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**INVESTIGATOR(S):**

**Name:** Ph.D. William Anderson

**Email:** william.anderson@utdallas.edu

**Phone Number:** 9728834660

**Principal:** Y

Organization: **University of Texas at Dallas**

Address: 800 West Campbell Road, AD15, Richardson, TX 750803021

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Email: william.anderson@utdallas.edu

Phone: (972) 883-4660

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**RPPR Final Report**  
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THE UNIVERSITY OF TEXAS AT DALLAS

1 **Annual Report for Grant # W911NF-15-1-0231, Project # 65843-CH**

2 **Numerical simulation of atmospheric boundary-layer flow over battlefield-scale complex**  
 3 **terrain: surface fluxes from resolved and subgrid scales**

4 *A final report submitted to the Army Research Office, Environmental Chemistry Program*  
 5 *(Program Manager: Dr. Robert A. Mantz)*

6 PI: William Anderson, Mechanical Engineering Department, The University of Texas at Dallas

7  
 8 **1. Background**

9 The atmospheric boundary layer (ABL) – the dense, thin layer of the atmosphere nearest to the Earth  
 10 – is dominated by vigorous turbulent mixing produced by aerodynamic drag (shear) and thermal  
 11 forcing (buoyancy). The depth of the ABL,  $H$ , varies with the diurnal cycle, where radiative surface  
 12 heating during daytime hours acts in conjunction with ambient shear to thicken the ABL, while  
 13 during nighttime hours the relatively cool surface reverses the affect of buoyancy and suppresses the  
 14 production of turbulence via shear. The relative contribution to turbulence production via shear  
 15 and buoyancy is commonly quantified via the Obukhov length. In the absence of buoyancy (neutral  
 16 conditions), turbulence is produced only via shear. For conceptual purposes, the neutrally-stratified  
 17 ABL may be divided vertically into two horizontal layers: the surface and inertial layer, where the  
 18 former occupies the lowest  $\approx 10\%$  of the ABL and is heavily influenced by attributes of the underlying  
 19 landscape, while the latter is presumed to be horizontally homogeneous for a “balanced” ABL [1] and  
 20 is affected by Coriolis accelerations. This vertical layering can be defined with the profiles [2]:

$$\frac{\langle u(z) \rangle_{xyt}}{u_*} = \frac{1}{\kappa} \ln \left( \frac{z-d}{z_0} \right), \quad (1)$$

21 and

$$\frac{\langle u(z) \rangle_{xyt} - U_G}{u_*} = \frac{1}{\kappa} \ln \left( \frac{zf}{u_*} \right) + C, \quad (2)$$

22 where, in this report, the first, second, and third components of any vector correspond with the  
 23 streamwise, spanwise, and vertical component, respectively (i.e.,  $\mathbf{x} = x\hat{i} + y\hat{j} + z\hat{k}$ ). Moreover,  
 24 averaging over dimension,  $a$ , is denoted by  $\langle \dots \rangle_a$ . Equation 1, thus, defines mean (Reynolds averaged)  
 25 streamwise wind in the surface layer, while Equation 2 defines the tendency of  $\langle u(z) \rangle_{xyt}$  towards the  
 26 outer flow associated with the geostrophic balance (i.e.,  $U_G = -(f\rho)^{-1} dp_0/dx$ , where  $f$  is the Coriolis  
 27 parameter,  $\rho$  is atmospheric density, and  $dp_0/dx$  is the aloft pressure gradient). In Equations 1 and  
 28 2,  $u_* = (\tau_w/\rho)^{1/2} = [(H/\rho) dp_0/dx]^{1/2}$ . In Equation 2, the additive constant is  $C \approx 4.3$  [2], while in  
 29 Equation 1,  $z_0$  is the elevation at which  $\langle u(z) \rangle_{xyt} = 0$ , otherwise known as the momentum roughness  
 30 length, and  $d$  is displacement height [3, 4]. The Project # 65843-CH research effort is directed  
 31 towards development of prognostic models for  $z_0$  in Equation 1 – especially for complex, multiscale  
 32 topographies that exhibit self similarity over a broad spectrum of modes (whether associated with  
 33 fractal landscapes [5, 6], vegetative canopies [7], urban geometries [8], or simply with understanding  
 34 how turbulent shear flows respond to canonical fractal geometries [9, 10]). Since  $z_0$  is the only input

35 argument required to define  $\langle u(z) \rangle_{xyt}$ , its importance to numerical weather prediction at battlefield  
 36 scales cannot be overstated.

37 Consider, for example, atmospheric flow over a representative urban landscape,  $h(x, y)$ . The  
 38 smallest scale at which  $h(x, y)$  exhibits variability,  $\delta_R$ , is expected to be  $\delta_R \sim 10^{-1}$  m. In large-eddy  
 39 simulations (LES) of the atmospheric boundary layer, where the computational domain extent is  
 40  $L \sim 10^3$  m, the number of grid points,  $N$ , sets the computational resolution,  $\delta_C$ . In recent times,  
 41 atmospheric LES with  $N \sim 10^3$  have been performed [11, 12], although these incur an immense  
 42 computational challenge, and in reality  $N \sim 10^2$  is more widespread. With this, one can deduce that  
 43 the computational mesh resolution,  $\delta_C$ , is typically  $\delta_C \sim 10^0$  to  $10^1$  m, and thus  $R \sim \delta_R/\delta_C \gtrsim 10^1$ .  
 44 ***This rudimentary deduction highlights the challenges inherent to modeling atmospheric***  
 45 ***surface layer turbulence over complex, natural terrains, and motivate the need for***  
 46 ***prognostic wall models.***

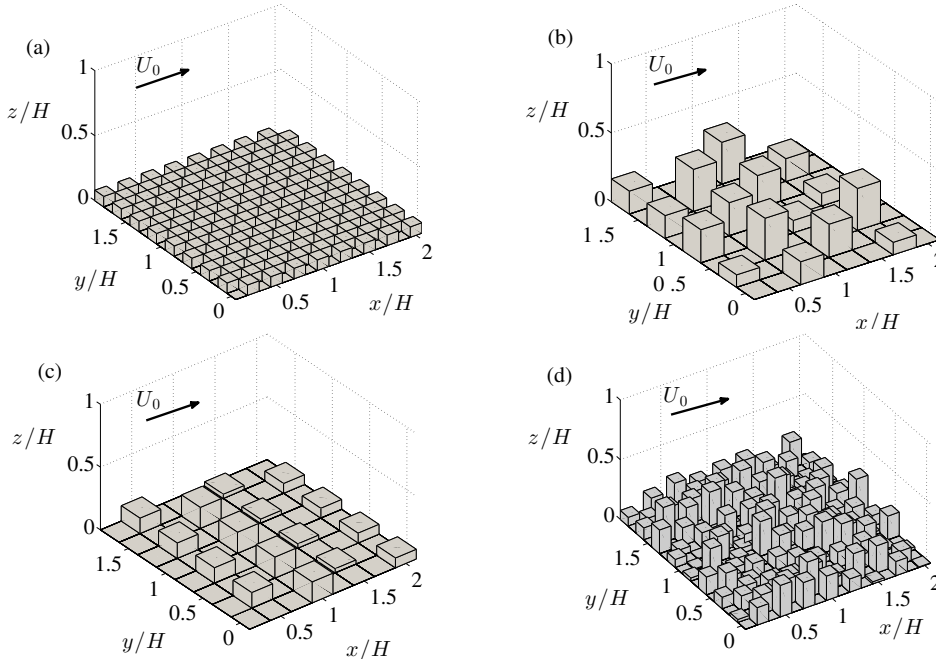


Figure 1: Illustration of synthetic, urban-like topographies considered in Zhu et al. [13] with freestream flow,  $U_0$ , added to denote forcing direction. Panel (a) shows a uniform height arrangement, Panels (b) and (c) show cases with bimodal heights, and Panel (c) shows a random case.

47 Given the physical definition of  $z_0$  (embodied in Eq. 1), it follows that a unique  $z_0$  exists for  
 48 each topography.  $z_0$  can be found *a posteriori* via experimental measurement, whereby the Reynolds-  
 49 averaged velocity profile is used to infer  $z_0$ . The practical limitations on this approach are obvious,  
 50 and this has motivated development of prognostic models. *A priori* estimation of  $z_0$  based only on  
 51 geometric details of the topography is an open research effort in boundary-layer meteorology. Models  
 52 for  $z_0$  play an important role in numerical weather prediction and generating realistic inflow conditions  
 53 for simulations of ABL turbulence [14–16].

54 During the Year # 1 research effort, we showed that models for  $z_0$  and  $d$  derived from the “engineer-  
 55 ing approach” can be leveraged in boundary-layer meteorology approaches [13, 16, 17]. Specifically,  
 56 we showed that:

$$d = d(\sigma_h) = \alpha_d \sigma_h, \quad (3)$$

57 where  $\alpha_d = 1.69$ , and  $\sigma_h$  is root-mean-square height, and

$$z_0(\sigma_h, sk) = \begin{cases} \alpha_G \sigma_h (1 + \beta sk), & \sigma_h / \langle h \rangle_{xy} + z_{0,p} < 1.15 \\ \alpha_G \sigma_h (1 + sk)^\beta, & \sigma_h / \langle h \rangle_{xy} + z_{0,p} \geq 2. \end{cases} \quad (4)$$

58 In Equation 4,  $\beta = 0.9$ ,  $\alpha_G = 0.1$ ,  $z_{0,p}$  is a “base” aerodynamic roughness to capture the influence  
 59 of small-scale elements, and  $s_k$  is the skewness of  $h(x, y)$ . In Zhu et al. [13], the efficacy of Eq.  
 60 4 was rigorously assessed via a comprehensive parametric study, all of which confirmed that: (1)  
 61 skewness is an important topographic momentum in setting the momentum roughness length; and  
 62 (2) only root-mean-square height is needed to model displacement height. Figure 1 illustrates sample  
 63 cases considered for the parametric testing. But, the cases considered in Zhu et al. [13] were “fully  
 64 resolved”, i.e.,  $R = 1$ . During the final years of the research effort, we considered fractal cities with  
 65  $R > 1$ , and leveraged the prognostic capabilities of Equations 3 and 4 to predict “subgrid-scale height”  
 66 effects.

## 67 2. Urban Fractal Geometries

68 In this study, we consider urban-like fractal geometries constructed via iterated function systems  
 69 (IFS) [18]. An IFS topology is, by definition, fractal, and is built via successive inner product  
 70 computations,  $\{f_i : \mathcal{X} \rightarrow \mathcal{X} | i = 1, 2, \dots, N_g\}$ , where  $\mathcal{X}$  is the underlying vector space. The resultant  
 71 function,  $\mathcal{F} = (\mathcal{X} : f_1, f_2, \dots, f_M)$ , is an IFS. For the present cases, we define  $f_i$  such that the side  
 72 length and height of each successive generation is  $l_{g+1} = rl_g$  and  $h_{g+1} = rh_g$ , where  $r$  is a pre-  
 73 defined scale-reduction factor. In order to consider topographies based upon a self-similar, Sierpiński  
 74 carpet template, we necessarily adopt  $r = 1/3$  for all cases (values exceeding this are not possible,  
 75 while values less than this would result in highly skewed topographies that would present resolution  
 76 challenges and not resemble urban-like topographies). In order to consider more realistic cases, we  
 77 added a random horizontal displacement to each successive mapping, such that the blocks were self-  
 78 similar across generations but displaced laterally to promote randomness. Figure 2 shows a sample  
 79 fractal city, built with IFS.

80 By virtue of the IFS approach, fractal dimension is nominated *a priori* via  $D = \log(N_b)/\log(r^{-1})$ ,  
 81 where  $N_b$  is the repeating factor (the number of descendant elements added during each IFS iteration).  
 82 Since the present topographies are composed via superposition of self-similar blocks, the spectral  
 83 density of height will not exhibit a power law dependence on wavelength with spectral exponent,  $\beta_s$ ,  
 84 (unlike the landscapes considered by [6]). Instead, the spectral density of the skewed topographies  
 85 would feature distinct peaks across the range of constituent wavelengths. Thus, Orey’s formula may  
 86 not be used to relate spectral exponent and fractal dimension. In total, we consider 33 cases in this  
 87 article, where Case 1 is a surface without any blocks, and serves the role of a benchmark case against  
 88 which others are compared. In the IFS construction process, the first generation is a single block at  
 89 the geometric center (Case 2; Figure 2a); beyond this, we consider the five series with different fractal  
 90 dimension,  $D$ , and increasing number of generations,  $N_g$  (Figure 2b-d). Cases 18 to 22 are identical to  
 91 Case 2, but with an added drag for sub-generation elements (defined via the equilibrium logarithmic  
 92 law, where effective roughness length is defined with Equation 4).

93 To assess the role of the random spatial shift added during each successive generation for Cases 3  
 94 to 17, we have also modeled flow over Cases 23 to 28, which are “pure” Sierpiński carpet topographies  
 95 with no spatial shift. Cases 23 to 25 are built with  $D = 1.2619$  (i.e., a comparison against Cases 3  
 96 to 5), while Cases 26 to 27 are built with  $D = 1.8928$  (i.e., a comparison against Cases 15 to 17).  
 97 Finally, as a matter of completeness, results of grid-sensitivity testing are presented in Appendix A.  
 98 Cases 29, 30, 31, 32, and 33 are geometrically identical to Cases 5, 8, 11, 14, and 17, respectively,  
 99 but the simulations have been carried out at higher resolution. Results have indicated no sensitivity  
 100 between the high- and low-resolution cases.

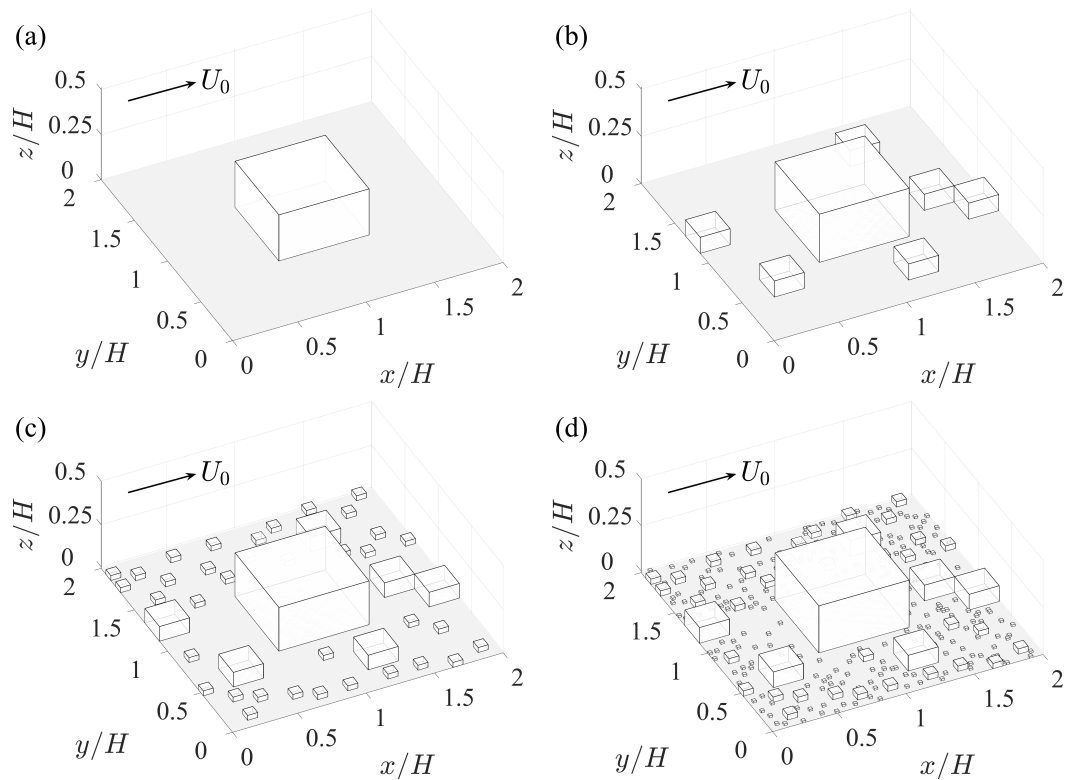


Figure 2: Illustration of multiscale, fractal-like urban topographies of Series 3 ( $D = 1.6309$ ): Panels (a) to (d) show Case 2, Case 9, Case 10, and Case 11, respectively. The topographies built in this figure are based on the Sierpiński carpet template with random horizontal shift, so they are fractals with self-similarity among generations. For example, in Panel (b), the six descendant blocks are self similar with the large central block in Panel (a), with the relation:  $l_2 = r \times l_1$  and  $h_2 = r \times h_1$ , where  $r = 1/3$  is the scale-reduction factor. Annotation,  $U_0$ , added to denote streamwise forcing direction.

Case	Series	$N_b$	$D$	$N_g$	$(\hat{z}_0/H)^*$	$d/H$	$z_{0,\text{Eff.}}/H^*$	$\alpha_g(N_g, D)$	$\alpha(D)$
1	-	-	-	0	$10^{-3}$	0	0.001	-	-
2	-	-	-	1	$10^{-3}$	0.135	4.173	0.016	0.050
3				2	$10^{-3}$	0.116	5.922	0.026	0.069
4	1	4	1.2619	3	$10^{-3}$	0.097	7.173	0.031	0.084
5				4	$10^{-3}$	0.097	7.795	0.034	0.091
6				2	$10^{-3}$	0.107	6.504	0.028	0.076
7	2	5	1.4650	3	$10^{-3}$	0.089	7.797	0.034	0.091
8				4	$10^{-3}$	0.089	8.327	0.037	0.097
9				2	$10^{-3}$	0.110	7.237	0.032	0.084
10	3	6	1.6309	3	$10^{-3}$	0.086	8.998	0.040	0.104
11				4	$10^{-3}$	0.085	9.795	0.044	0.113
12				2	$10^{-3}$	0.101	7.275	0.032	0.084
13	4	7	1.7712	3	$10^{-3}$	0.084	9.516	0.043	0.109
14				4	$10^{-3}$	0.082	10.117	0.045	0.114
15				2	$10^{-3}$	0.106	7.571	0.034	0.087
16	5	8	1.8928	3	$10^{-3}$	0.080	9.505	0.044	0.108
17				4	$10^{-3}$	0.078	10.216	0.047	0.116
18	$1^M$	4	1.2619	1	2.710	0.061	7.689	-	-
19	$2^M$	5	1.4650	1	2.959	0.054	8.671	-	-
20	$3^M$	6	1.6309	1	3.544	0.051	10.267	-	-
21	$4^M$	7	1.7712	1	3.662	0.046	10.653	-	-
22	$5^M$	8	1.8928	1	3.800	0.041	11.068	-	-
23				2	$10^{-3}$	0.106	6.184	0.027	0.072
24	1	4	1.2619	3	$10^{-3}$	0.094	7.361	0.032	0.086
25				4	$10^{-3}$	0.091	7.508	0.033	0.089
26				2	$10^{-3}$	0.108	7.804	0.035	0.089
27	5	8	1.8928	3	$10^{-3}$	0.085	10.135	0.047	0.115
28				4	$10^{-3}$	0.083	10.602	0.049	0.121
29	1	4	1.2619	1	$10^{-3}$	0.096	9.024	-	-
30	2	5	1.4650	1	$10^{-3}$	0.087	9.873	-	-
31	3	6	1.6309	1	$10^{-3}$	0.084	10.164	-	-
32	4	7	1.7712	1	$10^{-3}$	0.085	10.513	-	-
33	5	8	1.8928	1	$10^{-3}$	0.083	11.205	-	-

Table 1: Topographies used to determine  $z_{0,\text{Eff.}}$  and evaluate correlation parameters. Cases 1 to 17 are randomly distributed Sierpiński carpet, Cases 18 to 22 are Case 2, accompanied by a model for unresolved generations, Cases 23 to 28 are “pure” Sierpiński carpet templates (i.e., without the horizontal shift). Cases 29 to 33 are equivalent to Cases 5, 8, 11, 14, and 17, respectively, but the simulations run with relatively higher resolution (in order to assess resolution dependence).  $r$  is the scale-reduction ratio,  $N_b$  is the repeating factor (IFS input argument),  $D$  is the fractal dimension,  $N_g$  is the generation number,  $d/H$  and  $z_{0,\text{Eff.}}/H$  is zero-plane displacement height and effective aerodynamic roughness length, respectively, determined *a posteriori* and normalized by the flow depth. In addition,  $\alpha$  and  $\alpha_g$  are roughness model parameters outlined in Equation 4 and Section 3.1. \*denotes multiplication by  $10^3$ .

### 101 3. Results

102 This section is composed of two complementary parts. In the first part (Section 3.1), results of  
103 *a posteriori*-determined model parameters from Equation 4 are shown. We show that Equation 4  
104 can reproduce the momentum deficit for unresolved generations, i.e., the “bulk” parameter exhibits  
105 striking similarity. In Section 3.2, we report vertical profiles of turbulence statistics. Averages are  
106 based upon many thousand large-eddy turnovers (a description of the LES code is not provided in this  
107 report, for brevity; the code has been summarized in numerous preceding articles). Note also that

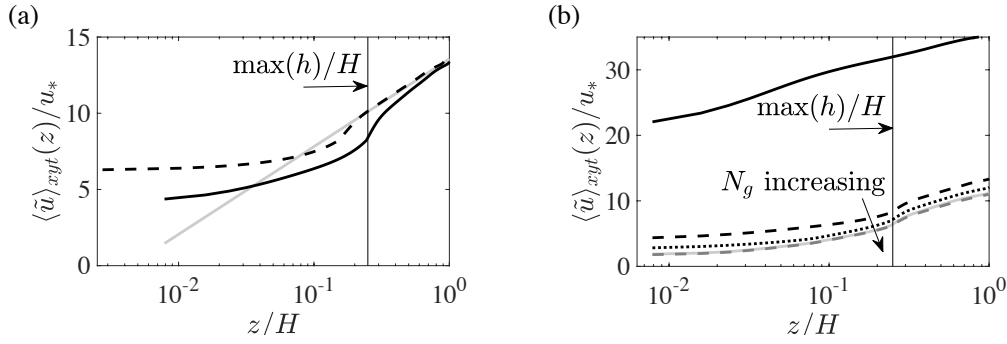


Figure 3: Vertical profiles of plane- and time-averaged streamwise velocity. Panel (a): Case 2 profile from LES (solid black), d-compensated velocity profile (dashed black), and logarithmic profile (solid gray). Panel (b): Case 1 (solid black), Case 2 (dashed black), Case 15 (dotted black), Case 16 (solid gray), and Case 17 (dashed gray). For reference, the maximum height of the topography is shown,  $\max(h)/H$ , which correlates closely with the inflection elevation; Panel (b) includes annotation for the direction of increasing  $N_g$ .

108 spatial (plane) averaging,  $\langle \dots \rangle_{xy}$ , is performed on a “fluid averaging” basis; that is, the cumulative  
 109 velocity at each level is divided only by the number of nodes occupied by fluid [19]. It is not instructive  
 110 to show results for all cases in Table 1; instead, we have considered “limiting cases”, on the premise  
 111 that turbulence statistics vary monotonically with intermediate topographic arrangements.

### 112 3.1. Parameter evaluation, model efficacy

113 Figure 3 shows vertical profiles of plane- and time-averaged streamwise velocity. Figure 3a shows the  
 114 velocity profile for Case 2, before (solid black) and after (dashed black) correction for the displacement  
 115 height,  $d = \int_0^h z f_{b,x}(z) dz / \int_0^h f_{b,x}(z) dz$ , where  $f_{b,x}$  is the streamwise component of the body force  
 116 imposed via IBM. For completeness,  $d/H$  is summarized in Table 1. With this,  $d/H$  is subtracted  
 117 from elevation,  $z/H$ , to generate the dashed black line, and this profile is extrapolated to determine  
 118  $z_{0,\text{Eff}}$  *a posteriori*. We remind the reader that spectral discretization in the horizontal directions  
 119 renders all cases physically equivalent to flow over an infinitely-repeating array of blocks.

120 Figure 3b shows vertical profiles of plane- and time-averaged streamwise velocity for Case 1, Case  
 121 2, and Cases 15 to 17. Thus, for a given fractal dimension (in Figure 3b,  $D = 1.8928$ ), the figure  
 122 reveals how momentum penalty increases monotonically with number of generations,  $N_g$ . Due to  
 123 the physics of the problem and the observation of diminishing influence of higher  $N_g$ , we find that  
 124 a convergence may exist even with additional IFS iterations. This is due to the homogenization of  
 125 additional blocks within the additional background roughness that is defined *a priori* within Equation  
 126 4 ( $z_{0,p}$ ).

127 For Cases 2 to 17, Figure 4a shows  $\alpha_g$  from Equation 4, determined *a posteriori* by fitting logarithmic  
 128 profiles to the d-compensated plane- and time-averaged velocity profiles (for completeness,  $z_{0,\text{Eff}}$   
 129 is summarized in Table 1). Annotation of the direction of increasing fractal dimension is also shown  
 130 on Figure 4a. One can deduce that  $\alpha_g = \alpha_g(N_g, D)$ , although the  $N_g$ -dependence attenuates over the  
 131 range considered. Physically, one can envision a scenario of “diminishing returns” with increasing  
 132  $N_g$ , to the extent that additional iterations will become homogenized with the “background”  $z_{0,p}$  in  
 133 Equation 4. Moreover, there are practical limitations on the requisite number of iterations, since nat-  
 134 ural urban topographies do not continue self-replicating *ad infinitum*. Note that the dependence and  
 135 trends in  $\alpha_g(N_g, D)$  are roughly equivalent for the “pure” and randomly-shifted fractal topographies  
 136 (symbols denoted in Figure 4a caption).

137 Given the ostensible  $\alpha_g$  tendency towards convergence at  $N_g = 4$ , beyond which  $\alpha_g = \alpha_g(N_g, D) \rightarrow$

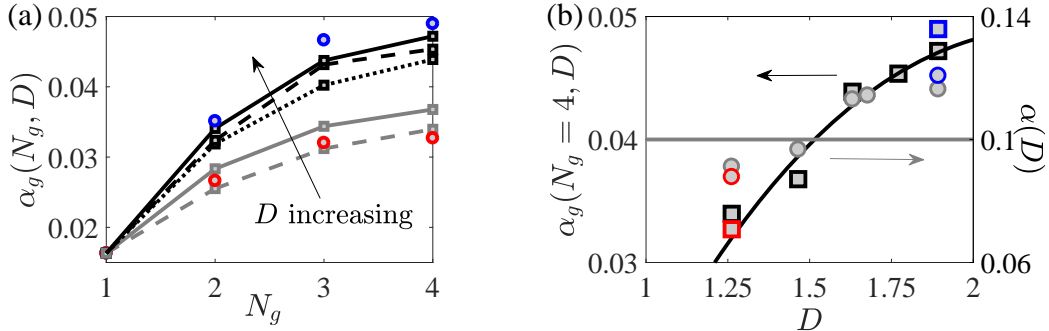


Figure 4: (Color) Equation 4 parameters. Panel (a) shows  $\alpha_g$  with respect to  $N_g$ , for the five values of  $D$  considered (Series 1, 2, 3, 4, and 5 denoted by dashed gray line, solid gray line, dotted black line, dashed black line, and solid black line, respectively). The squares correspond with the randomly-shifted cases (Cases 2, 3-17), while the circles correspond with a “pure” Sierpiński carpet template, with red for Cases 23 to 25, and blue for Cases 26 to 28, respectively. Panel (b) shows the coefficient  $\alpha_g(N_g = 4, D)$  (left ordinate, filled squares) and  $\alpha = \alpha_g(N_g = 4, D)(1 + sk)^\beta$  (right ordinate, filled circles) with respect to fractal dimension. Red circle and square correspond with Case 25, blue circle and square correspond with Case 28.

138  $\alpha_g(D)$ , we selected  $\alpha_g(N_g = 4, D)$  as the “steady state” value. These values are summarized in Figure  
 139 4b, alongside the aggregate prefactor,  $\alpha(D) = \alpha_g(N_g = 4, D)(1 + sk)^\beta$ , in Equation 4. It is important  
 140 to note that while  $\alpha_g(N_g = 4, D) \sim D^2$  (solid black),  $\alpha(D) \rightarrow \alpha \approx 10^{-1}$ , as the  $D$ -dependence  
 141 effectively vanishes; inclusion of the values for the “pure” and randomly-shifted Sierpiński carpet  
 142 template reveals that the layout is essentially negligible. Note that either  $\alpha = 10^{-1}$  or  $\alpha_g(1 + sk)^\beta$   
 143 can be adopted for modeling unresolved generations.

144 It is worthwhile noting that  $\alpha = 10^{-1}$  has a strong precedent [6]. Many prior studies — based on  
 145 simulations, experiments, and field observations of the atmospheric surface layer, and in applications  
 146 to a variety of dramatically-different rough boundaries — have reported  $\alpha \approx 10^{-1}$ . Although it is  
 147 beyond the scope of this article to survey all preceding studies, we direct the interested reader to  
 148 articles including Garratt [20], Zagarola and Smits [21], Tsai and Tsuang [22], Schultz and Flack [23],  
 149 Flack and Schultz [16], Anderson and Meneveau [6], Anderson [24], Zhu et al. [13], and references  
 150 therein. For completeness,  $\alpha(D)$  and  $\alpha_g(N_g, D)$  are summarized in Table 1.

151 The prognostic efficacy of the model can be readily assessed by considering a series of additional  
 152 LES cases, wherein the Case 2 simulation is repeated, but the sub-generation roughness model is  
 153 used where  $h(x, y) = 0$ , that is,  $z_{0,p}$  in Equation 4 becomes  $z_{0,p} \rightarrow \alpha_g(D)(1 + sk)^\beta \sigma_h^\Delta$ , where  $\sigma_h^\Delta$  is  
 154 the root-mean-square associated with sub-generation iterations (i.e.,  $2 \leq N_g \leq 4$ ). In this sense, the  
 155 aerodynamic loading associated with sub-generation scale iterations is imposed entirely through the  
 156 equilibrium logarithmic law, and the approach is valid since the model parameters have been tuned *a*  
 157 *priori* from the preceding exercise. For  $D = 1.2619$ , the results of this approach are shown in Figure  
 158 5, where the solid black and solid gray profiles correspond with Cases 2 and 5, respectively, where the  
 159 velocity shift between the profiles is attributed to unresolved descendant generations.

160 When  $z_{0,p} \rightarrow z_0$  (in Equation 4), however, the profiles for Case 5 and Case 18 (Case 2 with  
 161 Equation 4 sub-generation model) collapse in the inertial layer; some deviation in the roughness  
 162 sublayer,  $z/H \lesssim \max(h)/H$ , is guaranteed since actual elements are not being resolved. Although  
 163 there are some differences in the roughness sublayer, they are confined to the lowest 10 % of the flow,  
 164 and thus do not undermine the prognostic capabilities of the model. To further emphasize this, Figure  
 165 5 includes an inset for error,

$$\varepsilon(D) = \left\langle \left[ \frac{[M(D, z_R) - O(D, z_R)]^2}{O(D, z_R)^2} \right]^{1/2} \right\rangle_{z_R}, \quad (5)$$

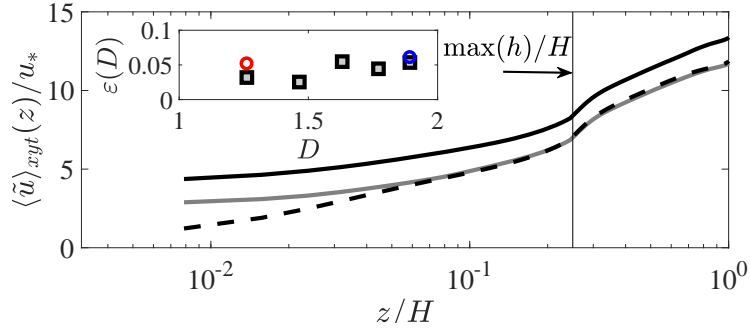


Figure 5: (Color) Vertical profile of plane- and time-averaged streamwise velocity for fully resolved Case 5 (solid gray), Case 2 without sub-generation model (solid black), and Case 18 (dashed black), which is identical to Case 2 but with sub-generation roughness model (Equation 4). Inset panel shows error,  $\varepsilon(D)$ , as per Equation 5 and accompanying text; squares correspond with randomly-shifted Sierpiński carpet templates, while red and blue symbols correspond with “pure” Sierpiński carpet template. Wall-normal position corresponding with maximum height,  $\max(h)/H$ , is shown for reference.

166 where  $M(D, z_R) = \langle \tilde{u} \rangle_{xyt}(z_R)$  for Cases 18 to 22.  $O(D, z_R) = \langle \tilde{u} \rangle_{xyt}(z_R)$  for Cases 5, 8, 11, 14, 17  
 167 (filled squares with black border), Case 25 (filled red circle), and Case 28 (filled blue circle) in the  
 168 range  $0.25 \leq z_R/H \leq 0.5$ . Other ranges could have been used, but we found this sufficient to quantify  
 169 differences between the modeled and actual cases. It is clear that  $\varepsilon(D)$  is small, and insensitive to  
 170  $D$ . These errors are reasonable and, again, do not undermine the prognostic capabilities embodied in  
 171 this approach.

172 The results presented in Figure 5 indicate that Equation 4 is a promising approach to the pa-  
 173 rameterization of small-scale, unresolved generations. The results also reveal that — in a first-order,  
 174 plane- and time-averaged sense — the roughness model parameters are insensitive to the spatial lay-  
 175 out of the underlying topography. However, these results do not illustrate how the dynamics of aloft  
 176 mixing responds to the omission (or random spatial shift) of descendant generations. In Section 3.2,  
 177 we address precisely this point via presentation of higher-order turbulence statistics.

### 178 3.2. Turbulence statistics and Sierpiński carpet templates

179 To assess the aerodynamic signature of the “pure” Sierpiński carpet template (Cases 23 to 28), the  
 180 randomly-shifted Sierpiński template (Cases 3 to 17), and cases with models for unresolved sub-  
 181 generations (Cases 18 to 22), we show vertical profiles of time- and plane-averaged first- to third-  
 182 order turbulence statistics. We show only plane-averaged statistics, aware that spatial heterogeneities  
 183 associated with emanating wakes, etc., may undermine the statistics. In results not shown here, we  
 184 have assessed vertical profiles from virtual towers, and the results differ only within the canopy (i.e.,  
 185 for  $z/\max(h) \lesssim 1$ ). This is due, mainly, to the overarching role played by the first generation (Case  
 186 2), where additional blocks appear to impart only a modest influence (at least as observed with the  
 187 turbulence statistics considered here).

188 Figure 6 shows vertical profiles of time- and plane-averaged streamwise velocity for cases corre-  
 189 sponding with the upper and lower limit of fractal dimension (Figure 6a:  $D = 1.2619$ ; Figure 6b:  
 190  $D = 1.8928$ ). The figures provide a means to assess: (1) how the outer (inertial) layer statistics re-  
 191 spond to parameterized descendant generations; and (2) how the outer layer responds to the random  
 192 spatial shift of successive generations, relative to the “pure” Sierpiński carpet template. There is  
 193 generally very close agreement between the different profiles for  $z/H \gtrsim 0.1$ , which falls beneath the  
 194 height of the large central block. This result is partially reflected in the Figure 5 inset, but the results  
 195 in Figure 6 would also suggest that the choice of “pure” or randomly-shifted successive generations  
 196 during IFS construction has negligible impact on the flow.

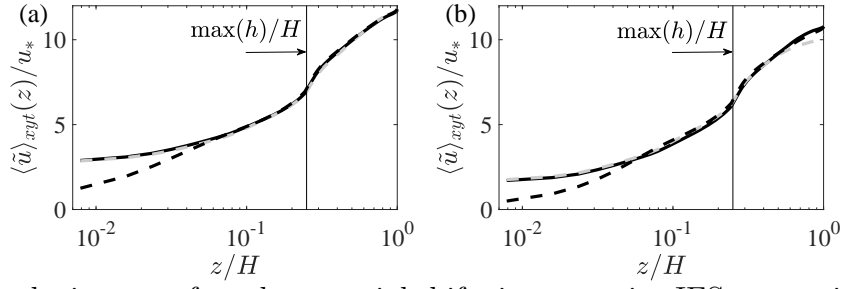


Figure 6: Assessing the impact of random spatial shifts in successive IFS generations. Vertical profiles of plane- and time-averaged streamwise velocity. On Panel (a), profiles correspond with Cases 5 (solid black), 18 (dashed black), and 25 (dashed gray). On Panel (b), profiles correspond with Cases 17 (solid black), 22 (dashed black), and 28 (dashed gray).

197 We present vertical profiles of Reynolds-averaged second- and third-order turbulence moments.  
 198 Beginning firstly with vertical profiles of the total turbulent stresses:

$$\langle \mathbf{u}' \otimes \mathbf{u}' \rangle_t = \underbrace{\langle \tilde{\mathbf{u}}' \otimes \tilde{\mathbf{u}}' \rangle_t}_{\text{Resolved}} + \underbrace{\langle \boldsymbol{\tau} \rangle_t}_{\text{Subgrid}} \quad (6)$$

199 where  $\langle \tilde{\mathbf{u}}' \otimes \tilde{\mathbf{u}}' \rangle_t$  and  $\langle \boldsymbol{\tau} \rangle_t$  denote contributions to the total stresses as depicted by the underbraces  
 200 [25]. Vertical profiles of the streamwise–wall-normal Reynolds stress,  $\langle u'w' \rangle_{x,y,t}(z)$ , are shown in  
 201 Figure 7(a,e). In addition, we compute the contribution from Quadrants 1 to 4 events to the resolved  
 202 streamwise–wall-normal turbulent stress based on,

$$\langle \tilde{u}' \tilde{w}' \rangle(\mathbf{x}, t; \mathcal{H})|_Q = \frac{1}{N_t} \sum_{j=1}^{N_t} \tilde{u}'(\mathbf{x}, t_j) \tilde{w}'(\mathbf{x}, t_j) I_Q(\mathbf{x}; \mathcal{H}), \quad (7)$$

where  $\mathcal{H}$  is the so-called hole size and used to nominate a threshold on the magnitude of terms contributing to the turbulent stresses,  $N_t$  is the number of instants over which the conditional stress is to be recorded, and  $I_Q(\mathbf{x}; \mathcal{H})$  is an indicator function used to isolate the role of different events based on their magnitude (or some predefined criteria):

$$I_Q(\mathbf{x}; \mathcal{H}) = \begin{cases} 1 & \text{if } |\tilde{u}'(\mathbf{x}, t) \tilde{w}'(\mathbf{x}, t)|_Q \geq T, \text{ and} \\ 0 & \text{if } |\tilde{u}'(\mathbf{x}, t) \tilde{w}'(\mathbf{x}, t)|_Q < T, \end{cases}$$

203 where  $T = \mathcal{H} = 0$  in the present study, though alternative criteria can be used to define the threshold.  
 204 For boundary layer turbulence, the ejection ( $Q2$ ) and sweep ( $Q4$ ) events are most relevant to the shear  
 205 stress, while the contributions from  $Q1$  and  $Q3$  events are known to be rather modest [26]. Existing  
 206 studies frequently report vertical profiles of the ratio of ejections to sweeps, i.e.,  $Q2/Q4$  [4, 19], and  
 207 we adopt this approach and show the results in Figure 7(b,f).

208 Finally, in Figure 7(c,d,g,h), we show vertical profiles of skewness of resolved streamwise and  
 209 vertical velocity fluctuation:

$$sk_{\tilde{u}'} = \frac{\langle (\tilde{u}')^3 \rangle_t}{\langle (\tilde{u}')^2 \rangle_t^{3/2}}, \quad \text{and} \quad sk_{\tilde{w}'} = \frac{\langle (\tilde{w}')^3 \rangle_t}{\langle (\tilde{w}')^2 \rangle_t^{3/2}}. \quad (8)$$

210 In the interest of generality, we show  $\langle u'w' \rangle_t$ ,  $Q2/Q4$ ,  $sk_{\tilde{u}'}$ , and  $sk_{\tilde{w}'}$  for Cases with  $D = 1.262$  (Figure  
 211 7a-d) and  $D = 1.893$  (Figure 7e-h) – the upper and lower limit of fractal dimension considered in this  
 212 article. In addition, we show results for the randomly-shifted Sierpiński templates, pure Sierpiński  
 213 templates, cases with modeled descendant generations, and Case 2, which serves as a comparative  
 214 benchmark (description of the figure profiles provided in caption).

215 A series of key deductions are made from inspection of Figure 7. First, note that for the modeled  
 216 cases (Cases 18 and 22, both denoted by light gray on the panels), the domain includes only the  
 217 first-generation block, while the *a posteriori*-determined roughness model parameters are used in  
 218 Equation 1 to model the effects of descendant generations during LES. To this extent, Cases 18 and  
 219 22 are low-pass filtered versions of Cases 5 and 17 (dark gray on Figure 7 panels), respectively, while  
 220 Cases 25 and 28 are “pure” Sierpiński carpet analogs of the fully-resolved, randomly-shifted cases.

221 Above the canopy,  $z/\max(h) > 1$  (horizontal black line), the profiles all rapidly converge. This  
 222 is especially true of the Reynolds stresses (Figure 7a,e), while effects of descendant blocks and/or  
 223 block layout are slightly more persistent in other statistics. For all turbulence quantities, however,  
 224 by  $z/h \gtrsim 2$  the profiles are indistinguishable. Recall that Case 2 (red) is a central block, but spectral  
 225 discretization in the horizontal directions renders this case equivalent to an infinitely-repeating array  
 226 of large blocks. To this extent, the arrangement may be viewed as flow over an array of blocks, and  
 227 thus the roughness sublayer will be occupied by a mixing layer-like structure (obstructed shear flow  
 228 [27]). Other flow metrics (spectral density, for example), may reveal persistent flow perturbations  
 229 over the depth of the domain [28], although such considerations are not the focus of this article.

230 It is certainly clear that differences emerge within the canopy,  $z/h \lesssim 1$ , although the differences  
 231 are modest. Perhaps most striking are differences in  $Q2/Q4$  (Figures 7b,f) and  $sk_{\tilde{u}'}$  (Figure 7c,g)  
 232 for  $z/h \lesssim 1$ .  $Q2/Q4$  is larger for the benchmark (Case 2, red), indicating that sweeps contribute  
 233 relatively less to the formation of Reynolds stresses. This can be readily understood via geometric  
 234 considerations: Case 2 is an isolated block, and so the contribution from sweeps due to Kelvin-  
 235 Helmholtz-like eddies in the roughness sublayer must be relatively weaker for this case since there  
 236 are a lesser number of elements available to initiate their production. The result is also reflected in  
 237 the skewness, where  $\tilde{u}'$  exhibits a greater tendency towards negative values within the canopy. The  
 238 vertical velocity skewness appears least sensitive to building layout.

239 The vertical profiles for  $Q2/Q4$  (Figure 7b,f) and  $sk_{\tilde{u}'}$  (Figure 7c,g) exhibit a distinct undulation  
 240 beneath the elevation of the first block,  $z/h \lesssim 1$ . A substantial effort was dedicated to isolating  
 241 the source of this undulation, since it was at first presumed to be a post-processing error. However,  
 242 the undulation remained insensitive to differing averaging techniques: for the purpose of this article,  
 243 plane averaging is used during LES to compute fluctuations, and the resultant turbulence quantities  
 244 are time averaged *a posteriori*; we also loaded a Reynolds-averaged flow,  $\langle \mathbf{u} \rangle_t$ , and used this to  
 245 accumulate fluctuations during LES, and the resultant turbulence quantities were indistinguishable to  
 246 those attained with the former approach. In additional testing (results not shown here) to demonstrate  
 247 that the undulation is a physical outcome associated with the urban topology considered – and not  
 248 the LES code or averaging procedures – we have modeled flow over an urban arrangement considered  
 249 previously by Coceal et al. [4]. They performed direct numerical simulation (DNS) of flow over an  
 250 array of uniform height, staggered blocks, under inertial-dominated conditions (i.e., dynamic similarity  
 251 exists between their DNS and our LES). The  $Q2/Q4$  profiles, and the contribution to the resolved  
 252 streamwise–wall-normal turbulent stresses from both, agree closely with the Coceal et al. [4] data.

253 Finally, it must be emphasized that the choice of building spatial template – pure or randomly-  
 254 shifted Sierpiński – or the use of models for descendant generation, has a minimal influence on the  
 255 inertial layer flow attributes (again, based on the turbulence statistics considered here).

#### 256 4. Concluding remarks

257 Multiscale, fractal-like geometries are commonly encountered by the atmospheric surface layer, with  
 258 urban topographies posing particular closure challenges owing to the disparate range of scales and  
 259 dramatic aerodynamic signature. The research effort enabled via ARO Project # 65843-CH has  
 260 addressed precisely this challenge. The PI and a supporting doctoral student have introduced a new  
 261 closure for modeling aerodynamic drag – used previously for “engineering”-type flows, but capable

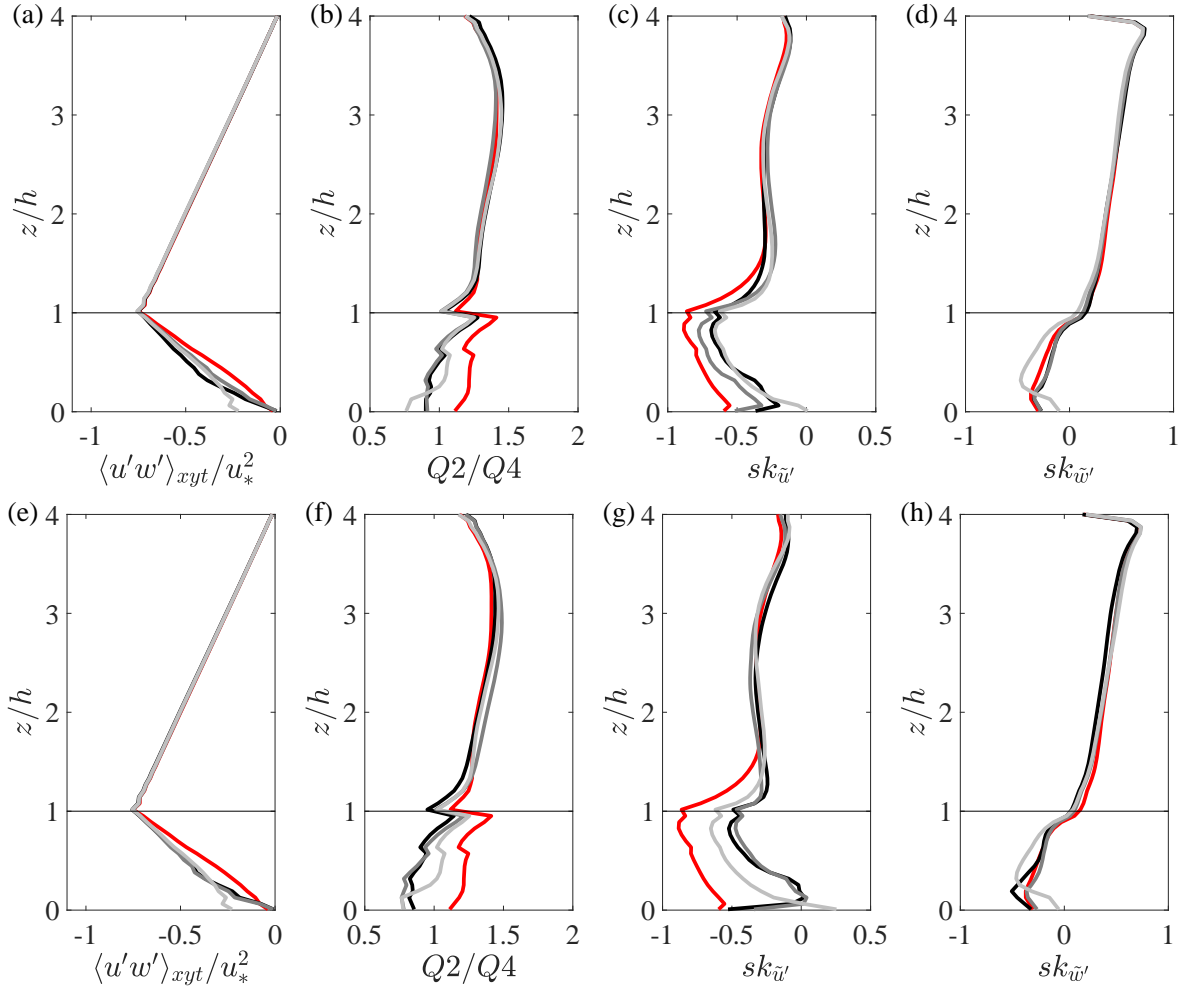


Figure 7: (Color) Vertical profiles of plane- and Reynolds-averaged turbulence statistics. Panels (a,e): streamwise–wall-normal Reynolds (total) stresses (Equation 6). Panels (b,f): ratio of contribution to production of resolved streamwise–wall-normal total stresses from sweeps, relative to ejections (Equations 7 and 3.2). Panels (c,g): skewness of resolved streamwise velocity (Equation 8). Panels (d,h): skewness of resolved vertical velocity (Equation 8). On Panels (a) to (d), line colors correspond with Case 25 (black), Case 5 (gray), Case 18 (light gray), and Case 2 (red); On Panels (e) to (h), line colors correspond with Case 28 (black), Case 17 (gray), Case 22 (light gray), and Case 2 (red).

262 of wider applicability – and have demonstrated excellent performance. In dissemination from this  
 263 research effort, the investigative team have demonstrated strong performance of the model in a variety  
 264 of contexts.

265 Towards the conclusion of the research effort, the PI and doctoral student turned attention to  
 266 the additional challenge of scalar flux prescription – indeed, in urban micrometeorology, scalar fluxes  
 267 have widespread importance – although unfortunately the project concluded before this work could  
 268 be completed. With this, it is the opinion of the PI that additional work should focus on application  
 269 of an Equation 4-style model, but for the application of scalar fluxes.

## 270 5. Doctoral Student: Travel Award and Princeton Visit

271 The research effort has been directed by PI Anderson, while a doctoral student, Xiaowei Zhu, has  
 272 performed simulations and post-processing in support of the research effort. Xiaowei is reaching  
 273 the end of his third year at UT Dallas, and will soon graduate from the program. He has been  
 274 supported by the ARO throughout his tenure at UT Dallas. In addition to his scholarly activities,

275 this student received a UT Dallas travel award in 2016. This award not only reflects university  
276 support for this ARO project, but it also provided Xiaowei with another travel opportunity (beyond  
277 attendance at conferences). The travel award funds were used to send Xiaowei for a one-week visit  
278 to Princeton University, where he worked with collaborator, Elie Bou-Zied. At Princeton, Xiaowei  
279 gained experience with modeling flows over heated cubes – resembling daytime conditions in the urban  
280 roughness sublayer. This is a new capability for the group, and it addresses a long-term interest on  
281 the need prognostic scalar roughness length models in flows over complex, multiscale environments  
282 [29]. As noted in the preceding section, this research will be the topic of our Year # 3 research effort.

## 283 6. Dissemination

284 The work described above (from Years # 1 and 2 of the research effort) has been presented through  
285 a series of invited seminars, conference talks/proceedings, peer-review publications, and outreach ac-  
286 tivities. The below summarizes dissemination of research during the research effort.

### 287 *Invited Talks:*

- 289 • November, 2018: Center for Astrophysics, Space Physics, and Engineering Research, Baylor  
290 University, Waco, Texas.
- 291 • October, 2018: Department of Mechanical Engineering, University of Houston, Houston, Texas.
- 292 • July, 2018: Invited Speaker in session “Scale-Resolving Turbulence Simulation”, World Congress  
293 on Computational Mechanics, New York, NY.
- 294 • June, 2018: Kavli Institute for Theoretical Physics, University of California at Santa Barbara  
295 – invited speaker in program “Planetary Boundary Layers in Atmospheres, Oceans, and Ice on  
296 Earth and Moons”.
- 297 • May, 2018: Department of Mechanical Engineering, Queensland University of Technology, Bris-  
298 bane, Australia.
- 299 • May, 2018: Department of Mechanical Engineering, Monash University, Melbourne, Australia.
- 300 • May, 2018: Department of Mechanical Engineering, The University of Melbourne, Melbourne,  
301 Australia.
- 302 • May, 2018: Department of Mechanical Engineering, Monash University, Melbourne, Australia.
- 303 • March, 2018: Invited Seminar, Department of Mechanical and Aerospace Engineering, Arizona  
304 State University, Tempe, Arizona.
- 305 • November, 2017: Guest lecture, “Turbulence modeling with large-eddy simulation”, Southern  
306 Methodist University, Dallas, Texas.
- 307 • November, 2017: Invited Speaker at Joint Workshop of Mathematicians, Biogeochemists and  
308 Atmospheric Scientists, Helsinki, Finland.
- 309 • November, 2017: Invited Seminar at Southwest Research Institute, University of Colorado at  
310 Boulder, Denver, Colorado.
- 311 • September, 2017: Invited Speaker at Symposium on Meteorology and Climate – Modeling for  
312 Air Quality, Davis, California.
- 313 • May, 2017: Department of Mechanical Engineering, The University of Utah.
- 314 • June, 2017: Invited Speaker at Engineering Mechanics Institute (EMI) Conference, University  
315 of California at San Diego, San Diego, Calif.
- 316 • May, 2017: School of Environment and Society, Tokyo Institute of Technology, Tokyo, Japan.

- 317 ● April, 2017: *Graduate Seminar Series, Department of Mechanical and Aerospace Engineering,*  
318 *University of Texas at Arlington, Arlington, TX.*
- 319 ● April, 2017: *Graduate school and fluid mechanics. Dallas County Community College District,*  
320 *STEM Foundation, “STEM Summit”, Dallas, TX.*
- 321 ● October, 2016: *Perot Museum, Social Sciences Evening, “Fluid flows in engineering and the*  
322 *environment: math, mystery . . . and a few anecdotes”, Dallas, TX.*
- 323 ● September, 2016: *Departmental Seminar Series, Department of Mechanical Engineering, Uni-*  
324 *versity of Texas at Dallas, Dallas, TX.*
- 325 ● October, 2016: *Departmental Seminar Series, Department of Mechanical Engineering, Univer-*  
326 *sity of Texas at Dallas, Dallas, TX.*

327 **Conference Proceedings:**

328 \*ARO-supported graduate student; †ARO-supported post-doctoral scholar; †Presenting author

- 329 ● Oral presentation: Zhu X<sup>\*,†</sup>, **Anderson W**, 2018: Proceedings of American Physical Society,  
330 Division of Fluid Dynamics, Atlanta, Georgia.
- 331 ● Oral presentation: Zhu X<sup>\*,†</sup>, **Anderson W**, 2018: Proceedings of the American Meteorological  
332 Society, Symposium on Boundary Layers and Turbulence, Oklahoma City, OK.
- 333 ● Oral presentation: Zhu X<sup>\*</sup>, **Anderson W**<sup>†</sup>, 2018: Proceedings of the 8<sup>th</sup> International Symposi-  
334 um on Environmental Hydraulics, South Bend, IN.
- 335 ● Oral presentation: Zhu X<sup>\*,†</sup>, Anderson W, 2017: Proceedings of American Physical Society,  
336 Division of Fluid Dynamics, Denver, Colorado.
- 337 ● Oral presentation: Zhu X<sup>\*,†</sup>, Anderson W, 2017: Proceedings of Bluebonnet Symposium on  
338 Thermal-Fluid Sciences, Dallas, Texas.
- 339 ● Oral presentation: Zhu X<sup>\*,†</sup>, **Anderson W**, 2016: Proceedings of American Physical Society,  
340 Annual Meeting of the Division of Fluid Dynamics, Portland, OR.
- 341 ● Oral presentation: Zhu X<sup>\*,†</sup>, **Anderson W**, 2016: Proceedings of American Meteorological  
342 Society, Symposium on Boundary Layers and Turbulence, Salt Lake City, UT.
- 343 ● Oral presentation: Li Q<sup>†</sup>, Bou-Zeid E, Anderson W, 2016: Proceedings of American Meteorolo-  
344 gical Society, Symposium on Boundary Layers and Turbulence, Salt Lake City, UT.

345 **Outreach:**

- 346 ● April, 2017: *Graduate school and fluid mechanics. Dallas County Community College District,*  
347 *STEM Foundation, “STEM Summit”, Dallas, TX.*
- 348 ● October, 2016: *Perot Museum, Social Sciences Evening, “Fluid flows in engineering and the*  
349 *environment: math, mystery . . . and a few anecdotes”, Dallas, TX.*

350 **Peer-Review Journal:**

351 \*ARO-supported graduate student; †Corresponding author

- 352 ● Li Q, Bou-Zeid E<sup>†</sup>, **Anderson W**, Grimmond S, Hultmark M, 2016: Quality and reliability of  
353 LES of convective scalar transfer at high Reynolds numbers. *International Journal of Heat and*  
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- 355 ● Zhu X<sup>+</sup>, Iungo V, Leonardi S, **Anderson W**<sup>†</sup>, 2016: Large-eddy simulation study of urban-  
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## 7. Mentoring

Grant # W911NF-15-1-0231 supported UT Dallas doctoral candidate, Xiaowei Zhu. Xiaowei joined the PI at UT Dallas in Fall, 2014, and completed three years of doctoral research under the PI’s supervision. He graduated in August, 2018, at the end of Project # 65843-CH. This was the PI’s first graduated doctoral student. This student has presented his research at numerous conferences, and is first author of two peer-review articles (one published in *Boundary-Layer Meteorology*, the second was submitted to the *Journal of Turbulence*, and we are currently revising the article based on reviewer critiques; see also Section 6). Following graduation from UT Dallas, Xiaowei moved to Mechanical Engineering Department at The Johns Hopkins University, where is a post-doctoral scholar.

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