

# Validation of Atmospheric Profile Retrievals from the SNPP NOAA-Unique Combined Atmospheric Processing System. Part 2: Ozone

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**Abstract**—This paper continues an overview of the validation of operational profile retrievals from the Suomi National Polar-Orbiting Partnership (SNPP), with focus here given to the infrared (IR) ozone profile environmental data record (EDR) product. The SNPP IR ozone profile EDR is retrieved using the cross-track IR sounder (CrIS), a Fourier transform spectrometer that measures high-resolution IR earth radiance spectra containing atmospheric state information, namely, vertical profiles of temperature, moisture, and trace gas constituents. The SNPP CrIS serves as the U.S. low earth orbit (LEO) satellite IR sounding system and will be featured on future Joint Polar Satellite System (JPSS) LEO satellites. The operational sounding algorithm is the National Oceanic and Atmospheric Administration-Unique Combined Atmospheric Processing System (NUCAPS), a legacy sounder science team algorithm that retrieves atmospheric profile EDR products, including ozone and carbon trace gases, with optimal vertical resolution under non-precipitating (clear to partly cloudy) conditions. The NUCAPS ozone profile product is assessed in this paper using extensive global *in situ* truth data sets, namely, ozonesonde observations launched from ground-based networks and from ocean-based intensive field campaigns, along with numerical weather predic-

tion model output. Based upon rigorous statistical analyses using these data sets, the NUCAPS ozone profile EDRs are determined to meet the JPSS Level 1 global performance requirements.

**Index Terms**—Algorithms, atmospheric measurements, geophysical measurements, infrared (IR) measurements, measurement errors, ozone, radiosondes, remote sensing, satellite applications.

## I. INTRODUCTION

THE operational U.S. Suomi National Polar-Orbiting Partnership (SNPP) satellite features the hyperspectral infrared (IR) cross-track IR sounder (CrIS) and advanced technology microwave sounder (ATMS) sounding system. The follow-on Joint Polar Satellite System (JPSS) is a U.S. National Oceanic and Atmospheric Administration (NOAA) operational satellite mission will feature CrIS/ATMS sounders onboard four satellites launched in the same orbit over the next two decades beginning in late 2017. The CrIS instrument is an advanced Fourier transform spectrometer that measures well-calibrated sensor data records (SDRs) consisting of high-resolution IR spectra in 1305 channels over three bands spanning  $\nu \approx [650, 2550] \text{ cm}^{-1}$ . The CrIS spectra allow for retrieval of atmospheric vertical profile environmental data records (EDRs) with the highest possible vertical resolution ( $\approx 2\text{--}5 \text{ km}$ ) comparable with predecessor sounding systems, namely, the *MetOp-A* and *-B* IR atmospheric sounding interferometer [1], [2] and the EOS-Aqua atmospheric IR sounder (AIRS) [3], [4].

Although sounder SDRs (radiance) have come to be directly assimilated into global numerical weather prediction (NWP) models via variational analysis schemes, they continue to be directly inverted operationally to retrieve orbital atmospheric profile EDRs in near real time, as originally envisioned by satellite sounding pioneers [5]–[10]. One advantage of direct inversion is the ready capability of inverting for numerous state parameters beyond atmospheric vertical temperature profile (AVTP) and atmospheric vertical moisture profile (AVMP), namely, trace gases, clouds, aerosols, and surface emissivity, among others.

The operational EDR retrieval algorithm for CrIS/ATMS is the NOAA-Unique Combined Atmospheric Processing System (NUCAPS) [11], [12]. The NUCAPS algorithm is based upon the heritage methodology developed for the EOS-Aqua

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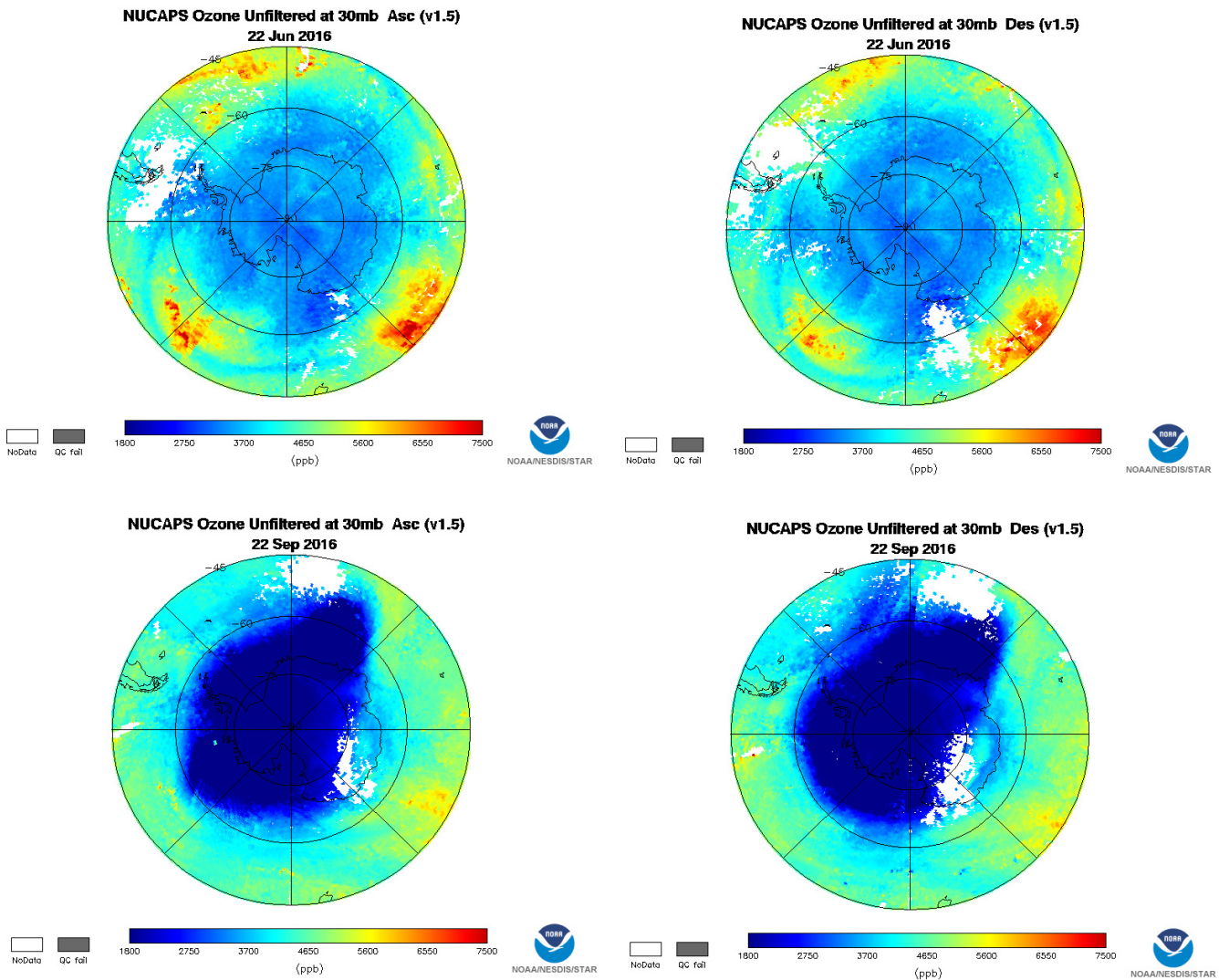


Fig. 1. NUCAPS retrieved 30-hPa layer ozone for ascending and descending orbits on (Top) June 22, 2016 and (Bottom) September 22, 2016 illustrating the seasonal depletion of ozone from SH winter to SH spring.

AIRS and is a modular implementation of the multistep AIRS Science Team retrieval algorithm version 5 [13], [14]. For more details on the NUCAPS algorithm, the reader may refer to [12], [13], or the algorithm theoretical basis document available online. The multistep NUCAPS physical retrieval module retrieves individual parameters in a step-by-step fashion, using only channels rigorously determined to be sensitive to that parameter [15], beginning with temperature and water vapor profiles, followed by ozone ( $O_3$ ) and trace gases, with the result output on the radiative transfer algorithm (RTA) 100 layers (AVTP is output on layer boundaries). The NUCAPS IR ozone profile EDR is currently used by the NOAA total ozone analysis using SBUV/2 and TOVS (TOAST), as well as in basic science applications.

Because of the multistep retrieval method, the quality of the ozone profile retrieval (and the other trace gases) will depend to some extent on the quality of the AVTP and AVMP retrievals. Thus, the performances of the temperature and moisture EDRs were first overviewed in [16], where it was demonstrated that the operational SNPP NUCAPS AVTP and AVMP EDRs comply with JPSS Level 1 requirements

(and declared validated as of September 2014). In this paper, the profile EDR validation is extended to the SNPP NUCAPS IR ozone profile EDR using ozonesonde collocations from land-based networks and ocean-based dedicated launches, along with numerical model comparisons.

## II. NUCAPS IR OZONE PROFILE EDR OVERVIEW

As mentioned above, users of the NUCAPS IR ozone profile EDR include the TOAST, in addition to science users interested in atmospheric chemistry and air quality [17], [18]. Satellite sounder EDR data sets are generally invaluable for numerous global environmental research studies [19]. To illustrate, Fig. 1 shows NUCAPS ozone retrievals for the 30-hPa RTA layer for June 22 and September 22, 2016, these being roughly the southern hemisphere (SH) winter solstice and spring equinox, respectively. As will be seen in Section III, the CrIS sensor has very good sensitivity to this layer, and as a result, the seasonal depletion of ozone from SH winter to spring, commonly referred to as the Antarctic “ozone hole” [20], is clearly observed by the NUCAPS ozone soundings.

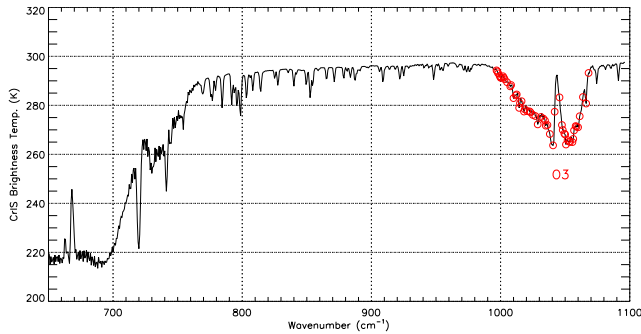


Fig. 2. Hamming apodized CrIS longwave IR brightness temperature spectrum (unapodized nominal spectral resolution  $0.625 \text{ cm}^{-1}$ ) for a marine nighttime case (10:22 UTC 9 June 2015,  $6.5^\circ\text{N}$ ,  $130.0^\circ\text{W}$ ) showing ozone channels (red circles) used in the NUCAPS multistep physical retrieval.

As also mentioned above, the NUCAPS physical retrieval algorithm utilizes information contained within the CrIS-measured IR earth radiance spectra to retrieve ozone. The NUCAPS ozone retrieval step applies an optimal estimation (OE) method to retrieve ozone using sensitive channels [15] [see Fig. 2] and an *a priori* background state consisting of a tropopause-based climatology [21].

Retrieval sensitivity to state parameters (e.g., ozone concentration) can be inferred from the averaging kernels (AKs) defined by [22]–[24]

$$\mathbf{A} \equiv \frac{\partial \hat{\mathbf{x}}}{\partial \mathbf{x}} \quad (1)$$

where the AK matrix  $\mathbf{A}$  is dimensioned  $m \times m$  ( $m$  being the number of RTA layers), and  $\hat{\mathbf{x}}$  and  $\mathbf{x}$  denote the retrieved and true states, respectively. The NUCAPS algorithm computes “effective” AKs,  $\mathbf{A}_e$ , for each retrieval that account for the trapezoidal basis functions used in the physical retrieval, the details of which can be found in [25]. Fig. 3 shows zonal-mean NUCAPS profiles taken from a global Focus Day, February 17, 2015, for the tropics, northern hemisphere (NH) and SH midlatitude, and polar zones. The left plot shows the RTA layer-averaged mean effective AKs for the ozone channels shown in Fig. 2, where it can be seen that the layer and magnitude of peak sensitivity increase from the poles to the tropics. Polar sensitivity peaks at around 100 hPa, whereas midlatitude and tropical sensitivity peak higher in the upper troposphere to lower stratosphere (UT/LS),  $\approx 50$  hPa, with a sharper peak exhibited in the tropics along with a secondary peak below the tropopause (middle plot) at around 300 hPa, which when combined with the primary peak shows UT/LS sensitivity of the NUCAPS ozone product over the tropics [21]. The greater sensitivity seen in the NH polar cap ( $60^\circ\text{--}90^\circ\text{N}$ ) versus the SH is related to the relatively higher ambient LS ozone concentration found in the NH over the SH (right plot) during late boreal winter. The ability of the CrIS to provide information about the ozone profile is also demonstrated by considering the NUCAPS algorithm degrees of freedom (DoF) for the ozone retrieval, which are shown for the February 17, 2015 Focus Day in Fig. 4. Generally speaking, DoF greater than unity is an indicator that more than one independent piece of information is contained within the measurements, thus enabling the retrieval to contribute

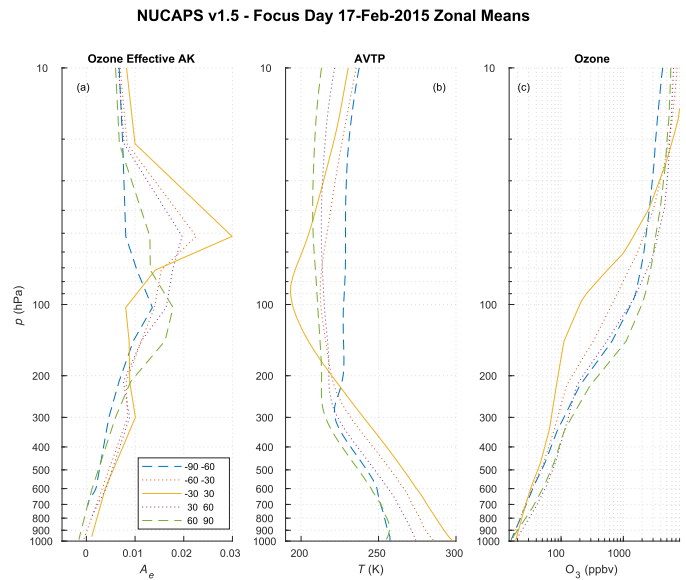


Fig. 3. Zonal-mean NUCAPS profiles calculated from a global Focus Day, February 17, 2015 ( $n = 2686$  granules). (a) RTA layer-averaged effective AKs  $A_e$  for nominal spectral-resolution CrIS ozone channels shown in Fig. 2. (b) Atmospheric vertical temperature profile retrievals. (c) IR ozone profile retrievals (log-log plot). The solid lines are tropics ( $30^\circ\text{S}$  to  $30^\circ\text{N}$ ), dotted lines are midlatitudes ( $30^\circ\text{--}60^\circ\text{S}$  and  $^\circ\text{N}$ ), and dashed lines are polar ( $60^\circ\text{--}90^\circ\text{S}$  and  $^\circ\text{N}$ ).

vertical profile information to the *a priori*. In Fig. 4, it can be seen that NUCAPS ozone DoF are generally  $\gtrsim 1$  globally speaking, with larger values  $\gtrsim 2$  found in midlatitude to polar zones, and smaller values  $\approx 1$  in regions of the tropics (possibly associated with deep convective clouds within the intertropical convergence zone) as well as over central Antarctica.

### III. IR OZONE PROFILE EDR PERFORMANCE ASSESSMENT

The JPSS Level 1 requirements for the CrIS IR ozone profile EDR are given in Table I, which are defined for global nonprecipitating cases on broad atmospheric layers made up of coarse layers. In the case of ozone, there is only one tropospheric layer (a consequence of the CrIS ozone sensitivity as evidenced in the AKs) and six spanning from the upper troposphere to the stratosphere that are to be computed as the average of coarse statistical layers. As described in [26], to avoid undesirable skewing of the sample distribution, we weight each deviation by the ozone layer mass abundance squared (i.e.,  $W^2$  weighting) in the computation of coarse-layer root-mean-square error (RMSE), bias (mean), and standard deviation ( $\sigma$ ).

#### A. CrIS Nominal Spectral Resolution

The operational NUCAPS algorithm (version 1.5) has run on nominal spectral-resolution (NSR) CrIS SDRs at  $\Delta\nu \approx 0.625, 1.25,$  and  $2.5 \text{ cm}^{-1}$  for the longwave, midwave, and shortwave IR bands, respectively [27], [28]. This section presents the validation of the operational ozone profile EDR based upon an offline v1.5 emulation.

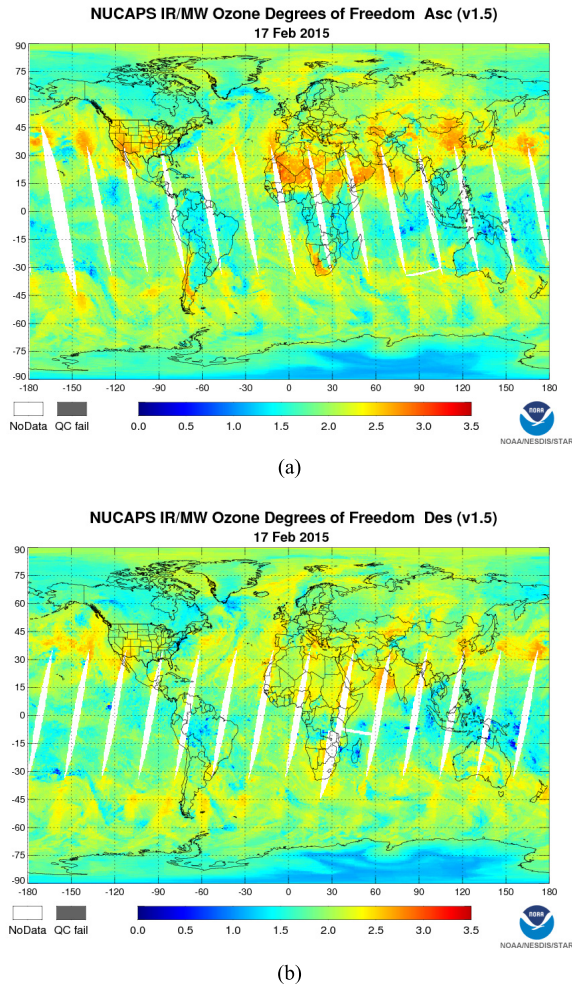


Fig. 4. NUCAPS ozone DoF for the global Focus Day, February 17, 2015. (a) Ascending orbit. (b) Descending orbit.

1) *Data*: Validation of the operational ozone profile EDR is primarily based upon collocations of truth data sets consisting of *in situ* ozone soundings obtained from electrochemical concentration cell (ECC) ozonesondes along with global output from the European Centre for Medium-Range Weather Forecasts (ECMWF) NWP model. Ozonesondes used in the analyses were acquired from land-based World Ozone and Ultraviolet Radiation Data Center (WOUDC) and Southern Hemisphere Additional Ozonesonde (SHADOZ) [29] network sites, along with unique SNPP-dedicated ECC ozonesondes launched during ship-based intensive cal/val campaigns [16], namely, NOAA Aerosols and Ocean Science Expeditions (AEROSE) [17], [30] and the 2015 CalWater ARM Cloud Aerosol Precipitation Experiment (ACAPEX) [31]–[33]. We have accumulated ozonesonde truth data sets collocated with SNPP CrIS spanning the period of early 2012–2015; the locations of these sites are shown in Fig. 5.

ECC ozonesondes typically measure ozone partial pressure in mPa with high vertical resolution (e.g., 1 s). These must be converted to fine layer abundances (molecules/cm<sup>2</sup>) and then reduced to 100 RTA layer abundances to yield correlative data for the NUCAPS ozone retrieval [26]. Ozonesonde partial pressures are first converted to number densities  $N_x$

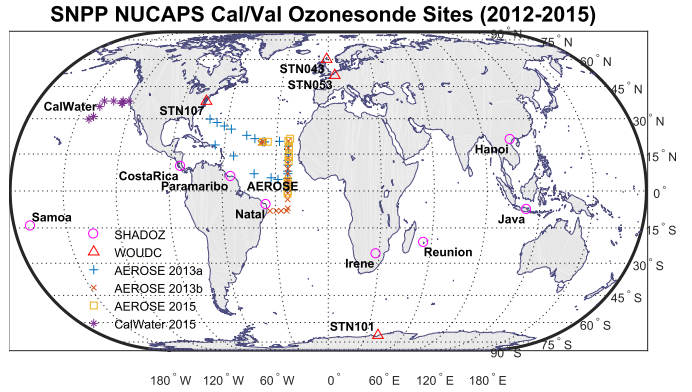


Fig. 5. Ozonesonde truth sites used for SNPP NUCAPS IR ozone profile EDR cal/val over the sampling period 2012–2015. Magenta circles denote SHADOZ sites, red triangles denote WOUDC sites, and blue +, red ×, gold □, and purple \* denote SNPP-dedicated ozonesondes launched from ship-based intensive campaigns (AEROSE and CalWater/ACAPEX). Map projection is equal area.

TABLE I  
JPSS LEVEL 1 REQUIREMENTS\* FOR IR OZONE PROFILE EDR

IR Ozone Profile (CrIS) Layer Average Proportional Error		
Atmospheric Broad-Layer	Threshold	Objective
<i>Precision (random error, <math>\sigma</math>)</i>		
Surface to 260 hPa (6 statistic layers)	20%	10%
260 hPa to 4 hPa (1 statistic layer)	20%	10%
<i>Accuracy (systematic error, bias)</i>		
Surface to 260 hPa (6 statistic layers)	±10%	±5%
260 hPa to 4 hPa (1 statistic layer)	±10%	±5%
<i>Combined Uncertainty (RMSE)</i>		
Surface to 260 hPa (6 statistic layers)	25%	15%
260 hPa to 4 hPa (1 statistic layer)	25%	15%

\*Source: Joint Polar Satellite System (JPSS) Program Level 1 Requirements Supplement — Final, Version 2.9, 27 June 2013, NOAA/NESDIS, p. 49.

(molecules/cm<sup>3</sup>) using the formula (in centimeter-gram-second units)

$$N_x(p_{x,\ell}, T_\ell) = 10^{-2} \left( \frac{p_{x,\ell}}{kT_\ell} \right) \quad (2)$$

where  $p_{x,\ell}$  is the partial pressure (in mPa) for constituent  $x \equiv \text{O}_3$  at ozonesonde level  $\ell$ ,  $T_\ell$  is the radiosonde temperature at level  $\ell$ ,  $k$  is the Boltzmann constant (ergs), and the factor  $10^{-2}$  converts partial pressure from mPa to dPa. Equation (2) is then integrated from the balloon burst level down and interpolated to RTA layer boundaries (i.e., “levels”) to enable calculation of RTA layer abundances [26].

Although the NUCAPS effective AKs (discussed in Section II) can be applied to “smooth” the correlative truth data and remove null-space source error implicit to the retrieval algorithm (thus yielding improved statistics), the primary focus of this paper is to evaluate the product’s performance

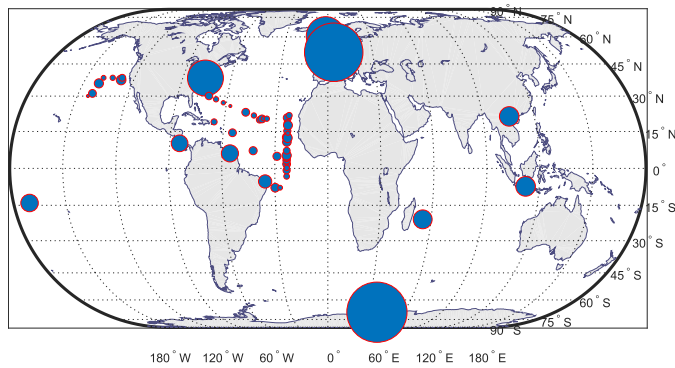


Fig. 6. Geographic histogram of SNPP NUCAPS FOR-ozonesonde collocation data used in the global land/sea statistical error analysis. Circle sizes depict the relative SNPP-ozonesonde collocation sample sizes for each ozonesonde launch location. Map projection is equal area.

against the metrics defined by the JPSS Level 1 requirements summarized in Table I. The JPSS requirements are applicable to the total system error, which includes the null-space error, thus precluding the use of AKs in demonstrating the product meets requirements. Thus, a more detailed breakdown of algorithm error sources, including null-space error using AKs, falls outside the scope of the current effort and will be the subject of future work (e.g., the JPSS-1 NUCAPS validation effort).

2) *Error Analysis*: As in the collocation methodology described in [16], we have imposed space–time collocation criteria to the NUCAPS-ozonesonde collocation data set, striking a balance between collocation mismatch uncertainty and sample size. In this case, FORs are included within  $\delta x \leq 125$ -km radius and  $-240 < \delta t < +120$  min of launches (note that the selected ozonesonde sites, including the dedicated ozonesonde launches, favored ozonesondes being launched prior to overpasses). Fig. 6 shows the corresponding geographic histogram of the distribution of the ozonesonde collocation sample on an equal-area map projection, where it can be seen that the combination of the ozonesonde sites described above provide adequate representation of global climate zones (tropics, midlatitudes, and polar) along with land and ocean surfaces.

The resulting global profile error statistics are given in Fig. 7, along with those separated by polar, midlatitude, and tropical zones given in Figs. 8–10, respectively. In Figs. 7–10, blue lines show the results of the NUCAPS retrievals (IR accepted cases, clear to partly cloudy) and magenta lines show the results of the *a priori* (climatological background) used in the physical retrieval. The left and right plots show the coarse-layer RMSE and bias  $\pm 1\sigma$  statistics, respectively. The JPSS Level 1 global specification requirements (Table I) for RMSE and bias are shown with dashed gray lines in the plots. The corresponding broad-layer averages for these statistics are depicted with asterisks in the plots, with the global results summarized in Table II. It should be noted that although we have included the JPSS global requirement lines and broad-layer averages in the zonal plots (Figs. 8–10) for reference, JPSS requirements are specified

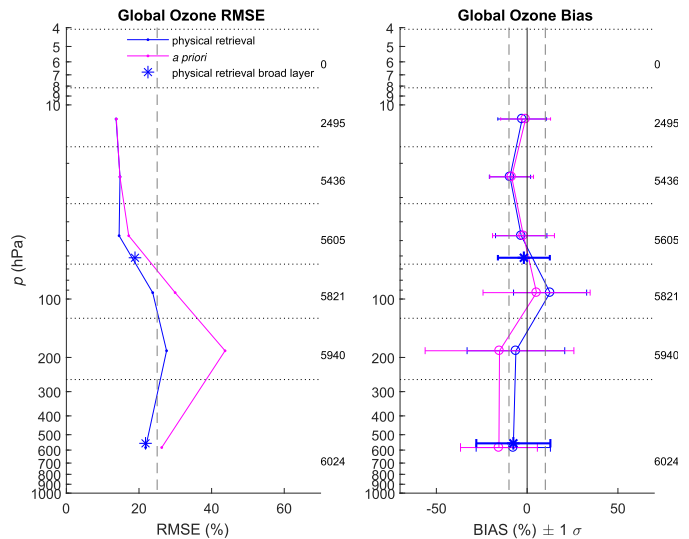


Fig. 7. Coarse-layer statistical assessment of the NUCAPS IR ozone profile EDR (offline v1.5 operational emulation, blue lines) versus collocated ozonesondes for retrievals accepted by the quality flag within space–time collocation criteria of  $\delta x \leq 125$ -km radius and  $-240 \leq \delta t \leq +120$  min of launches over a sampling period of April 4, 2012 to December 12, 2015. (Left) RMSE results. (Right) Bias  $\pm 1\sigma$  results. NUCAPS IR physical retrieval (under clear to partly cloudy conditions) and *a priori* (climatological background) performances are given as blue and magenta lines, respectively, with collocation sample size for each coarse-layer given in the right margins. The gray dashed lines designate the JPSS Level 1 global performance requirements for two broad atmospheric layers defined in Table I, with asterisks denoting the calculated broad-layer averages for the physical retrievals.

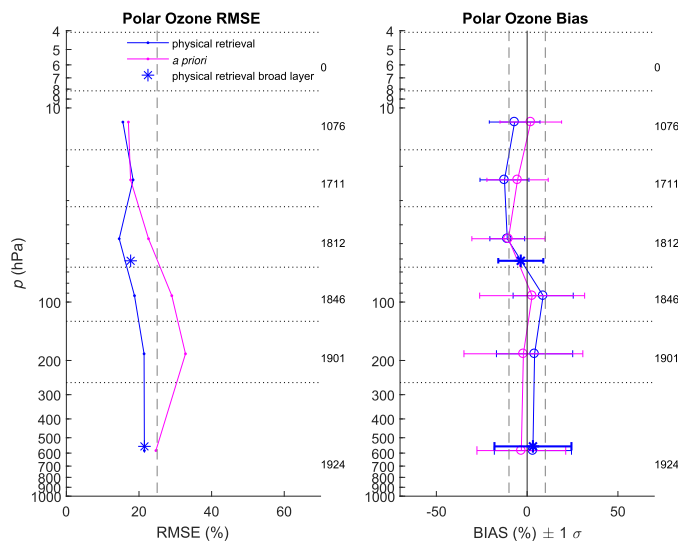


Fig. 8. As Fig. 7 except for NUCAPS retrievals collocated with ozonesondes in the NH and SH polar caps.

for global ensembles only; thus, they do not have any direct bearing on results obtained for any type of subsample binning (e.g., latitude zones).

A scatterplot of NUCAPS versus ozonesonde layer-averaged  $O_3$  molecular abundances for the two broad atmospheric layers is shown in Fig. 11. The majority of the data fall along the one-to-one line with the exception of a region between the two layers, where a small number of NUCAPS retrievals in the 260–4 hPa layer (red + symbols) are seen to significantly overestimate the ozone concentration relative to the ozonesondes. The region in question roughly corresponds

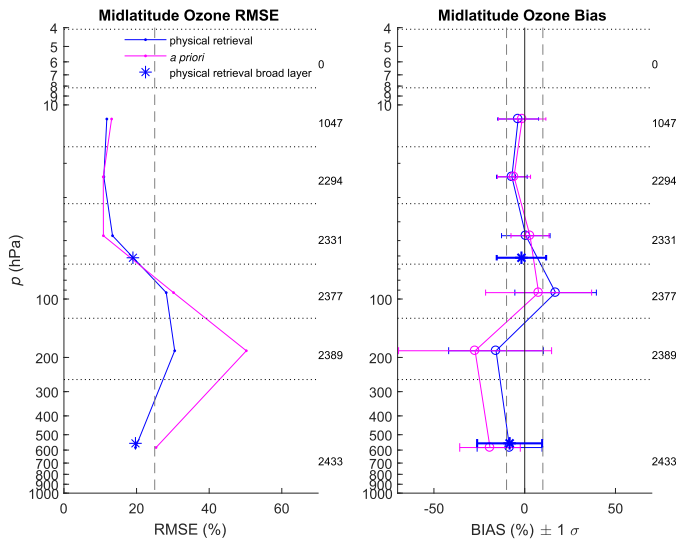


Fig. 9. As Fig. 7 except for NUCAPS retrievals collocated with ozonesondes in the midlatitude zones.

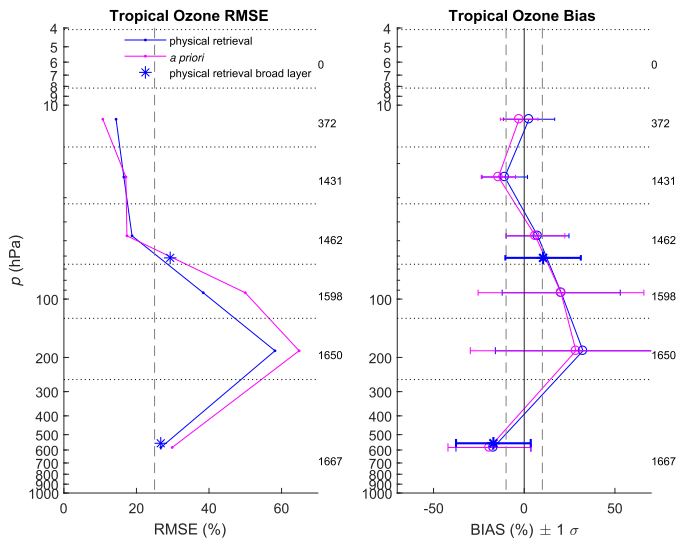


Fig. 10. As Fig. 7 except for NUCAPS retrievals collocated with ozonesondes in the tropical zone.

TABLE II  
VALIDATED GLOBAL IR OZONE PROFILE EDR  
MEASUREMENT UNCERTAINTY

Atmospheric Broad-Layer	Observed Uncertainty				
	RMSE	bias	$\sigma$	$r$	$p$
Surface to 260 hPa	23.2%	-9.4%	21.2%	0.69	0
260 hPa to 4 hPa	18.9%	-1.8%	14.3%	0.77	0

to the tropopause region, where two potential sources of error would include *a priori* and null-space errors. Null-space errors result from the limitations in the CrIS instrument’s vertical resolution and sensitivity (e.g., Fig. 3); this issue will be explored using the NUCAPS effective AKs in a future paper. The correlation coefficients,  $r$ , along with corresponding  $p$ -values, are included in Table II, where it is seen that the broad-layer correlations between NUCAPS and ozonesondes are  $\approx +0.7$ .

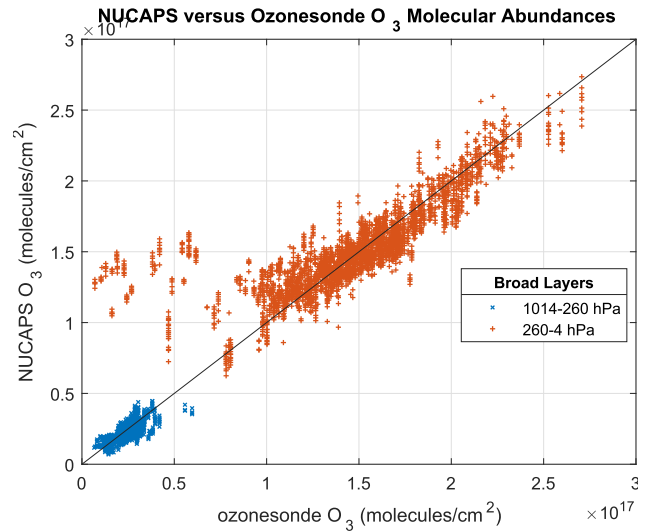


Fig. 11. Scatterplot of NUCAPS versus ozonesonde layer-averaged  $O_3$  molecular abundances ( $\text{molecules}/\text{cm}^2$ ) for the two broad atmospheric layers defined in this paper: 1014–260 hPa (blue  $\times$ ) and 260–4 hPa (red  $+$ ). Correlation coefficients  $r$  and  $p$ -values are given in Table II.

In discussing further the results presented in Figs. 7–10, it is first recalled that the NUCAPS ozone physical retrieval step uses an OE method that relies on a formal *a priori* derived based upon a climatological background state [21]. Figs. 7–10 demonstrate the ability of the retrieval (blue lines) to move the *a priori* state (magenta lines) toward the ozonesonde-observed state as evidenced by the significantly reduced  $\sigma$  and RMSE for layers where the CrIS channels have sensitivity [Fig. 3(a)]. Because the *a priori* (magenta) is based upon climatology, it is not surprising that it exhibits very little global bias, making further improvement from the retrieval difficult to achieve (Fig. 7, right plot). Thus, the NUCAPS OE ozone retrieval uses the observed IR spectral information to measure deviations from the *a priori* (i.e., mean) state, resulting in the reduction of random errors ( $\sigma$  and RMSE), but not necessarily the systematic error.

We find that the global ozone profile EDR meets the JPSS requirements, with the only exception being the precision ( $\sigma$ ) for the tropospheric broad layer (surface to 260 hPa), which falls somewhat outside of the 20% requirement for this layer. However, referring back to the AKs shown in Fig. 3, it is noted that the CrIS instrument possesses little sensitivity in the troposphere, thereby requiring the algorithm to relax to the *a priori*. The overall results for SNPP NUCAPS presented here are comparable with those reported previously for the *Aqua* AIRS version 5 ozone product [34]. Therefore, based on our findings (Fig. 7 and Table II), it is concluded that the NUCAPS ozone profile EDR generally meets the JPSS Level 1 requirements.

Similar performance patterns (both RMSE and bias) are observed in the three climate zones, with overall profile performances improving with latitude zone from the tropics to the poles. The diminished performance in the tropics (Fig. 10) is associated with what may potentially be a suboptimal *a priori* (magenta lines) combined with reduced ozone DoF (Fig. 4) and ozone AK sensitivity at higher altitudes (Fig. 3). The physical retrieval significantly improves the *a priori* in

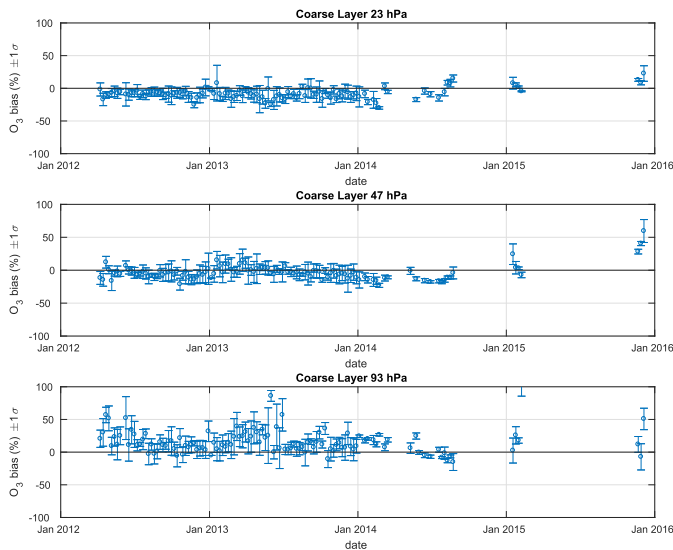


Fig. 12. Weekly statistical time series (bias  $\pm 1\sigma$ ) for NUCAPS v1.5 IR ozone profile EDR retrievals versus the collocated ozonesondes (Fig. 6) acquired over the sampling period of 2012–2015 for three UT/LS coarse layers. (Top)  $\approx 23$  hPa. (Middle)  $\approx 47$  hPa. (Bottom)  $\approx 93$  hPa.

UT/LS in both the polar and midlatitude zones (Figs. 8 and 9, respectively), whereas the improvement is reduced, but nevertheless still evident, for the tropical cases (Fig. 10). Global seasonal stability in the retrievals for three UT/LS coarse layers (23, 47, and 93 hPa) over the ozonesonde acquisition period is demonstrated in Fig. 12. Weekly biases generally fall within  $-20\%$  to  $0\%$  for the 23-hPa layer,  $\pm 20\%$  for the 47-hPa layer, and  $-10\%$  to  $+40\%$  at 93 hPa, with very little seasonal variability or long-term trends. Note that two short acquisition periods at the beginning and ending of 2015 correspond to dedicated ozonesondes acquired over ocean during the 2015 CalWater/ACAPEX and AEROSE campaigns, the former obtained under inclement weather conditions in the Pacific [32], [33] and the latter obtained over the tropical Atlantic (see Fig. 5).

### B. CrIS Full Spectral Resolution

As discussed in [16], the operational SNPP NUCAPS v1.5 has previously run on CrIS spectra with reduced resolution in the midwave and shortwave bands due to truncated interferograms in those bands during operational processing; these reduced-resolution spectra have been referred to as “nominal resolution” as this was the original (nominal) resolution of the operational SDRs. However, offline production of SNPP full spectral resolution (FSR) CrIS SDRs ( $\Delta\nu \approx 0.625 \text{ cm}^{-1}$  in all three bands) began in December 2014 [35], with operational Interface Data Processing Segment (IDPS) production starting in March 2017. Given that CrIS FSR SDRs will be produced operationally going forward (i.e., for the remainder of the SNPP lifetime as well as the follow-on JPSS satellite series, with the JPSS-1 launch tentatively scheduled for November 2017), a preliminary experimental offline NUCAPS version was developed to run on CrIS FSR data for demonstration studies [36]. A completed version (v2.0.5), representing the operational delivery of the

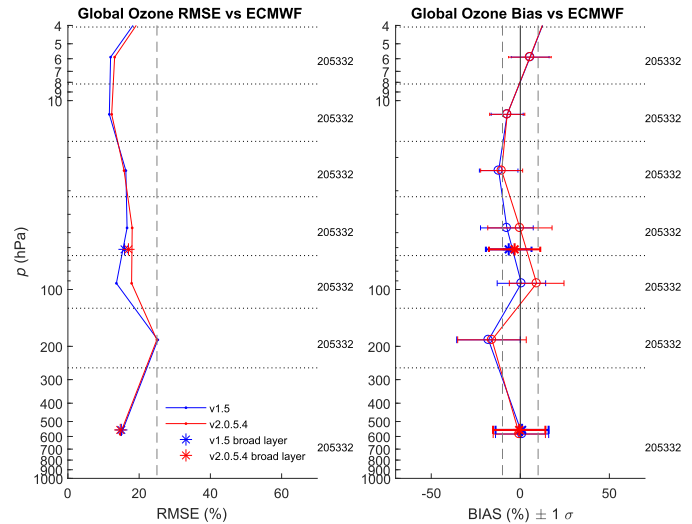


Fig. 13. As Fig. 7 except statistical assessment of offline NUCAPS v2.0.5 (CrIS FSR, red lines) and v1.5 (CrIS nominal resolution, blue lines) IR physical retrievals versus collocated ECMWF model output (analysis or forecast nearest in time, red lines) for retrievals accepted by the quality flag for a global Focus Day, February 17, 2015. Global yields for v2.0.5 and v1.5 accepted cases are 83.4% and 63.5%, respectively, indicating a marked improvement in the v2.0.5 quality acceptance yield.

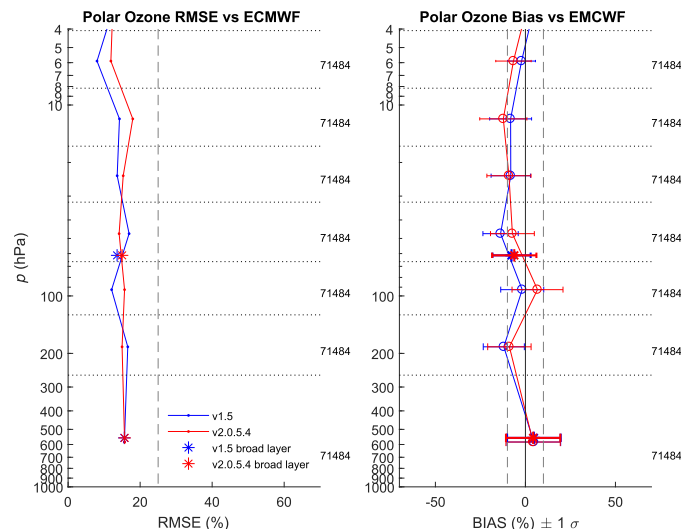


Fig. 14. As Fig. 13 except for NUCAPS retrievals collocated with ECMWF within the NH and SH polar caps.

NUCAPS system in FSR mode, was demonstrated and delivered for operational implementation in July to August 2017.

Because CrIS FSR SDRs were not operationally available during the ozonesonde acquisition period, a preliminary assessment of the NUCAPS FSR algorithm has been performed versus numerical forecast model output (viz., ECMWF) for a global Focus Day (February 17, 2015) [26] where the CrIS FSR SDRs were made available offline. As in Section III-A2, Fig. 13 shows the global results, Figs. 14–16 show the breakdowns by latitude zones. In Figs. 13–16, the red lines show the FSR v2.0.5 NUCAPS results with blue lines showing the v1.5 NSR results for comparison. The patterns are similar (but not identical) to those obtained when using ozonesondes as the baseline (see Figs. 7–10), with improved performance occurring with latitude zone

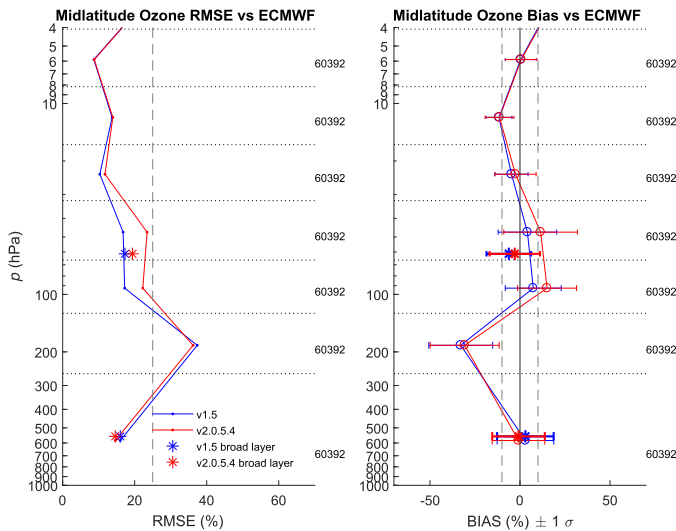


Fig. 15. As Fig. 13 except for NUCAPS retrievals collocated with ECMWF within the midlatitude zones.

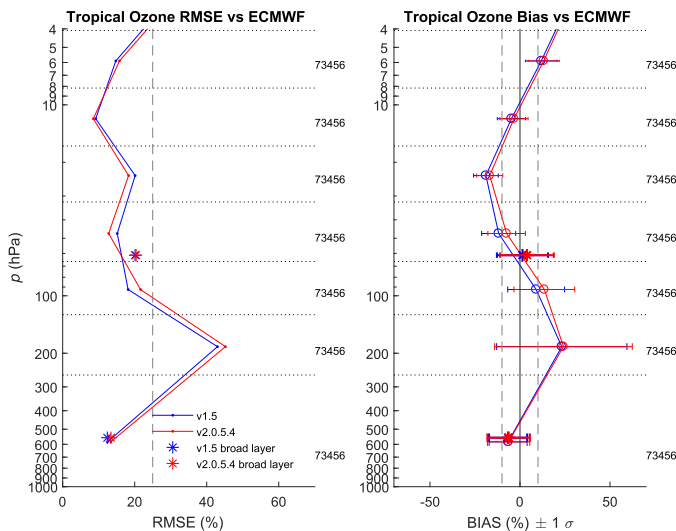


Fig. 16. As Fig. 13 except for NUCAPS retrievals collocated with ECMWF within the tropical zone.

from tropical (Fig. 16) to midlatitude (Fig. 15) to polar zones (Fig. 14). Of particular note, the NUCAPS v2.0.5 FSR algorithm demonstrates a significant improvement over the v1.5 NSR algorithm in the IR + MW retrieval quality acceptance yield, from 63.5% to 83.4%, while demonstrating comparable performance. Rejected cases typically occur under environmental conditions that present challenges to passive IR retrievals but are otherwise of meteorological interest (e.g., cloudiness associated with convection). In spite of this, it is seen that the NUCAPS FSR (v2.0.5) algorithm otherwise performs comparably with the fully validated NUCAPS NSR (v1.5), with the broad-layer averages (denoted with asterisks) generally meeting the JPSS Level 1 requirements relative to ECMWF.

#### IV. CONCLUSION

This paper has presented the formal validation of the SNPP NUCAPS IR ozone profile EDR in continuation of the validation of atmospheric vertical temperature and moisture profile

EDRs described in [16]. Based upon a globally representative sample of collocated ozonesondes and ECMWF model output, it has been shown that the NUCAPS v1.5 IR ozone profile EDR (CrIS-FSR) meets JPSS Level 1 broad-layer global performance requirements (Tables I and II) and has thus attained validated maturity. It is noted that the ozonesonde sites used in this analysis (Fig. 5) include those from all global climate zones (tropical, midlatitude, and polar), as well as unique marine-based data sets obtained from ship over both the Pacific and Atlantic Oceans (i.e., AEROSE and CalWater/CAPEX campaigns). The NUCAPS OE physical retrieval was shown to improve upon the climatological *a priori* in UT/LS layers (Figs. 7 and 13) where CrIS has sensitivity (Fig. 3). Results vary somewhat depending on latitude zone (tropical, midlatitude, and polar), with a general improvement seen at higher latitudes as would be expected given the variation in ozone DoF (Fig. 4) and in vertical sensitivity (Fig. 3). The algorithm has been successfully implemented for SNPP CrIS-FSR SDRs (v2.0.5), these being produced for future JPSS satellites and operationally from SNPP since March 2017, with increased yield and comparable performance versus the validated NUCAPS v1.5 algorithm (Fig. 13). Full validation of the JPSS-1 NUCAPS-FSR algorithm (including future upgrades) versus global ensembles of collocated ozonesondes (including dedicated ozonesondes) will be the subject of future work.

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#### REFERENCES

- [1] F. R. Cayla, "IASI infrared interferometer for operations and research," in *High Spectral Resolution Infrared Remote Sensing for Earth's Weather and Climate Studies* (NATO ASI Series), vol. 19, A. Chedin, M. T. Chahine, and N. A. Scott, Eds. Berlin Germany: Springer-Verlag, 1993, pp. 9–19.
- [2] F. Hilton *et al.*, "Hyperspectral earth observation from IASI: Five years of accomplishments," *Bull. Amer. Meteorol. Soc.*, vol. 93, no. 3, pp. 347–370, 2012.
- [3] H. H. Aumann *et al.*, "AIRS/AMSU/HSB on the Aqua mission: Design, science objectives, data products, and processing systems," *IEEE Trans. Geosci. Remote Sens.*, vol. 41, no. 2, pp. 253–264, Feb. 2003.

- [4] M. T. Chahine *et al.*, "AIRS: Improving weather forecasting and providing new data on greenhouse gases," *Bull. Amer. Meteorol. Soc.*, vol. 87, no. 7, pp. 911–926, Jul. 2006.
- [5] W. L. Smith, "An iterative method for deducing tropospheric temperature and moisture profiles from satellite radiation measurements," *Monthly Weather Rev.*, vol. 95, no. 6, pp. 363–369, 1967.
- [6] W. L. Smith, "An improved method for calculating tropospheric temperature and moisture from satellite radiometer measurements," *Monthly Weather Rev.*, vol. 96, no. 6, pp. 387–396, 1968.
- [7] W. L. Smith, "Iterative solution of the radiative transfer equation for the temperature and absorbing gas profile of an atmosphere," *Appl. Opt.*, vol. 9, pp. 1993–1999, Sep. 1970.
- [8] M. T. Chahine, "Determination of the temperature profile in an atmosphere from its outgoing radiance," *J. Opt. Soc. Amer.*, vol. 58, no. 12, pp. 1634–1637, 1968.
- [9] M. I. Chahine, "Inverse problems in radiative transfer: Determination of atmospheric parameters," *J. Atmos. Sci.*, vol. 27, no. 6, pp. 960–967, 1970.
- [10] M. I. Chahine, "A general relaxation method for inverse solution of the full radiative transfer equation," *J. Atmos. Sci.*, vol. 29, no. 4, pp. 741–747, 1972.
- [11] A. Gambacorta *et al.*, "The NOAA unique CrIS/ATMS processing system (NUCAPS): First light retrieval results," in *Proc. 18th ITSC*, Toulouse, France, 2012.
- [12] A. Gambacorta, C. D. Barnet, and M. Goldberg, "Status of the NOAA unique CrIS ATMS processing system (NUCAPS): Algorithm development and lessons learned from recent field campaigns," in *Proc. 20th ITSC*, Lake Geneva, WI, USA, 2015.
- [13] J. Susskind, C. D. Barnet, and J. M. Blaisdell, "Retrieval of atmospheric and surface parameters from AIRS/AMSU/HSB data in the presence of clouds," *IEEE Trans. Geosci. Remote Sens.*, vol. 41, no. 2, pp. 390–409, Feb. 2003.
- [14] J. Susskind, J. M. Blaisdell, L. Iredell, and F. Keita, "Improved temperature sounding and quality control methodology using AIRS/AMSU data: The AIRS science team version 5 retrieval algorithm," *IEEE Trans. Geosci. Remote Sens.*, vol. 49, no. 3, pp. 883–907, Mar. 2011.
- [15] A. Gambacorta and C. D. Barnet, "Methodology and information content of the NOAA NESDIS operational channel selection for the cross-track infrared sounder (CrIS)," *IEEE Trans. Geosci. Remote Sens.*, vol. 51, no. 6, pp. 3207–3216, Jun. 2013.
- [16] N. R. Nalli *et al.*, "Validation of atmospheric profile retrievals from the SNPP NOAA-unique combined atmospheric processing system. Part 1: Temperature and moisture," *IEEE Trans. Geosci. Remote Sens.*, vol. 55, no. 12, Dec. 2017, doi: 10.1109/TGRS.2017.2744558.
- [17] N. R. Nalli *et al.*, "Multiyear observations of the tropical atlantic atmosphere: Multidisciplinary applications of the NOAA Aerosols and Ocean Science Expeditions," *Bull. Amer. Meteorol. Soc.*, vol. 92, pp. 765–789, Jun. 2011.
- [18] J. W. Smith and N. R. Nalli, "Ozone sonde and NOAA-unique sounder measurements of AEROSE 2010 ozone via lightning-induced nitrogen oxides and biomass burning emission over the Tropical Atlantic Ocean," in *Proc. AMS Annu. Meet.*, Atlanta, GA, USA, Feb. 2014.
- [19] T. Pagano, *AIRS Project Status* (Sounder Science Team Meeting). Pasadena, CA, USA: NASA/JPL, May 2013. [Online]. Available: [http://airs.jpl.nasa.gov/documents/science\\_team\\_meeting\\_archive/2013\\_05/slides/Pagano\\_AIRS\\_Proj\\_Status\\_May\\_2013.pdf](http://airs.jpl.nasa.gov/documents/science_team_meeting_archive/2013_05/slides/Pagano_AIRS_Proj_Status_May_2013.pdf)
- [20] S. Solomon, D. J. Ivey, D. Kinnison, M. J. Mills, R. R. Neely, III, and A. Schmidt, "Emergence of healing in the Antarctic ozone layer," *Science*, vol. 353, no. 6296, pp. 269–274, 2016.
- [21] J. C. Wei *et al.*, "Ozone profile retrieval from an advanced infrared sounder: Experiments with tropopause-based climatology and optimal estimation approach," *J. Atmos. Ocean. Tech.*, vol. 27, pp. 1123–1139, Jul. 2010.
- [22] G. Backus and F. Gilbert, "Uniqueness in the inversion of inaccurate gross earth data," *Philos. Trans. Roy. Soc. London A, Math. Phys. Sci.*, vol. 266, no. 1173, pp. 123–192, 1970.
- [23] B. J. Conrath, "Vertical resolution of temperature profiles obtained from remote radiation measurements," *J. Atmos. Sci.*, vol. 29, no. 7, pp. 1262–1271, 1972.
- [24] C. D. Rodgers and B. J. Connor, "Intercomparison of remote sounding instruments," *J. Geophys. Res.*, vol. 108, no. D3, p. 4116, 2003.
- [25] E. S. Maddy and C. D. Barnet, "Vertical resolution estimates in version 5 of AIRS operational retrievals," *IEEE Trans. Geosci. Remote Sens.*, vol. 46, no. 8, pp. 2375–2384, Aug. 2008.
- [26] N. R. Nalli *et al.*, "Validation of satellite sounder environmental data records: Application to the cross-track infrared microwave sounder suite," *J. Geophys. Res. Atmos.*, vol. 118, no. 24, pp. 13,628–13,643, 2013.
- [27] M. D. Goldberg, H. Kilcoyne, H. Cikanek, and A. Mehta, "Joint Polar Satellite System: The United States next generation civilian polar-orbiting environmental satellite system," *J. Geophys. Res. Atmos.*, vol. 118, no. 24, pp. 13,463–13,475, 2013.
- [28] Y. Han *et al.*, "Suomi NPP CrIS measurements, sensor data record algorithm, calibration and validation activities, and record data quality," *J. Geophys. Res. Atmos.*, vol. 118, no. 22, pp. 12,734–12,748, 2013.
- [29] A. M. Thompson, J. C. Witte, S. J. Oltmans, and F. J. Schmidlin, "SHADOZ—A tropical ozonesonde-radiosonde network for the atmospheric community," *Bull. Amer. Meteorol. Soc.*, vol. 85, no. 10, pp. 1549–1564, 2004.
- [30] V. Morris *et al.*, "Measuring Trans-Atlantic aerosol transport from Africa," *Eos Trans. AGU*, vol. 87, no. 50, pp. 565–571, 2006.
- [31] F. M. Ralph *et al.*, "CalWater field studies designed to quantify the roles of atmospheric rivers and aerosols in modulating U.S. West Coast precipitation in a changing climate," *Bull. Amer. Meteorol. Soc.*, vol. 97, no. 7, pp. 1209–1228, 2016.
- [32] N. R. Nalli *et al.*, "Satellite sounder observations of contrasting tropospheric moisture transport regimes: Saharan air layers, Hadley cells, and atmospheric rivers," *J. Hydrometeorol.*, vol. 17, no. 12, pp. 2997–3006, 2016.
- [33] P. J. Neiman *et al.*, "An analysis of coordinated observations from NOAA's Ronald H. Brown Ship and G-IV aircraft in a landfalling atmospheric river over the North Pacific during CalWater-2015," *Monthly Weather Rev.*, vol. 145, no. 11, pp. 3647–3669, 2017.
- [34] M. Divakarla *et al.*, "Evaluation of atmospheric infrared sounder ozone profiles and total ozone retrievals with matched ozonesonde measurements, ECMWF ozone data, and ozone monitoring instrument retrievals," *J. Geophys. Res.*, vol. 113, p. D15308, Aug. 2008.
- [35] Y. Han, Y. Chen, X. Xiong, and X. Jin, "S-NPP CrIS full spectral resolution SDR processing and data quality assessment," in *Annu. Meet.*, Phoenix, AZ, USA, Jan. 2015. [Online]. Available: <https://ams.confex.com/ams/95Annual/webprogram/Paper261524.html>
- [36] A. Gambacorta *et al.*, "An experiment using high spectral resolution CrIS measurements for atmospheric trace gases: Carbon monoxide retrieval impact study," *IEEE Geosci. Remote Sens. Lett.*, vol. 11, no. 9, pp. 1639–1643, Sep. 2014.



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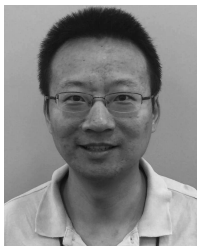
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