAD(1)=(CM(1)*PRAD)**ONETRD TES14400
  SUMA=CN(1)*CM(1) TES14500
  WRITE(1,1000) TES14600
1000 FORMAT(1H1) TES14700
  WRITE(1,1100) SUMA TES14800
1100 FORMAT(/,20X,26HTOTAL MASS OF PARTICLES = ,E11.4,/) TES14900
  I=1 TES15000
  WRITE(1,1200) I,CM(I),I,RAD(I) TES15100
1200 FORMAT(20X,3HCM(,I2,4H) = ,E10.4,20X,4HRAD(,I2,4H) = ,E10.4,/) TES15200
  DO 50 I=2,NUMPAR TES15300
  CM(I)=CM(I-1)*RATIO TES15400
  RAD(I)=(CM(I)*PRAD)**ONETRD TES15500
  WRITE(1,1200) I,CM(I),I,RAD(I) TES15600
  50 CONTINUE TES15700
  CALL SETUP TES15800
  WRITE(1,1000) TES15900
  DO 10 I=1,NSTEP TES16000
  IF(TIM.GT.TIMMAX) STOP TES16100
  SUMA=C.0 TES16200
  SUMH=C.0 TES16300
  CALL AEROSOL TES16400
  TIM=TIM+DELTIM TES16500
  WRITE(1,1300) I,TIM,DELTIM TES16700
1300 FORMAT(///,20X,12HITERATION = ,I3,5X,7HTIME = ,E10.4,5X TES16800

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112HTIME STEP = ,E10.4)
WRITE(1,1400)
1400 FORMAT(///,20X,5HCLASS,5X,13HCONCENTRATION,5X,
114HCHANGE IN MASS,10X,5HDM/DT,/,20X,
25H-----,5X,13(1H-),5X,14(1H-),10X,5H-----)
DO 5 J=1,NUMPAK
WRITE(1,1500) J,CN(J),DELGM(J),DMDT(J)
1500 FORMAT(20X,I3,8X,E12.4,6X,E12.4,7X,E12.4)
IF(CN(J).LT.0.0) CN(J)=0.0
C=====
C TOTAL PARTICLE MASS CONCENTRATION
C=====
SUMA=SUMA+CN(J)*CM(J)
C
C=====
C HEAT RELEASED IN TIME STEP
C=====
SUMH=SUMH+DLLHA(J)
C
5 CONTINUE
WRITE(1,1100) SUMA
WRITE(1,1600) SUMH
1600 FORMAT(/,20X,49HTOTAL HEAT RELEASED TO ATMOSPHERE IN TIME STEP = ,
1 E11.4)
WRITE(1,1700)
1700 FORMAT(/,5X,110(1H*))
10 CONTINUE
STOP
END
TES16900
TES17000
TES17100
TES17200
TES17300
TES17400
TES17500
TES17600
TES17700
TES17800
TES17900
TES18000
TES18100
TES18200
TES18300
TES18400
TES18500
TES18600
TES18700
TES18800
TES18900
TES19000
TES19100
TES19200
TES19300
TES19400
TES19500
TES19600
TES19700

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	SUBROUTINE AERUSOL	AER0000
C	COMMON/FLAGS/FIRST,IMPURE,INTRCT,INSTCN,FROZEN,PSMALL,NUM1,NUM2,	AER00100
	1TFIX,VFSIG,CDHC,CWF,NAGLOM,SKIP,INSTFM	AER00200
	LOGICAL FIRST,IMPURE,INTRCT,INSTCN,FROZEN,PSMALL,NUM1,NUM2,	AER00300
	1TFIX,VFSIG,CDHC,CWF,NAGLOM,SKIP	AER00400
	COMMON/CONSTN/C,CK,CLF,CLL,CM(50),CMW,CSW,D,DELTIN,HL,PI,	AER00500
	1RG,RHOW,DNFTRD,CVM,CVH,DIFCM(50),NUMPAR,RHOP,CP,	AER00600
	2SPEMIN,BK,TNOT,PNOT,DNOT,AF,BF,TF,AL,BL,CSI,ESO	AER00700
	COMMON/VARIABLES/DELTIM,RHOVA,TAMB,TIM,CNU,S,A,B,	AER00800
	1RHUES,P	AER00900
	COMMON/ARRAYS/CN(50),CSD(50),DELHA(50),DELWV(50),HD(50),	AER01000
	1RHUVS(50),RAD(50),TSUR(50),V(50),DELCM(50),	AER01100
	2CF(50,50),DMDT(50),STATE(50),VG(50)	AER01200
C	*****	AER01300
C		AER01400
C	DEFINITION OF VARIABLES USED --	AER01500
C		AER01600
C	I. FLAGS (TRUE OR FALSE VALUES ONLY) -- INDICATE THE FOLLOWING	AER01700
C	CDHC ---- CHANGE IN DROP HEAT CONTENT CONSIDERED	AER01800
C	CWF ----- CONDENSED WATER PRESENT	AER01900
C	FIRST --- USED ON CLASS ONE ONLY	AER02000
C	FROZEN -- A FROZEN PARTICLE (FREEZING CONSIDERED)	AER02100
C	IMPURE -- PRESENCE OF IMPURITIES IN PARTICLE	AER02200
C	INTRCT -- CALCULATION INTERACTIVE WITH ATMOSPHERE	AER02300
C	INSTCN -- INSTANT CONDENSATION IS ASSUMED	AER02400
C	INSTFM -- INSTANT FREEZING/MELTING IS ASSUMED	AER02500
C	NAGLOM -- OMIT AGGLOMERATION	AER02600
C	NCDHC --- NEGATIVE OF CDHC	AER02700
C	NINTRC -- NEGATIVE OF INTRCT	AER02800
C	NSPMFG -- NEGATIVE OF SPEMFG	AER02900
C	NUM1 ---- NUMPAR = 1	AER03000
C	NUM2 ---- NUMPAR = 2	AER03100
C	PSMALL -- PARTICLE SMALL ENOUGH FOR SURFACE TENSION OR CURVATURE	AER03200
C	TO AFFECT WATER VAPOR DENSITY OVER SURFACE	AER03300
C	SKIP ---- DO NOT SKIP CONDENSATION/EVAPORATION LOGIC	AER03400
C	SPEMFG -- SATURATION RATIO EXCEEDS SPEMIN	AER03500
C	TFIX ---- FIXED TIME STEP	AER03600
C	TIMF ---- CHANGE IN TIME IS AN INPUT	AER03700
C	VFSIG --- SIGNIFICANCE OF VENTILATION FACTORS	AER03800
C		AER03900
C		AER04000
C	II. REAL VALUED VARIABLES	AER04100
C	C ----- MAXIMUM PERMISSABLE CHANGE PER TIME STEP	AER04200
C	CK ----- CONDUCTIVITY OF HEAT IN AIR	AER04300
C	CLF ----- LATENT HEAT OF FREEZING	AER04400
C	CLL ----- LATENT HEAT OF VAPORIZATION OF WATER AT TEMPERATURE	AER04500
C	ZERO DEGREES CENTIGRADE	AER04600
C	CLM ----- LATENT HEAT OF MELTING (ICE-LIQUID TRANSITION)	AER04700
C	CM ----- MASS OF PARTICLE OF CLASS(I)	AER04800
C	CMW ----- MOLECULAR WEIGHT OF WATER	AER04900
C	CN ----- PARTICLE CONCENTRATION	AER05000
C	CNU ----- KINEMATIC VISCOSITY OF AIR AT RHOVA AND TAMB	AER05100
C	CP ----- SPECIFIC HEAT OF AIR AT CONSTANT PRESSURE	AER05200
C	CSI ----- SPECIFIC HEAT OF ICE	AER05300
C	CSW ----- SPECIFIC HEAT OF WATER	AER05400
C	CVH ----- HEAT VENTILATION FACTOR COEFFICIENT	AER05500

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C      CVM ----- MASS VENTILATION FACTOR COEFFICIENT          AER05600
C      D ----- DIFFUSIVITY OF WATER VAPOR IN AIR             AER05700
C      DN ----- TOTAL D(CN)/DT                                AER05800
C      DNDTAG -- D(CN)/DT FOR AGGLOMERATION                     AER05900
C      DNPTLC -- D(CN)/DT FOR EVAPORATION/CONDENSATION         AER06000
C      DELCM --- CHANGE IN MASS OF PARTICLE                     AER06100
C      DELH ---- HEAT RELEASED BY CONDENSATION                 AER06200
C      DELHA --- HEAT RELEASED TO ATMOSPHERE                    AER06300
C      DLLHD --- CHANGE IN DROP HEAT CONTENT                   AER06400
C      DELTIM -- TIME STEP                                       AER06500
C      DELTII -- STANDARD INITIAL VALUE FOR DELTIM             AER06600
C      DELWV --- CHANGE IN WATER VAPOR MASS                     AER06700
C      DMDT ---- DM/DT                                          AER06800
C      ES ----- SATURATION VAPOR PRESSURE OF WATER AT TEMPERATURE TAMB AER06900
C      HD ----- HEAT CONTENT OF DROP                           AER07000
C      NUNPAR -- NUMBER OF PARTICLE CLASSES                      AER07100
C      P ----- AMBIENT AIR PRESSURE                            AER07200
C      PI ----- 3.14159 . . .                                  AER07300
C      PR ----- PRANDTL NUMBER OF THE GAS (AIR)               AER07400
C      RAD ----- PARTICLE RADIUS                               AER07500
C      RE ----- REYNOLDS NUMBER OF PARTICLE                   AER07600
C      RG ----- GAS CONSTANT                                   AER07700
C      RHGDS --- SATURATION WATER VAPOR DENSITY AT TEMPERATURE TAMB AER07800
C      RHGP ---- PARTICLE DENSITY                               AER07900
C      RHOVA --- WATER VAPOR DENSITY                            AER08000
C      RHOVS --- WATER VAPOR DENSITY OVER DROP SURFACE         AER08100
C      RHOw ---- DENSITY OF WATER                              AER08200
C      S ----- SATURATION RATIO                               AER08300
C      SC ----- SCHMIDT NUMBER OF THE GAS (AIR)               AER08400
C      SPEMIN -- SPECIFIED MINIMUM VALUE OF S                   AER08500
C      TAMB ---- AMBIENT AIR TEMPERATURE                        AER08600
C      TIM ---- TIME                                           AER08700
C      TSUR ---- DROP SURFACE TEMPERATURE                      AER08800
C      V ----- VELOCITY OF PARTICLE RELATIVE TO AIR           AER08900
C      VISC ---- DYNAMIC VISCOSITY OF AIR                       AER09000
C      VH ----- VENTILATION FACTOR FOR HEAT TRANSFER         AER09100
C      VM ----- VENTILATION FACTOR FOR MASS TRANSFER          AER09200
C                                                                 AER09300
C*****AER09400
      LOGICAL NINTRC,NCDHC,TIME,SPEMFG                          AER09500
      TFFLAG=TAMB.LT.TF                                         AER09600
      IF(FROZEN.OR.TFFLAG) GO TO 300                             AER09700
C=====AER09800
C  SET CONSTANTS TO LIQUID VALUES                             AER09900
C=====AER10000
      WRITE(1,1000)                                             AER10100
1000 FORMAT(////,25X,10(1H*),8H LIQUID ,10(1H*))             AER10200
      1 A=AL                                                    AER10300
      B=BL                                                       AER10400
      CLN=CLL                                                    AER10500
      CSU=CSW                                                    AER10600
C                                                                 AER10700
C=====AER10800
C  THTENS EQUATIONS FOR SATURATION VAPOR PRESSURE IN CGS     AER10900
C=====AER11000
      2 ES=ES1*EXP(A*(TAMB-273.16)/(TAMB-B))                   AER11100

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```

C
WRITE(1,1100) ES
1100 FORMAT(/,25X,37HSATURATION VAPOR PRESSURE OF WATER = ,E10.4)
C
RHOES=ES*CMW/(RG*TAMB)
RHOESF=RHOES*.99
S=RHOVA/RHOES
WRITE(1,1200) S
1200 FORMAT(/,25X,19HSATURATION RATIO = ,E10.4)
SPEMFG=S.GF.SPEMIN
SKIP=SPEMFG.OR.CWF
C=====
C IF NO CONDENSED WATER AND S BELOW CONDENSATION THRESHOLD
C THEN BYPASS EVAPORATION/CONDENSATION COMPUTATION
C=====
IF(.NOT.SKIP) GO TO 950
FIRST=.TRUE.
C=====
C 'INSTCN' LOGIC NOT YET DEFINED
C=====
IF(INSTCN) GO TO 200
C=====
C '.NOT.INSTFM' LOGIC NOT YET DEFINED
C=====
IF(.NOT.INSTFM) GO TO 700
NCDHC=.NOT.CDHC
3 DRHOVA=0.0
DO 900 I=1,NUMPAR
DRHO=DRHOVA-RHOVS(I)
C=====
C 'IMPURE' AND 'PSMALL' LOGIC NOT YET DEFINED
C=====
IF(IMPURE) GO TO 400
IF(PSMALL) GO TO 500
IF(VFSIG) GO TO 70
VM=1.
VH=1.
C=====
C DRHO = RHOVA - RHOVS
C=====
5 DRHO=RHOES*(S-1.)/(1.+(D*CLN*CLN*RHOES*CMW*VM/
1(RG*TAMB*TAMB*CK*VH)))
FORPIR=4.*PI*RAD(I)
C=====
C GROWTH RATE OF PARTICLE
C=====
DMDTI=FORPIR*D*DRHO*VM
C
C=====
C HEAT RELEASE RATE OF PARTICLE
C=====
DHDT=DMDTI*CLN
C
C
TSUR(I)=DHDT*(1./(FORPIR*CK))*(1./(1.+VH))+TAMB
10 DELCMI=DMDTI*DELTIM
IF(TFIX) GO TO 15

```

```

AER11200
AER11300
AER11400
AER11500
AER11600
AER11700
AER11800
AER11900
AER12000
AER12100
AER12200
AER12300
AER12400
AER12500
AER12600
AER12700
AER12800
AER12900
AER13000
AER13100
AER13200
AER13300
AER13400
AER13500
AER13600
AER13700
AER13800
AER13900
AER14000
AER14100
AER14200
AER14300
AER14400
AER14500
AER14600
AER14700
AER14800
AER14900
AER15000
AER15100
AER15200
AER15300
AER15400
AER15500
AER15600
AER15700
AER15800
AER15900
AER16000
AER16100
AER16200
AER16300
AER16400
AER16500
AER16600
AER16700

```

```

      IF(FIRST) GO TO 50
      15 CMI=CM(I)+DELDMI
C=====
C TO TREAT CLASS 1 PARTICLES AS DRY NUCLEI, SET
C NEGATIVE DELDMI TO ZERO FOR I = 1
C=====
      CNI=CN(I)
      20 DELH=CLN*DELDMI
      DELHDI=0
      IF(NCLHC) GO TO 30
C=====
C IF FROZEN USE ICF VALUE FOR CSW
C=====
      DELHDI=CSW*DELDMI*TSUR(I)
C
C=====
C TOTAL HEAT RELEASED BY CLASS
C=====
      30 DELHA(I)=(DELH-DELHDI)*CNI
      DRHOVA=DRHOVA-CNI*DELDMI
      GO TO 80
C=====
C TEST TIME STEP
C=====
      50 DLCMCM=DELDMI/CM(I)
      IF(ABS(DLCMCM).LE.C) GO TO 52
      DELTIM=DELTIM/2.
      GO TO 10
      52 IF(ABS(DLCMCM).LE.C/2.0) GO TO 60
C=====
C TIME STEP ALL RIGHT
C=====
      FIRST=.FALSE.
      GO TO 15
      60 DELTIM=DELTIM*2.
      GO TO 10
C=====
C COMPUTE VENTILATION FACTORS
C=====
      70 SQRTRE=SQRT(2.*RAD(I)*V(I)/CNU)
      SC=CNU/D
      PR=CP*CNU/(CK*RHOVA)
      VM=1.+CVM*SQRTRE*SC**CNETRD
      VH=1.+CVH*SQRTRE*PR**CNETRD
      GO TO 5
C
      80 CONTINUE
      DELCM(I)=DELDMI
      DMDT(I)=DMDTI
      900 CONTINUE
      IF(INTRCT) RHOVA=RHOVA+DRHOVA
C=====
C ALSO UPDATE P AND TAMB IF INTERACTIVE
C USING TOTAL DELHA AND CONSTANT-VOLUME SPECIFIC HEAT
C=====
      IF(RHOVA.LT.0.0) RHOVA=0.0

```

```

AER16800
AER16900
AER17000
AER17100
AER17200
AER17300
AER17400
AER17500
AER17600
AER17700
AER17800
AER17900
AER18000
AER18100
AER18200
AER18300
AER18400
AER18500
AER18600
AER18700
AER18800
AER18900
AER19000
AER19100
AER19200
AER19300
AER19400
AER19500
AER19600
AER19700
AER19800
AER19900
AER20000
AER20100
AER20200
AER20300
AER20400
AER20500
AER20600
AER20700
AER20800
AER20900
AER21000
AER21100
AER21200
AER21300
AER21400
AER21500
AER21600
AER21700
AER21800
AER21900
AER22000
AER22100
AER22200
AER22300

```

```

C=====AER22400
C CALL CONCEN OR NCONCN                                AER22500
C=====AER22600
  950 CALL NCONCN
C                                AER22700
  RETURN                                                AER22800
  200 STOP                                              AER22900
C=====AER23000
C FREEZE/MELT LOGIC                                    AFR23100
C=====AER23200
  300 IF(FROZEN.AND..NOT.TFFLAG) GO TO 370             AER23300
        IF(.NOT.FROZEN.AND.TFFLAG) GO TO 340           AER23400
C                                AER23500
        WRITE(1,1010)                                   AER23600
  1010 FORMAT(////,25X,10(1H*),8H FROZEN ,10(1H*))    AER23700
C=====AER23800
C SET CONSTANTS TO ICE VALUES                         AER23900
C=====AER24000
  310 CLN=CLF                                          AER24100
        A=AF                                           AER24200
        B=BF                                           AER24300
        CS0=CSI                                        AER24400
        GO TO 2                                         AER24500
C                                AER24600
  340 FROZEN=.TRUE.                                    AER24700
        CLN=CLL                                        AER24800
        WRITL(1,1020)                                   AER24900
  1020 FORMAT(////,25X,10(1H*),18H SWITCH TO FROZEN ,10(1H*))
        IF(.NOT.CWF) GO TO 310                         AER25000
        TLHF=C.O                                       AER25100
        DO 350 I=1,NUMPAR                              AER25200
  350 TLHF=TLHF+CN(I)*CM(I)                            AER25300
        TLHF=TLHF*CLN                                  AER25400
        WRITE(1,1030) TLHF                             AER25500
  1030 FORMAT(/,10X,16HHEAT RELEASED = ,E11.4)        AER25600
C=====AER25700
C UPDATE P AND TAMB IF INTERACTIVL                   AER25800
C=====AER25900
        GO TO 310                                       AER25900
C                                AER26000
C=====AER26100
C RESET TO LIQUID VALUES                             AER26200
C=====AER26300
  370 FROZEN=.FALSE.                                   AER26400
        CLN=CLF                                        AER26500
        WRITE(1,1040)                                   AER26600
  1040 FORMAT(////,25X,10(1H*),18H SWITCH TO LIQUID ,10(1H*))
        IF(.NOT.CWF) GO TO 1                           AER26700
        TLHF=C.C                                       AER26800
        DO 380 I=1,NUMPAR                              AER26900
  380 TLHF=TLHF-CN(I)*CM(I)                            AER27000
        TLHF=TLHF*CLN                                  AER27100
C=====AER27200
C                                AER27300
C                                AER27400
C=====AER27500
C UPDATE P AND TAMB IF INTERACTIVL                   AER27600
C=====AER27700
        WRITE(1,1030) TLHF                             AER27700
        GO TO 1                                         AER27800
C                                AER27900

```

C
400 STOP
500 STOP
700 STOP
END

AER 2800J
AER 28100
AER 28200
AER 28300
AFR 28400

```

SUBROUTINE SETUP
C
COMMON/FLAGS/FIRST,IMPURE,INTRCT,INSTCN,FROZEN,PSMALL,NUM1,NUM2,
ITFIX,VFSIG,CDHC,CWF,NAGLUM,SKIP,INSTFM
LOGICAL FIRST,IMPURE,INTRCT,INSTCN,FROZEN,PSMALL,NUM1,NUM2,
ITFIX,VFSIG,CDHC,CWF,NAGLUM,SKIP
COMMON/CONSTN/C,CK,CLF,CLL,CM(50),CMW,CSW,D,DELTIN,HL,PI,
IRG,RHOW,ONETRD,CVM,CVH,DIFCM(50),NUMPAR,RHOP,CP,
2SPEMIN,BK,TNOT,PNOT,DNOT,AF,BF,TF,AL,BL,CSI,ESO
COMMON/VARIAB/DELTIM,RHOVA,TAMB,TIM,CNU,S,A,B,
IRHOES,P
COMMON/ARRAYS/CN(50),CSD(50),DELHA(50),DELWV(50),HD(50),
IRHOVS(50),RAD(50),TSUR(50),V(50),DELCM(50),
2CF(50,50),DMDT(50),STATE(50),VG(50)
C*****SET01400
C
C THE FOLLOWING ARE CONSTANT ASSIGNMENTS --
C*****SET01800
C-----SET02000
C  CONSTANTS FOR SATURATION VAPOR PRESSURE OVER LIQUID COMPUTATION
C  DATA AL / 17.27 / SET02100
C  DATA BL / 35.68 / SET02200
C-----SET02400
C  CONSTANTS FOR SATURATION VAPOR PRESSURE OVER ICE COMPUTATION
C  DATA AF / 21.87 / SET02500
C  DATA BF / 7.66 / SET02600
C-----SET02800
C  BOLTZMANN'S CONSTANT
C  DATA BK / 1.38E-16 / SET02900
C  ERGS / DEGREE KELVIN SET03000
C-----SET03200
C  CONDUCTIVITY OF HEAT IN AIR
C  DATA CK / 2.42E3 / SET03300
C  ERGS / CENTIMETER - SECOND - DEGREE KELVIN SET03400
C-----SET03600
C  LATENT HEAT OF FREEZING OF WATER AT ZERO DEGREES CENTIGRADE
C  DATA CLF / 2.83E10 / SET03700
C  ERGS / GRAM SET03800
C-----SET04000
C  LATENT HEAT OF VAPORIZATION OF WATER AT ZERO DEGREES CENTIGRADE
C  DATA CLL / 2.5E10 / SET04100
C  ERGS / GRAM SET04200
C-----SET04400
C  MOLECULAR WEIGHT OF WATER
C  DATA CMW / 18.0 / SET04500
C-----SET04700
C  SPECIFIC HEAT OF AIR AT CONSTANT PRESSURE AND TEMPERATURE TAMB
C  DATA CP / 1.01E7 / SET04800
C-----SET05000
C  SPECIFIC HEAT OF WATER (ICE VALUE)
C  DATA CSI / 2.09E7 / SET05100
C  ERGS / GRAM - DEGREE KELVIN SET05200
C-----SET05400
C  SPECIFIC HEAT OF WATER (LIQUID VALUE) SET05500

```

```

DATA CSW / 4.18E07 / SET05600
C ERGS / GRAM - DEGREE KELVIN SET05700
C-----SET05800
C CONSTANTS FOR COMPUTING VENTILATION FACTOR SET05900
DATA CVH / 0.3 / SET06000
DATA CVM / 0.3 / SET06100
C-----SET06200
C STANDARD VALUE FOR DIFFUSIVITY OF WATER SET06300
DATA DNOT / .220 / SET06400
C CENTIMETER 2 / SECOND SET06500
C-----SET06600
C COEFFICIENT OF TETENS EQUATION FOR VAPOR PRESSURE SET06700
DATA ES0 / 6108. / SET06800
C-----SET06900
C STANDARD PRESSURE SET07000
DATA PNOT / 1.013E6 / SET07100
C DYNES / CENTIMETER 2 SET07200
C-----SET07300
C DENSITY OF WATER SET07400
DATA RHDW / 1.0 / SET07500
C GRAMS / CENTIMETER 3 SET07600
C-----SET07700
C STANDARD TEMPERATURE SET07800
DATA TNOT / 273. / SET07900
C DEGREES KELVIN SET08000
C-----SET08100
C GAS CONSTANT SET08200
DATA RG / 8.314E7 / SET08300
C ERGS / MOL - DEGREE KELVIN SET08400
C-----SET08500
C FREEZING TEMPERATURE SET08600
DATA TF / 273.16 / SET08700
C DEGREES KELVIN SET08800
C-----SET08900
C SET09000
C=====SET09100
C DIFFUSIVITY OF WATER VAPOR - ICT EQUATION SET09200
C=====SET09300
D=DNOT*(TAMB/TNOT)**1.81*(PNOT/P) SET09400
C SET09500
NUM1=NUMPAR.EQ.1 SET09600
NUM2=NUMPAR.EQ.2 SET09700
CLM=CLF-CLL SET09800
PI=4.0*ATAN(1.0) SET09900
IF(NUM1) GO TO 60 SET10000
DO 50 I=2,NUMPAR SET10100
DIFCM(I-1)=CM(I)-CM(I-1) SET10200
50 CONTINUE SET10300
60 DIFCM(NUMPAR)=1.0 SET10400
RETURN SET10500
END SET10600

```

```

SUBROUTINE COLLIS
C
COMMON/FLAGS/FIRST,IMPURE,INTRCT,INSTCN,FROZEN,PSMALL,NUM1,NUM2,
ITFIX,VFSIG,CDHC,CWF,NAGLOM,SKIP,INSTFM
LOGICAL FIRST,IMPURE,INTRCT,INSTCN,FROZEN,PSMALL,NUM1,NUM2,
ITFIX,VFSIG,CDHC,CWF,NAGLOM,SKIP
COMMON/CONSTN/C,CK,CLF,CLL,CM(50),CMW,CSW,D,DELTIN,HL,PI,
IRG,RHOW,ONETRD,CVM,CVH,DIFCM(50),NUMPAR,RHOP,CP,
2SPEMIN,BK,TNOT,PNOT,DNOT,AF,BF,TF,AL,BL,CSI,ESO
COMMON/VARIAB/DELTIM,RHOVA,TAMB,TIM,CNU,S,A,B,
IRHOES,P
COMMON/ARRAYS/CN(50),CSD(50),DELHA(50),DELWV(50),HD(50),
IRHOVS(50),RAD(50),TSUR(50),V(50),DELCM(50),
2CF(50,50),DMDT(50),STATE(50),VG(50)
DIMENSION DB(50)
C=====
C VISCOSITY OF AIR - SUTHERLAND EQUATION - CGS UNITS
C=====
VISC=TAMB**1.5*145.8E-07/(110.4+TAMB)
C=====
C COMPUTE BROWNIAN COLLISION FREQUENCY FB
C=====
TEMPD=BK*TAMB/(6.0*PI*VISC)
DO 100 I=1,NUMPAR
DB(I)=TEMPD/RAD(I)
100 CONTINUE
DO 300 J=1,NUMPAR
DO 200 I=1,NUMPAR
RIJ=RAD(I)+RAD(J)
FB=4.0*(DB(I)+DB(J))
C=====
C ADD GRAVITATIONAL COLLISION FREQUENCY FGT
C=====
CF(I,J)=FB+FGT(I,J)
C
200 CONTINUE
300 CONTINUE
WRITE(1,1000)VISC
1000 FORMAT(/,35X,12HVISCOSITY = ,F10.4)
RETURN
END
COL00000
COL00100
COL00200
COL00300
COL00400
COL00500
COL00600
COL00700
COL00800
COL00900
COL01000
COL01100
COL01200
COL01300
COL01400
COL01500
COL01600
COL01700
COL01800
COL01900
COL02000
COL02100
COL02200
COL02300
COL02400
COL02500
COL02600
COL02700
COL02800
COL02900
COL03000
COL03100
COL03200
COL03300
COL03400
COL03500
COL03600
COL03700
COL03800
COL03900
COL04000
COL04100
COL04200

```

```

FUNCTION FGT(I,J)
C
COMMON/FLAGS/FIRST,IMPURE,INTRCT,INSTCN,FROZEN,PSMALL,NUM1,NUM2,
1TFIX,VFSIG,CDHC,CWF,NAGLOM,SKIP,INSTFM
LOGICAL FIRST,IMPURE,INTRCT,INSTCN,FROZEN,PSMALL,NUM1,NUM2,
1TFIX,VFSIG,CDHC,CWF,NAGLOM,SKIP
COMMON/CONSTN/C,CK,CLF,CLL,CM(50),CMW,CSW,D,DELTIN,HL,PI,
1RG,RHOW,DNETRD,CVM,CVH,DIFCM(50),NUMPAR,RHOP,CP,
2SPEMIN,BK,TNOT,PNOT,DNOT,AF,BF,TF,AL,BL,CSI,ESQ
COMMON/VARIAB/DELTIM,RHOVA,TAMB,TIM,CNU,S,A,B,
1RHOES,P
COMMON/ARRAYS/CN(50),CSD(50),DELHA(50),DELWV(50),HD(50),
1RHOVS(50),RAD(50),TSUR(50),V(50),DELICM(50),
2CF(50,50),DMDT(50),STATE(50),VG(50)
FGT=C.0
C=====FGT01500
C DUMMY VALUE OF FGT WHEN V(I) AVAILABLE REPLACE BY FGT01600
C RIJ = RAD(I) + RAD(J) FGT01700
C FGT = PI * RIJ * RIJ * ABS(V(I) - V(J)) FGT01800
C OR USE FALLING RATES VG INSTEAD OF V FGT01900
C=====FGT02000
RETURN FGT02100
END FGT02200

```

```

SUBROUTINE NCONCN
C
COMMON/FLAGS/FIRST,IMPURE,INTRCT,INSTCN,FROZEN,PSMALL,NUM1,NUM2,
1TFIX,VFSIG,CDHC,CWF,NAGLOM,SKIP,INSTFM
LOGICAL FIRST,IMPURE,INTRCT,INSTCN,FROZEN,PSMALL,NUM1,NUM2,
1TFIX,VFSIG,CDHC,CWF,NAGLOM,SKIP
COMMON/CONSTN/C,CK,CLF,CLL,CM(50),CMW,CSW,D,DELTIN,HL,PI,
1RG,RHOW,ONETRD,CVM,CVH,DIFCM(50),NUMPAR,RHOP,CP,
2SPEMIN,BK,TNUT,PNOT,DNOT,AF,BF,TF,AL,BL,CSI,ESQ
COMMON/VARIAB/DELTIM,RHOVA,TAMB,TIM,CNU,S,A,B,
1RHOES,P
COMMON/ARRAYS/CN(50),CSD(50),DELHA(50),DELWV(50),HD(50),
1RHOVS(50),RAD(50),TSUR(50),V(50),DELCM(50),
2CF(50,50),DMDT(50),STATE(50),VG(50)
DIMENSION TN(50)
LOGICAL NSPMFG
IF(NAGLOM) GO TO 50
CALL COLLIS
50 NSPMFG=S.LT.SPEMIN
WRITE(1,1000)
1000 FORMAT(/,76X,7HRATE OF,
1/,30X,5HCLASS,5X,23HCHANGE IN CONCENTRATION,5X,
223HCHANGE IN CONCENTRATION,/,42X,18HDUE TO EVAPORATION,11X,
320HDUE TO AGGLOMERATION,/,44X,15HOR CONDENSATION,/,30X,5H-----,
45X,23(1H-),5X,23(1H-))
C=====
C COMPUTE CHANGE OF CLASS 1
C=====
CN11=CN(1)
DNDTAG=0.0
IF(NAGLOM) GO TO 104
DO 100 J=1,NUMPAR
100 DNDTAG=DNDTAG-CF(1,J)*CN(J)
IF(TFIX) GO TO 103
C=====
C TO PERMIT VARIABLE DELTIM WITH '.SKIP.', REVISE
C AEROSOL TO COMPUTE DELHA AND DRHOVA AFTER
C CALL OF NCONCN
C=====
101 IF(SKIP.OR.DNDTAG*DELTIM.LE.C) GO TO 102
DELTIM=DELTIM/2.
GO TO 101
102 IF(SKIP.OR.DNDTAG*DELTIM.GE.C/2.) GO TO 103
DELTIM=DELTIM*2.
GO TO 102
103 DNDTAG=DNDTAG*CN11
IF(SKIP) GO TO 104
DELCNI=C.0
GO TO 105
104 IF(DELCM(1).GT.0.0) DELCNI=-CN11*DELCM(1)/DIFCM(1)+AMAX1(0.0,
1-CN(2)*DELCM(2)/DIFCM(1))
C=====
C CLASS 1 CANNOT LOSE MASS TO SMALLER PARTICLES
C=====
IF(DELCM(1).LE.0.0) DELCNI=AMAX1(0.0,
1-CN(2)*DELCM(2)/DIFCM(1))

```

```

105 TN(1)=CNI1+DNDTAG*DELTIM+DELCNI
    I=1
    WRITE(1,1200) I,DELCNI,DNDTAG
1200 FORMAT(30X,I3,10X,E12.4,11X,E12.4)
C
    IF(NUM1) GO TO 205
C=====
C COMPUTE CHANGE OF CLASS 2
C=====
    CNI=CN(2)
    DNDTAG=0.0
    IF(NAGLOM) GO TO 113
    DNDTAG=-CF(2,1)*CNI1*.5
    DO 110 J=2,NUMPAR
110  DNDTAG=DNDTAG-CF(2,J)*CN(J)
    DNDTAG=DNDTAG*CNI+CF(1,1)*CNI1*CNI*.5
    IF(SKIP) GO TO 113
    DELCNI=0.0
    GO TO 115
113  IF(DELCM(2).GT.0.0) DELCNI=-CNI*DELCM(2)/DIFCM(2)+AMAX1(0.0,
    1CN(1)*DELCM(1)/DIFCM(1))+AMAX1(0.0,-CN(3)*DELCM(3)/DIFCM(2))
    IF(DELCM(2).LE.0.0) DELCNI=CNI*DELCM(2)/DIFCM(1)+AMAX1(0.0,
    1CN(1)*DELCM(1)/DIFCM(1))+AMAX1(0.0,-CN(3)*DELCM(3)/DIFCM(2))
115  TN(2)=CNI+DNDTAG*DELTIM+DELCNI
    I=2
    WRITE(1,1200) I,DELCNI,DNDTAG
C
    IF(NUM2) GO TO 205
C=====
C COMPUTE CHANGE OF CLASSES 3 THROUGH NUMPAR
C=====
    DO 200 I=3,NUMPAR
    CNI1=CNI
    CNI=CN(I)
    AGLOM1=0.0
    AGLOM2=0.0
    DNDTAG=0.0
    I1=I-1
    I2=I-2
    IF(NAGLOM) GO TO 131
    DO 120 J=1,NUMPAR
120  DNDTAG=DNDTAG-CF(I,J)*CN(J)
    TWOJI=2.0**(I-1)
    DO 130 J=1,I2
    AGLOM1=AGLUM1-CF(I,J)*CN(J)*TWOJI
    TWOJI=TWOJI*2.0
    AGLOM2=AGLUM2+CF(I1,J)*TWOJI*CN(J)
130  CONTINUE
    AGLOM1=AGLUM1-CF(I,I1)*CN(I1)*.5
    AGLOM2=AGLUM2+CF(I1,I1)*.5*CNI1
    DNDTAG=(DNDTAG+AGLUM1)*CNI+AGLUM2*CNI1
    IF(SKIP) GO TO 131
    DELCNI=0.0
    GO TO 135
131  IF(I.LT.NUMPAR) GO TO 133
    DELCNI=AMAX1(0.0,CN(I1)*DELCM(I1)/DIFCM(I1))

```

```

NC005600
NC005700
NC005800
NC005900
NC006000
NC006100
NC006200
NC006300
NC006400
NC006500
NC006600
NC006700
NC006800
NC006900
NC007000
NC007100
NC007200
NC007300
NC007400
NC007500
NC007600
NC007700
NC007800
NC007900
NC008000
NC008100
NC008200
NC008300
NC008400
NC008500
NC008600
NC008700
NC008800
NC008900
NC009000
NC009100
NC009200
NC009300
NC009400
NC009500
NC009600
NC009700
NC009800
NC009900
NC010000
NC010100
NC010200
NC010300
NC010400
NC010500
NC010600
NC010700
NC010800
NC010900
NC011000
NC011100

```

GO TO 135	NC011200
133 IF (DELCM(I).GT.0.0) DELCNI=-CN(I)*DELCM(I)/DIFCM(I)+AMAX1(0.0,	NC011300
1CN(I1)*DELCM(I1)/DIFCM(I1))+	NC011400
2AMAX1(0.0,-CN(I+1)*DELCM(I+1)/DIFCM(I))	NC011500
IF (DELCM(I).LE.0.0) DELCNI=CN(I)*DELCM(I)/DIFCM(I)+AMAX1(0.0,	NC011600
1CN(I1)*DELCM(I1)/DIFCM(I1))+	NC011700
2AMAX1(0.0,-CN(I+1)*DELCM(I+1)/DIFCM(I))	NC011800
135 TN(I)=CNI+DNNTAG*DELTIM+DELCNI	NC011900
WRITE(1,1200) I,DELCNI,DNNTAG	NC012000
200 CONTINUE	NC012100
C	NC012200
C=====	NC012300
C UPDATE ALL CN	NC012400
C=====	NC012500
205 DO 210 I=1,NUMPAR	NC012600
210 CN(I)=TN(I)	NC012700
C	NC012800
C=====	NC012900
C TEST FOR APPEARANCE/DISAPPEARANCE OF CONDENSED WATER	NC013000
C=====	NC013100
IF(CWF) GO TO 160	NC013200
C=====	NC013300
C CLASS 1 PARTICLES TREATED AS DRY NUCLEI	NC013400
C=====	NC013500
IF(NSPMFG) RETURN	NC013600
IF(DELCM.GT.0.0) GO TO 150	NC013700
140 CONTINUE	NC013800
RETURN	NC013900
150 CWF=.TRUE.	NC014000
RETURN	NC014100
160 DO 170 I=2,NUMPAR	NC014200
IF(CN(I).GT.0.0) RETURN	NC014300
170 CONTINUE	NC014400
C	NC014500
IF(NSPMFG) CWF=.FALSE.	NC014600
RETURN	NC014700
END	NC014800

```

SUBROUTINE CONCEN
C
C=====CON00000
C-----CON00100
C WARNING -- CONCEN MAY ONLY BE VALID FOR
C CONDENSING NOT EVAPORATING PARTICLES
C=====CON00200
C-----CON00300
C-----CON00400
C-----CON00500
C-----CON00600
COMMON/FLAGS/FIRST,IMPURE,INTRCT,INSTCN,FROZEN,PSMALL,NUM1,NUM2,
ITFIX,VFSIG,CDHC,CWF,NAGLOM,SKIP,INSTFM
CON00700
LOGICAL FIRST,IMPURE,INTRCT,INSTCN,FROZEN,PSMALL,NUM1,NUM2,
CON00800
ITFIX,VFSIG,CDHC,CWF,NAGLOM,SKIP
CON00900
COMMON/CONSTN/C,CK,CLF,CLL,CM(50),CMW,CSW,D,DELTIN,HL,PI,
CON01000
IRG,RHOW,UNETRD,CVM,CVH,DIFCM(50),NUMPAR,RHUP,CP,
CON01100
2SPEMIN,BK,TNOT,PNOT,DNOT,AF,BF,TF,AL,BL,CSI,ESO
CON01200
COMMON/VARIABLES/DELTIM,RHOVA,TAMB,TIM,CNU,S,A,B,
CON01300
IRHOES,P
CON01400
COMMON/ARRAYS/CN(50),CSD(50),DELHA(50),DELWV(50),HD(50),
CON01500
IRHOVS(50),RAD(50),TSUR(50),V(50),DELCM(50),
CON01600
2CF(50,50),DMDT(50),STATE(50),VG(50)
CON01700
DIMENSION DN(50)
CON01800
LOGICAL SPEMFG
CON01900
IF(NAGLOM) GO TO 50
CON02000
CALL CCLLIS
CON02100
WRITE(1,1000)
CON02200
1000 FORMAT(//,48X,7HRATE OF,
CON02300
1/,30X,5HCLASS,5X,23HCHANGE IN CONCENTRATION,
CON02400
2/,41X,20HDUE TO AGGLOMERATION,/,30X,5H-----,
CON02500
35X,23(1H-))
CON02600
50 NSPEMFG=S.LT.SPEMIN
CON02700
C COMPUTE CHANGE OF CLASS 1
CON02800
CN12=CN(1)
CON02900
DN12AG=0.
CON03000
IF(NAGLOM) GO TO 102
CON03100
DO 90 J=1,NUMPAR
CON03200
DN12AG=DN12AG+CF(1,J)*CN(J)
CON03300
90 CONTINUE
CON03400
92 IF(ITFIX) GO TO 97
CON03500
IF(SKIP.OR.DN12AG*DELTIM.LE.C) GO TO 97
CON03600
C=====CON03700
C-----CON03800
C TO PERMIT VARIABLE DELTIM WITH '.SKIP.', REVISE
CON03900
C AFROSOL TO COMPUTE DELHA AND DRHOVA AFTER
CON04000
C CALL OF CONCEN
CON04100
C=====CON04200
DELTIM=DELTIM/2.
CON04300
GO TO 92
CON04400
97 IF(SKIP) GO TO 102
CON04500
DN(1)=-CN12*DN12AG
CON04600
GO TO 103
CON04700
102 DN(1)=-CN12*DMDT(1)/DIFCM(1)-CN12*DN12AG
CON04800
C-----CON04900
C-----CON05000
C-----CON05100
C=====CON05200
C-----CON05300
C XDND IS RATE OF CHANGE OF CN(1) DUE TO AGGLOMERATION
CON05400
C=====CON05500
XDND=-CN12*DN12AG
CON05500

```

```

      IJK=1
      WRITE(1,1100) IJK,XDND
1100 FORMAT(30X,I3,10X,E12.4)
C=====
C COMPUTE CHANGE OF CLASS 2
C=====
      DNDTAG=0.
      IF(NAGLOM) GO TO 106
      DO 105 J=1,NUMPAR
      DNDTAG=DNDTAG+CF(2,J)*CN(J)
105 CONTINUE
C=====
C XDND IS RATE OF CHANGE OF CN(2) DUE TO AGGLOMERATION
C=====
      XDND=-CNI1*DNDTAG+.5*CF(1,1)*CNI2*CNI2
      IJK=2
      WRITE(1,1100) IJK,XDND
      IF(SKIP) GO TO 106
      DN(2)=-CNI1*DNDTAG+.5*CF(1,1)*CNI2*CNI2
      GO TO 107
106 DN(2)=-CNI1*DMDT(2)/DIFCM(2)+CNI2*DMDT(1)/DIFCM(1)-CNI1*DNDTAG
      1+.5*CF(1,1)*CNI2*CNI2
107 IF(NUM2) GO TO 125
C=====
C COMPUTE CHANGE OF CLASSES 3 THROUGH NUMPAR
C=====
      DO 120 I=3,NUMPAR
      CNI=CNI(I)
      IF(I.EQ.NUMPAR) CNI=0.0
      I1=I-1
      I2=I-2
      IF(SKIP) GO TO 108
      DNDTEC=0.0
      GO TO 109
108 DNDTEC=CNI1*DMDT(I1)/DIFCM(I1)-CNI*DMDT(I)/DIFCM(I)
      DNDTAG=0.0
      IF(NAGLOM) GO TO 117
109 DNDTAG=.5*CF(I1,I1)*CNI1*CNI1+.75*CF(I1,I2)*CNI1*CNI2
      SUMA=0.0
      TWOJ1=2.0**((1-I))
      DO 110 J=1,I2
      SUMA=SUMA+CF(I,J)*CN(J)*TWOJ1
      TWOJ1=TWOJ1*2.
110 CONTINUE
      SUMB=0.0
      DO 115 J=I1,NUMPAR
      SUMB=SUMB+CF(I,J)*CN(J)
115 CONTINUE
      IF(I.EQ.NUMPAR) CNI=CNI(I)
      DNDTAG=DNDTAG+CNI*(SUMA-SUMB)
C=====
C DN IS TOTAL RATE OF CHANGE OF CN
C=====
117 DN(I)=DNDTEC+DNDTAG
      CNI2=CNI1
      CNI1=CNI

```

```

      WRITE(1,1100) I,DNDTAG                                CON11200
120 CONTINUE                                              CON11300
C=====CON11400
C  UPDATE ALL CN                                          CON11500
C=====CON11600
      125 DO 130 I=1,NUMPAR                                CON11700
      130 CN(I)=CN(I)+DN(I)*DELTIM                        CON11800
C                                                        CON11900
C=====CON12000
C  TEST FOR APPEARANCE/DISAPPEARANCE OF CONDENSED WATER CON12100
C=====CON12200
      IF(CWF) GO TO 160                                    CON12300
      IF(NSPMFG) RETURN                                    CON12400
      DO 140 I=1,NUMPAR                                    CON12500
      IF(DELCM(I).GT.C.C) GO TO 150                      CON12600
140 CONTINUE                                             CON12700
      RETURN                                              CON12800
150 CWF=.TRUE.                                           CON12900
      RETURN                                              CON13000
160 DO 170 I=2,NUMPAR                                    CON13100
      IF(CN(I).GT.C.C) RETURN                            CON13200
170 CONTINUE                                             CON13300
C                                                        CON13400
      IF(NSPMFG) CWF=.FALSE.                             CON13500
      RETURN                                              CON13600
      END                                                CON13700

```

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