

REPORT DOCUMENTATION PAGE**Form Approved
OMB No. 0704-0188**

Public reporting burden for this collection of information is estimated to average 1 hour per response, including the time for reviewing instructions, searching data sources, gathering and maintaining the data needed, and completing and reviewing the collection of information. Send comments regarding this burden estimate or any other aspect of this collection of information, including suggestions for reducing this burden to Washington Headquarters Service, Directorate for Information Operations and Reports, 1215 Jefferson Davis Highway, Suite 1204, Arlington, VA 22202-4302, and to the Office of Management and Budget, Paperwork Reduction Project (0704-0188) Washington, DC 20503.

PLEASE DO NOT RETURN YOUR FORM TO THE ABOVE ADDRESS.**1. REPORT DATE (DD-MM-YYYY)**
01-07-2022**2. REPORT TYPE**
Journal Article**3. DATES COVERED (From - To)****4. TITLE AND SUBTITLE**
Ionospheric Boundaries Derived from Auroral Images**5a. CONTRACT NUMBER****5b. GRANT NUMBER****5c. PROGRAM ELEMENT NUMBER****6. AUTHOR(S)**
Gareth Chisham (1), Angeline G. Burrell (2), Evan G. Thoman (3), Yun-Ju Chen (4)**5d. PROJECT NUMBER****5e. TASK NUMBER****5f. WORK UNIT NUMBER****7. PERFORMING ORGANIZATION NAME(S) AND ADDRESS(ES)**
1) British Antarctic Survey, Cambridge, UK; 2) U.S. Naval Research Laboratory, Washington, D.C., USA;
3) Dartmouth College, Hanover, NH, USA; 4) University of Texas at Dallas, Richardson, TX, USA**8. PERFORMING ORGANIZATION
REPORT NUMBER****9. SPONSORING/MONITORING AGENCY NAME(S) AND ADDRESS(ES)****10. SPONSOR/MONITOR'S ACRONYM(S)****11. SPONSORING/MONITORING
AGENCY REPORT NUMBER****12. DISTRIBUTION AVAILABILITY STATEMENT**
DISTRIBUTION STATEMENT A. Approved for public release: distribution unlimited.**13. SUPPLEMENTARY NOTES****14. ABSTRACT**

This paper presents updated methods for locating the Poleward and Equatorward Auroral Luminosity Boundaries (PALB and EALB) directly from IMAGE Far UltraViolet (FUV) images of the Northern Hemisphere auroral oval. Separate boundaries are determined from images measured at different FUV wavelengths. In addition, new methods for indirectly estimating the Open-Closed magnetic field line Boundary (OCB) and the Equatorward Precipitation Boundary (EPB) locations are presented; these new boundaries are derived from a combination of the auroral luminosity boundary estimates with statistical latitudinal offsets derived from comparisons with low-altitude spacecraft Particle Precipitation Boundaries (PPBs). Subsequently, we derive new circle model fits for all these boundary data sets, as well as new quality control criteria for these model fits. The suitability of circle fits for each of the data sets is discussed, and the OCB and PALB circle fits are validated against the Convection Reversal Boundary (CRB), as measured by low-altitude in situ spacecraft. All the new boundary data sets, covering the epoch May 2000 to October 2002, are freely available online.

15. SUBJECT TERMS
Ionosphere, polar cap, auroral oval**16. SECURITY CLASSIFICATION OF:**a. REPORT
Ab. ABSTRACT
Ac. THIS PAGE
A**17. LIMITATION OF
ABSTRACT**
Dist A**18. NUMBER
OF PAGES**
52**19a. NAME OF RESPONSIBLE PERSON**
Angeline G. Burrell**19b. TELEPHONE NUMBER (Include area code)**
(202) 404-4065

1 **Ionospheric Boundaries Derived from Auroral Images**

2 **G. Chisham ¹, A. G. Burrell ², E. G. Thomas ³, and Y.-J. Chen ⁴**

3 ¹British Antarctic Survey, Cambridge, UK

4 ²Space Science Division, US Naval Research Laboratory, Washington, DC, USA

5 ³Thayer School of Engineering, Dartmouth College, Hanover, New Hampshire, USA

6 ⁴University of Texas at Dallas, Richardson, Texas, USA

7 **Key Points:**

- 8 • Improved poleward and equatorward auroral luminosity boundaries are determined
9 from IMAGE FUV measurements.
10 • New OCB and EPB data sets are derived from these observations.
11 • New circle model fits are derived for all the boundary data sets.

Corresponding author: Gareth Chisham, gchi@bas.ac.uk

12 Abstract

13 This paper presents updated methods for locating the Poleward and Equatorward Au-
 14 roral Luminosity Boundaries (PALB and EALB) directly from IMAGE Far UltraVio-
 15 let (FUV) images of the Northern Hemisphere auroral oval. Separate boundaries are de-
 16 termined from images measured at different FUV wavelengths. In addition, new meth-
 17 ods for indirectly estimating the Open-Closed magnetic field line Boundary (OCB) and
 18 the Equatorward Precipitation Boundary (EPB) locations are presented; these new bound-
 19 aries are derived from a combination of the auroral luminosity boundary estimates with
 20 statistical latitudinal offsets derived from comparisons with low-altitude spacecraft Partic-
 21 le Precipitation Boundaries (PPBs). Subsequently, we derive new circle model fits for
 22 all these boundary data sets, as well as new quality control criteria for these model fits.
 23 The suitability of circle fits for each of the data sets is discussed, and the OCB and PALB
 24 circle fits are validated against the Convection Reversal Boundary (CRB), as measured
 25 by low-altitude in situ spacecraft. All the new boundary data sets, covering the epoch
 26 May 2000 to October 2002, are freely available online.

27 Plain Language Summary

28 The ability to measure and model near-Earth space, its response to driving forces
 29 in the interplanetary solar wind, and the impact of these forces on the Earth’s upper at-
 30 mosphere is the basis for the understanding and forecasting of space weather. Because
 31 there is a strong physical connection between the Earth’s magnetosphere (the region of
 32 near-Earth space dominated by the Earth’s magnetic field) and ionosphere (the ionised
 33 region of the upper atmosphere) near the poles, observations of the ionosphere provide
 34 information about the magnetospheric state. Dynamic processes within the magneto-
 35 sphere accelerate electrons and ions within the space environment giving them the en-
 36 ergy to precipitate in the upper atmosphere. These high-energy charged particles excite
 37 atmospheric atoms and molecules, creating aurora. The light from the aurora can be ob-
 38 served at ultraviolet wavelengths by space-based imagers. In this study we identify the
 39 higher and lower latitude edges of the aurora and determine their relationship to impor-
 40 tant particle boundaries. Circles are then fitted to the spatial boundary variations to make
 41 sure they are defined across all longitudes. These data sets can be used to identify and
 42 track the spatial extent and dynamics of magnetospheric regions and boundaries.

43 1 Introduction

44 The Earth’s magnetosphere covers a vast region of near-Earth space, but in situ
 45 measurements of magnetospheric dynamics are limited to a small number of Earth-orbiting
 46 spacecraft that can only provide single-point measurements at a particular time. Extend-
 47 ing these measurements to describe the whole system relies on interpolation, extrapo-
 48 lation, speculation, or the use of magnetospheric models. However, at high latitudes the
 49 terrestrial ionosphere couples with the outer regions of the magnetosphere with progres-
 50 sively higher latitudes coupling to regions further from the Earth (Kivelson et al., 1996;
 51 S. W. H. Cowley, 2000). This Magnetosphere-Ionosphere (MI) coupling means that iono-
 52 spheric dynamics and properties can provide insights about the state of the magneto-
 53 sphere.

54 Magnetospheric regions can be categorised in different ways, such as defining re-
 55 gions by the local particle populations or currents. One of the simplest divisions is the
 56 separation of the magnetosphere into regions defined by their different magnetic field topolo-
 57 gies (Dungey, 1961; Milan et al., 2017):

- 58 1. ‘Open’ magnetic field lines connect directly from the geomagnetic field to the In-
 59 terplanetary Magnetic Field (IMF) that emanates from the Sun, and expands through-

60 out the solar system. These field lines converge into the geomagnetic polar regions,
61 forming the polar caps.

62 2. ‘Closed’ magnetic field lines are geomagnetic field lines that exist wholly within
63 the geospace domain, connecting directly from the Southern to the Northern Hemi-
64 sphere ionosphere, equatorward of the polar caps, with no direct connection to the
65 IMF.

66 The boundary between these topologically-different magnetospheric regions is commonly
67 known as the Open-Closed magnetic field line Boundary (OCB) (Lockwood, 1998).

68 One way of identifying different magnetospheric regions and boundaries is through
69 auroral precipitation into the ionosphere (Newell et al., 2004; Kilcommons et al., 2017).
70 The auroral region is immediately equatorward of the ionospheric signature of the OCB
71 and is characterised by a range of different auroral forms that provide information about
72 associated magnetospheric dynamics. At the poleward edge of the auroral region, au-
73 roral emissions change as a result of the transition from closed to open field lines at the
74 OCB. At the equatorward edge of the auroral region the disappearance of auroral emis-
75 sions results from the reduction in precipitating particles to low levels, although precip-
76 itation can often be measured equatorward of the auroral emission boundary. The Equ-
77 atorward Precipitation Boundary (EPB) can be measured by low altitude spacecraft and
78 matches well with the low latitude boundary of solar wind-driven ionospheric plasma con-
79 vection (Greenwald et al., 2002).

80 The magnetospheric system is highly dynamic, as it (and consequently, the cou-
81 pled ionosphere) responds to changes in driving forces in the solar wind, as well as to
82 atmospheric losses (such as Joule heating) and losses to the interplanetary medium (such
83 as substorm plasmoid ejections) (Milan et al., 2003; Chisham et al., 2008). Both the po-
84 lar cap and the auroral regions expand and contract in response to magnetic reconnec-
85 tion driven by both solar wind and internal magnetospheric dynamics. Consequently, the
86 OCB, the EPB, and the auroral region between them are in continual motion (Siscoe
87 & Huang, 1985; S. Cowley & Lockwood, 1992; Milan et al., 2003, 2012). Hence, many
88 aspects of MI science, as well as space weather applications, require the regular measure-
89 ment or estimation of the locations of these boundaries (Newell et al., 2014; Zhu et al.,
90 2020).

91 It is possible to identify different regions and boundaries in the ionosphere as the
92 characteristics of the precipitation of particles from the magnetosphere into these dif-
93 ferent ionospheric regions (polar cap, auroral region, sub-auroral region) are distinct. This
94 is due to differences in the plasma in the magnetospheric regions they map to (Sotirelis
95 & Newell, 2000). These changes in precipitation are best measured by plasma detectors
96 on spacecraft in low-Earth orbit, which provide highly accurate and precise boundary
97 identifications at single locations (e.g., Newell et al., 1991, 1996). The drawback of such
98 observations is that they provide infrequent and sparse estimates of a boundary, which
99 makes them unsuitable for providing comprehensive spatio-temporal pictures of the bound-
100 aries and their motion. Such comprehensive boundary definitions are required both to
101 improve the high-latitude specifications for space weather modelling, and to facilitate
102 statistical studies of the polar region (Newell et al., 2004, 2014; Redmon et al., 2010; Chisham,
103 2017a; Kilcommons et al., 2017; Burrell et al., 2020a).

104 Fortunately, there are other ground- and space-based instruments that can observe
105 larger regions of the high-latitude ionosphere, allowing a more complete mapping of the
106 regions and boundaries. From the ground, both coherent and incoherent scatter radars
107 (Blanchard et al., 2001; Chisham & Freeman, 2003; Aikio et al., 2006; Chisham et al.,
108 2007), as well as photometers and all-sky cameras (Blanchard et al., 1995; X.-C. Chen
109 et al., 2017) have been used to locate boundaries in the ionosphere. On a larger scale,
110 the Active Magnetosphere and Planetary Electrodynamics Response Experiment (AM-
111 PERE) (Anderson et al., 2014; Coxon et al., 2018) and the SuperMAG archive of global

112 ground-based magnetometer data (Gjerloev, 2012) provide full-hemispheric coverage of
 113 high-latitude magnetic field variability, providing maps of MI current systems. Space-
 114 based auroral images, through their extensive coverage of the auroral oval, are proba-
 115 bly the best tool for identifying the spatial extent of the auroral region, and provide the
 116 ability to monitor both the OCB and EPB simultaneously. There is a long history of de-
 117 termining the poleward edge of the auroral oval from space-based auroral imager data
 118 (Kauristie et al., 1999; Baker et al., 2000; Carbary et al., 2003; Boakes et al., 2008; Long-
 119 den et al., 2010). This boundary has regularly been used as a proxy for the OCB in order
 120 to study processes occurring there, like magnetic reconnection (Hubert et al., 2006;
 121 Chisham et al., 2008; Hubert et al., 2010).

122 In this paper, we present boundaries derived from the Far UltraViolet (FUV) im-
 123 agers on the Imager for Magnetopause-to-Aurora Global Exploration (IMAGE) space-
 124 craft (Mende et al., 2000a). The Northern Hemisphere IMAGE epoch has proved an ex-
 125 tremely good data set for studies of space weather (Frey et al., 2004; Boakes et al., 2009,
 126 2011; Milan, 2009; Milan et al., 2009; Mooney et al., 2020, 2021). Longden et al. (2010)
 127 presented a reliable, automated method to estimate the location of Poleward Auroral Lu-
 128 minosity Boundaries (PALBs) and Equatorward Auroral Luminosity Boundaries (EALBs)
 129 in IMAGE FUV measurements. Here, we introduce additional selection criteria that re-
 130 move potentially unreliable boundaries from the Longden et al. (2010) data sets. We also
 131 present new methods for estimating the OCB and EPB from the measured PALBs and
 132 EALBs, respectively. We also model all these boundary data sets with circles using the
 133 method outlined by Chisham (2017a), with the important addition of a figure of merit
 134 that allows the reliability of the data to be more easily assessed. These fitted circle ver-
 135 sions of the boundaries allow easier implementation into models and certain data anal-
 136 ysis applications. Finally, the new OCB circle data set is validated in a method similar
 137 to the IMAGE OCB validation performed in Burrell et al. (2020a).

138 2 Instrumentation

139 2.1 IMAGE FUV

140 The IMAGE spacecraft (Burch, 2000; Gibson et al., 2000) was launched on 25 March
 141 2000. One of the key aims of the mission was to simultaneously image both proton and
 142 electron aurora using three FUV imagers. The data set described in this paper was com-
 143 piled from IMAGE FUV images of the Northern Hemisphere auroral regions extending
 144 from May 2000 to October 2002, during which time the spacecraft was located within
 145 a highly elliptical polar orbit with an inclination of 90° . The orbit apogee during this
 146 time was positioned over the Northern Hemisphere polar ionosphere at $\sim 7 R_E$; the or-
 147 bit perigee was located in the upper regions of the Southern Hemisphere ionosphere, at
 148 ~ 1000 km altitude. The orbital period was ~ 13.5 hrs, allowing imaging of the North-
 149 ern Hemisphere auroral oval for most of this time.

150 The FUV instrument on the IMAGE spacecraft (Mende et al., 2000a) comprised
 151 three imagers: The Spectrographic Imagers (SI), termed SI12 and SI13 (Mende et al.,
 152 2000c), and the Wideband Imaging Camera (WIC) (Mende et al., 2000b). All three im-
 153 agers had a temporal resolution of ~ 2 min. The SI12 imager measured Doppler-shifted
 154 Lyman- α emissions at 121.8 nm that occur as a result of proton precipitation. The SI13
 155 imager measured oxygen emissions at 135.6 nm that occur as a result of energetic elec-
 156 tron precipitation. The WIC imager measured emissions in the N_2 Lyman-Birge-Hopfield
 157 band (140-190 nm) that occur as a result of electron precipitation. Images from both the
 158 SI imagers had a resolution of 128×128 pixels, whereas those from the WIC imager had
 159 a resolution of 256×256 pixels.

160 The high spatial resolution of the imagers and the near-perpendicular pointing an-
 161 gles around apogee resulted in good images of the complete auroral oval for much of the

162 orbit. However, as the spacecraft approached perigee the images were taken from increas-
 163 ingly oblique angles. As a consequence of this, the entirety of the auroral oval was not
 164 captured in these images and the auroral features are increasingly smeared. Thus, bound-
 165 ary determinations from the images taken at these points in the orbits are less accurate.
 166 Additional uncertainties in boundary determinations also result from minor inaccura-
 167 cies in the spacecraft pointing direction and from the challenges of dealing with dayglow.
 168 The impacts of these issues on the boundary locations are discussed in full detail by Longden
 169 et al. (2010).

170 2.2 DMSP

171 The Defense Meteorological Satellite Program (DMSP) consists of a constellation
 172 of polar orbiting, sun-synchronous satellites in low Earth orbit (~ 830 km altitude). Data
 173 from the Special Sensor J (SSJ)/4 instruments (Hardy et al., 1984) on each spacecraft
 174 allowed the accurate identification of the locations of ionospheric Particle Precipitation
 175 Boundaries (PPBs) along the DMSP orbits. These PPBs occur where the boundaries
 176 between magnetospheric regions characterised by different ion and electron distributions
 177 map to the ionosphere. The orbital period of the DMSP spacecraft was ~ 101 min, and
 178 ion and electron energy spectra were measured every second, which corresponds to a po-
 179 tential spatial resolution of the boundary determinations along the orbit of ~ 7.5 km ($\sim 0.1^\circ$
 180 latitude).

181 The algorithms used to identify PPBs in the DMSP data are different for measure-
 182 ments made in the dayside and nightside portions of the ionosphere, due to the local time
 183 differences in the particle precipitation regions. The algorithm used for automated identi-
 184 fication of dayside PPBs from the DMSP energy spectra is detailed in Sotirelis and Newell
 185 (2000), and is based on boundaries between the dayside precipitation regions classified
 186 by Newell et al. (1991). The ‘doc’ boundary, which represents an unambiguous transi-
 187 tion between open and closed precipitation regions, is taken as the best estimate of the
 188 OCB on the dayside. The ‘deq’ boundary, which is located at the equatorward edge of
 189 diffuse precipitation, is taken as the EPB on the dayside.

190 The algorithm used for automated identification of nightside PPBs from the DMSP
 191 energy spectra is detailed in Newell et al. (1996) and Sotirelis and Newell (2000). The
 192 ‘b6’ boundary on the nightside, which represents the poleward edge of the subvisual au-
 193 roral drizzle that occurs poleward of the main auroral oval, is taken as the best proxy
 194 for the OCB on the nightside. The most equatorward of the ‘b1e’ boundary (the ‘zero-
 195 energy’ electron convection boundary) and the ‘b2i’ boundary (the ion isotropy bound-
 196 ary) is taken as the EPB on the nightside.

197 Following Sotirelis and Newell (2000), if multiple crossings of the same magneto-
 198 spheric precipitation region are identified on a single spacecraft overpass, then the most
 199 poleward boundary for each region is used in the determination of the OCB and the most
 200 equatorward boundary for each region is used in the determination of the EPB. Where
 201 clear transitions between precipitation regions cannot be made, the PPBs are not used
 202 in this analysis. These post-processing checks identify failures of the boundary algorithm,
 203 and remove ambiguous boundaries.

204 Data from the Ion Velocity Meter (IVM) on the DMSP spacecraft (Heelis & Han-
 205 son, 1998) are also used to identify Convection Reversal Boundaries (CRBs) that are used
 206 here for validation of the poleward boundaries (Y.-J. Chen et al., 2015; Burrell et al.,
 207 2020a). These CRBs indicate where the ionospheric plasma drifts change from moving
 208 sunward to anti-sunward (or vice versa). The locations of the CRBs have lower uncer-
 209 tainties when the plasma density is high and a high-fidelity signal is returned. Hence,
 210 the CRB data set is biased to local summer months (May-August) in the Northern Hemi-
 211 sphere.

212 During the epoch of the IMAGE FUV observations, PPBs were available from four
 213 DMSP spacecraft: F12, F13, F14, and F15. CRBs were available from F13 and F15. The
 214 Northern Hemisphere orbit of DMSP F13 was approximately dawn to dusk in AACGM
 215 co-ordinates, whereas the DMSP F12, F14, and F15 orbits were approximately in the
 216 pre-noon to post-dusk direction. These restrictions in the available orbits mean that there
 217 are no DMSP measurements in the pre-midnight and early morning sectors in the North-
 218 ern Hemisphere.

219 **3 Pre-existing IMAGE FUV boundary determination methods and** 220 **data sets**

221 The new auroral luminosity boundary data sets presented here are refinements of
 222 those determined by Longden et al. (2010). The method of Longden et al. (2010) was
 223 based on that of Boakes et al. (2008), who first developed an automated method for iden-
 224 tifying boundaries in the IMAGE FUV data. Longden et al. (2010) improved this method-
 225 ology in several ways, the most important of which was adding the ability to deal with
 226 the presence of bifurcated auroral ovals (which they showed can comprise up to 50% of
 227 the data base in the midnight sector during active geomagnetic times). This method re-
 228 quires no prior knowledge of the presence of bifurcations in the oval; it involves the fit-
 229 ting of two different functions to the data to deal with the possibility of both single and
 230 double ovals. The Longden et al. (2010) methodology also allows the estimation of the
 231 PALBs and EALBs with no prior knowledge of the level of auroral activity (i.e., the method
 232 is not based on fixed intensity thresholds as some previous boundary determination meth-
 233 ods were).

234 Here, we present a brief overview of the Longden et al. (2010) methodology:

- 235 1. The data in each auroral image are divided into 24 regional segments in the Altitude-
 236 Adjusted Corrected GeoMagnetic (AACGM) latitude and Magnetic Local Time
 237 (MLT) coordinate system (Shepherd, 2014), each segment covering 1 h of MLT.
- 238 2. For each 1 h MLT segment, a latitudinal intensity profile is constructed by find-
 239 ing the average pixel intensity in bins of 1° AACGM latitude in the range 50° to
 240 90°.
- 241 3. Two separate model functions are fitted to each latitudinal intensity profile:
 - 242 (a) a function with a single Gaussian component and a quadratic background, and
 - 243 (b) a function with two Gaussian components and a quadratic background.
- 244 4. Of the two fitted functions, the better model for that profile is chosen by deter-
 245 mining the reduced χ^2 goodness-of-fit statistic for both functions. This is performed
 246 separately for the latitudinal intensity profiles measured at each MLT.
- 247 5. When the single Gaussian function provides the better fit, the location of the PALB
 248 is estimated as being offset poleward from the location of the centre of the Gaus-
 249 sian peak by the Full Width at Half Maximum (FWHM) of the Gaussian profile,
 250 whereas the EALB is estimated as being offset equatorward from the centre of the
 251 Gaussian peak by the FWHM. When the double Gaussian function provides the
 252 better fit, the location of the PALB is estimated as being offset poleward from the
 253 location of the poleward peak of the two Gaussian components by the FWHM of
 254 that peak, whereas the EALB is estimated as being offset equatorward from the
 255 location of the equatorward peak of the two Gaussian components by the FWHM
 256 of that peak.
- 257 6. Criteria are applied both during and after the fitting process to discard bound-
 258 aries where the fitting has been poor, or the data incomplete. However, as with
 259 any automated technique it has its limitations, and high-latitude sun-aligned arcs,
 260 and subauroral features such as detached arcs and patches may occasionally re-

261 sult in inaccurate modelling of the boundaries (see Longden et al. (2010) for full
262 details).

263 The analyses of Boakes et al. (2008) and Longden et al. (2010) showed that the PALBs
264 determined from the three different FUV imagers are not co-located with each other; nei-
265 ther are they co-located with the PPB estimates of the OCB, which are generally accepted
266 as the most accurate markers of the OCB. The offsets between the different measure-
267 ments are a result of the separate behaviour of ions and electrons on magnetic field lines
268 close to the boundary. In order to estimate the most probable OCB locations from the
269 measured PALB locations, Longden et al. (2010) determined statistical latitudinal off-
270 sets between the measured PALBs and DMSP estimates of the OCB. They also deter-
271 mined the variation of these statistical offsets with MLT. These offsets for the original
272 PALBs from each FUV imager were published by Longden et al. (2010). The original
273 Longden PALB and EALB data sets are freely available (Chisham, 2017b).

274 Subsequent studies have shown the usefulness of this original data set. Longden
275 et al. (2014) used these data to study the local time variation and scaling of poleward
276 auroral boundary dynamics. Mooney et al. (2020) used the data set to study the spatio-
277 temporal variation of the OCB through the substorm cycle. They followed up this work
278 by using the IMAGE boundary data to evaluate auroral forecasts (Mooney et al., 2021).
279 Recently, Hoque et al. (2021) have used the data set to evaluate convection reversal bound-
280 aries measured by the SuperDARN radars.

281 Following the work of Longden et al. (2010), Chisham (2017a) showed that by as-
282 suming that the PALBs are well modelled by circles (in AACGM co-ordinates), then these
283 data sets can be more practically used in the development of new ionospheric climatolo-
284 gies and empirical models. In this new methodology, empirical data are binned in a co-
285 ordinate system that adapts relative to the position of the boundary. Chisham (2017a)
286 presented a demonstration of how this methodology can be used to improve the clima-
287 tological representation of ionospheric vorticity. The derived PALB circle fits form a key
288 part of the OCBpy python data analysis package (Burrell et al., 2020b).

289 4 New boundary determination methods and data sets

290 4.1 Changes to existing boundary acceptance criteria

291 The new PALB and EALB data bases are a subset of the Longden et al. (2010) data
292 sets. In the creation of the new PALB and EALB data bases, changes have been made
293 to the previously published methodology to deal with two separate issues.

294 4.1.1 *Edge of field of view issues*

295 The original boundary-determination algorithm developed by Longden et al. (2010)
296 didn't adequately consider problems that might be introduced owing to the location of
297 the edge of the instrument field of view. In Longden et al. (2010), acceptance criterion
298 2 in section 3.2 of that paper stated - 'The Gaussian centre(s) μ must fall within the AACGM
299 latitude range of the given intensity profile.' This criterion required that more than half
300 of the latitudinal auroral intensity variation had been observed, but did not require that
301 the estimated boundary location itself was within the field of view. This criterion oc-
302 casionally introduced issues when there were anomalous or missing intensity values at
303 the edge of the field of view which resulted in the appearance of a false peak in the lat-
304 itudinal intensity profile. It also introduced biases if the observable portion of the au-
305 roral profile at the edge of the field of view was not strictly Gaussian. As a consequence
306 of these issues, the Longden algorithm often placed boundaries equatorward of the im-
307 age field of view without having observed the whole profile.

Figure 1. Auroral images measured by the WIC FUV imager at 14:10 UT on 15 December 2001 (day 349). PALBs and EALBs are shown on the figure as yellow squares using (a) the original Longden boundary data set, and (b) applying the new data selection criteria. The yellow solid lines show circle fits to the PALBs and EALBs.

308 Figure 1a presents an example WIC image where some of the boundaries (yellow
 309 squares) have been evaluated to be equatorward of the image field of view. On the day-
 310 side, where dayglow dominates the image, variations in the dayglow intensity have re-
 311 sulted in false EALBs equatorward of the field of view. In the main auroral oval on the
 312 nightside, EALBs have also been placed equatorward of the edge of the field of view. Al-
 313 though there are often instances where this boundary placement may be accurate, in this
 314 instance partially observed latitudinal intensity profiles that are not strictly Gaussian
 315 have led to the EALBs being placed further equatorward than they are likely to be. Not
 316 requiring the observation of the full auroral profile within the field of view has led to these
 317 erroneous boundary placements. The solid yellow circles show circle fits to the PALB
 318 and EALB locations (see Section 5). The EALB circle fit is heavily affected by the er-
 319 roneous EALB estimates.

320 We have amended the original criterion such that boundaries in the new data set
 321 instead need to be in a 1° AACGM latitude bin that contains good data and lies at least
 322 2 pixels inward from the edge of the IMAGE field of view. This minimises the impact
 323 of these edge effects in most cases. Figure 1b presents the boundaries from the revised
 324 data set for the same time interval; the EALBs that were originally placed outside of the
 325 field of view have been removed from the image. Consequently, the EALB fitted circle
 326 is now more representative of the true equatorward edge of the auroral oval. As shown
 327 in this example, this problem is more of an issue for the EALBs than the PALBs as the
 328 EALBs are typically closer to the edge of the image field of view. This issue is most per-
 329 tinent for images taken when the IMAGE spacecraft was close to perigee; at these times
 330 the spatial area of the ionosphere contained within the field of view of each image was
 331 reduced and the auroral oval often extended outside of the image field of view.

332 *4.1.2 Anomalous mapping issues*

333 Within the Longden et al. (2010) IMAGE FUV data base there are occasional or-
 334 bits for which the image mapping to the ionosphere for particular imagers is in error.
 335 These anomalies in the Longden data base were identified through the comparison of the
 336 parameters of circle fits to the PALBs and EALBs (see Section 5) for the three FUV im-
 337 agers along a complete polar orbit. This anomalous mapping is also clear in individual
 338 comparisons of simultaneous images from all three FUV imagers.

Figure 2. Auroral images measured by the IMAGE FUV imagers at 06:52 UT on 30 September 2000 (day 274), from (a) the SI12 imager, (b) the SI13 imager, and (c) the WIC imager. The yellow squares indicate the locations of the PALBs and EALBs determined using the original Longden methodology.

339 Figure 2 presents an example that shows this anomalous mapping; it presents im-
 340 ages from 06:51:53 UT on 30 September 2000 (day 274) as measured by (a) SI12, (b) SI13,
 341 and (c) WIC. Although the auroral ovals viewed by SI12 and WIC show slightly differ-
 342 ent features (as expected for proton and electron aurora), the PALBs and EALBs de-
 343 termined from the auroral profiles are very similar, and the oval is centred close to the

344 AACGM pole. However, the auroral oval in the SI13 image, whilst showing similar fea-
 345 tures to the WIC aurora (they both represent electron aurora), is shifted significantly
 346 toward dawn (by $\sim 5^\circ$ - 7°). This anomalous shift in the mapped SI13 auroral position per-
 347 sists throughout this orbit.

348 Boundaries determined from data in the orbits in the Longden data base that are
 349 affected by this anomalous mapping have been removed from the new PALB and EALB
 350 data sets. This issue resulted in the removal of $\sim 2\%$ of the data base, predominantly SI12
 351 and SI13 data. It is possible that future analysis may provide the opportunity to cor-
 352 rect this mapping problem and for these intervals to be included in future versions of the
 353 IMAGE FUV boundary data.

354 4.2 The Poleward Auroral Luminosity Boundary (PALB)

355 A summary of the new PALB data base is presented in Figure 3. The black solid
 356 line shows the variation with MLT of the median PALB latitude for each imager: (a) SI12,
 357 (b) SI13, and (c) WIC. The thick error bars at each MLT show the latitude range from
 358 the lower to the upper quartile of the measured PALB latitudes, whereas the thinner bars
 359 show the latitude range from the 10% to 90% percentiles. Although the variations for
 360 the three imagers are similar, there are subtle differences in the median latitudes of the
 361 different boundaries (as discussed above), particularly in the SI12 variation, which ex-
 362 tends to lower latitudes in the afternoon sector. This is probably owing to differences
 363 in the relationship of electron and proton precipitation to the OCB in this sector.

Figure 3. Statistics of PALB locations identified in data from the IMAGE FUV imagers (a)
 SI12, (b) SI13, and (c) WIC, in AACGM and MLT co-ordinates. The solid black line describes
 the MLT variation of the median PALB latitude. The thick black error bars present the quartile
 latitude variation at each MLT. The thin black error bars at each MLT present the extent of
 the distribution from the 10% to 90% percentile. The orange histograms present the number of
 PALB observations in each MLT sector.

364 The orange histograms in each map highlight the number of boundaries measured
 365 in each MLT sector across the whole measurement epoch. This shows that for all the im-
 366 agers a larger number of boundaries are measured in MLT sectors in the nightside iono-
 367 sphere ($\sim 10^5$), where the auroral emissions are strongest and where there is less contam-
 368 ination of the images by dayglow. The number of boundaries decreases significantly to-
 369 ward noon, where the number of observed boundaries is lower ($\sim 10^4$). The variation shown
 370 in Figure 3 is similar to that of the original PALB data set as was presented in Figure
 371 4a of Longden et al. (2010). The main difference here is the reduction in the number of
 372 WIC boundaries from dawn, through noon, to dusk. This is because the new criterion
 373 now excludes potentially erroneous boundaries in the original data set that were com-
 374 plicated by the effects of dayglow. Whereas the number of WIC boundaries in the orig-
 375 inal data set peaked around dawn and dusk, now the peak is near midnight, consistent
 376 with the other two imagers.

377 These data bases are located in the files PALB_SI12_v2, PALB_SI13_v2, and PALB_WIC_v2
 378 in the directory ‘raw_boundaries’ at [https://doi.org/10.5285/fa592594-93e0-4ee1-8268-](https://doi.org/10.5285/fa592594-93e0-4ee1-8268-b031ce21c3ca)
 379 [b031ce21c3ca](https://doi.org/10.5285/fa592594-93e0-4ee1-8268-b031ce21c3ca) (Chisham, 2022).

380 4.3 The Open-Closed magnetic field line Boundary (OCB)

381 In the original studies of Boakes et al. (2008) and Longden et al. (2010) the PALBs
 382 were compared with DMSP OCB proxies to determine statistical latitudinal offsets be-

383 tween the boundaries. In these studies, the latitudinal offsets were determined separately
 384 for each imager in each MLT sector (see the methods presented in those papers). How-
 385 ever, with three imagers this methodology leads to three distinct estimates of the OCB
 386 rather than one. This can lead to difficulties when explaining the differences between dis-
 387 parate OCB estimates from different data sets (e.g., Mooney et al., 2020).

388 Here, we present a new methodology that combines boundary data from all the im-
 389 agers and their relationships with the DMSP OCB proxy to make a single best estimate
 390 of the OCB location at each time. The statistical comparisons between the IMAGE PALBs
 391 and the DMSP OCB estimates are here treated differently to the methods used by Boakes
 392 et al. (2008) and Longden et al. (2010). First, we are more selective about which DMSP
 393 data are used in the comparison. This is important because the quality of the DMSP
 394 precipitation boundaries can be affected by a variety of different sources of uncertainty,
 395 including the presence of nuisance signals and magnetic perturbations (Kilcommons et
 396 al., 2017). Figure 4 displays as orange dots the locations of all the OCB estimates from
 397 the four different DMSP spacecraft used over the epoch of the study (DMSP F12, F13,
 398 F14, and F15). For each spacecraft the boundaries appear in fixed bands dictated by their
 399 orbits. The DMSP orbits cross the polar regions at fixed local times in geographic co-
 400 ordinates, but these spread into bands when the orbits are transformed into the AACGM
 401 and MLT co-ordinate system. The black error bars and solid line in Figure 4 represent
 402 the variation of the median and quartile SI13 PALBs with MLT for reference, as presented
 403 in Figure 3.

Figure 4. The locations of DMSP OCB proxies (orange dots) through the measurement epoch, in AACGM and MLT co-ordinates. The black line and error bars show the median and quartile variation of the SI13 PALB data set for reference. The red lines define areas within which the DMSP OCB estimates are most reliable.

404 In Figure 4 we have split the observed DMSP OCBs into two groups; those mea-
 405 sured at MLTs and latitudes within the enclosed red regions and those measured at MLTs
 406 and latitudes outside of these regions. We define the areas within the red regions as those
 407 where the DMSP OCB estimates are most reliable for use in our statistical comparisons.
 408 There are two reasons why we only use the observed OCBs from within the red regions:

- 409 1. In the MLT sectors within the red regions, the DMSP orbits cover the full AACGM
 410 latitude range where the OCB is likely to be. Outside of the red regions, due to
 411 the restricted spacecraft orbits, the latitudinal distribution of the observed OCBs
 412 is truncated on the equatorward side. Comparing the OCBs with the PALBs in
 413 these MLT sectors leads to a bias in the statistical offsets owing to a biased dis-
 414 tribution of uncertainties in the OCB observations. This unreliability in the sta-
 415 tistical offset is greatest in the MLT sectors where the latitudinal distribution is
 416 most truncated.
- 417 2. Boundaries are sharper and more easily identified in the DMSP data when the space-
 418 craft orbit is such that it crosses the boundary close to the perpendicular direc-
 419 tion. Comparing the spread of the orbits with the SI13 median PALB variation
 420 in Figure 4 shows that for those orbits that pass closer to the AACGM pole, the
 421 DMSP orbit passes closer to perpendicular to the median PALB location. These
 422 DMSP OCB estimates will tend to be the most accurate. For the OCBs determined
 423 in MLT sectors outside of the red regions, the DMSP orbits are close to being par-
 424 allel with the median PALB location. In some cases these orbits will be ‘skimming’
 425 the boundary, and for many of the boundaries determined in this region, the bound-
 426 ary transition will be spread over a greater time interval in the data making it more
 427 difficult to unambiguously determine the boundary location, and increasing the

428 uncertainty in its estimate. In addition, if the boundary latitude is fluctuating dur-
 429 ing the spacecraft pass, then multiple, complex crossings may be observed in these
 430 regions. This is another reason why we have not used these DMSP OCB estimates.

431 Following the methodology of Longden et al. (2010), for each DMSP OCB latitude
 432 estimate at time t ($\lambda_{OCB}(MLT, t)$) we identify the latitude of the closest PALB ($\lambda_P(i, MLT, t)$)
 433 for each FUV imager i that is within ± 5 min UT and within ± 0.5 h MLT. The latitu-
 434 dinal difference between these two boundaries ($\Delta\lambda_P(i, MLT, t)$) is determined from

$$435 \quad \Delta\lambda_P(i, MLT, t) = \lambda_P(i, MLT, t) - \lambda_{OCB}(MLT, t) \quad (1)$$

436 These latitudinal differences are not constant at a particular MLT throughout the data
 437 set, partly due to the presence of instrumental and algorithmic noise.

438 Subsequently, for each MLT sector the distributions of the estimates of $\Delta\lambda_P(i, MLT, t)$
 439 for the complete data set are compiled. From these difference distributions, both Boakes
 440 et al. (2008) and Longden et al. (2010) took the modes of the distributions to provide
 441 the best estimates for the statistical offsets in each MLT sector. These difference distri-
 442 butions are now typically Gaussian (due to the removal of unreliable DMSP OCB bound-
 443 ary estimates), which suggests that the uncertainties are random, and that the average
 444 value provides the best estimate of the typical offset. As the distributions are broadly
 445 symmetrical, we choose here to use the median value of the offset in each MLT sector,
 446 as it is more accurately determined than the mode. Median values are only determined
 447 for MLT sectors where the difference distribution is comprised of at least 100 values. Hence,
 448 for each imager, in each MLT sector, the statistical offset is defined as

$$449 \quad \Delta\lambda_P(i, MLT) = \text{Median}(\Delta\lambda_P(i, MLT, t)) \quad (2)$$

450 In Figure 5 the values of $\Delta\lambda_P(i, MLT)$ are presented. The black, orange, and red
 451 squares represent the values for the SI12, SI13, and WIC imagers, respectively. The error
 452 bars represent the quartile range of the distributions at each MLT. Owing to the or-
 453 bits of the DMSP spacecraft in the northern hemisphere (see Figure 4), the median off-
 454 sets are restricted to two MLT sectors covering the late morning (0600 to 1200 MLT)
 455 and dusk (1500 to 2200 MLT) sectors. In the morning sector the median offsets for the
 456 three imagers are within 1° of each other, and in the late morning the OCB lies on av-
 457 erage $\sim 2^\circ$ equatorward of the PALBs. This is because aurora associated with the cusp
 458 region typically exist poleward of the OCB around noon. In the late afternoon the me-
 459 dian offset for the SI12 proton aurora boundary diverges significantly from those for SI13
 460 and WIC by ~ 2 - 3° . Consequently, the SI12 PALB lies on average equatorward of the
 461 OCB in this region, whereas the other PALBs lie on average poleward of the OCB.

Figure 5. The variation of the median offsets between the IMAGE FUV PALBs and the DMSP OCBs in each 1-hr MLT sector, for the SI12 (black), SI13 (orange), and WIC (red) imagers. The error bars present the quartile range of the offsets (separated slightly in MLT for clarity). The solid lines present the harmonic function fits to the median values for each imager.

462 In order to provide the best estimate of the statistical offsets for each imager at all
 463 MLTs, we model the median offsets as a function of MLT using a second-order harmonic
 464 function (as in Carbary et al., 2003; Boakes et al., 2008; Longden et al., 2010) given by

$$465 \quad L(\phi) = C_0 + C_1 \cos \phi + D_1 \sin \phi + C_2 \cos 2\phi + D_2 \sin 2\phi \quad (3)$$

466 where ϕ represents MLT as an angle ($MLT \times 15^\circ$), starting with 0° at midnight, and in-
 467 creasing toward dawn with increasing MLT, and C_0 , C_1 , C_2 , D_1 , and D_2 are the coef-
 468 ficients of the fit. The best fits, and hence the coefficients for each imager, are determined

469 using the Levenberg-Marquardt least squares fitting routine. In Figure 5 the black, or-
 470 ange, and red lines represent the resulting fits for the SI12, SI13, and WIC imagers, re-
 471 spectively. The coefficients of the fits for each imager are presented in Table 1. These
 472 coefficients allow the determination of the poleward boundary offset $L_P(i, \phi)$ at all MLTs
 473 for each imager.

Table 1. The harmonic function fit coefficients for the OCB offsets from the imager PALBs (Equation 3). These define L_P for each imager.

Imager	C_0	C_1	D_1	C_2	D_2
SI12	0.0405	-1.5078	0.5095	0.9371	0.0708
SI13	0.5797	-0.6837	-0.5020	0.2971	-0.4173
WIC	1.0298	-1.1249	-0.7380	0.1838	-0.6171

474 As discussed previously, the earlier studies of Boakes et al. (2008) and Longden et
 475 al. (2010) produced separate estimates of the OCB for each imager, based on a combi-
 476 nation of similarly estimated offsets from the DMSP OCBs and statistical intercompar-
 477 isons between the imager data sets. Here, we combine the estimated PALB measurements
 478 from all the available imagers from any one time to provide a single OCB estimate for
 479 that time. We use the fitted offset curves for each imager, $L_P(i, \phi)$, in combination with
 480 the observed PALBs to determine our OCB estimate using

$$481 \lambda_{OCB}(\phi, t) = \sum_{i=1}^n \frac{\lambda_P(i, \phi, t) - L_P(i, \phi)}{n} \quad (4)$$

482 where n represents the number of available imagers. For times and MLTs where data
 483 from not all PALBs are available, we just use those that are.

484 A summary of the new OCB data base is presented in Figure 6. As with the PALB
 485 statistics, the solid black line shows the variation of the median OCB latitude with MLT.
 486 The thick black error bars at each MLT show the latitude range from the lower to the
 487 upper quartile of the OCB latitudes, whereas the thinner bars show the latitude range
 488 from the 10% to 90% percentiles. Again, the orange histogram represents the number
 489 of boundaries derived in each MLT sector across the whole measurement epoch. As with
 490 the PALBs from which these boundaries are derived, there are a larger number of bound-
 491 aries in the nightside ionosphere ($\sim 10^5$), where the auroral emissions are strongest and
 492 dayglow weakest.

Figure 6. Statistics of OCB locations derived from the IMAGE FUV PALBs in AACGM co-ordinates. The solid black line describes the MLT variation of the median OCB latitude. The thick black error bars present the quartile latitude variation at each MLT. The thin black error bars at each MLT present the extent of the distribution from the 10% to 90% percentile. The orange histograms present the number of OCB estimations in each MLT sector.

493 This new data set is labelled OCB_v2 in the directory ‘raw_boundaries’ at <https://doi.org/10.5285/fa592594-93e0-4ee1-8268-b031ce21c3ca> (Chisham, 2022). Those who wish to use estimates of the
 494 OCB determined from a single imager only rather than from the combined data sets can
 495 do so using the relevant PALB and the offsets described by Equation 3 and Table 1.
 496

4.4 The Equatorward Auroral Luminosity Boundary (EALB)

A summary of the new EALB data base is presented in Figure 7. As with the PALB statistics, the solid black line shows the variation of the median EALB latitude with MLT for each imager: (a) SI12, (b) SI13, and (c) WIC. The thick black error bars at each MLT show the latitude range from the lower to the upper quartile of the measured EALB latitudes, whereas the thinner bars show the latitude range from the 10% to 90% percentiles. The variations for the three imagers show greater differences than with the PALB median latitude variation, with the shape of the median SI12 (proton aurora) EALB, being distinctly different to that of the median SI13 and WIC (electron aurora) EALBs (which are similar). The median SI12 boundary extends further equatorward in the early afternoon sector, whereas the median SI13 and WIC boundaries extend further equatorward in the late morning sector. This is likely owing to the different equatorward extents of the electron and proton precipitation in the morning and afternoon sectors.

Figure 7. Statistics of EALB locations from the IMAGE FUV imagers (a) SI12, (b) SI13, and (c) WIC, in AACGM and MLT co-ordinates. The solid black line describes the MLT variation of the median EALB latitude. The thick black error bars present the quartile latitude variation at each MLT. The thin black error bars at each MLT present the extent of the distribution from the 10% to 90% percentile. The orange histograms present the number of EALB observations in each MLT sector.

The orange histograms in each panel represent the number of boundaries measured in each MLT sector across the whole measurement epoch. As with the PALBs a larger number of boundaries are measured in the nightside ionosphere where the auroral emissions are strongest. Again, the number of boundaries decreases significantly toward noon, where the effect of dayglow on the ability to identify EALBs is greater than on the PALBs. These variations were not presented for the original data set in Longden et al. (2010), although they were similar to those for the PALBs. Again, the new criteria employed here better identify potentially erroneous boundaries within the original Longden et al. (2010) data set.

The data bases are located in the files EALB_SI12_v2, EALB_SI13_v2, and EALB_WIC_v2 in the directory ‘raw_boundaries’ at <https://doi.org/10.5285/fa592594-93e0-4ee1-8268-b031ce21c3ca> (Chisham, 2022).

4.5 The Equatorward Precipitation Boundary (EPB)

This is the first time that a proxy for the EPB has been estimated indirectly from auroral imager data. This is achieved using statistical latitudinal offsets that have been determined through a comparison of the measured EALBs with the DMSP ‘b1e’/‘b2s’ and ‘deq’ EPBs discussed above. As with the derivation of the OCB proxy presented earlier, we are selective about which DMSP EPB data are used in the comparison with the EALBs. Figure 8 displays as orange dots the locations of all the EPB estimates identified by the four different DMSP spacecraft used in this study (DMSP F12, F13, F14, and F15). As with the DMSP OCBs, the boundaries again appear in fixed bands across the polar region dictated by the spacecraft orbits and the co-ordinate transform from geographic to AACGM and MLT. The black solid line and error bars show the variation of the median and quartile SI13 EALBs with MLT, as presented previously in Figure 7b. The red sectors highlight where the DMSP EPB estimates are most reliable, following the rationale described for the DMSP OCBs.

Figure 8. The locations of DMSP EPB proxies (orange dots) through the measurement epoch, in AACGM and MLT co-ordinates. The black line and error bars show the median and quartile variation of the SI13 EALB data set for reference. The red lines define areas within which the DMSP EPB estimates are most reliable.

536 We follow the same methodology as with the OCB boundary determination in order
 537 to provide estimates of the statistical EPB offsets from the EALBs measured by each
 538 imager at all MLTs, modelling the median offsets as a function of MLT using the second-
 539 order harmonic function given in Equation 3. Figure 9 presents the median offsets with
 540 the associated quartile ranges, and the estimated fits for the model offsets $L_E(i, \phi)$; the
 541 black, orange, and red symbols and lines represent the variations for the SI12, SI13, and
 542 WIC imagers, respectively. The coefficients of the fits for each of the different imagers
 543 are presented in Table 2

Table 2. The harmonic function fit coefficients for the EPB offsets from the imager EALBs (Equation 3). These define L_E for each imager.

Imager	C_0	C_1	D_1	C_2	D_2
SI12	-0.1447	-1.9779	2.6799	0.5778	-1.2297
SI13	0.2500	-2.9931	0.8818	0.8511	-0.6300
WIC	-0.4935	-2.1186	0.3188	0.5749	-0.3118

544 Following the same methodology as with the determination of the OCB proxy, we
 545 combine the estimated EALB measurements from the different imagers from any one time
 546 to provide a single EPB estimate for that time, using an equatorward boundary version
 547 of Equation 4,

$$548 \lambda_{EPB}(\phi, t) = \sum_{i=1}^n \frac{\lambda_E(i, \phi, t) - L_E(i, \phi)}{n} \quad (5)$$

549 Again, for times and MLTs where boundaries from not all the imagers are available, we
 550 just use those that are.

Figure 9. The variation of the median offsets between the IMAGE FUV EALBs and the DMSP EPBs in each 1-hr MLT sector, for the SI12 (black), SI13 (orange), and WIC (red) imagers. The error bars present the quartile range of the offsets (separated slightly in MLT for clarity). The solid lines present the harmonic function fits to the median values for each imager.

551 A summary of the new EPB data base is presented in Figure 10. As with previ-
 552 ous figures, the black solid line shows the variation of the median EPB latitude with MLT.
 553 The thick black error bars at each MLT show the latitude range from the lower to the
 554 upper quartile of the EPB latitudes, whereas the thinner bars show the latitude range
 555 from the 10% to 90% percentiles. Again, the orange histogram represents the number
 556 of boundaries derived in each MLT sector across the whole measurement epoch. The most
 557 striking difference between the EPB and EALB variations is that the poleward excu-
 558 rsion of the EALBs around noon has been significantly reduced due to the offset between
 559 the two boundaries that is typically observed there (see Figure 9). The resulting median
 560 EPB variation is therefore more circular in shape.

Figure 10. Statistics of EPB locations derived from the IMAGE FUV EALBs in AACGM co-ordinates. The solid black line describes the MLT variation of the median EPB latitude. The thick black error bars present the quartile latitude variation at each MLT. The thin black error bars at each MLT present the extent of the distribution from the 10% to 90% percentile. The orange histograms present the number of EPB estimations in each MLT sector.

561 This data set is labelled EPB_v2 in the directory ‘raw_boundaries’ at [https://doi.org/10.5285/fa592594-](https://doi.org/10.5285/fa592594-93e0-4ee1-8268-b031ce21c3ca)
 562 93e0-4ee1-8268-b031ce21c3ca (Chisham, 2022). Those who wish to use estimates of the
 563 EPB determined from a single imager only rather than from the combined data sets can
 564 do so using the relevant EALB and the offsets described by Equation 3 and Table 2.

565 5 Fitted circle boundary methodology and data sets

566 Here, we present new circle fits for all the boundary data sets described above, in-
 567 cluding new improved criteria for helping to identify acceptable fits, an assessment of the
 568 suitability of circle fits for each data set, and a validation of the OCB circle fits.

569 5.1 Methodology

570 As discussed earlier, boundaries modelled as circles can be very useful in the de-
 571 velopment of high-latitude ionospheric climatologies and empirical models (Chisham, 2017a).
 572 Circles represent simple and practical model functions, and an easily scalable represen-
 573 tation of the boundary locations at all MLTs. These model circles also provide a method
 574 for filling gaps in those MLT sectors where data is missing, particularly on the dayside
 575 where dayglow is a major issue; fitting circles to the MLT variations of the observed bound-
 576 aries at any one time provides a model description of that boundary at all MLTs.

577 The circle fitting method employed here is the modified least squares method of
 578 Umbach and Jones (2003) as used by Chisham (2017a). This method does not require
 579 the location of the centre of the circle to be specified in advance. The fitted circles are
 580 subsequently defined by three parameters; two which define the co-ordinates of the cen-
 581 tre of the circle relative to the AACGM pole, and one that defines the circle radius. In
 582 the data sets, the fitted circles are described by the following three variables: r_{cent} is the
 583 co-latitude distance from the AACGM pole to the fitted circle centre; ϕ_{cent} is the MLT
 584 angle (MLT $\times 15^\circ$) of the line from the AACGM pole to the fitted circle centre; and r_{circ}
 585 is the radius of the fitted circle in co-latitude from the fitted circle centre.

586 The circle data sets are labelled PALB_si12_circle_v2, PALB_si13_circle_v2, PALB_wic_circle_v2,
 587 OCB_circle_v2, EALB_si12_circle_v2, EALB_si13_circle_v2, EALB_wic_circle_v2, and EPB_circle_v2,
 588 in the directory ‘circle_fit_boundaries’ at [https://doi.org/10.5285/fa592594-93e0-4ee1-](https://doi.org/10.5285/fa592594-93e0-4ee1-8268-b031ce21c3ca)
 589 8268-b031ce21c3ca (Chisham, 2022).

590 5.2 Circle fit acceptance criteria

591 In the original study of Chisham (2017a), criteria were developed to discard poor,
 592 unrealistic or erroneous circle fits of the PALBs and OCB. These criteria were derived
 593 from visual inspection of the probability distributions of the fitted parameters, and how
 594 these distributions varied with the number of 1-hr MLT sectors n where PALBs were con-
 595 tributing to the fit (from the minimum of $n = 4$ needed for a minimal circle fit to the
 596 maximum $n = 24$, when estimates were available at all MLTs). The criteria for accep-
 597 tance were set at $n > 7$, $r_{cent} \leq 8^\circ$, and $10^\circ < r_{circ} \leq 22^\circ$. (The last of these criteria
 598 is obviously only applicable for the poleward boundaries). However, as the PALB cir-
 599 cle centres are predominantly shifted anti-sunward of the AACGM pole, these criteria

600 failed to remove many potentially erroneous fits where the circle centres were shifted sig-
 601 nificantly sunward of the pole.

Figure 11. Contour map showing the level of occurrence of OCB fitted circle centre locations in AACGM co-ordinates for different numbers of MLT sectors contributing to the fit: (a) 8, (b) 11, (c) 14, (d) 17, (e) 20, and (f) 23. The red circles represent the chosen criteria circles within which the associated fitted circles are thought to be more reliable. The text outlines the number (and percentage) of fitted circle centres that are located inside and outside of the red criteria circle. The co-ordinate system is AACGM with 1200 MLT (noon) to the top of each panel.

602 In order to more effectively address this issue, here we recommend new criteria, based
 603 on the two-dimensional spatial distributions of the fitted circle centres, and how they vary
 604 with n . This is in addition to the $n > 7$ criterion. Figure 11 presents contour plots of
 605 the occurrence of the locations of the circle centres for the fitted circles derived from the
 606 OCB boundary data set for six different values of n (8, 11, 14, 17, 20, and 23), in AACGM
 607 co-ordinates. In each panel of the figure there is a preferred region for the location of the
 608 circle centres where the occurrence is high (dark orange/black region). The spread of circle
 609 centres outside of this region reduces with increasing n , suggesting that there is in-
 610 creasing uncertainty in the circle centre estimates away from this core region, especially
 611 for lower values of n .

612 To deal with the increased uncertainty associated with fitted circles with centres
 613 away from this core region, we define a circular region centred around this core of ob-
 614 servations (shown in red in Figure 11). We propose that fitted circles with centres out-
 615 side of this red ‘criteria circle’ should be viewed with suspicion. These criteria circles have
 616 been chosen for each separate boundary data set by visual inspection of the circle centre
 617 contour maps (such as those in Figure 11) for all values of n . The criteria circles are
 618 defined by the co-ordinates of their centre (x_{crit}, y_{crit}) on a Cartesian grid, where the
 619 x -axis is positive towards dawn (0600 MLT) and the y -axis is positive towards noon (1200
 620 MLT), and the radius of the criteria circle (r_{crit}). The Cartesian values of the fitted circle
 621 centres (x_{cent}, y_{cent}) are determined from r_{cent} and ϕ_{cent} using:

$$622 \quad x_{cent} = r_{cent} \sin \phi_{cent} \quad (6)$$

623 and

$$624 \quad y_{cent} = -r_{cent} \cos \phi_{cent}. \quad (7)$$

625 To be accepted, we define a merit value r_{merit} for each fitted circle that must sat-
 626 isfy:

$$627 \quad r_{merit}^2 = (x_{cent} - x_{crit})^2 + (y_{cent} - y_{crit})^2 \leq r_{crit}^2. \quad (8)$$

628 The value of r_{merit} for each circle fit also forms part of the data set. Each panel in Fig-
 629 ure 11 details the number (and percentage) of fitted circle centres that are located both
 630 inside ($r_{merit} < r_{crit}$), and outside ($r_{merit} > r_{crit}$), of the criteria circle. The percent-
 631 age of values located outside of the criteria circle reduces with increasing n .

632 The values of x_{crit} , y_{crit} , and r_{crit} estimated for each of the boundary data sets are
 633 presented in Table 3. The value of r_{crit} is only a recommendation and can be adjusted
 634 to suit the needs of the data user.

635 5.3 Assessment of suitability of circle fits

636 The study of Chisham (2017a) contained no assessment of how apt a circle fit was
 637 for the observed PALBs. Although the auroral oval is roughly circular when viewed in

Table 3. Recommended values of the parameters describing the criteria circles to help assess the fitted circle reliability for each data set. (see text for details).

Boundary	x_{crit}	y_{crit}	r_{crit}
SI12 Poleward	-1.5	-3.5	5.0
SI12 Equatorward	-1.0	-3.5	5.0
SI13 Poleward	0.0	-2.5	5.0
SI13 Equatorward	0.0	-5.0	5.0
WIC Poleward	0.0	-3.5	5.0
WIC Equatorward	1.0	-4.0	5.0
OCB	-0.5	-2.5	5.0
EPB	1.0	-2.5	5.0

638 geomagnetic co-ordinates, there are particular conditions where the poleward boundary
 639 is more elliptical than circular. In addition, the equatorward boundary typically diverges
 640 significantly from circular on the dayside. Hence, we accept that the circles should gener-
 641 ally be viewed as a simple first-order model of the boundaries. However, their simplic-
 642 ity makes them particularly useful for modelling work.

643 Here, we present a statistical assessment of the suitability of circle model fits for
 644 each of the boundary data sets, in order to provide a level of confidence in the model fits.
 645 Figure 12 presents, as black square symbols, the MLT variation of the mean difference
 646 between the fitted circle locations and the corresponding measured (in the case of the
 647 PALBs and EALBs) or derived (in the case of OCBs and EPBs) boundary locations. The
 648 error bars show the standard deviation of the boundary differences in each MLT sector.
 649 Hence, the dotted line at 0° represents the location of the measured/derived boundaries.
 650 The orange region defines where the differences are less than 1° , which we assess as be-
 651 ing an acceptable range for the mean offsets, within which the circle model fits are typ-
 652 ically a good assumption.

Figure 12. Comparison of circle fit locations with measured/derived boundary locations.
 Each panel is labelled with the corresponding boundary data set name. The black symbols and
 error bars represent the mean and standard deviation of the fitted circle latitudes from the mea-
 sured/derived boundary latitudes in each MLT sector, respectively. Hence, the dotted lines at
 0° offset represent the relative location of the measured/derived boundary. The orange shading
 defines the region within which the mean offset is less than 1° latitude.

653 Figure 12 shows that circle fits are particularly suitable for modelling the PALBs,
 654 particularly the SI12 and SI13 PALBs, for which the mean offsets are very close to 0°
 655 at all MLTs, and for which the standard deviations are approximately 1° . The match
 656 for the WIC PALB and the OCB are not as good, particularly around noon. However,
 657 all the mean offsets are still less than 1° , with most being less than 0.5° . The circle model
 658 fits for the EALBs are not as good, with the fitted circle locations typically lying $\sim 1^\circ$
 659 equatorward of the measured boundary locations around noon. This is as expected from
 660 the typical shape of the equatorward boundary observed on the dayside. This suggests
 661 that future studies should consider whether there is a more suitable function for mod-
 662 elling the shape of the EALBs. However, as a result of the large offset between the EALBs
 663 and the EPB reported in Figure 9, the EPB variation is well modelled by a circle, with
 664 almost all the mean offsets being less than 0.5° . Overall, we suggest that for most mod-

665 elling or analysis purposes, that the model circle fits provide a good representation of
 666 the boundaries.

667 5.4 Validation of poleward boundaries

668 The appropriateness of the new PALB and OCB circle fits is further assessed by
 669 comparing them to the Convection Reversal Boundary (CRB) identified by the plasma
 670 convection flow taken from DMSP (Y.-J. Chen et al., 2015; Y.-J. Chen & Heelis, 2018),
 671 as was done for the previous IMAGE data base by Burrell et al. (2020a). This valida-
 672 tion is limited to MLTs within 0500-0700 MLT and 1700-1900 MLT to avoid the influ-
 673 ence that the IMF has on the plasma convection pattern and the flow reversal bound-
 674 ary locations. The most significant influence is in the region close to the local noon and
 675 midnight sectors where magnetic reconnection occurs (see Burrell et al. (2020a) for de-
 676 tails). The validation paired the combined (OCB) and individual instrument (PALB)
 677 circle fit latitudes with r_{merit} below 5.0° with CRB values identified within 60 s of UT
 678 of the OCB/PALB observations. Figure 13 presents the locations of all the DMSP CRBs
 679 relative to the circle fit OCBs/PALBs (positioned in the figure at a nominal location of
 680 74°), while the differences between the two boundaries within 1 h of dawn and dusk are
 681 shown in Figure 14.

Figure 13. DMSP CRB locations relative to the IMAGE OCB and PALB circle fits (set at 74° , marked by an orange circle) for satellite F13 (top row, navy points) and F15 (bottom row, light blue points). The first column shows the paired CRB and combined OCB data set, while the subsequent columns show the paired boundaries for the SI12, SI13, and WIC PALBs.

Figure 14. Histogram of differences between the DMSP CRB locations from both F13 and F15 and the OCB/PALB for times within 05:00-07:00 MLT and 17:00-19:00 MLT. The first column shows the histogram and key statistics for the combined OCB data set, while the subsequent columns show the histograms and key statistics for the SI12, SI13, and WIC PALBs.

682 With the exception of the WIC PALB histogram (which does not have enough points
 683 to draw statistically significant conclusions from) the combined OCB, SI12 PALB, and
 684 SI13 PALB histograms all have means and medians that are roughly equal with simi-
 685 lar standard deviations. The combined OCB and SI12 PALB histograms are centered
 686 very close to zero, with no notable skew or kurtosis. The SI13 PALB histogram, how-
 687 ever, has a small negative offset and a negative skew. This is to be expected given the
 688 observed median offsets between the SI13 PALB and the OCB as shown in Figure 5.

689 Comparing these results to those presented in Burrell et al. (2020a) shows that the
 690 new circle fits and their selection criteria result in less pairings for SI12 and SI13 (the
 691 old boundaries had 744 pairings for SI12 and 434 pairings for SI13). However, the statis-
 692 tics are similar; SI12 still has a positive central bias while SI13 has a negative central
 693 bias. The mean and median values for the new SI12 and SI13 boundary differences are
 694 higher than for the old boundary differences, though not by a significant amount (pre-
 695 viously means were 0.32° for SI12 and -0.13° for SI13). However, the Burrell et al. (2020a)
 696 analysis used PALBs adjusted by the derived offset to the OCB in the comparison, whereas
 697 here we use the unadjusted PALBs.

698 The new combined OCB data set shows the best agreement with the DMSP CRBs
 699 when compared to the new individual PALBs and the old data set. It has the largest num-

700 ber of boundary pairings, with the central statistics closest to zero, and the smallest sig-
 701 nificant standard deviation. This improved agreement provides confidence in the com-
 702 bined OCB data set, which now matches the performance of the AMPERE OCB proxy
 703 (where the mean difference was -0.01° in the northern hemisphere and -0.02° in the south-
 704 ern hemisphere).

705 6 Summary and Conclusions

706 This study presents improved poleward and equatorward auroral luminosity bound-
 707 ary estimates from the IMAGE FUV instruments, as well as new estimates of the OCB
 708 and EPB. The newly derived OCBs and EPBs provided here have benefited from im-
 709 proved data selection when determining the IMAGE PALBs and EALBs, respectively.
 710 New data bases of model circle fits to the MLT variation of the measured/derived bound-
 711 aries have benefited from a new figure of merit (r_{merit}), that helps to assess the relia-
 712 bility of the fit. The new circle fit OCBs and PALBs have been validated against an in-
 713 dependent polar cap boundary proxy. The results of this validation show that the com-
 714 bined OCB data set performs better than previous IMAGE boundary data sets, and matches
 715 the performance of other existing OCB proxies.

716 Modelling and statistical studies of ionospheric, thermospheric, and magnetospheric
 717 processes should avoid mixing data from inside the polar cap, within the auroral oval,
 718 and equatorward of the auroral oval when determining averages, typical values, or typ-
 719 ical conditions. The boundaries provided here and the software that comprises OCBpy
 720 provide a set of tools that could be used to improve studies in these regions as well as
 721 developing tools for space weather modelling. Future missions with similar imaging ca-
 722 pabilities (such as SMILE), could contribute in this area by applying the methods pre-
 723 sented here to data from their own instruments.

724 Acknowledgments

725 GC was supported by the British Antarctic Survey Polar Science for Planet Earth Pro-
 726 gramme, funded by the UK Natural Environment Research Council (NERC) as part of
 727 UK Research and Innovation (UKRI) as well as the UK NERC directed grant NE/V002732/1.
 728 AGB was supported by the Office of Naval Research. EGT was supported by the Na-
 729 tional Science Foundation (NSF) under grant AGS-1934997. YJC was supported by AFOSR
 730 MURI grant (FA9559-16-1-0364) and NASA LWS grant (80NSSC20K177). The authors
 731 would like to thank those involved in producing the original Boakes and Longden bound-
 732 ary data sets from the IMAGE data: Peter Boakes, Nicola Longden, Mervyn Freeman,
 733 Steve Milan, and Gary Abel. The derivation of the original Longden boundaries was funded
 734 by UK STFC grant PP/E002110/1.

735 The new boundaries derived here can be found at <https://doi.org/10.5285/fa592594-93e0-4ee1-8268-b031ce21c3ca> (Chisham, 2022). The authors would like to thank the NASA
 736 Space Physics Data Facility and National Space Science Data Center. The IMAGE FUV
 737 data were provided courtesy of the instrument PI Stephen Mende (University of Cali-
 738 fornia, Berkeley). We thank the PI, the IMAGE mission, and the IMAGE FUV team
 739 for data usage and processing tools. The raw image data, and software, were acquired
 740 from <http://sprg.ssl.berkeley.edu/image/>. We thank Tom Sotirelis (Johns Hopkins Uni-
 741 versity, Applied Physics Laboratory) for providing the DMSP boundary data base.
 742

743 References

- 744 Aikio, A. T., Pitkanen, T., Kozlovsky, A., & Amm, O. (2006). Method to locate the
 745 polar cap boundary in the nightside ionosphere and application to a substorm
 746 event. *Ann. Geophys.*, *24*, 1905–1917.
 747 Anderson, B. J., Korth, H., Waters, C. L., Green, D. L., Merkin, V. G., Barnes,

- 748 R. J., & Dyrud, L. P. (2014). Development of large-scale Birkeland currents
749 determined from Active Magnetosphere and Planetary Electrodynamics Re-
750 sponse Experiment. *Geophys. Res. Lett.*, *41*, 3017–3025.
- 751 Baker, J. B., Clauer, C. R., Ridley, A. J., Papitashvili, V. O., Brittnacher, M. J.,
752 & Newell, P. T. (2000). The nightside poleward boundary of the auroral
753 oval as seen by DMSP and the Ultraviolet Imager. *J. Geophys. Res.*, *105*,
754 21267–21280.
- 755 Blanchard, G. T., Ellington, C. L., Lyons, L. R., & Rich, F. J. (2001). Incoherent
756 scatter radar identification of the dayside magnetic separatrix and measure-
757 ment of magnetic reconnection. *J. Geophys. Res.*, *106*, 8185–8195.
- 758 Blanchard, G. T., Lyons, L. R., Samson, J. C., & Rich, F. J. (1995). Locating the
759 polar cap boundary from observations of 6300 Å auroral emission. *J. Geophys.*
760 *Res.*, *100*, 7855–7862.
- 761 Boakes, P. D., Milan, S. E., Abel, G. A., Freeman, M. P., Chisham, G., & Hubert,
762 B. (2009). A statistical study of the open magnetic flux content of the magne-
763 tosphere at the time of substorm onset. *Geophys. Res. Lett.*, *36*, L04105. doi:
764 10.1029/2008GL037059
- 765 Boakes, P. D., Milan, S. E., Abel, G. A., Freeman, M. P., Chisham, G., & Hu-
766 bert, B. (2011). A superposed epoch investigation of the relation between
767 magnetospheric solar wind driving and substorm dynamics with geosyn-
768 chronous particle injection signatures. *J. Geophys. Res.*, *116*, A01214. doi:
769 10.1029/2010JA016007
- 770 Boakes, P. D., Milan, S. E., Abel, G. A., Freeman, M. P., Chisham, G., Hubert, B.,
771 & Sotirelis, T. (2008). On the use of IMAGE FUV for estimating the lati-
772 tude of the open/closed magnetic field line boundary in the ionosphere. *Ann.*
773 *Geophys.*, *26*, 2759–2769.
- 774 Burch, J. L. (2000). IMAGE mission overview. *Space Sci. Rev.*, *91*, 1–14.
- 775 Burrell, A. G., Chisham, G., Milan, S. E., Kilcommons, L., Chen, Y.-J., Thomas,
776 E. G., & Anderson, B. (2020a). AMPERE polar cap boundaries. *Annales*
777 *Geophysicae*, *38*. doi: 10.5194/angeo-38-481-2020
- 778 Burrell, A. G., Chisham, G., & Reistad, J. (2020b). *aburrell/ocbpy: v0.2.1 (0.2.1)*.
779 Zenodo. doi: 10.5281/zenodo.4289226
- 780 Carbary, J. F., Sotirelis, T., Newell, P. T., & Meng, C.-I. (2003). Auroral bound-
781 ary correlations between UVI and DMSP. *J. Geophys. Res.*, *108*(1018). doi: 10
782 .1029/2002JA009378
- 783 Chen, X.-C., Han, D.-S., Lorentzen, D. A., Oksavik, K., Moen, J. I., & Baddeley,
784 L. J. (2017). Dynamic properties of throat aurora revealed by simultaneous
785 ground and satellite observations. *J. Geophys. Res.*, *122*, 3469–3486.
- 786 Chen, Y.-J., & Heelis, R. A. (2018). Motions of the convection reversal boundary
787 and local plasma in the high-latitude ionosphere. *Journal of Geophysical Re-*
788 *search: Space Physics*, *123*, 2953–2963. doi: 10.1002/2017ja024934
- 789 Chen, Y.-J., Heelis, R. A., & Cumnock, J. A. (2015). Response of the ionospheric
790 convection reversal boundary at high latitudes to changes in the interplane-
791 tary magnetic field. *Journal of Geophysical Research: Space Physics*, *120*,
792 5022–5034. doi: 10.1002/2015ja021024
- 793 Chisham, G. (2017a). A new methodology for the development of high-latitude iono-
794 spheric climatologies and empirical models. *J. Geophys. Res.*, *122*, 932–947.
- 795 Chisham, G. (2017b). *Auroral boundary derived from image satellite mission data*
796 *(May 2000 - October 2002), version 1.1 [Dataset]*. Polar Data Centre, Nat-
797 ural Environment Research Council. doi: 10.5285/75aa66c1-47b4-4344-ab5d
798 -52ff2913a61e
- 799 Chisham, G. (2022). *Ionospheric boundaries derived from IMAGE satellite mission*
800 *data (May 2000 - October 2002), version 2.0 [Dataset]*. NERC EDS UK Polar
801 Data Centre. doi: 10.5285/fa592594-93e0-4ee1-8268-b031ce21c3ca
- 802 Chisham, G., & Freeman, M. P. (2003). A technique for accurately determining the

- 803 cusp-region polar cap boundary using SuperDARN HF radar measurements.
804 *Ann. Geophys.*, *21*, 983–996.
- 805 Chisham, G., Freeman, M. P., Abel, G. A., Lam, M. M., Pinnock, M., Coleman,
806 I. J., ... Villain, J.-P. (2008). Remote sensing of the spatial and temporal
807 structure of magnetopause and magnetotail reconnection from the ionosphere.
808 *Rev. Geophys.*, *46*(RG1004). doi: 10.1029/2007RG000223
- 809 Chisham, G., Lester, M., Milan, S. E., Freeman, M. P., Bristow, W. A., Grocott,
810 A., ... Walker, A. D. M. (2007). A decade of the Super Dual Auroral Radar
811 Network (SuperDARN): Scientific achievements, new techniques and future
812 directions. *Surv. Geophys.*, *28*, 33–109.
- 813 Cowley, S., & Lockwood, M. (1992). Excitation and decay of solar wind-driven flows
814 in the magnetosphere–ionosphere system. *Ann. Geophys.*, *10*, 103–115.
- 815 Cowley, S. W. H. (2000). Magnetosphere–ionosphere interactions: a tutorial re-
816 view. In S. Ohtani (Ed.), *Magnetospheric current systems* (Vol. 118, pp. 91–
817 106). American Geophysical Union. doi: 10.1029/GM118p0091.
- 818 Coxon, J. C., Milan, S. E., & Anderson, B. J. (2018). A review of Birke-
819 land current research using AMPERE. In *Electric currents in geospace*
820 *and beyond* (Vol. 235, pp. 259–278). American Geophysical Union. doi:
821 10.1002/9781119324522.ch16.
- 822 Dungey, J. W. (1961). Interplanetary magnetic field and the auroral zones. *Physical*
823 *Review Letters*, *6*, 47–48.
- 824 Frey, H. U., Mende, S. B., Angelopoulos, V., & Donovan, E. F. (2004). Substorm
825 onset observations by IMAGE-FUV. *J. Geophys. Res.*, *109*, A10304. doi: 10
826 .1029/2004JA010607
- 827 Gibson, W. C., Burch, J. L., Scherrer, J. R., Tapley, M. B., Killough, R. L., Volpe,
828 F. A., ... Hosten, S. J. (2000). The IMAGE observatory. *Space Sci. Rev.*, *91*,
829 15–50.
- 830 Gjerloev, J. (2012). The SuperMAG data processing technique. *J. Geophys. Res.*,
831 *117*, A09213.
- 832 Greenwald, R. A., Shepherd, S. G., Sotirelis, T. S., Ruohoniemi, J. M., & Barnes,
833 R. J. (2002). Dawn and dusk sector comparisons of small-scale irregularities,
834 convection, and particle precipitation in the high-latitude ionosphere. *Journal*
835 *of Geophysical Research: Space Physics*, *107*. doi: 10.1029/2001JA000158
- 836 Hardy, D. A., Schmitt, L. K., Gussenhoven, M. S., Marshall, F. J., Yeh, H. C., Shu-
837 maker, T. L., ... Pantazis, J. (1984). Precipitating electron and ion detectors
838 (SSJ/4) for the block 5D/flights 6-10 DMSP satellites: Calibration and data
839 presentation. Rep. AFGL-TR-84-0317, Air Force Geophys. Lab., Hanscom
840 AFB, Mass.
- 841 Heelis, R. A., & Hanson, W. B. (1998). Measurements of thermal ion drift ve-
842 locity and temperature using planar sensors. In R. F. Pfaff, J. Borovsky,
843 & T. D. Young (Eds.), *Measurement techniques in space plasmas: Particles*
844 (Vol. 102, pp. 61–71). American Geophysical Union.
- 845 Hoque, S. N. M. A., Alvi, A., Mahmud, M. R., Sakhawat, S., & Bhuiyan, M. I.
846 (2021). On the use of directional derivatives of SuperDARN convection flow
847 for estimating open-closed field line boundary during the intervals of substorm
848 activity on 14 october 2001. *Polar Sci.*, *28*, 100655.
- 849 Hubert, B., Aikio, A. T., Amm, O., Pitkanen, T., Kauristie, K., Milan, S. E., ...
850 Gerard, J.-C. (2010). Comparison of the open-closed field line boundary loca-
851 tion inferred using IMAGE-FUV SI12 images and EISCAT radar observations.
852 *Ann. Geophys.*, *28*, 883–892.
- 853 Hubert, B., Milan, S. E., Grocott, A., Blockx, C., Cowley, S. W. H., & Gerard, J.-C.
854 (2006). Dayside and nightside reconnection rates inferred from IMAGE-FUV
855 and Super Dual Auroral Radar Network data. *J. Geophys. Res.*, *111*, A03217.
- 856 Kauristie, K., Weygand, J., Pulkkinen, T. I., Murphree, J. S., & Newell, P. T.
857 (1999). Size of the auroral oval: UV ovals and precipitation boundaries com-

- 858 pared. *J. Geophys. Res.*, *104*, 2321–2331.
- 859 Kilcommons, L. M., Redmon, R. J., & Knipp, D. J. (2017). A new DMSP magne-
860 tometer and auroral boundary data set and estimates of field-aligned currents
861 in dynamic auroral boundary coordinates. *Journal of Geophysical Research:*
862 *Space Physics*, *122*, 9068–9079. doi: 10.1002/2016ja023342
- 863 Kivelson, M. G., Russell, C. T., & Brown, M. E. (1996). Introduction to space
864 physics. *Physics Today*, *49*. doi: 10.1063/1.2807586
- 865 Lockwood, M. (1998). Identifying the open-closed field line boundary. In J. Moen
866 (Ed.), *Polar cap boundary phenomena* (pp. 73–90). Kluwer Academic Publish-
867 ers.
- 868 Longden, N., Chisham, G., & Freeman, M. P. (2014). Magnetic local time vari-
869 ation and scaling of poleward auroral boundary dynamics. *J. Geophys. Res.*,
870 *119*, 10006–10022.
- 871 Longden, N., Chisham, G., Freeman, M. P., Abel, G. A., & Sotirelis, T. (2010).
872 Estimating the location of the open-closed magnetic field line boundary from
873 auroral images. *Ann. Geophys.*, *28*, 1659–1678.
- 874 Mende, S., Heeterdks, H., Frey, H., Lampton, M., Geller, S., Abiad, R., ... Trond-
875 sen, T. (2000b). Far ultraviolet imaging from the IMAGE spacecraft. 2.
876 wideband FUV imaging. *Space Sci. Rev.*, *91*, 271–285.
- 877 Mende, S., Heeterdks, H., Frey, H., Lampton, M., Geller, S., Habraken, S., ... Cog-
878 ger, L. (2000a). Far ultraviolet imaging from the IMAGE spacecraft. 1. system
879 design. *Space Sci. Rev.*, *91*, 243–270.
- 880 Mende, S., Heeterdks, H., Frey, H., Stock, J., Lampton, M., Geller, S., ... Lauche,
881 H. (2000c). Far ultraviolet imaging from the IMAGE spacecraft. 3. spectral
882 imaging of the Lyman- α and OI 135.6 nm. *Space Sci. Rev.*, *91*, 287–318.
- 883 Milan, S. E. (2009). Both solar wind-magnetosphere coupling and ring current in-
884 tensity control of the size of the auroral oval. *Geophys. Res. Lett.*, *36*, L18101.
885 doi: 10.1029/2009GL039997
- 886 Milan, S. E., Clausen, L. B. N., Coxon, J. C., Carter, J. A., Walach, M.-
887 T., Laundal, K., ... Anderson, B. J. (2017). Overview of solar
888 wind-magnetosphere-ionosphere-atmosphere coupling and the generation
889 of magnetospheric currents. *Space Sci. Rev.*, *206*, 547–573.
- 890 Milan, S. E., Gosling, J. S., & Hubert, B. (2012). Relationship between interplan-
891 etary parameters and the magnetopause reconnection rate quantified from
892 observations of the expanding polar cap. *J. Geophys. Res.*, *117*, A03226. doi:
893 10.1029/2011JA017082
- 894 Milan, S. E., Hutchinson, J., Boakes, P. D., & Hubert, B. (2009). Influences on the
895 radius of the auroral oval. *Ann. Geophys.*, *27*, 2913–2924.
- 896 Milan, S. E., Lester, M., Cowley, S. W. H., Oksavik, K., Brittnacher, M., Greenwald,
897 R. A., ... Villain, J.-P. (2003). Variations in the polar cap area during two
898 substorm cycles. *Ann. Geophys.*, *21*, 1121–1140.
- 899 Mooney, M. K., Forsyth, C., Rae, I. J., Chisham, G., Coxon, J. C., Marsh, M. S., ...
900 Hubert, B. (2020). Examining local time variations in the gains and losses of
901 open magnetic flux during substorms. *J. Geophys. Res.*, *125*(e2019JA027369).
- 902 Mooney, M. K., Marsh, M. S., Forsyth, C., Sharpe, C., Hughes, T., Bingham, S., ...
903 Chisham, G. (2021). Evaluating auroral forecasts against satellite observations.
904 *Space Weather*, *19*(e2020SW002688).
- 905 Newell, P. T., Burke, W. J., Sanchez, E. R., Meng, C.-I., Greenspan, M. E., &
906 Clauer, C. R. (1991). The low-latitude boundary layer and the boundary
907 plasma sheet at low altitude: Prenoon precipitation regions and convection
908 reversal boundaries. *J. Geophys. Res.*, *96*, 21013–21023.
- 909 Newell, P. T., Feldstein, Y. I., Galperin, Y. I., & Meng, C.-I. (1996). Morphology of
910 nightside precipitation. *J. Geophys. Res.*, *101*, 10737–10748.
- 911 Newell, P. T., Liou, K., Zhang, Y., Sotirelis, T., Paxton, L. J., & Mitchell, E. J.
912 (2014). Ovation prime-2013: Extension of auroral precipitation model to

- 913 higher disturbance levels. *Space Weather*, *12*. doi: 10.1002/2014SW001056
914 Newell, P. T., Ruohoniemi, J. M., & Meng, C.-I. (2004). Maps of precipitation
915 by source region, binned by IMF, with inertial convection streamlines. *J. Geo-*
916 *phys. Res.*, *109*, A10206. doi: 10.1029/2004JA010499
- 917 Redmon, R. J., Peterson, W. K., Andersson, L., Kihn, E. A., Denig, W. F.,
918 Hairston, M., & Coley, R. (2010). Vertical thermal O⁺ flows at 850 km in
919 dynamic auroral boundary coordinates. *Journal of Geophysical Research*, *115*,
920 A00J08. doi: 10.1029/2010JA015589
- 921 Shepherd, S. G. (2014). Altitude-adjusted corrected geomagnetic coordinates: Defi-
922 nition and functional approximations. *J. Geophys. Res.*, *119*, 7501–7521. doi:
923 10.1002/2014JA020264
- 924 Siscoe, G. L., & Huang, T. S. (1985). Polar cap inflation and deflation. *J. Geophys.*
925 *Res.*, *90*, 543–547.
- 926 Sotirelis, T., & Newell, P. T. (2000). Boundary-oriented electron precipitation
927 model. *J. Geophys. Res.*, *105*, 18655–18673.
- 928 Umbach, D., & Jones, K. N. (2003). A few methods for fitting circles to data. *IEEE*
929 *Trans. Instrum. Meas.*, *52*, 1881–1885.
- 930 Zhu, Q., Deng, Y., Richmond, A., Maute, A., Chen, Y.-J., Hairston, M., . . .
931 Mitchell, E. (2020). Impacts of binning methods on high-latitude electro-
932 dynamic forcing: static vs. boundary-oriented binning methods. *J. Geophys.*
933 *Res.*, *125*, e2019JA027270. doi: 10.1029/2019JA027270

Figure 1.

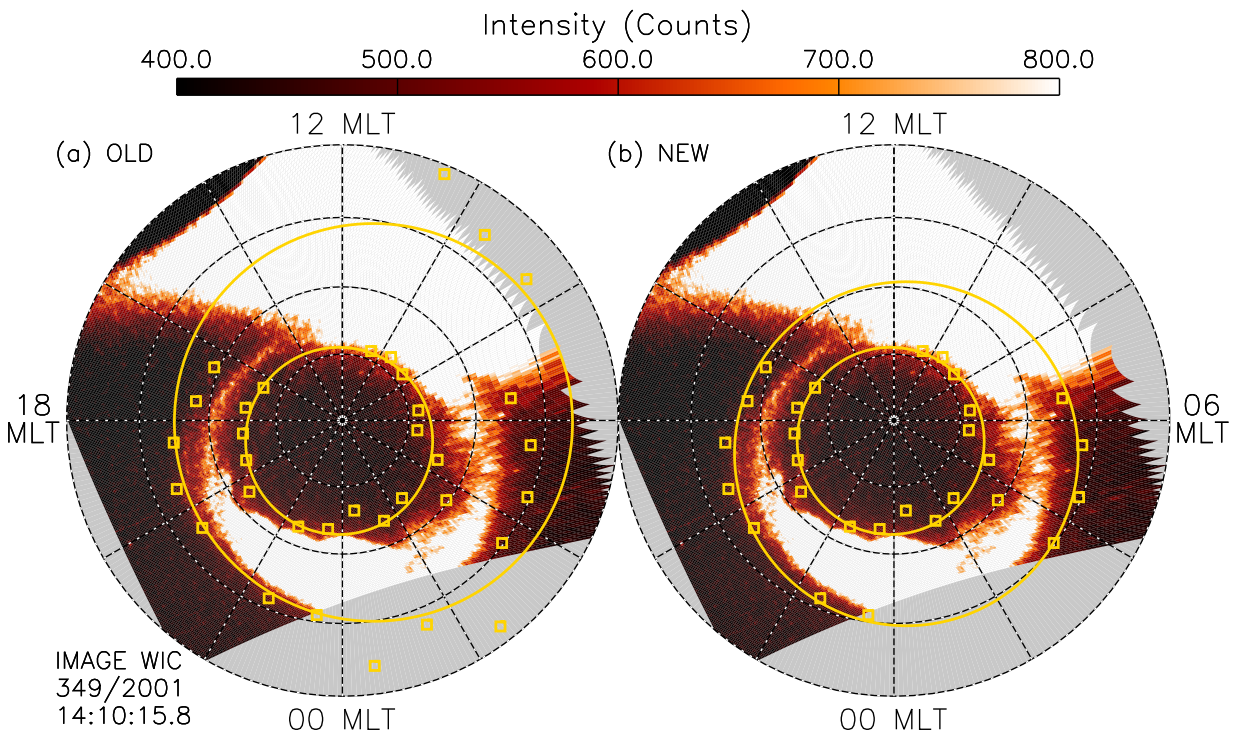


Figure 2.

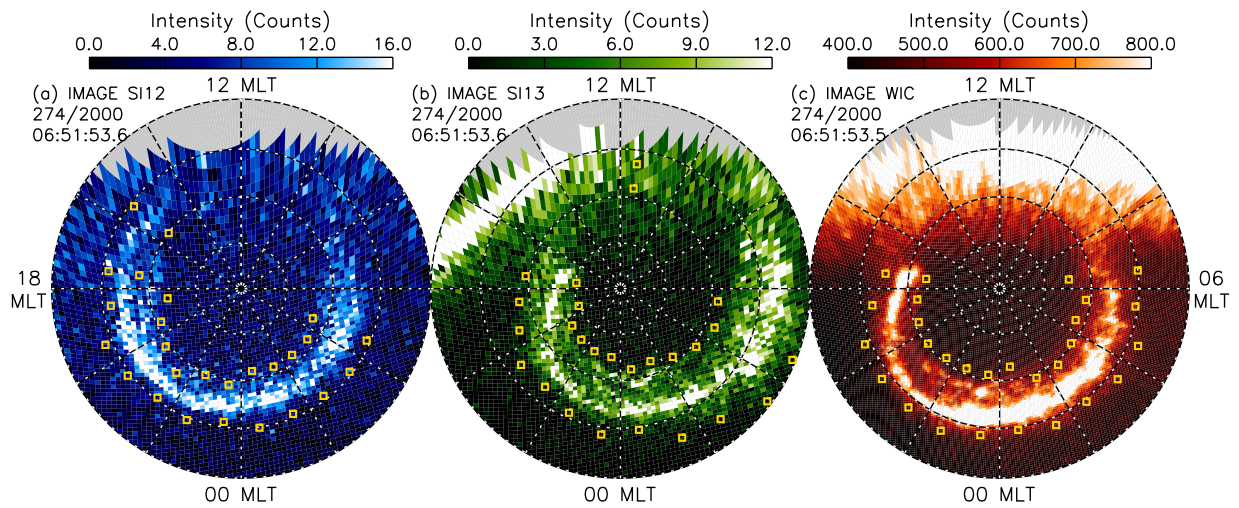
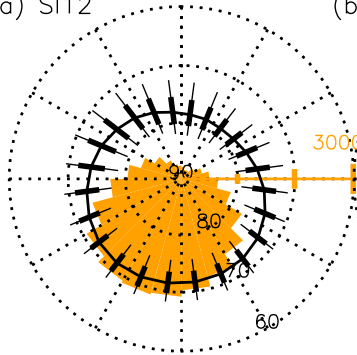
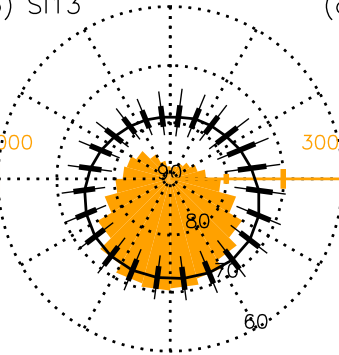


Figure 3.

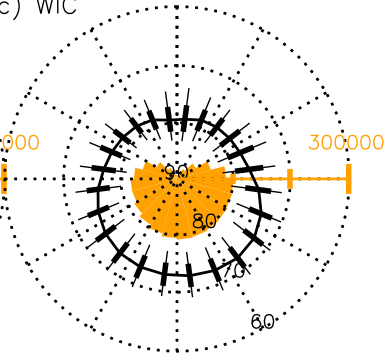
(a) SI12



(b) SI13



(c) WIC



300000

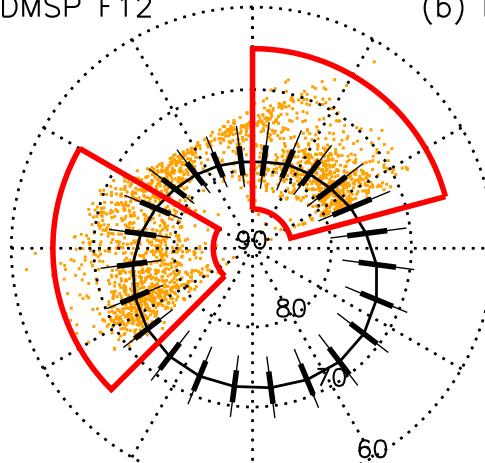
300000

300000

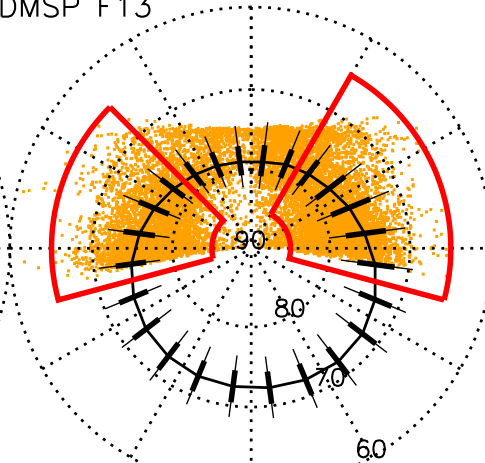
Figure 4.

Poleward Boundaries

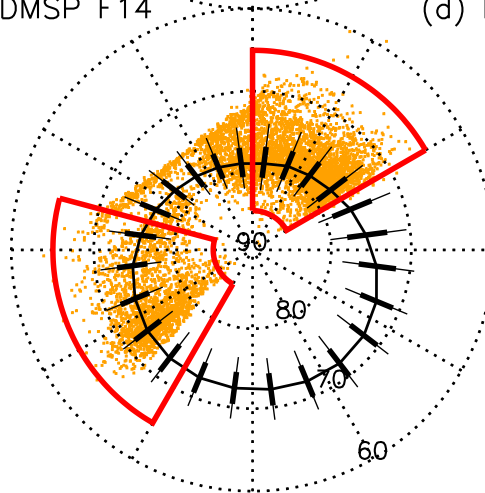
(a) DMSP F12



(b) DMSP F13



(c) DMSP F14



(d) DMSP F15

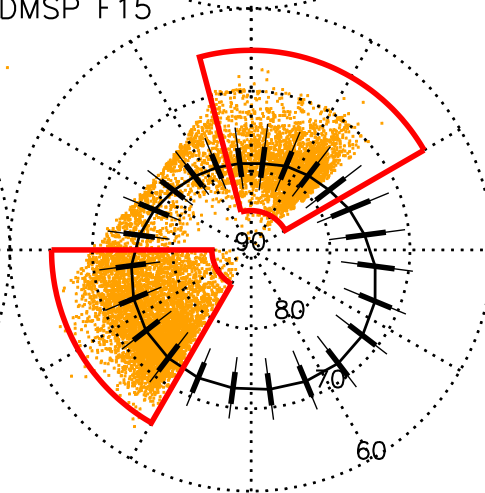


Figure 5.

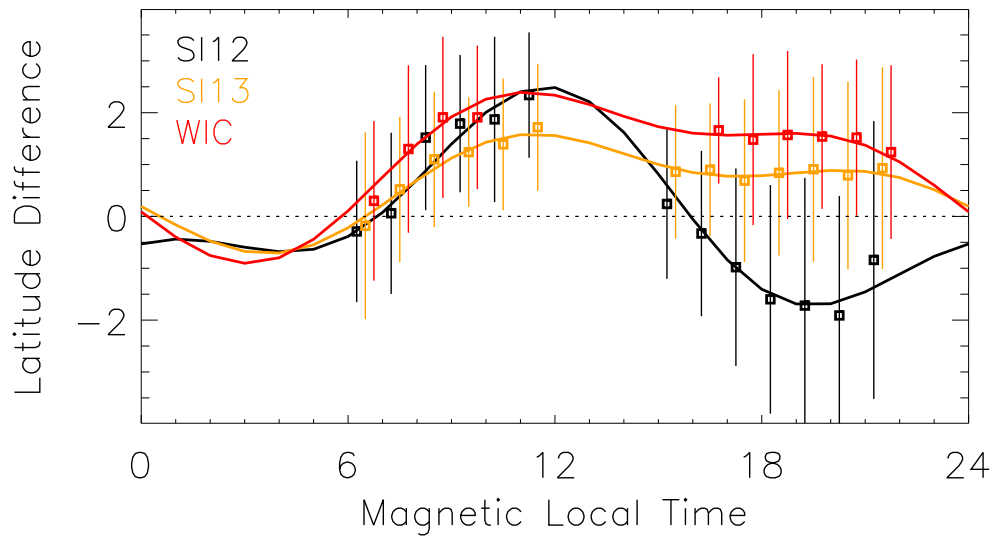


Figure 6.

OCB

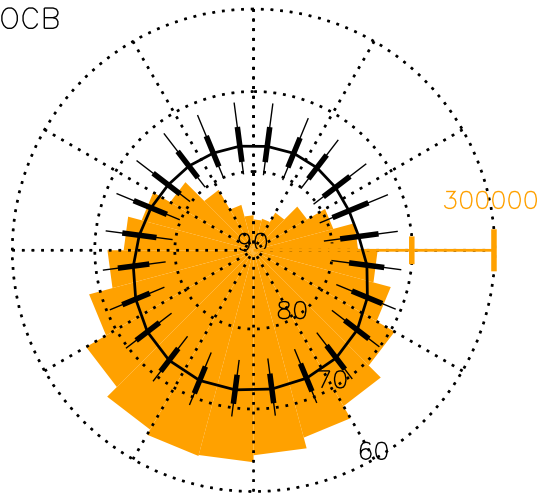
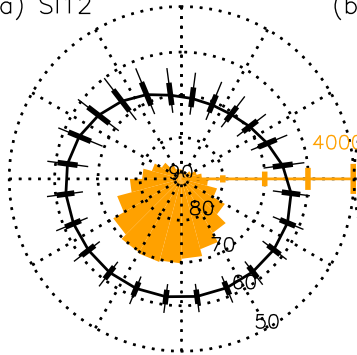
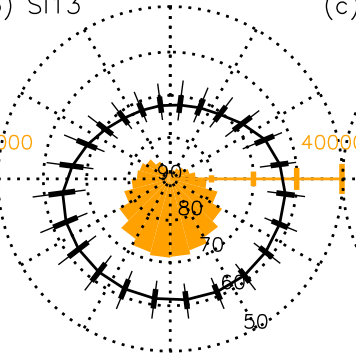


Figure 7.

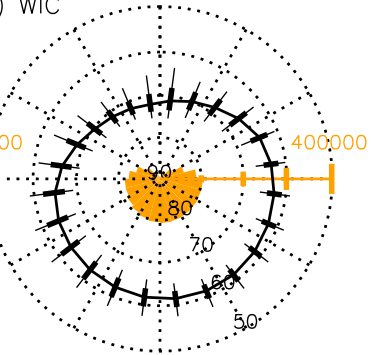
(a) SI12



(b) SI13



(c) WIC



400,000

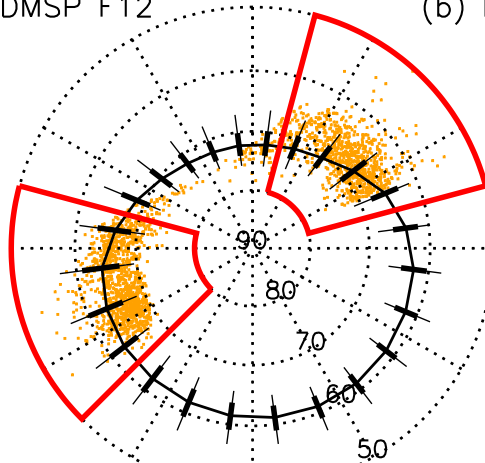
400,000

400,000

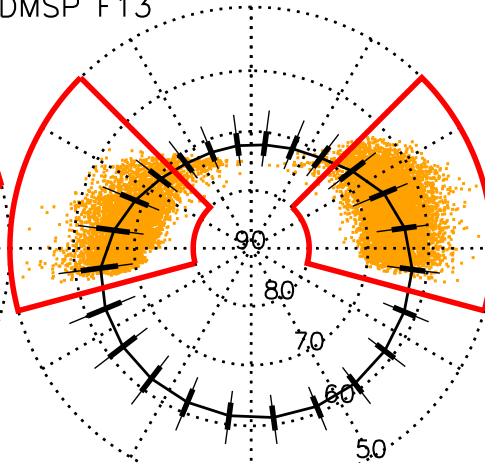
Figure 8.

Equatorward Boundaries

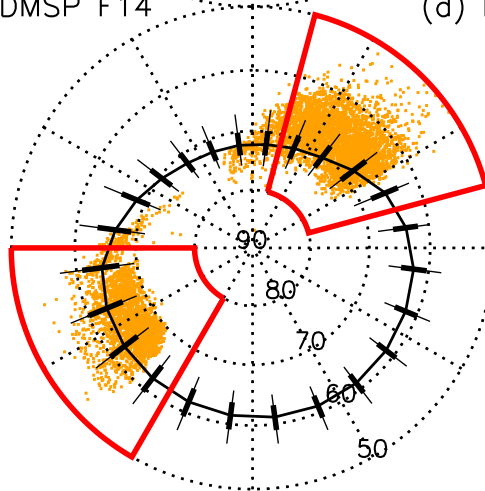
(a) DMSP F12



(b) DMSP F13



(c) DMSP F14



(d) DMSP F15

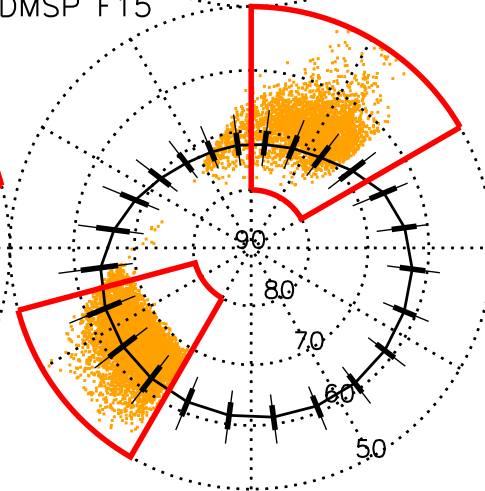


Figure 9.

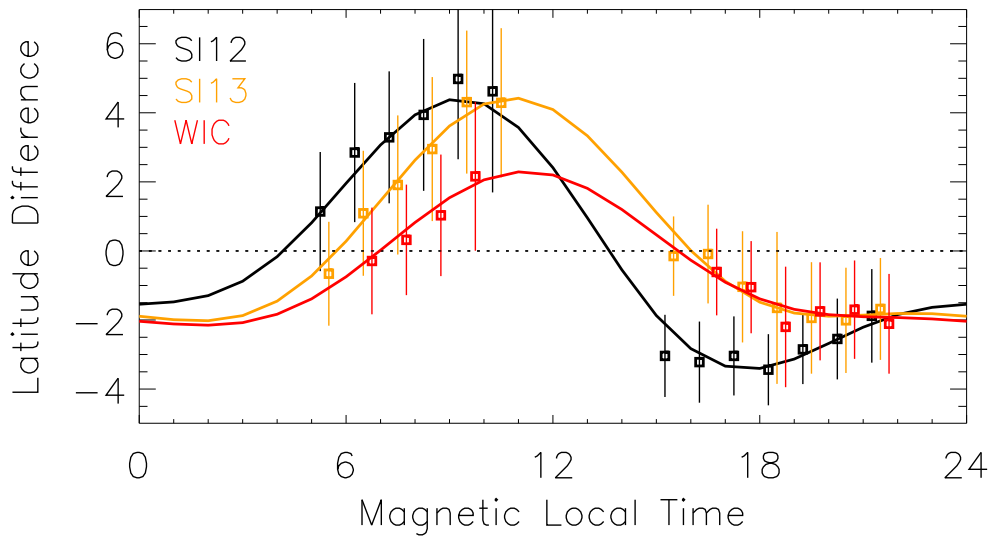


Figure 10.

EPB

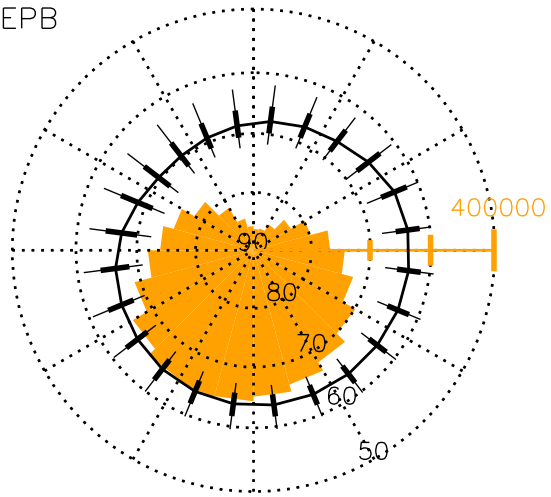


Figure 11.

IMAGE FUV-derived OCB
Fitted circle centre locations

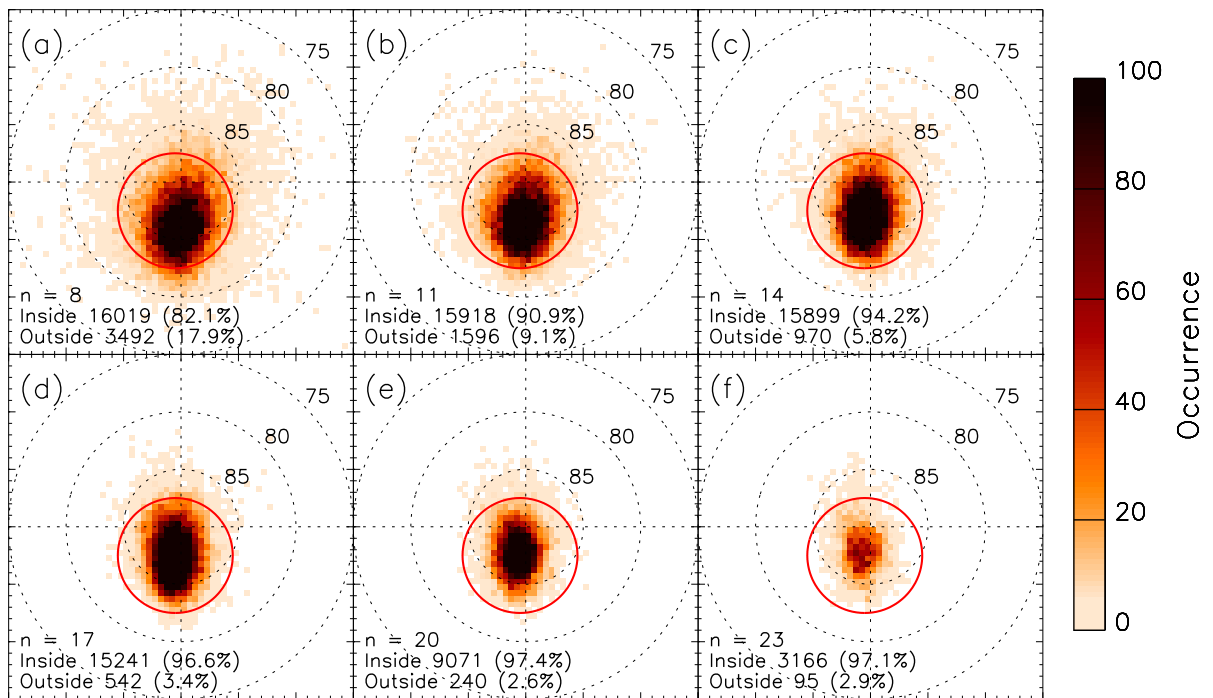


Figure 12.

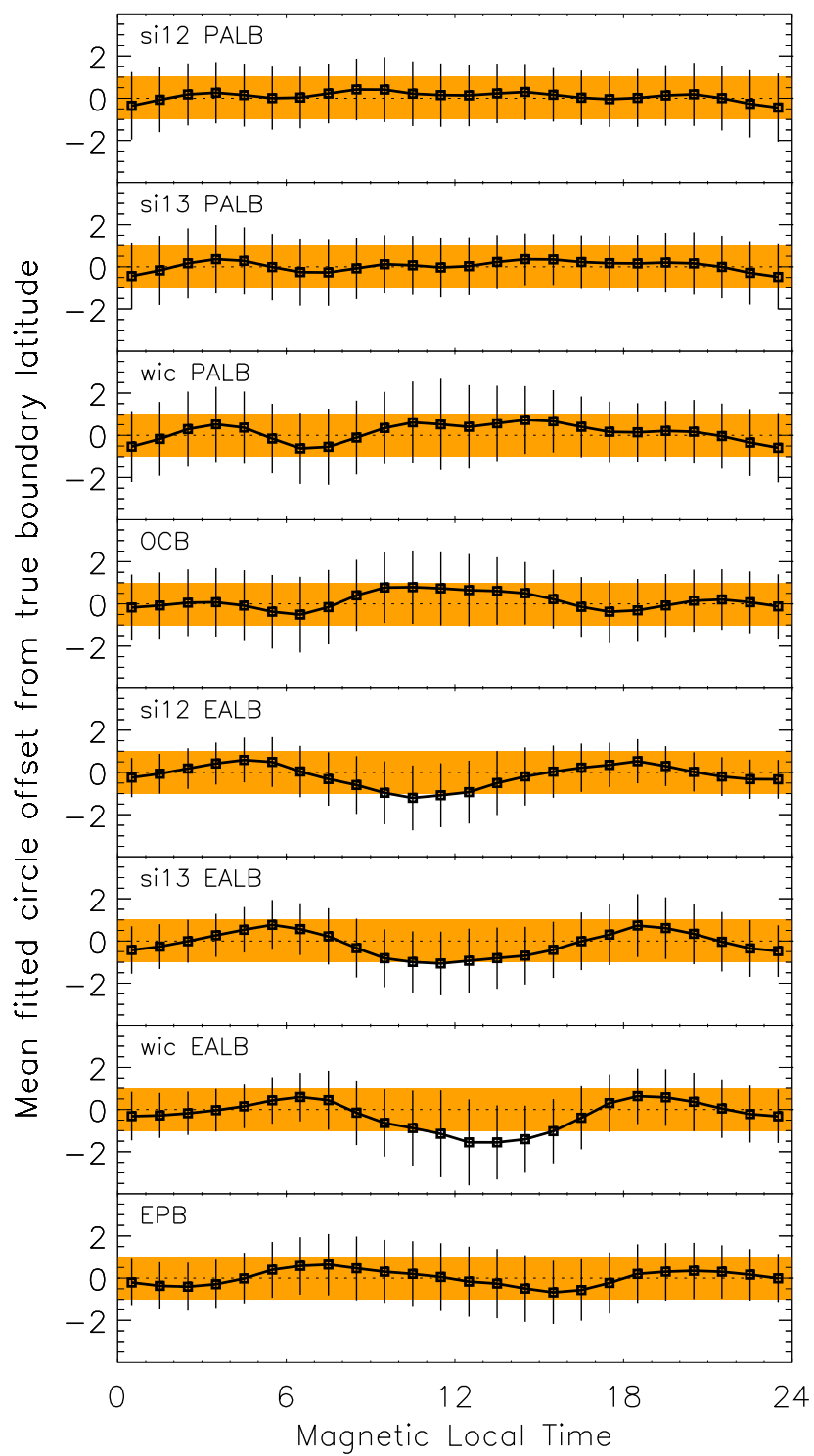


Figure 13.

Paired DMSP CRB locations in OCB coordinates

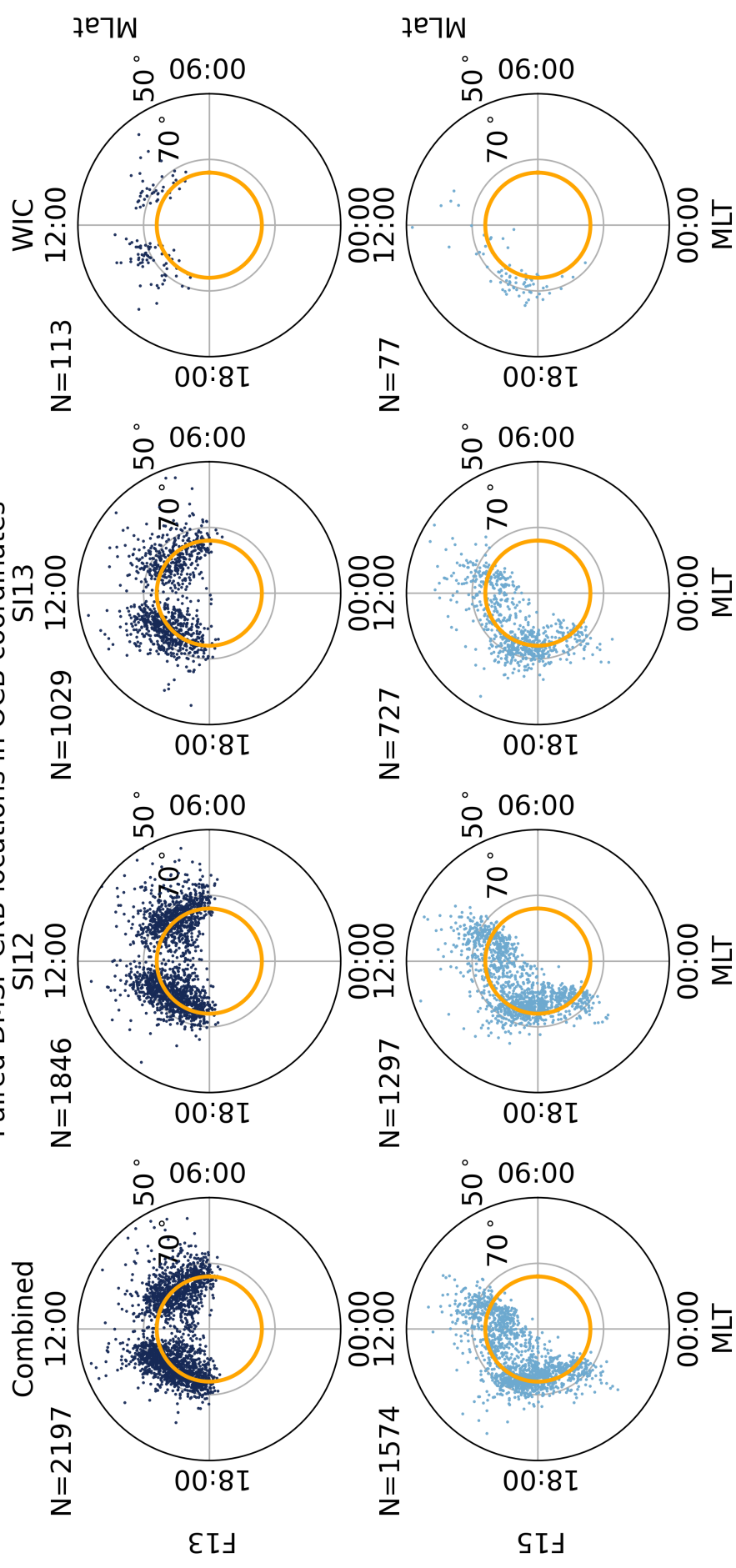


Figure 14.

Boundary differences within 1 h of 18:00 or 06:00 MLT

