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EVALUATION OF THE NOAA-4 VTPR THERMAL WINDS FOR NUCLEAR FALLOUT--ETC(U)
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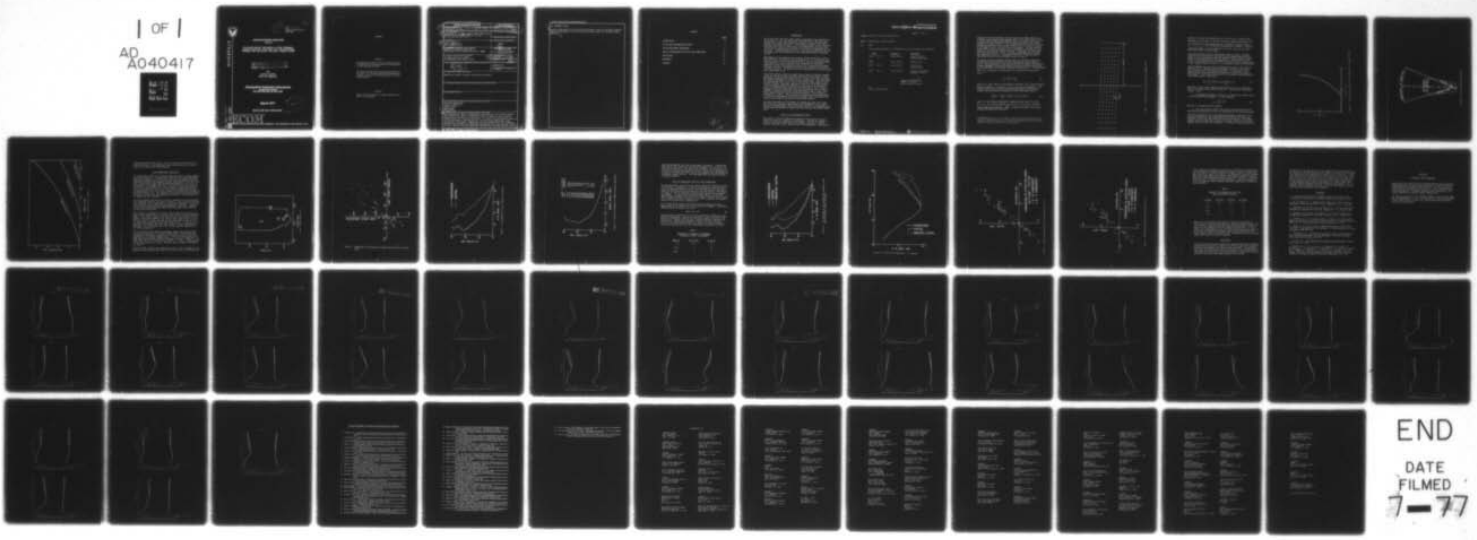

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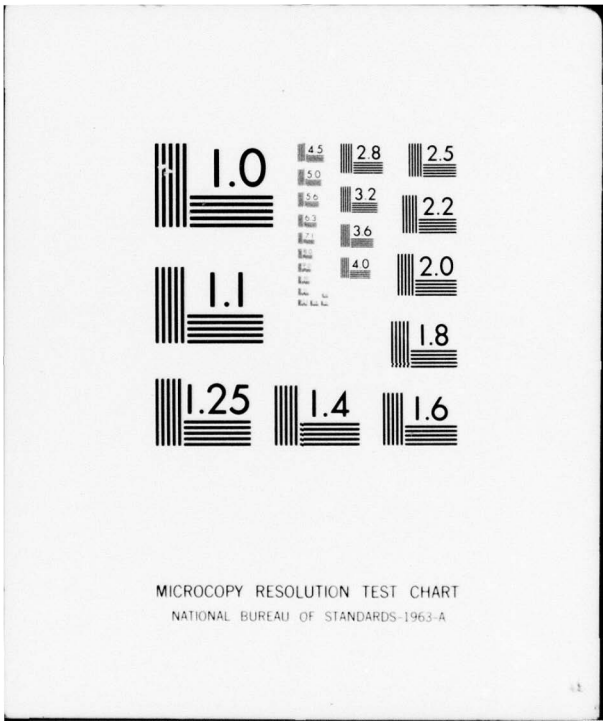
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RESEARCH AND DEVELOPMENT TECHNICAL REPORT
ECOM-5815

EVALUATION OF THE NOAA-4 VTPR THERMAL WINDS FOR NUCLEAR FALLOUT PREDICTIONS

By
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March 1977

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20. ABSTRACT *(Continue on reverse side if necessary and identify by block number)*

The procedure and theory for determining wind profiles from radiance measurements obtained by vertical temperature profile sounders on polar orbiting meteorological satellites are presented. The use of the wind profiles from 15 to 30 km altitude in the generation of nuclear fallout prediction on the battlefield is discussed. Comparisons of winds obtained with this technique with profiles obtained from radiosonde measurements have been made for 33 sets of data. Results of these comparisons indicate that the winds derived from

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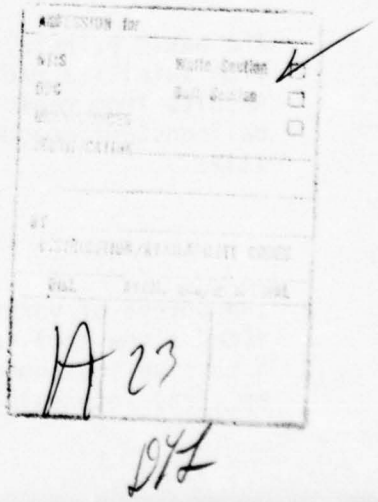
20. ABSTRACT (cont)

radiance measurements are sufficiently accurate to satisfy the Army's requirements for wind data, in the 15 to 30 km altitude range, for nuclear fallout prediction.

CONTENTS



INTRODUCTION

On 16 July 1945, the first atomic bomb was exploded in the Southern New Mexico desert. From this experience, and others which followed, man has begun to realize the necessity for protection against the harmful effects of an atomic explosion. These effects include initial radiation, heat and blast effects, shock waves, and contamination resulting from the fallout of radioactive particles carried into the atmosphere by the explosion. Protection against fallout is available through both the use of shelters and from predictions of the time and location of hazardous fallout.

Requirements for the generation of nuclear fallout predictions on the battlefield are discussed in several documents [1-4]. The primary meteorological input is the wind data obtained by the radio-tracked balloonsonde soundings (usually referred to as radiosonde or rawinsonde soundings) of the upper atmosphere made by the US Army Artillery Meteorological Sections. Soundings to at least 30 km altitude are required at 6-hour intervals. No other battlefield requirements for meteorological data extend above 15 km.

The nuclear fallout prediction capability has been on readiness standby status for the past 30 years and has exerted a distinct impact on Army battlefield meteorological operations as a result of the data collection and analysis necessary for nuclear fallout predictions. The requirement for balloonsonde data in the 15 to 30 km altitude range imposes a substantial increase (over that required for the 0 to 15 km observations) in meteorological equipment, operations, and man-hours on the battlefield. The Meteorological Satellite Technical Area of the US Army's Atmospheric Sciences Laboratory (ASL) has conducted an active research program to develop applications of meteorological data to upper level winds (above 15 km) for nuclear fallout prediction. This program is referred to as SATFAL. This research has demonstrated that meteorological satellite measurements can provide the upper level winds more efficiently than currently operational balloonsonde systems.

The basic principles and techniques for determining upper level winds from satellite radiometric measurements are presented in this report. Results from measurements made by the NOAA-4 satellite are compared with balloonsonde measurements to demonstrate and evaluate the SATFAL capabilities.

SATFAL WIND DETERMINATION THEORY

The advent of vertical temperature soundings from satellites has provided a new tool for atmospheric measurements. The thermal sounder is a passive instrument which senses the outgoing radiances in several spectral intervals as seen from the top of the atmosphere. The vertical

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Andrew G. Revoir

ANDREW G. REVOIR
Chief, Analysis Division

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Above listed employees

Temperature Profile Radiometer (VTPR)* flown on the NOAA satellites is a thermal sounder which scans perpendicular to the orbital path in order to obtain data extending approximately 1000 km to either side of the subsatellite track (Figure 1). The radiances measured by the VTPR are integrated signatures of the thermal structure in the vertical column of the atmosphere in the field of view of the sounder. The basic physical theory which relates the radiance to the temperature structure is contained in the radiative transfer equation. Algorithms have been developed which allow one to "invert" the radiance measurements and obtain the temperature profiles. A temperature profile, if desired, may be determined at each of the grid points shown in Figure 1.

The use of satellite measured radiance data to calculate a usable approximation of the wind field has been discussed by several authors [5-7]. These observations have been based upon well-known relationships between atmospheric thermal and dynamic structures with the wind given by the geostrophic wind through the thermal wind equation. The geostrophic wind, $\vec{V}_g$, is given by (cf Holton [8] for a derivation of this equation)

$$\vec{V}_g = \frac{1}{\rho f} \vec{k} \times \vec{\nabla}_h p \tag{1}$$

where $\vec{k}$ is a unit vector in the vertical direction, $\vec{\nabla}_h p$ is the horizontal pressure gradient, ρ is density, and f is the Coriolis parameter. The vertical shear (derivative) of the geostrophic wind is given by the thermal wind which may be expressed in finite difference form by

$$\vec{V}_g(P_2) - \vec{V}_g(P_1) = \frac{R}{f} \vec{k} \times \vec{\nabla}_p \bar{T} \ln (P_1/P_2) \tag{2}$$

where $\bar{T}$ is the average temperature between pressure levels P_1 and P_2 ($P_2 < P_1$), R is the gas coefficient for dry air, and $\vec{\nabla}_p \bar{T}$ is the horizontal temperature gradient. (NOTE: Pressure and height are related through the hydrostatic equation.)

A detailed discussion of the VTPR together with operational data processing procedures used by the National Environmental Satellite Service (NESS) is given by McMillian et al. [9] (1973).

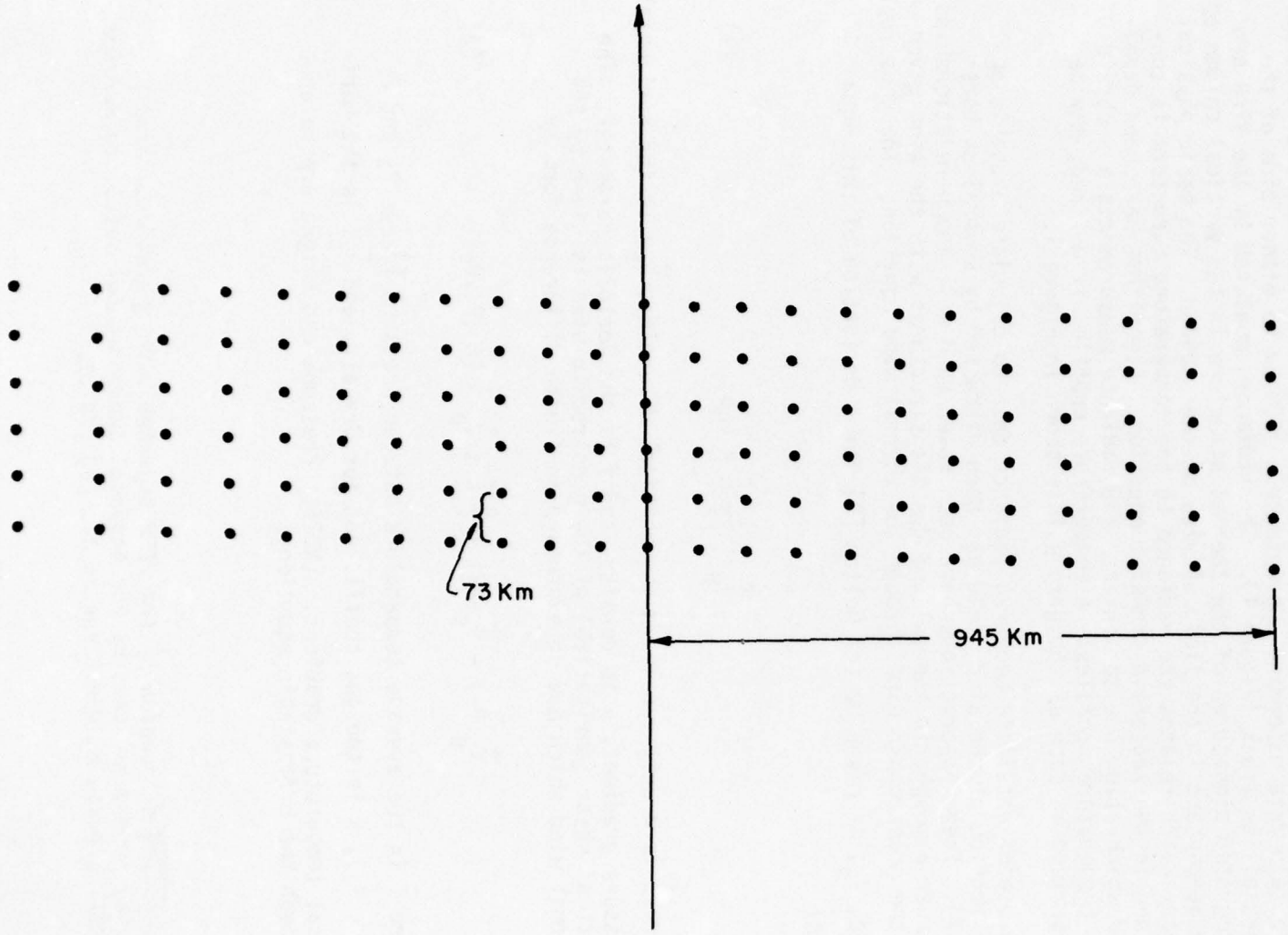

Figure 1. Distribution of Earth Scan Scene for the NOAA-4 VTPR.

Equation (2) shows that determination of the wind at P_2 requires only knowledge of the wind at P_1 and the horizontal temperature gradient $\vec{\nabla}_h \bar{T}$. This procedure can be extended through a sequence of pressure levels $P_1 > P_2 > \cdots > P_n$. This is the basic concept used in SATFAL. A balloonsonde measurement is used to provide the "tie-on" value at the lowest level (at or near 15 km altitude).

The relatively dense grid of observations shown in Figure 1 provides sufficient data for numerical estimation of the temperature gradients. However, as is typical with numerical differentiation problems, considerable care must be taken in this computation. After much trial and error, it was determined that sufficient accuracy could be obtained from a least squares planar fit over a 7 x 7 rectangular array of observations.

Complete details for the fallout predictions are given in TM 3-210 [3] and FM 3-22 [4]. Relevant physical parameters such as nuclear burst time, location, and yield are assumed known. (A discussion of the methods by which this information is obtained is outside the scope of this report.) The effect of wind drift is calculated from the effective wind profile which is defined by

$$\vec{V}(z) = \int_0^z [\vec{U}(\mu)/\tau(\mu)]d\mu \Big/ \int_0^z d\mu/\tau(\mu) \qquad (3)$$

where $\vec{U}(\mu)$ is the actual (measured) wind profile and $\tau(\mu)$ is the fall rate for a nominal size particle (Figure 2). The principal predictands (Figure 3) are:

a. The downrange distance, R, which is related to the nuclear yield, Y, and the effective windspeed, V, by the expression

$$R = AY^{2/5} V^{1/2} \qquad (4)$$

where A is a proportionality constant.

b. The right and left radial lines which are determined from the effective wind directions at the 2/3 stem height and the cloud top height.

The fallout prediction for a specified yield becomes a function of the effective windspeed at the cloud bottom height and the effective wind direction at the cloud top height and the 2/3 stem height. These three heights, which are functions of yield, are shown in Figure 4 for yields ranging from 10 kilotons to 10 megatons. (The data for constructing this

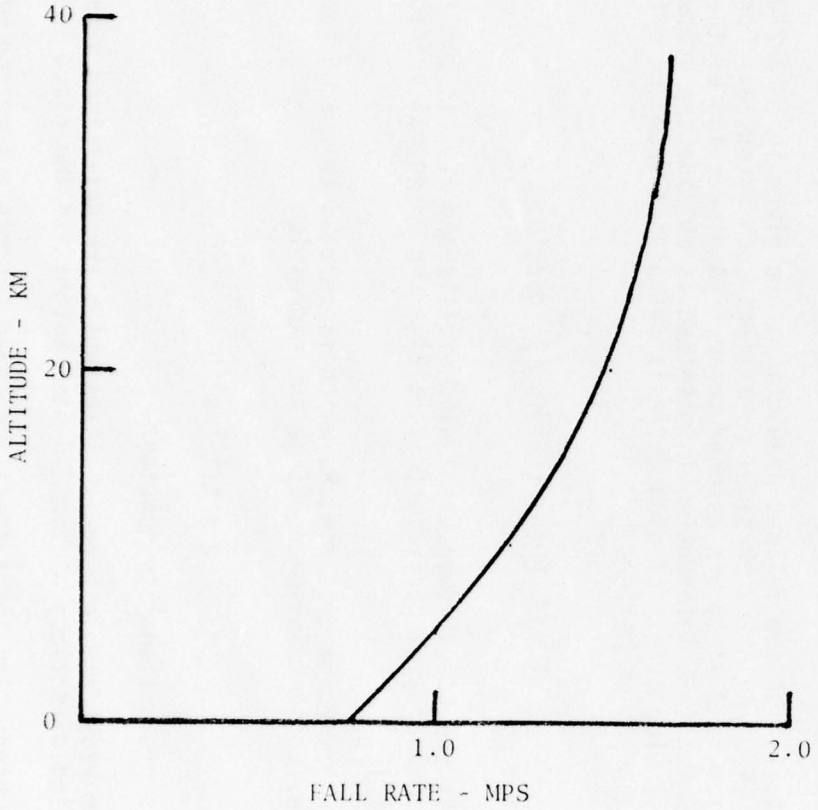

Figure 2. Fall rate for nominal fallout particles - source FM 3-22.

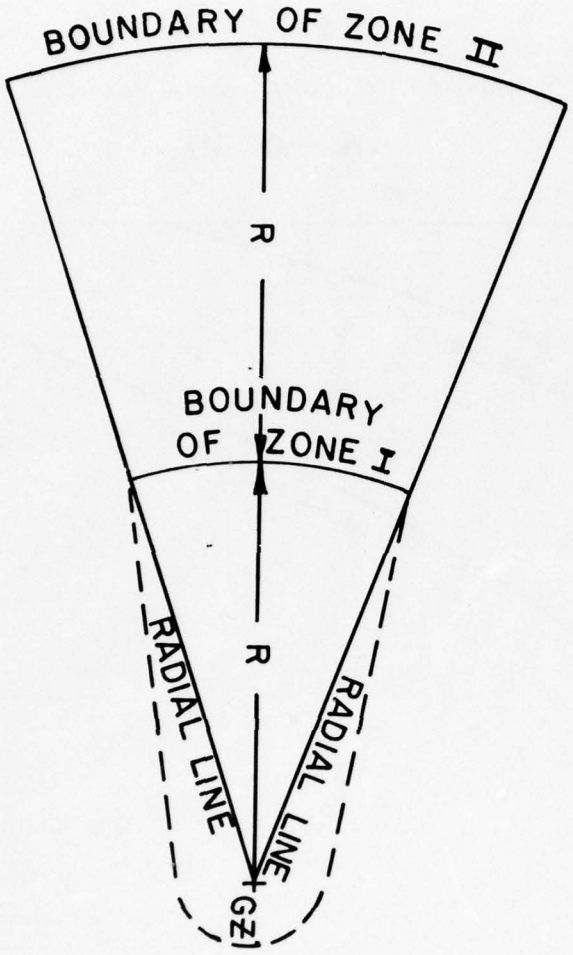

Figure 3. Fallout prediction diagram.

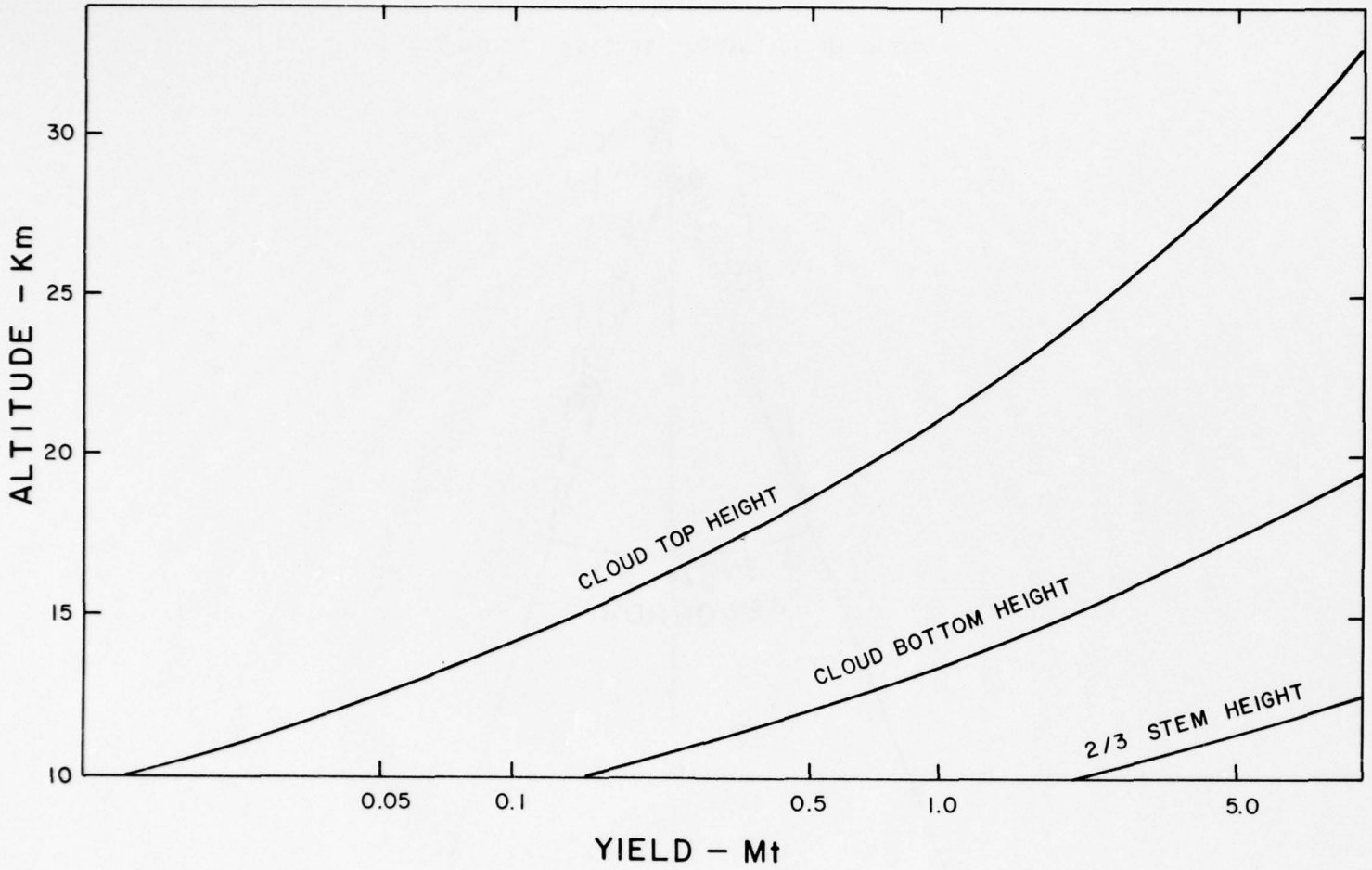

Figure 4. Nuclear cloud height parameters.

graph were taken from TM 3-210.) Only the direction of the effective wind is required for altitudes above the cloud bottom height, which is generally a rather large altitude range.

SATFAL/RADIOSONDE COMPARISONS

Six radiosonde stations are located on White Sands Missile Range (WSMR), New Mexico (Figure 5). A special data collection was conducted between February and December 1975 to obtain a data base for SATFAL evaluation. Radiosonde balloons were launched (nearly) simultaneously from the three lower range stations (WSD, SMR, and HMN) approximately 1 hour prior to the passover of the NOAA-4 satellite. (The release time was chosen to minimize the time variability of the wind data.) Eleven abutting pressure layers extending from 125 to 10 mb (with corresponding altitudes of approximately 15 to 31 km) were selected for thermal wind computations employing Equation (2). The average of the three measured wind profiles was used as "truth" data for comparative evaluation. Thirty-three sets of comparisons were collected.

The total thermal wind from 15 to 30 km is compared with the corresponding radiosonde measured wind shear in Figure 6. Dots represent the east-west component; crosses represent the north-south component. Except for a few isolated cases, generally good agreement is obtained. Computed correlation coefficients were 0.786 for the east-west winds and 0.743 for the north-south winds.

More insight is available from Figure 7 which shows the east-west component of the mean profiles. The total shear for the average radiosonde winds is 18.9, which compares well with the total thermal wind of 16.3. However, the thermal wind lags the typical large actual shear in the 15 to 20 km region, is about the same as the actual shear in the 20 to 25 km region, and exceeds the actual shear above 25 km. It should also be noted that the thermal shear for this average is almost constant with altitude. This is probably due to the limited vertical resolution in the VTPR measurements.

The observations of the previous paragraph suggest that perhaps better results are obtainable if the total thermal shear is distributed percentagewise in the manner suggested by climatology instead of as calculated for the individual pressure layers. The zonal wind component for WSMR taken from the IRIG reference atmosphere is shown in Figure 8. The total shear from 15 to 30 km for this profile is 19 mps. The percentage of this value occurring in each 1 km thick layer is shown in the table on Figure 8.

Modified SATFAL profiles were computed for each of the 33 samples by computing the total thermal wind from 15 to 30 km and then distributing the

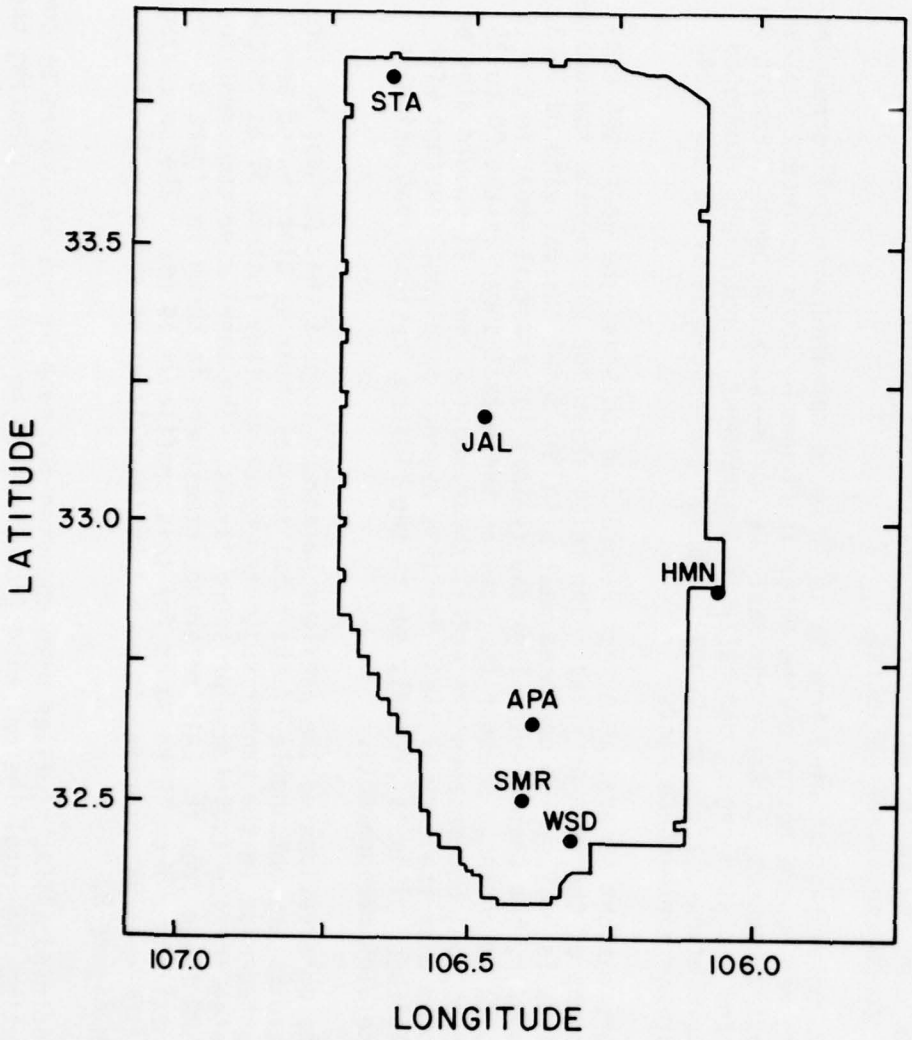

Figure 5. WSMR radiosonde sites.

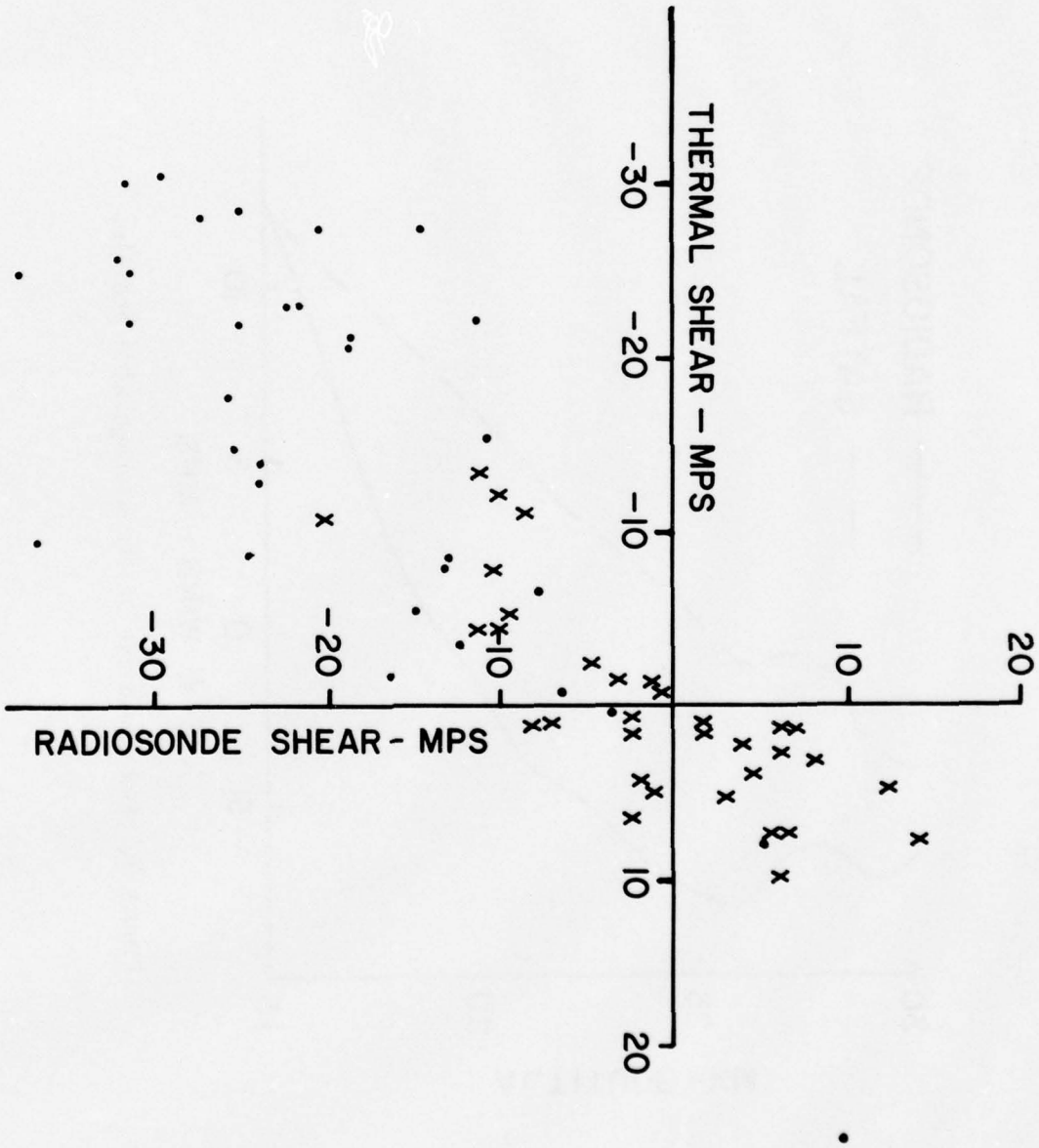

Figure 6. Comparison of radiosonde and thermal shear for the 15-30-km layer.

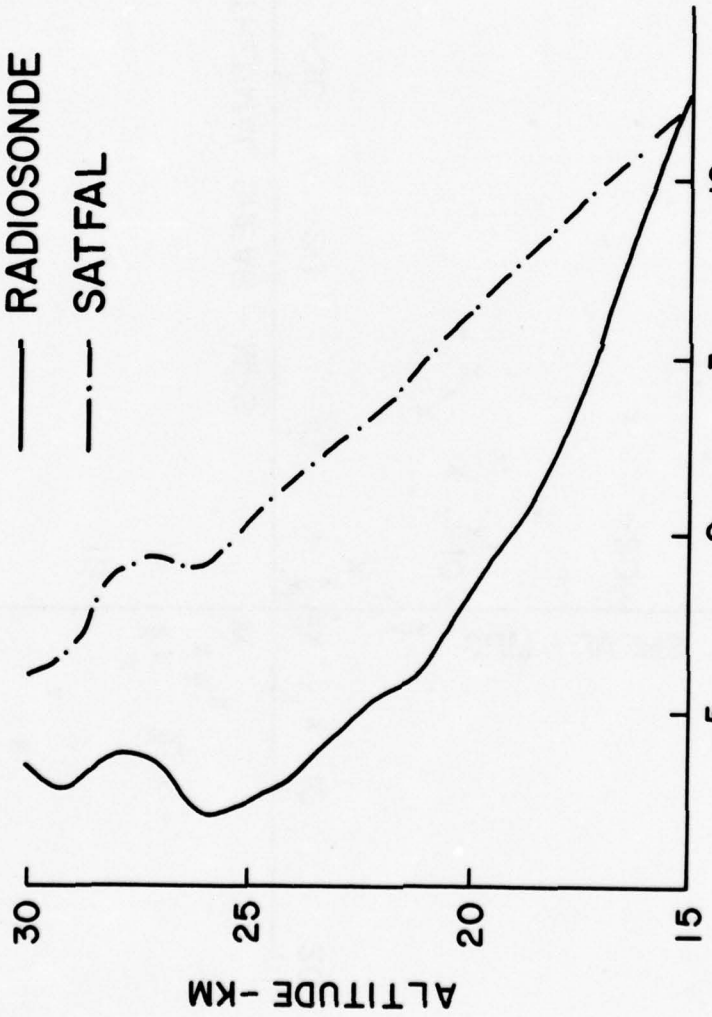

Figure 7. E-W component of the average wind profile.

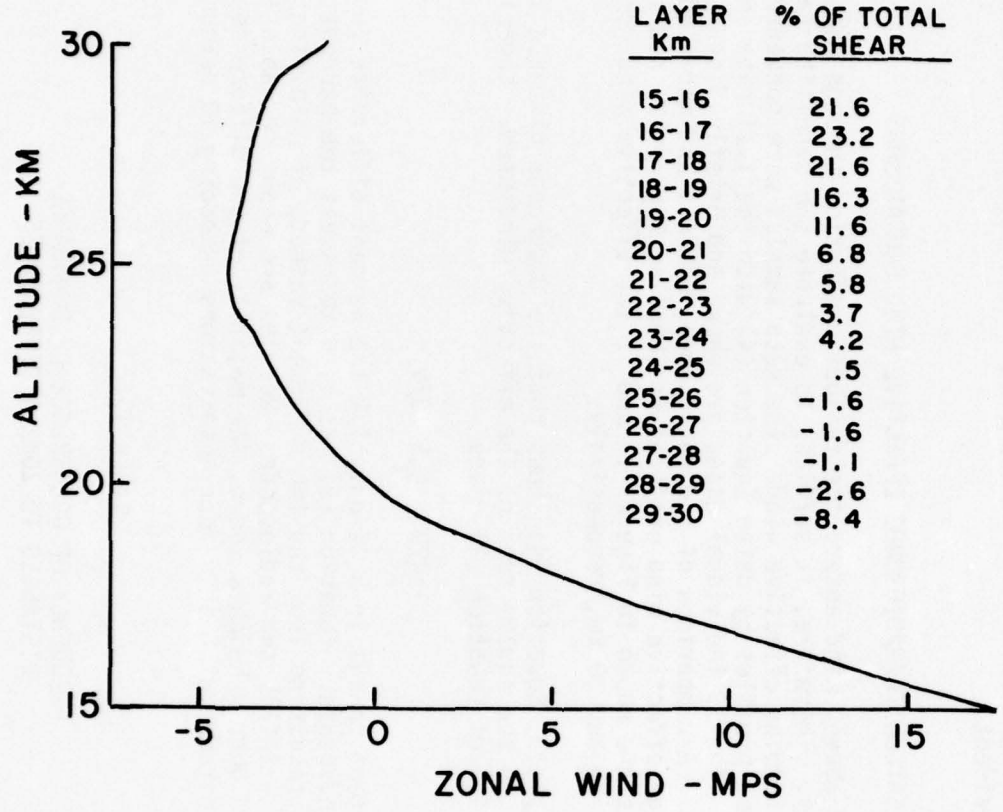

Figure 8. Zonal wind component of WSMR reference atmosphere.

total percentagewise using the values shown in Figure 8. (Through the remainder of this report the term "modified SATFAL" will be used to indicate wind obtained in this manner while "SATFAL" will be used to indicate profiles obtained by using the standard thermal wind computations.) In Figure 9 the modified SATFAL results are compared with the SATFAL and radiosonde average profiles. The improvement with the modified SATFAL is readily evident.

SATELLITE/RADIOSONDE EFFECTIVE WIND COMPARISONS

As discussed above, wind enters the fallout prediction through the effective velocity. Therefore, it suffices to evaluate the satellite capability to determine effective wind. The data samples were converted to effective wind profiles by using Equation (4) with the fall rate shown in Figure 2. The 33 individual cases are shown and briefly discussed in the appendix. A comparison of the radiosonde, SATFAL, and modified SATFAL average effective wind profiles is shown in Figure 10, and statistical results are shown in Figures 11 and 12 for effective wind comparisons at 22.5 km and 30 km, respectively.

For a fixed yield, Equation (4) shows that the downrange distance is proportional to the square root of the effective windspeed. Logarithmic differentiation of Equation (2) gives

$$\Delta R/R = 0.5\ \Delta V/V \qquad (5)$$

Thus the percent change in R is only 1/2 the percent difference in the effective windspeeds. Equation (5) was used to assess the percent difference in R resulting from the two different methods of obtaining the wind profile - SATFAL and radiosonde. Results are shown in Table 1. For the cloud bottom heights shown, the majority of the differences is less than 5 percent with only four observations exceeding 10 percent.

TABLE 1

FREQUENCY OF OCCURRENCE OF DOWNRANGE DISTANCES TO ZONE 1 DIFFERENCES

$\Delta R/R$ (%)	at 17.5 km	at 20 km
<5	27	15
5-10	5	14
10-15	1	4

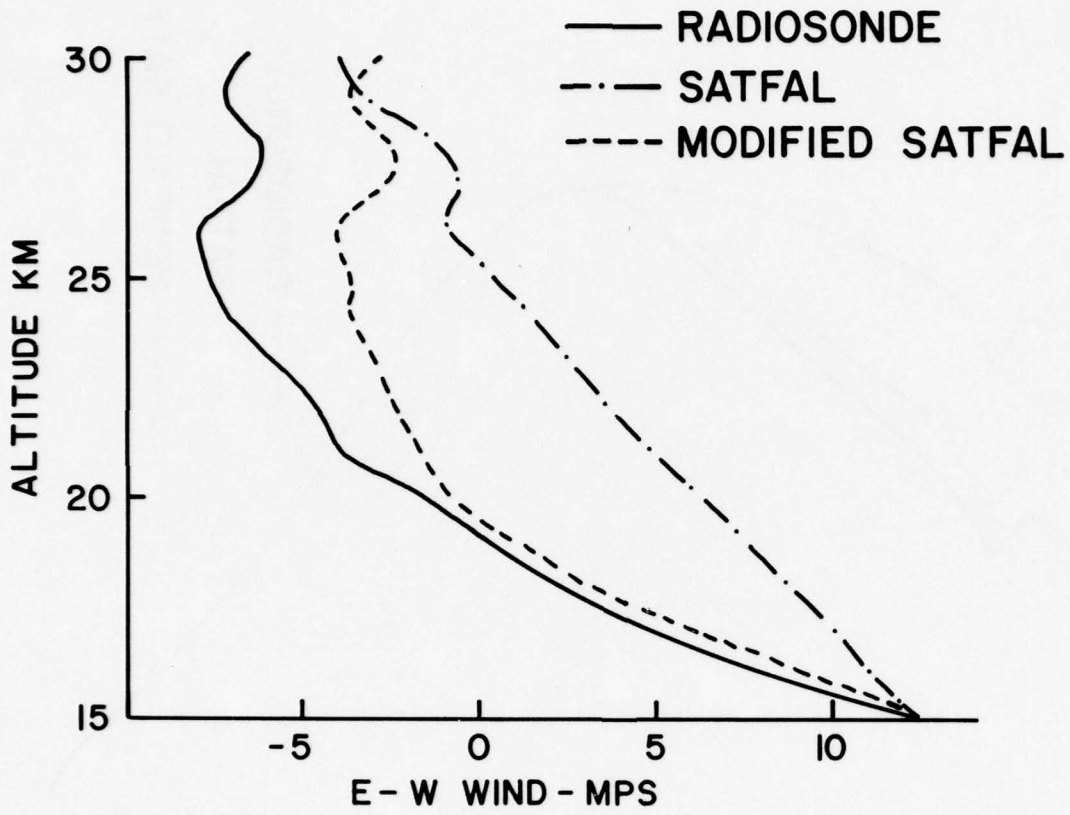

Figure 9. E-W component of the average wind component with modified SATFAL profile included.

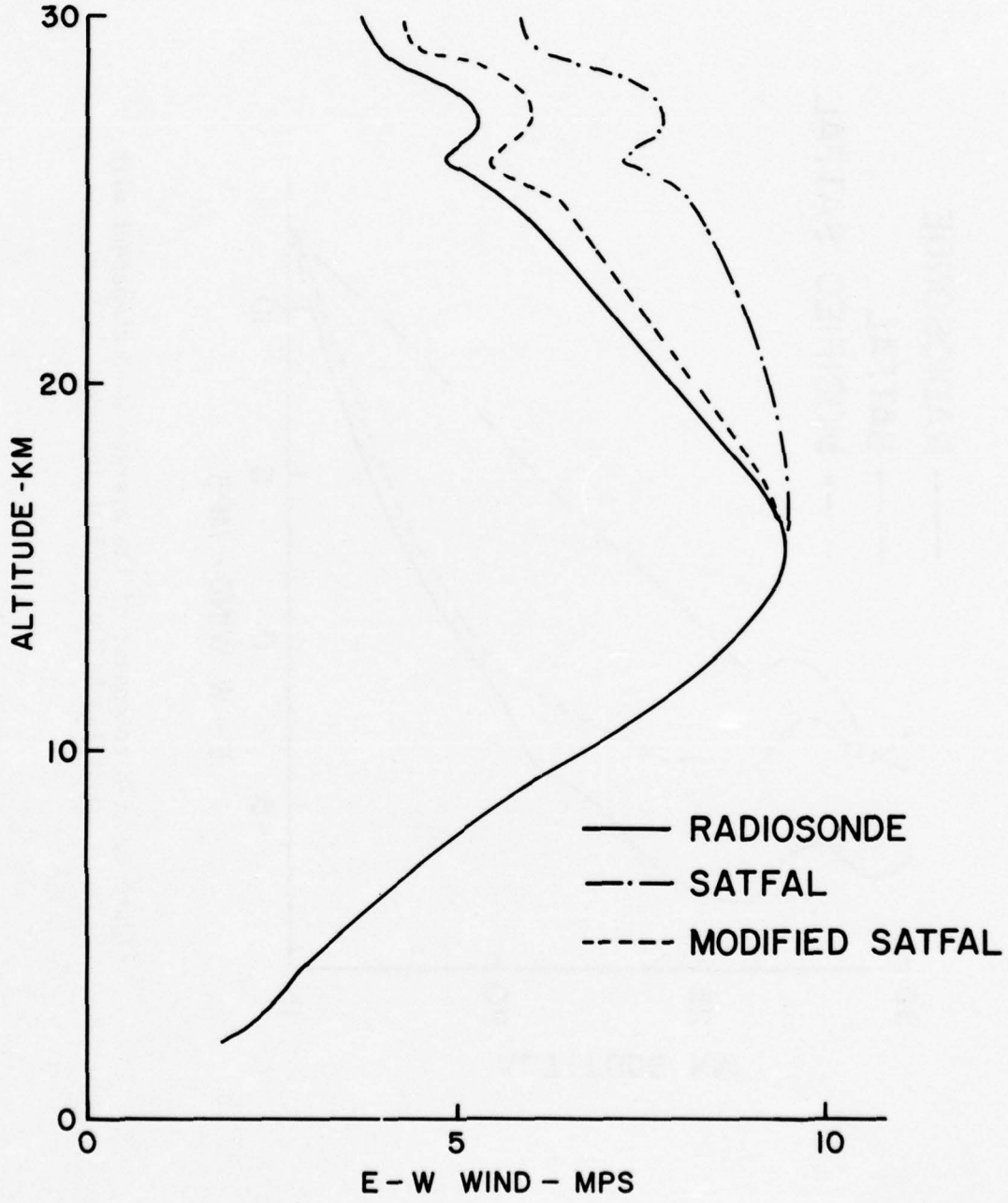

Figure 10. Effective wind comparisons - E-W component.

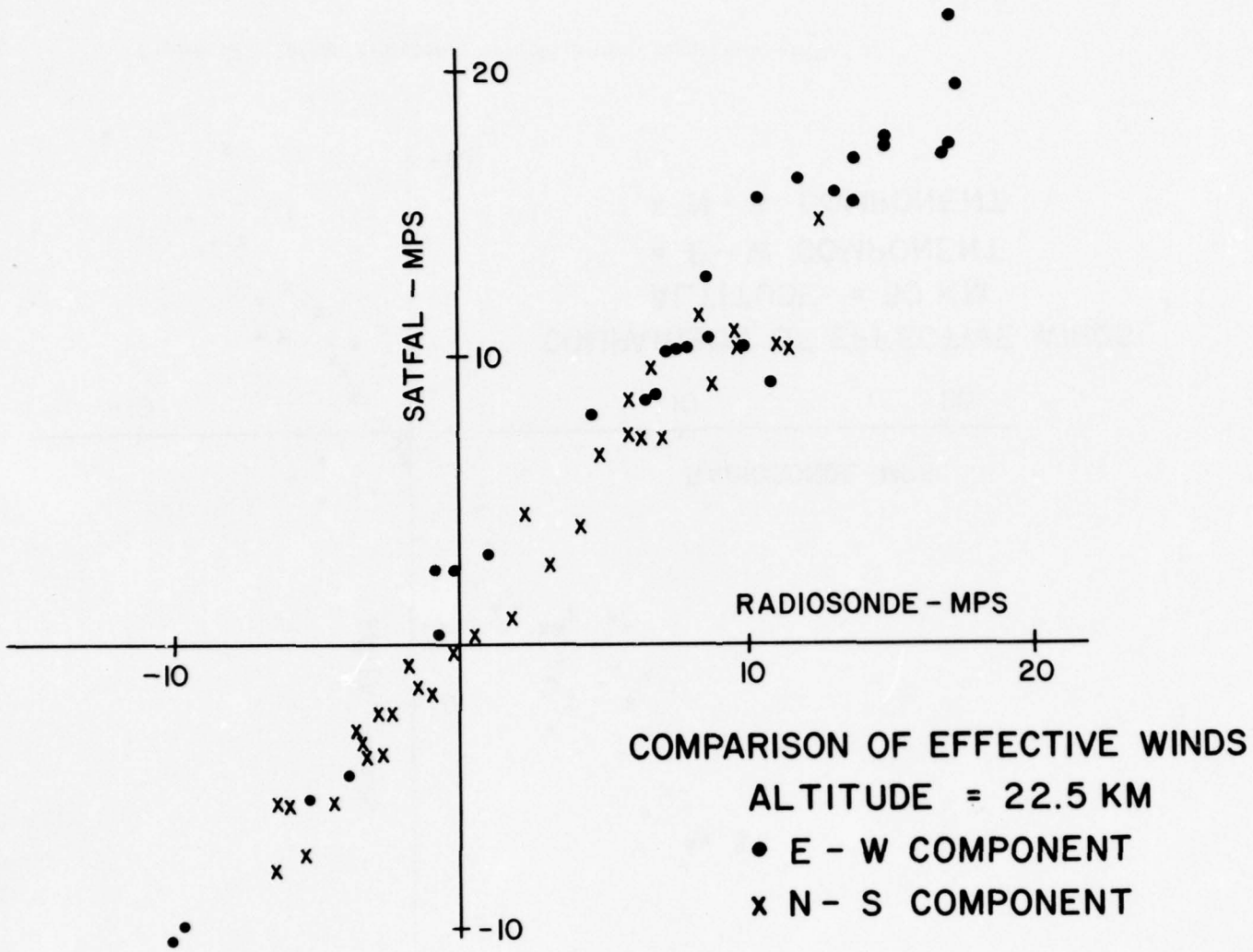

Figure 11. SATFAL/radiosonde effective wind comparison at 22.5 km.

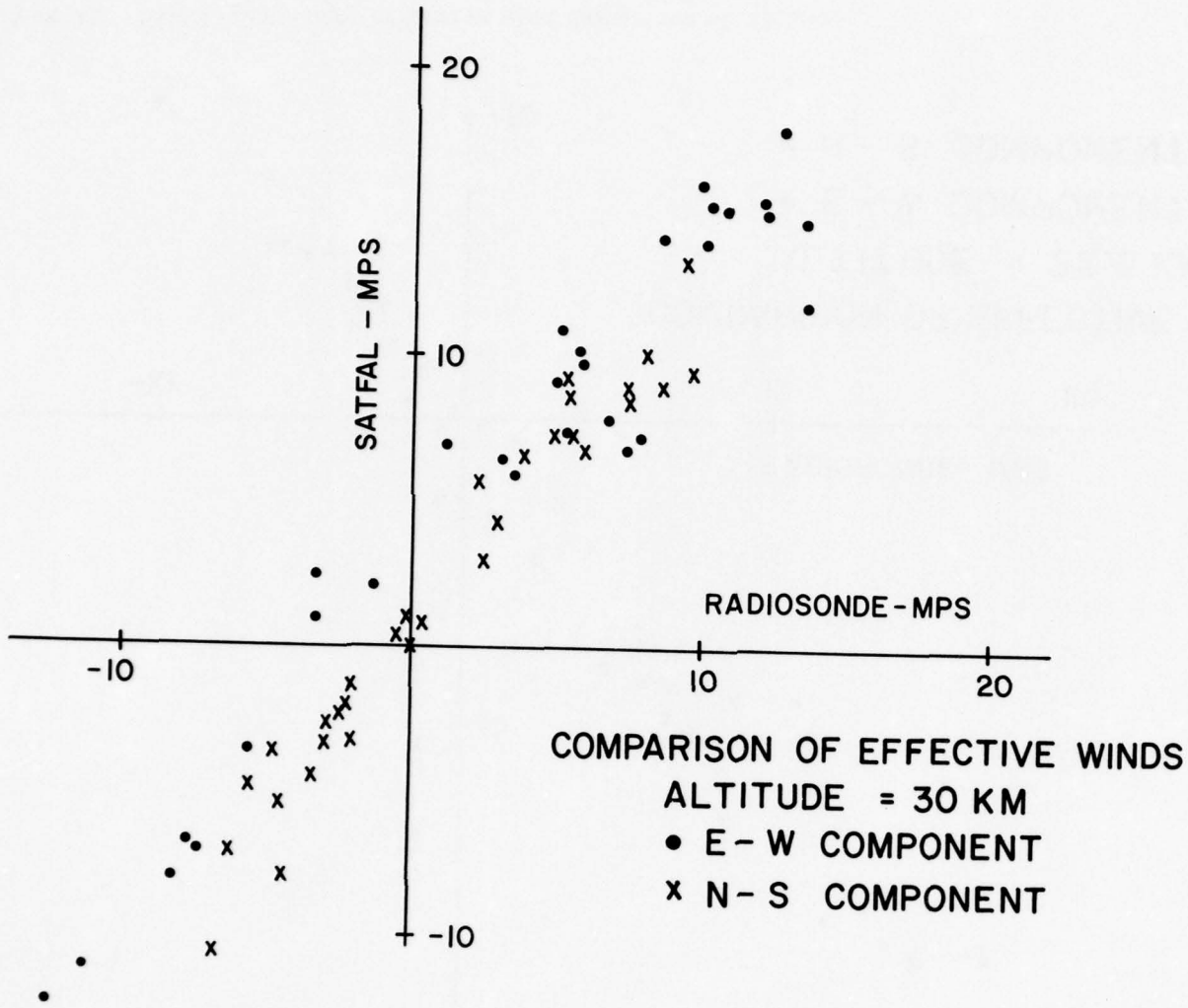

Figure 12. SATFAL/radiosonde effective wind comparison at 30 km.

Table 2 shows the frequency of occurrence of effective wind shear direction differences for three different altitudes. Direction differences are quite small, usually less than 10 and frequently less than 5 degrees. Most of the cases where wind direction differences exceeded 15 degrees were under light windspeed conditions where small changes in the vector wind can result in large direction differences. Very good agreement to 30 km altitude was found in SATFAL/radiosonde comparisons of effective wind direction.

TABLE 2

FREQUENCY OF OCCURRENCE OF EFFECTIVE
WIND DIRECTION DIFFERENCES

$\Delta\Theta$ (deg)	at 20 km	at 25 km	at 30 km
<5	23	15	14
5-10	7	8	6
10-15	2	4	5
>15	1	6	8

When used in conjunction with Figure 4, the results presented in Tables 1 and 2 show the difference which might be expected between fallout predictions computed from SATFAL and radiosonde winds. As an example, for an assumed yield of 2.5 megatons (or smaller): (1) the radial line corresponding to the 2/3 stem height would be the same, (2) the percent difference in downrange distance would probably be less than 5 percent, and (3) the radial lines corresponding to the cloud top height should agree to within 5 degrees.

CONCLUSIONS

The system discussed herein employs the advanced technology provided by meteorological satellites to provide an efficient means for satisfying the Army's requirements for upper altitude wind data. The procedure and theory for determination of the wind profile from radiance measurements obtained by vertical temperature sounders on polar orbiting meteorological satellites have been presented and discussed.

An analysis of 33 comparisons of winds obtained from radiosonde balloons and satellite radiances measured by the NOAA-4 VTPR shows that the satellite derived winds are adequate for nuclear fallout prediction. It was also shown that the thermal shear is significantly less than the actual shear in the 15 to 20 km level, and this is believed to be a result of the coarse vertical resolution of the satellite sounder. Modification of the computed thermal winds by the addition of climatological data provides improved agreement with the radiosonde measurements. This is apparent from the average case shown as well as from inspection of the individual cases presented in the appendix.

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APPENDIX

INDIVIDUAL DATA COMPARISONS

Comparison of the effective velocity determined from the balloonsonde wind measurements and the derived winds from the satellite radiance measurements are presented in Figure A1-A33. The three separate velocity profiles are identified as follows: A cross indicates radiosonde measurements; a triangle indicates SATFAL computations; and a square identifies the modified SATFAL computations.

The data are plotted at 1 km increments between 2 and 30 km MSL (WSMR MSL is approximately 1.25 km). In some cases the radiosonde data terminates below 30 km; in these cases, usable data were not obtained above the altitude shown.

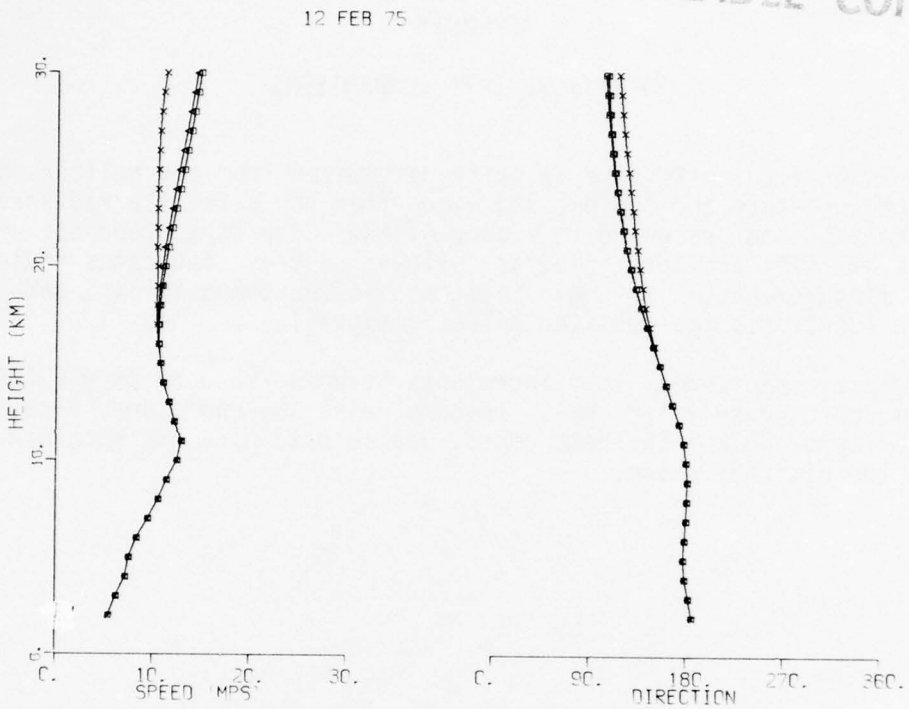


Figure A1. Satellite/radiosonde effective velocity comparison for 12 Feb 75.

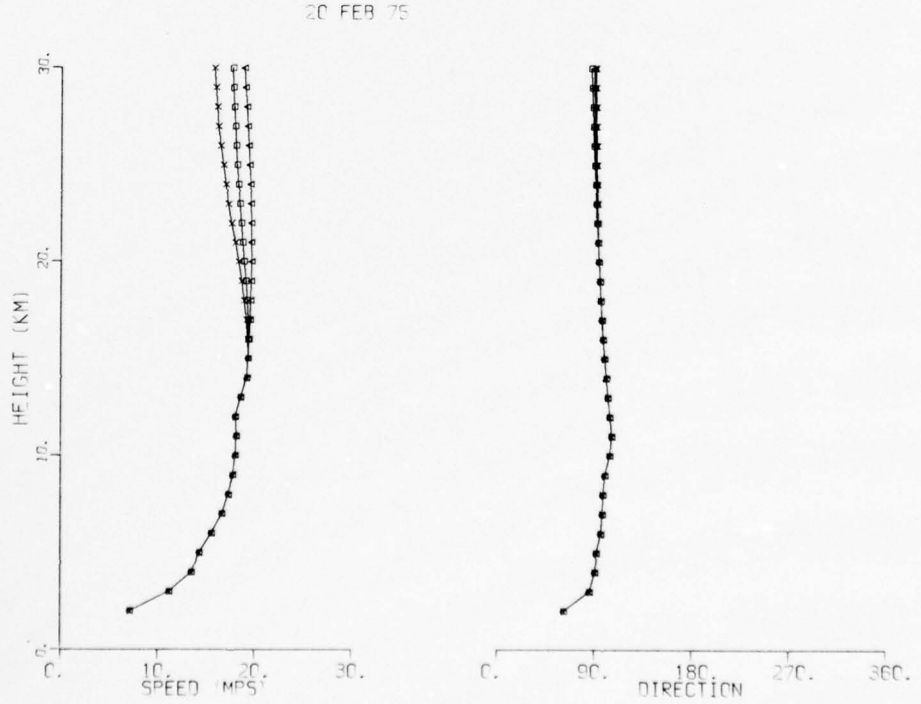


Figure A2. Satellite/radiosonde effective velocity comparison for 20 Feb 75.

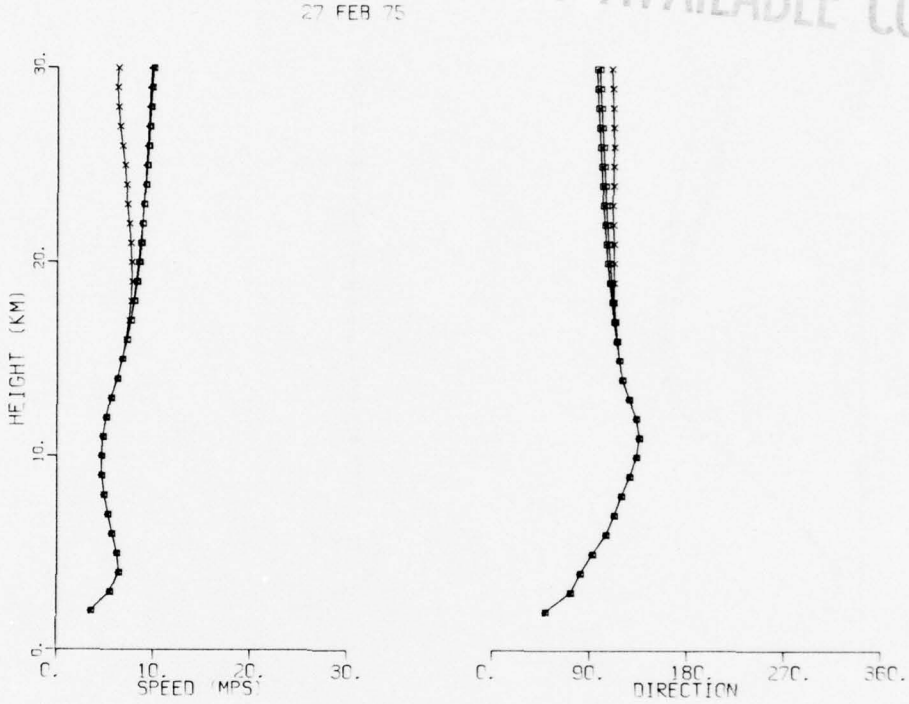


Figure A3. Satellite/radiosonde effective velocity comparison for 27 Feb 75.

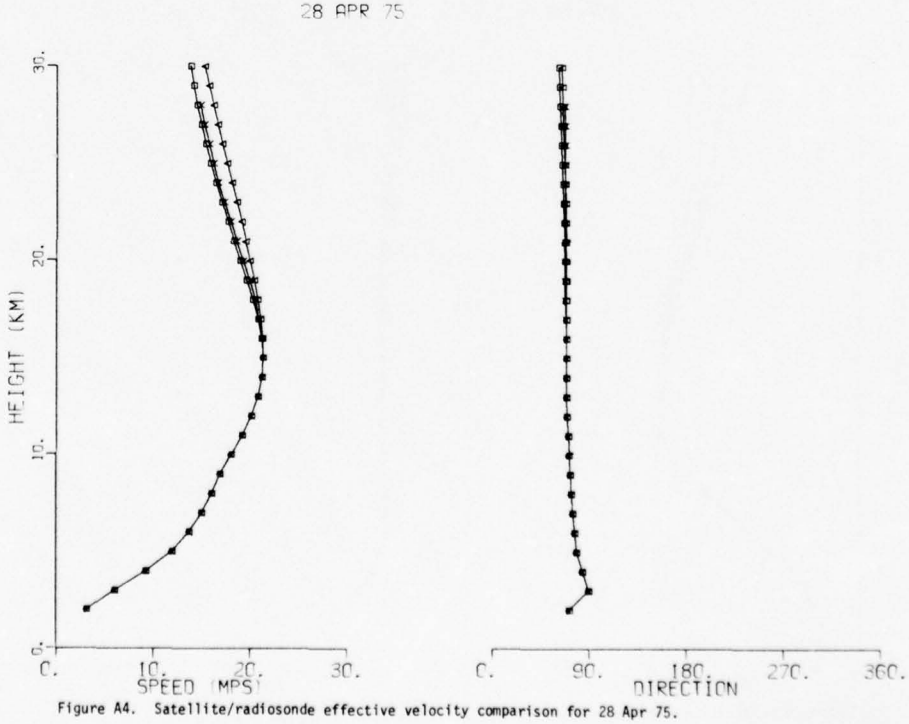


Figure A4. Satellite/radiosonde effective velocity comparison for 28 Apr 75.

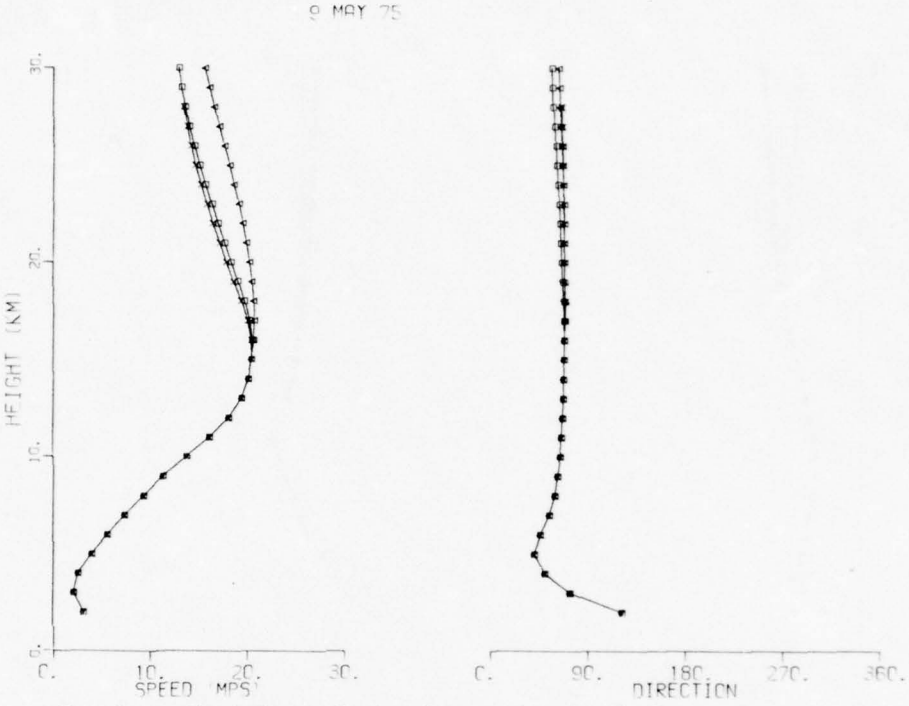
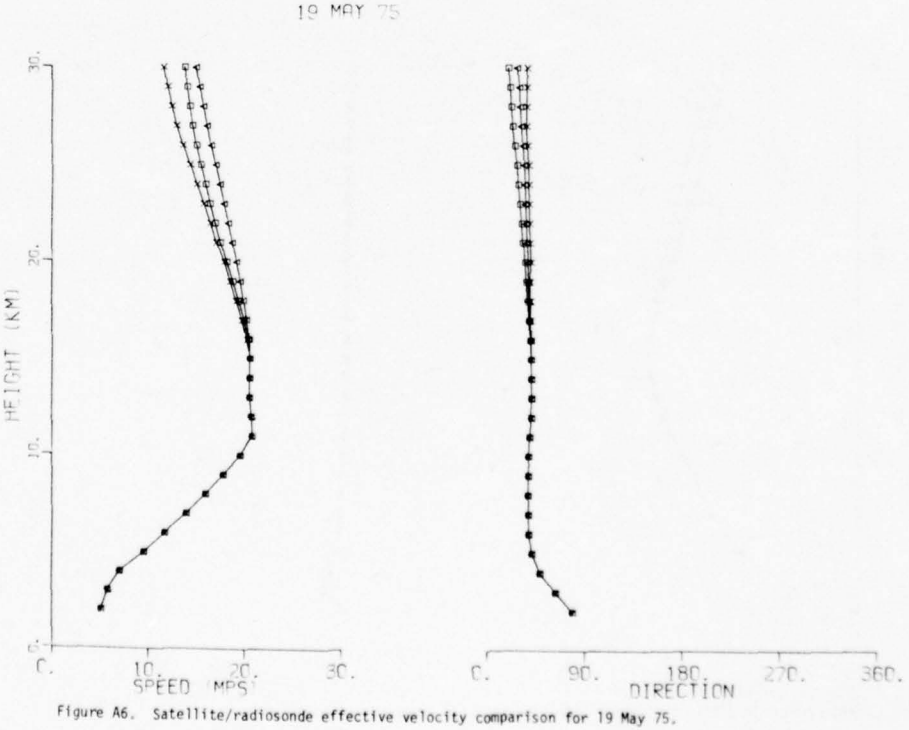


Figure A5. Satellite/radiosonde effective velocity comparison for 9 May 75.



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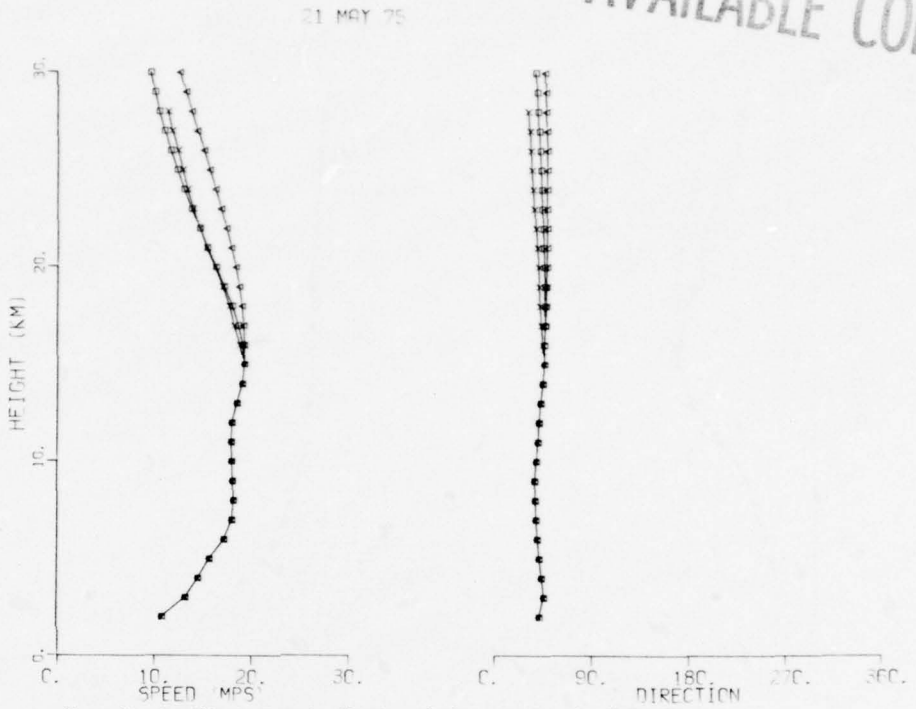

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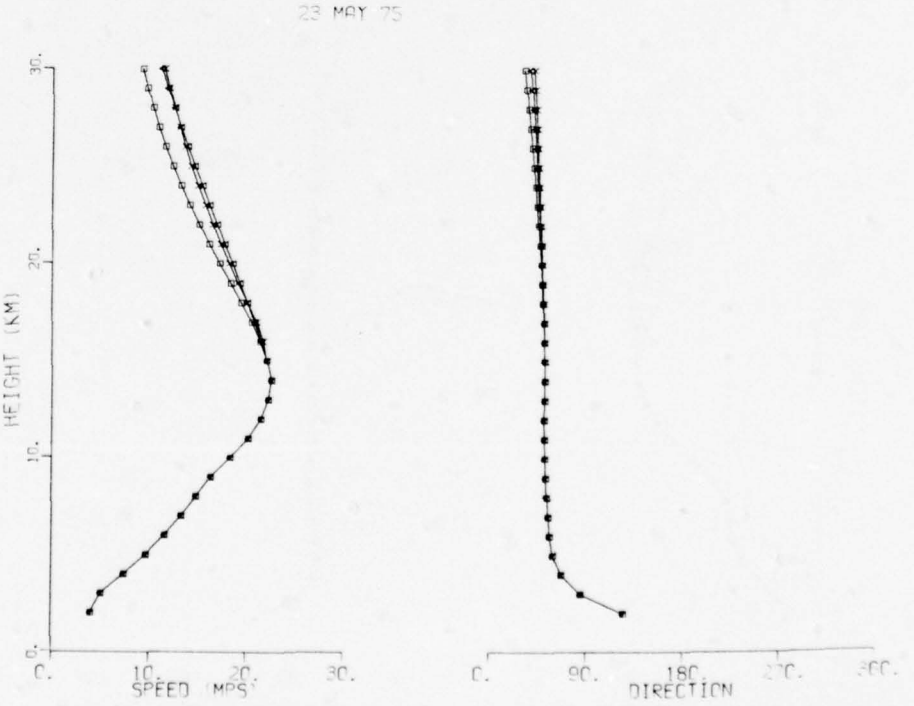

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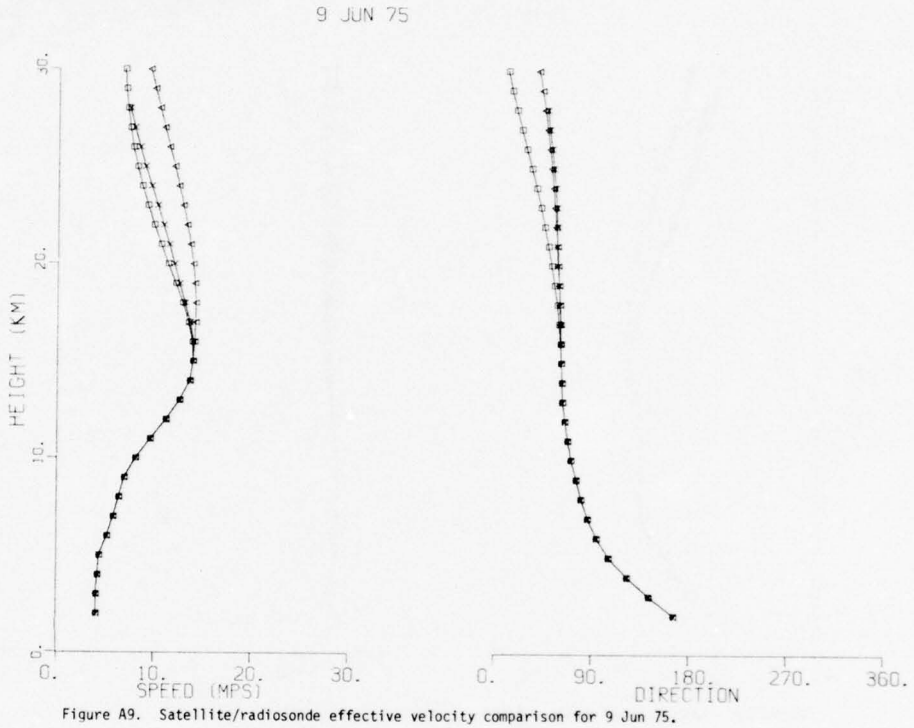
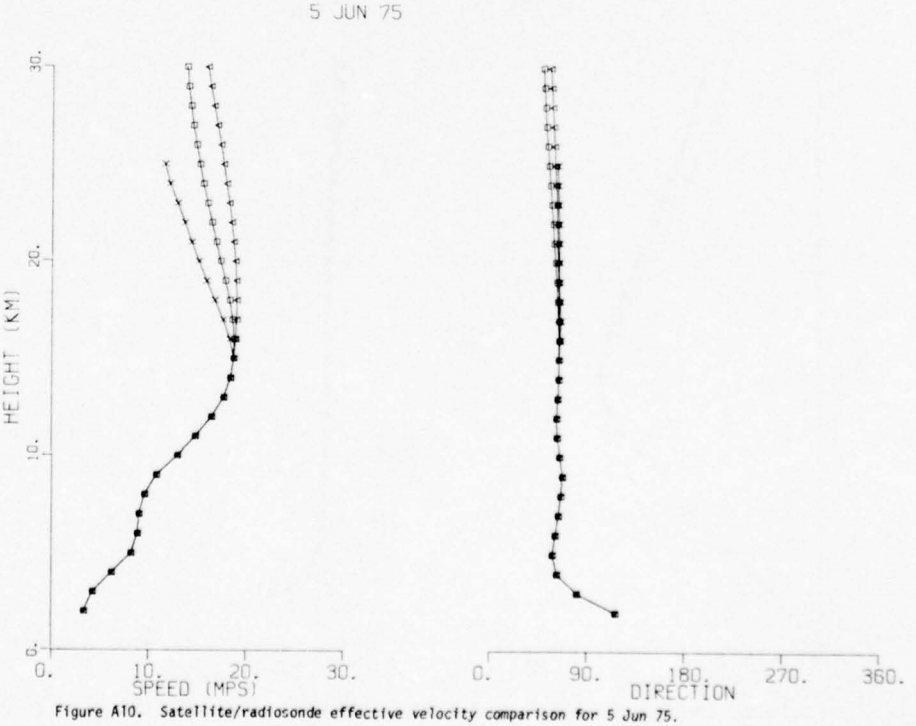


Figure A9. Satellite/radiosonde effective velocity comparison for 9 Jun 75.



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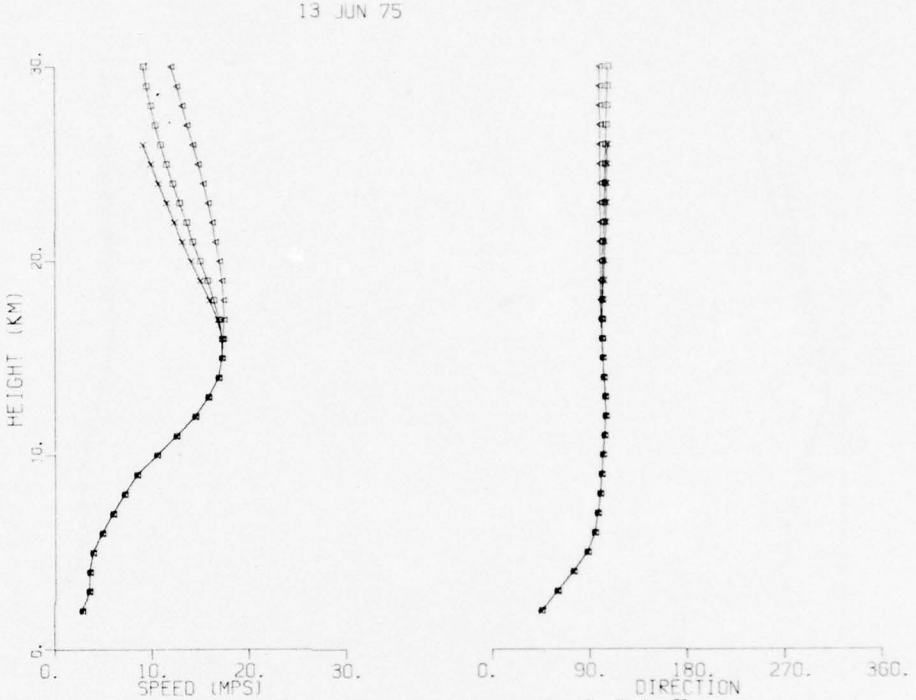

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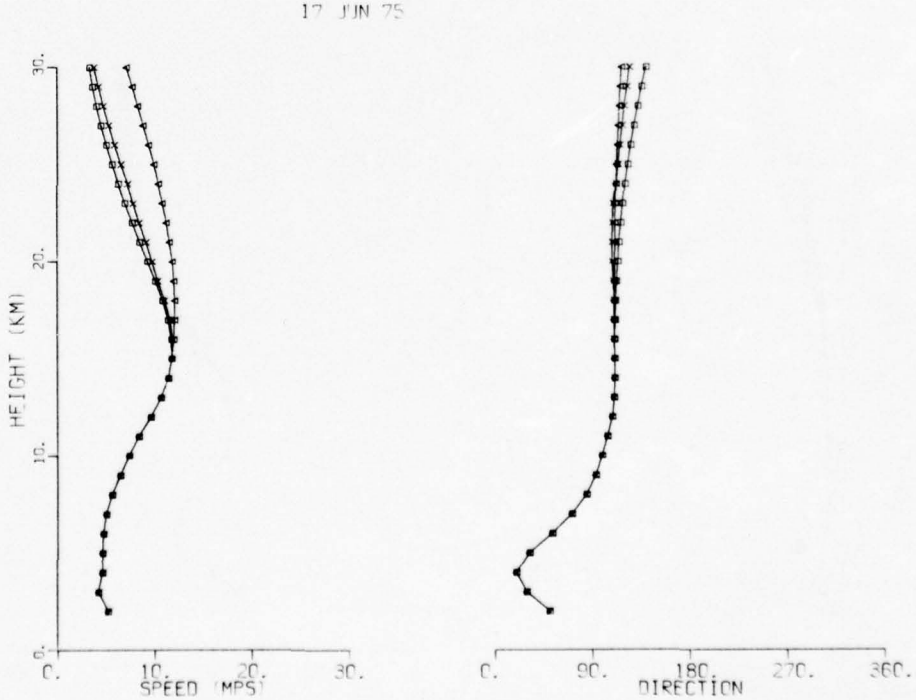

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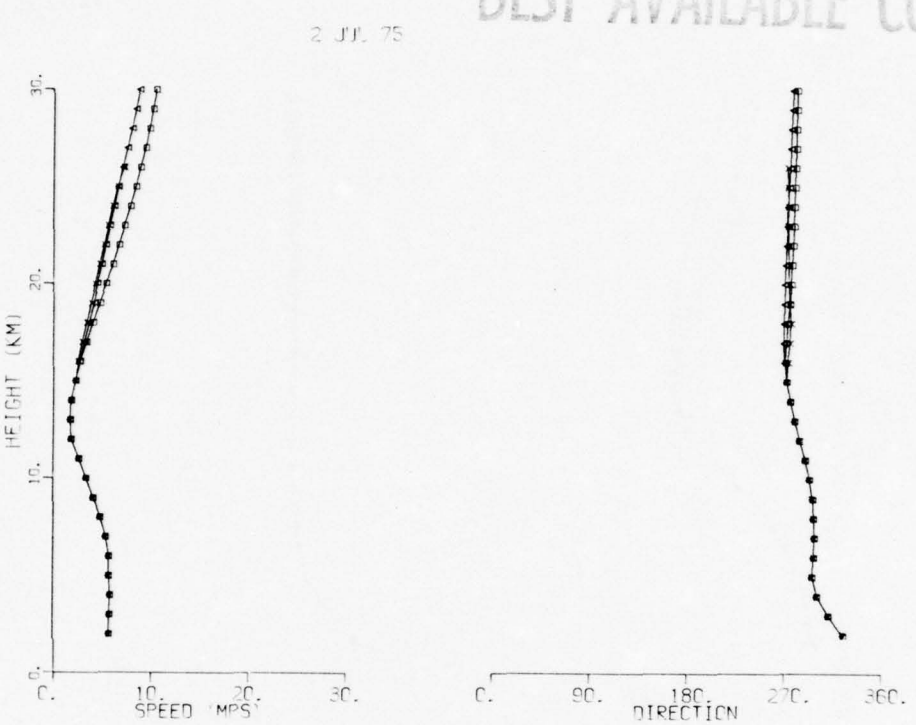


Figure A13. Satellite/radiosonde effective velocity comparison for 2 Jun 75.

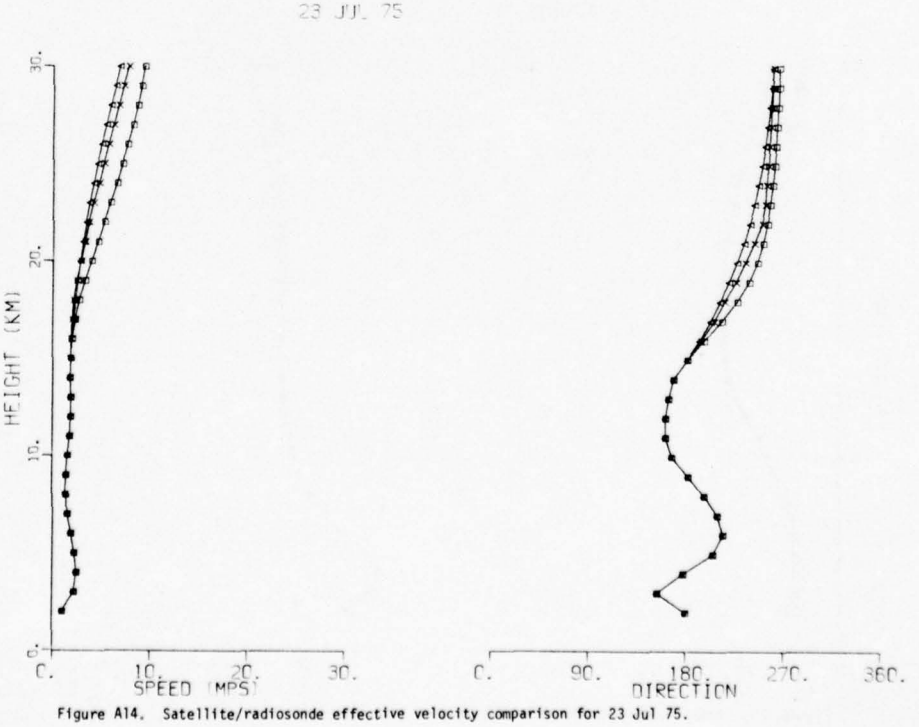


Figure A14. Satellite/radiosonde effective velocity comparison for 23 Jul 75.

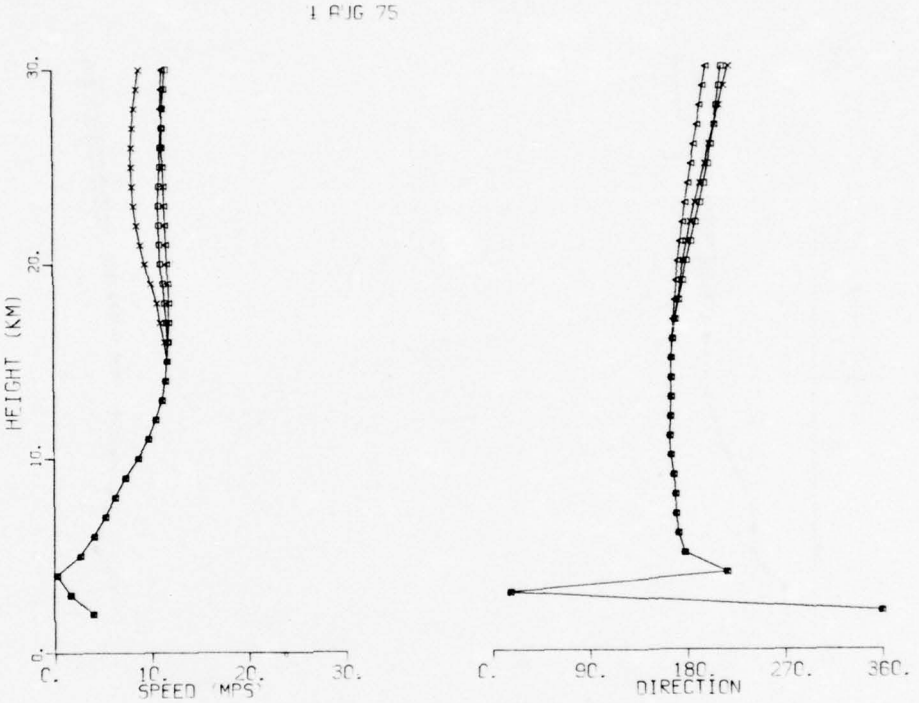


Figure A15. Satellite/radiosonde effective velocity comparison for 4 Aug 75.

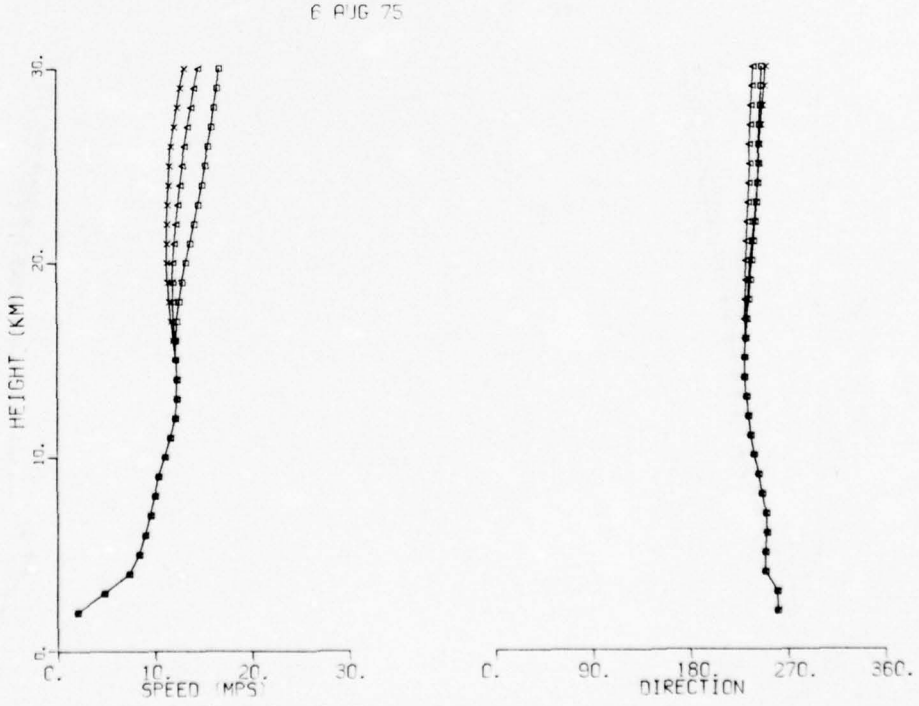


Figure A16. Satellite/radiosonde effective velocity comparison for 6 Aug 75.

7 AUG 75

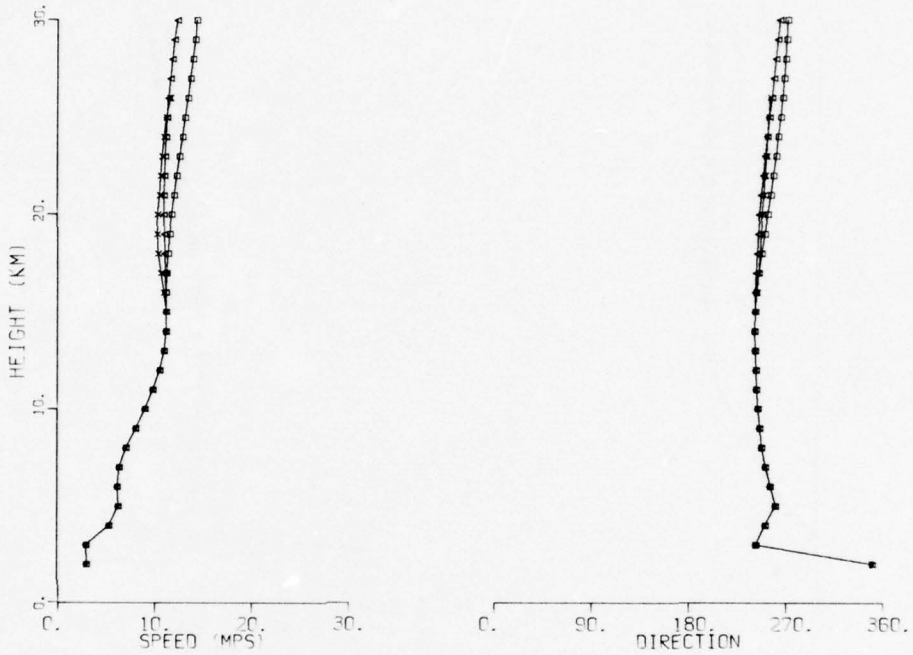
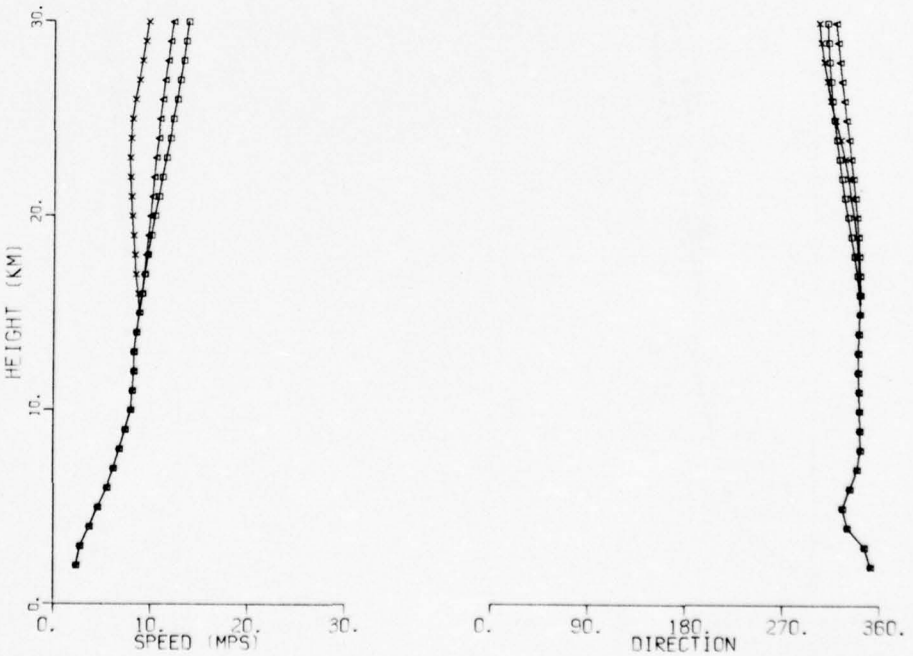


Figure A17. Satellite/radiosonde effective velocity comparison for 7 Aug 75.

19 AUG 75



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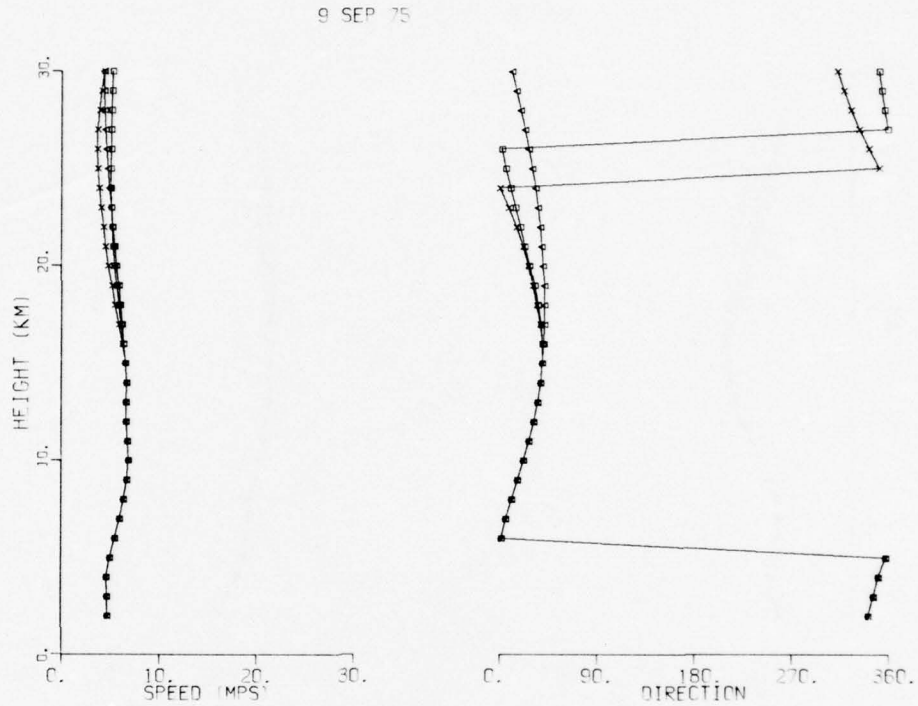


Figure A19. Satellite/radiosonde effective velocity comparison for 9 Sep 75.

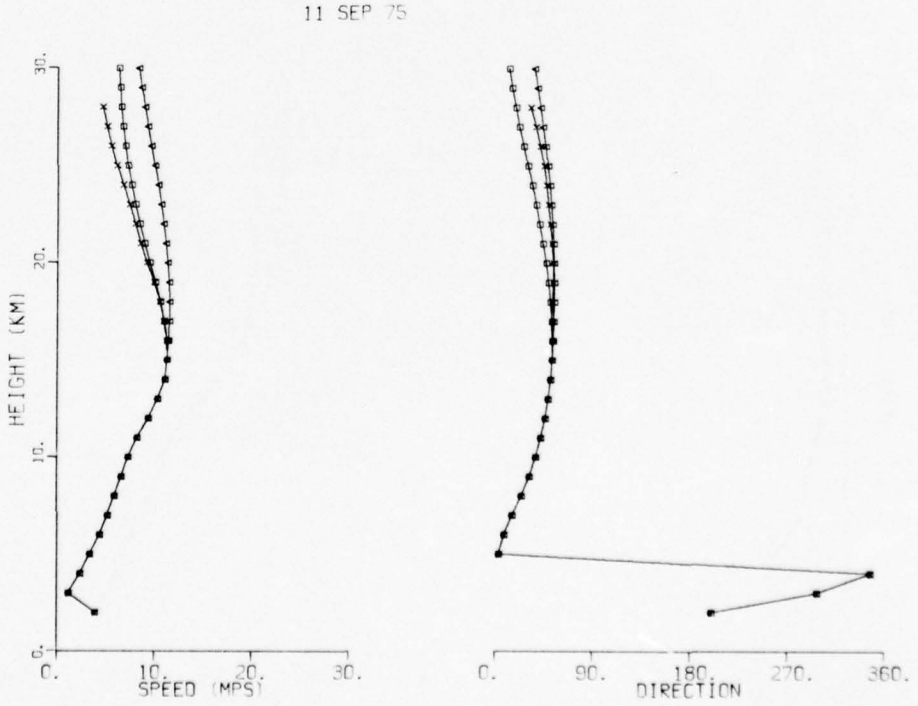
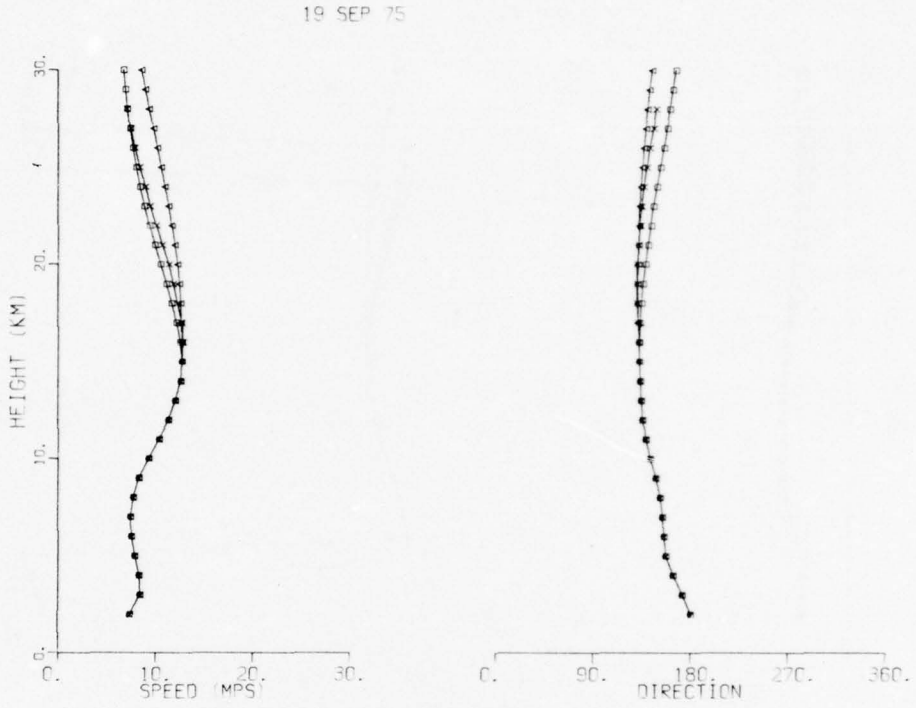


Figure A20. Satellite/radiosonde effective velocity comparison for 11 Sep 75.



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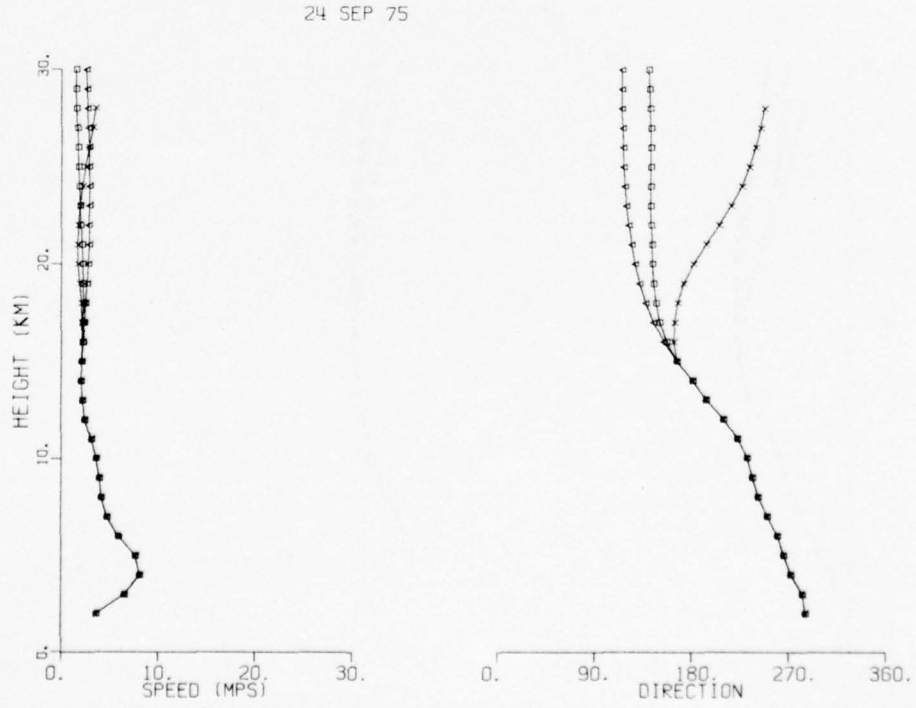

Figure A22. Satellite/radiosonde effective velocity comparison for 24 Sep 75.

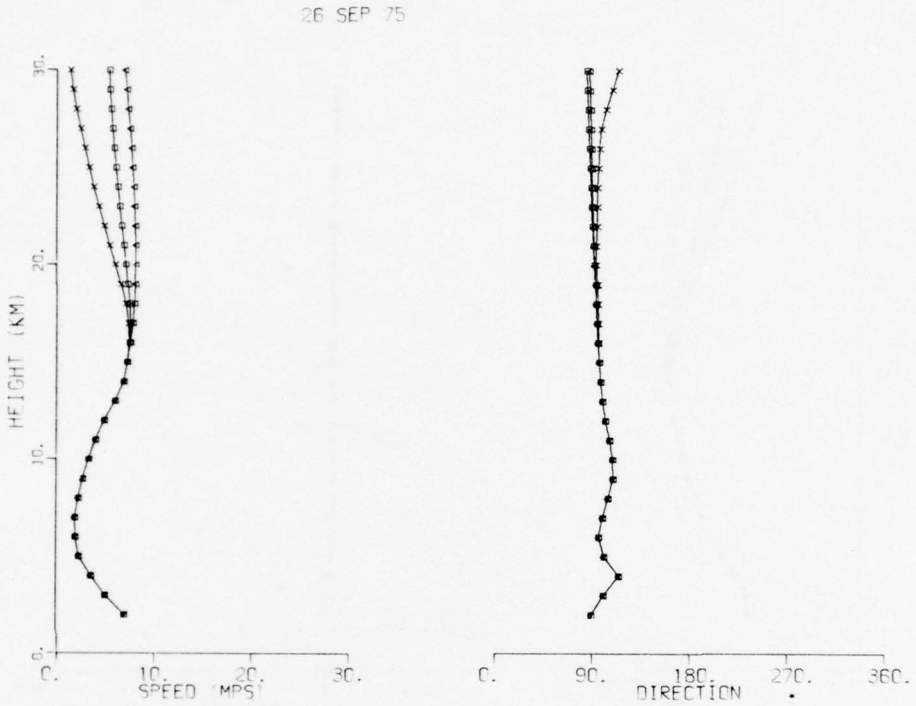

Figure A23. Satellite/radiosonde effective velocity comparison for 26 Sep 75.

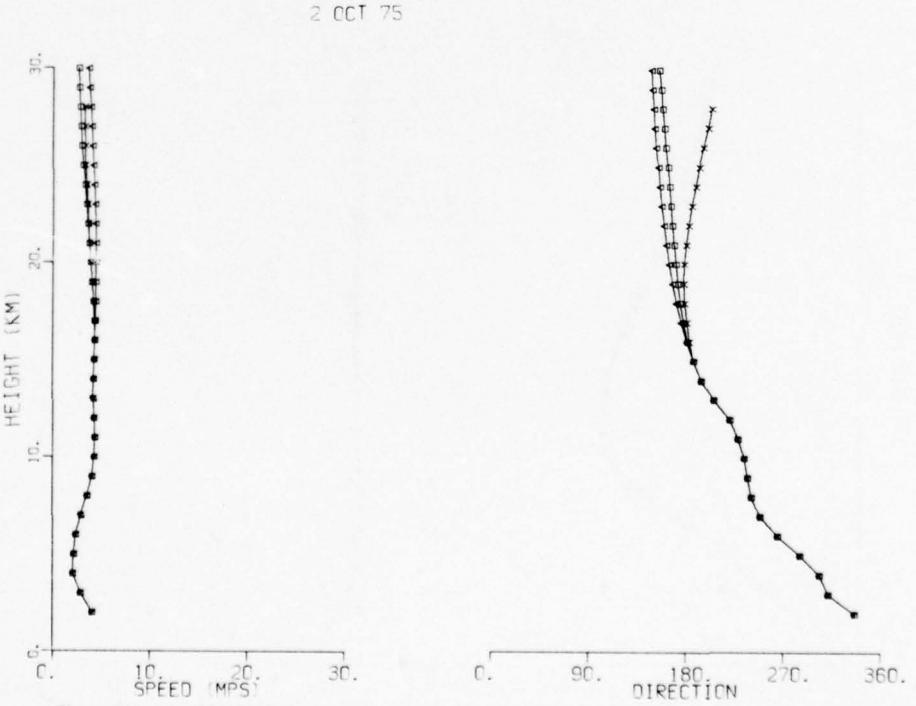

Figure A24. Satellite/radiosonde effective velocity comparison for 2 Oct 75.

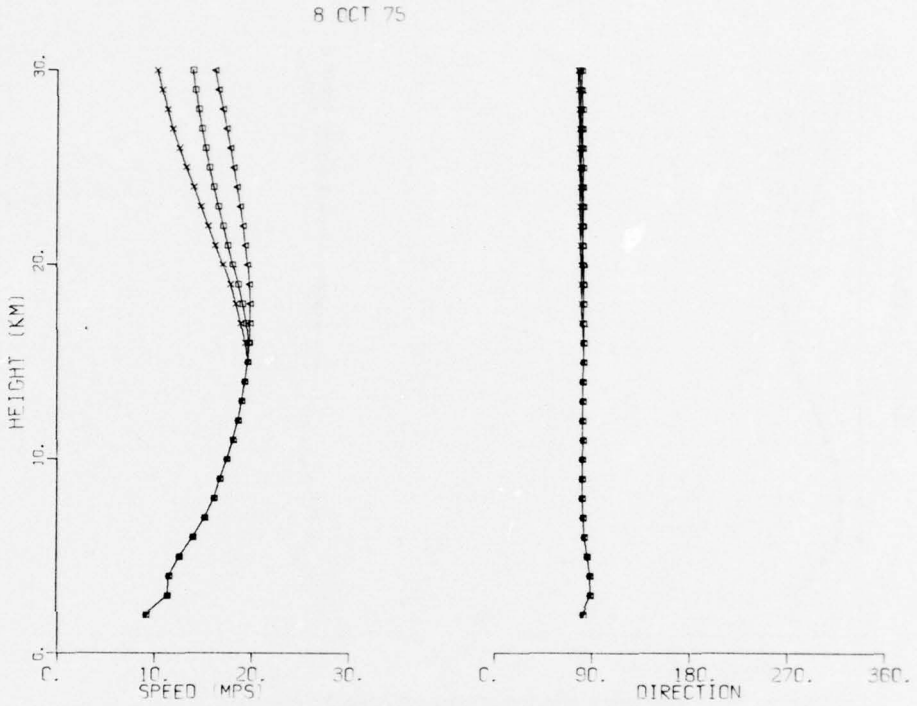

Figure A25. Satellite/radiosonde effective velocity comparison for 8 Oct 75.

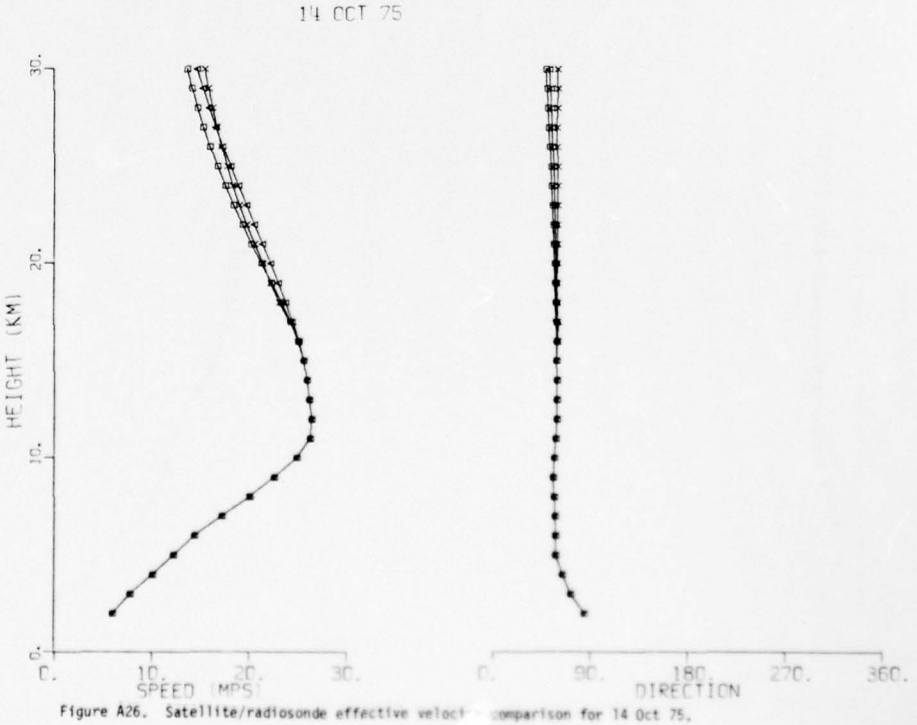

Figure A26. Satellite/radiosonde effective velocity comparison for 14 Oct 75.

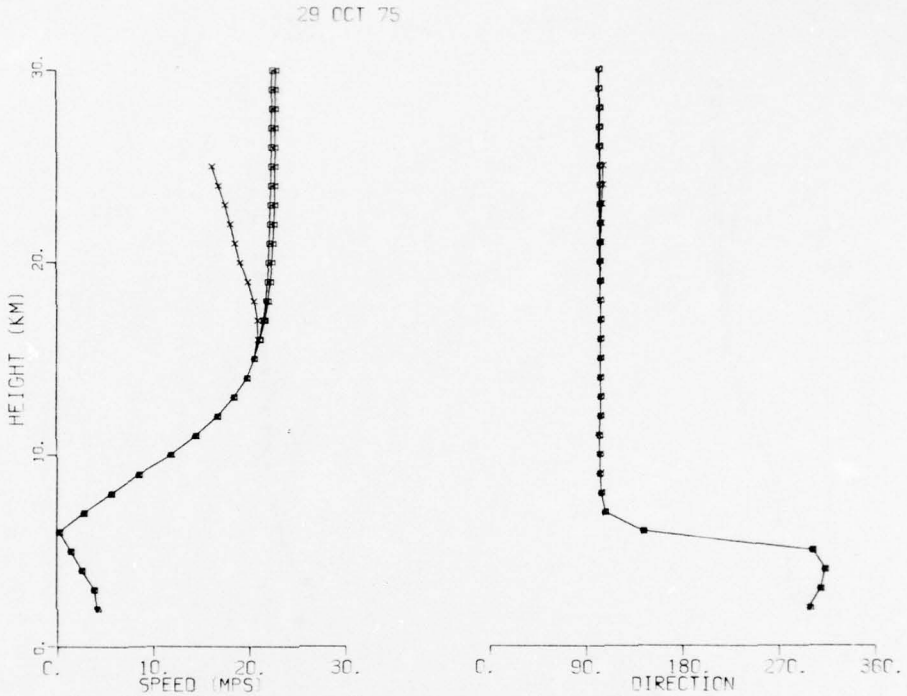

Figure A27. Satellite/radiosonde effective velocity comparison for 29 Oct 75.

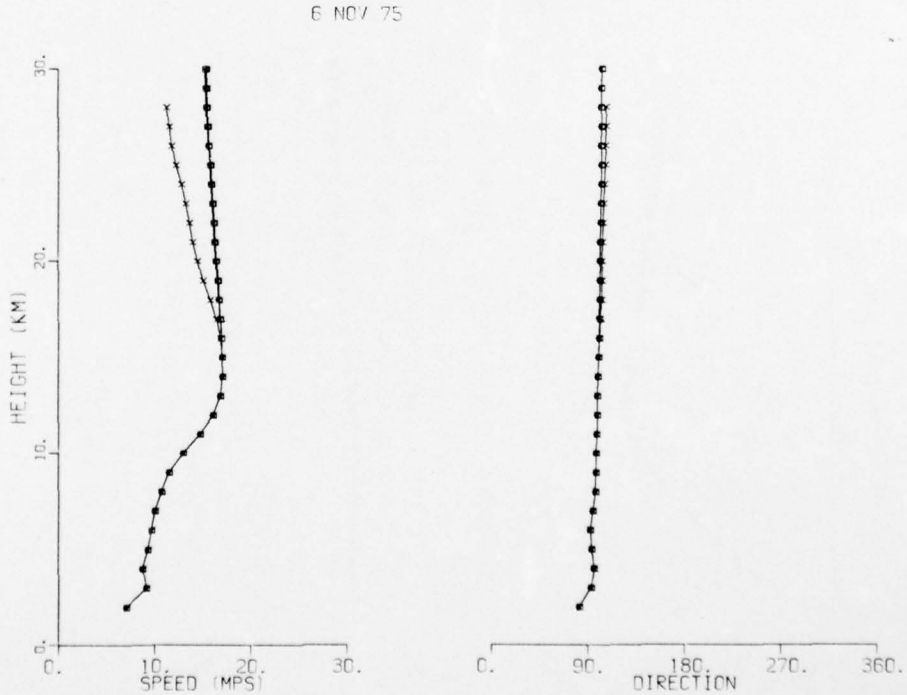

Figure A28. Satellite/radiosonde effective velocity comparison for 6 Nov 75.

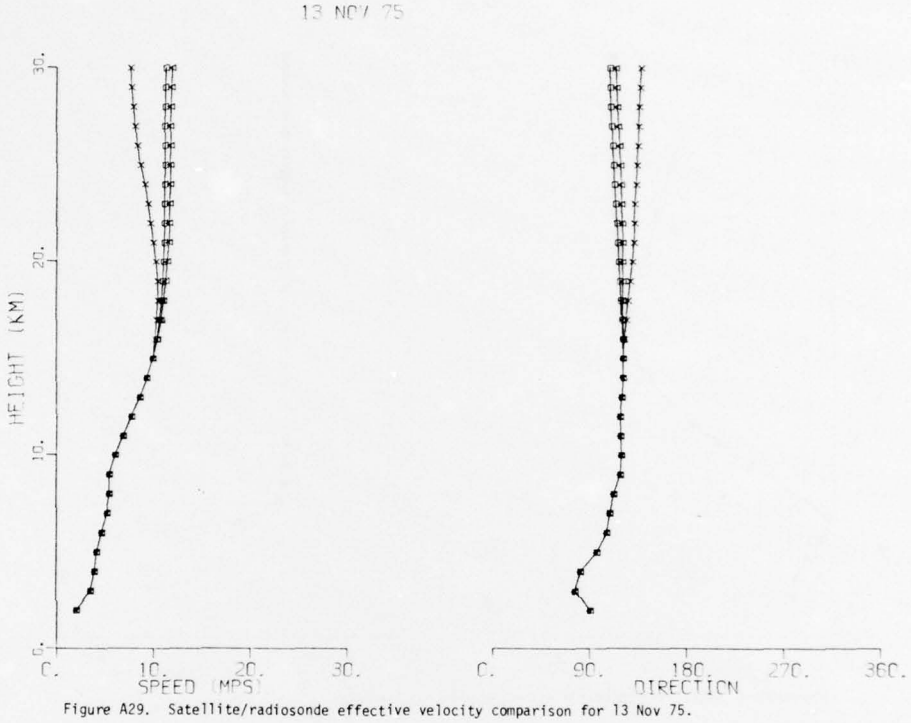

Figure A29. Satellite/radiosonde effective velocity comparison for 13 Nov 75.

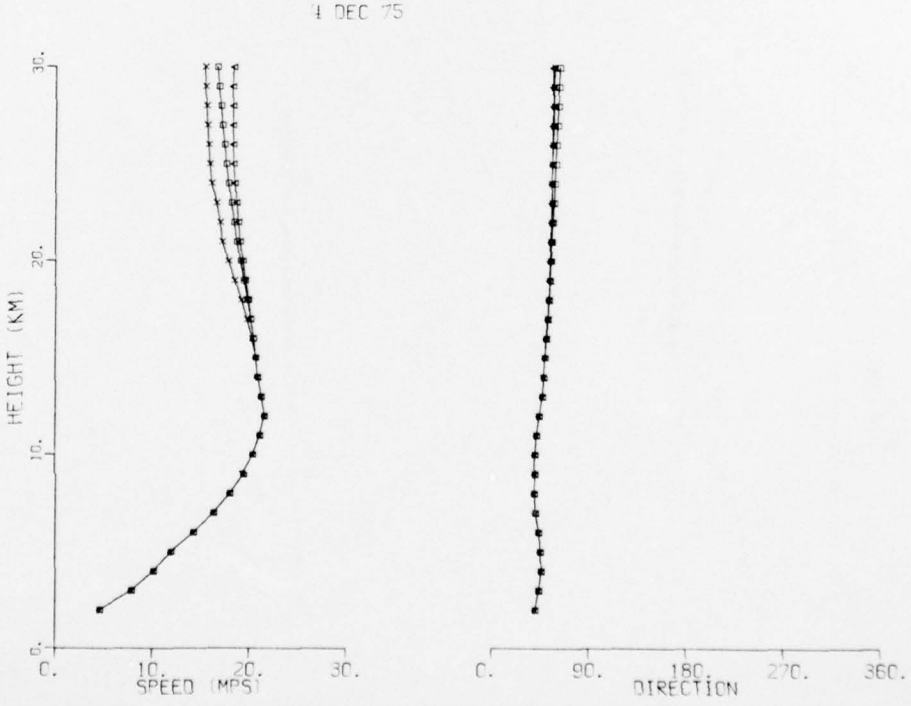

Figure A30. Satellite/radiosonde effective velocity comparison for 4 Dec 75.

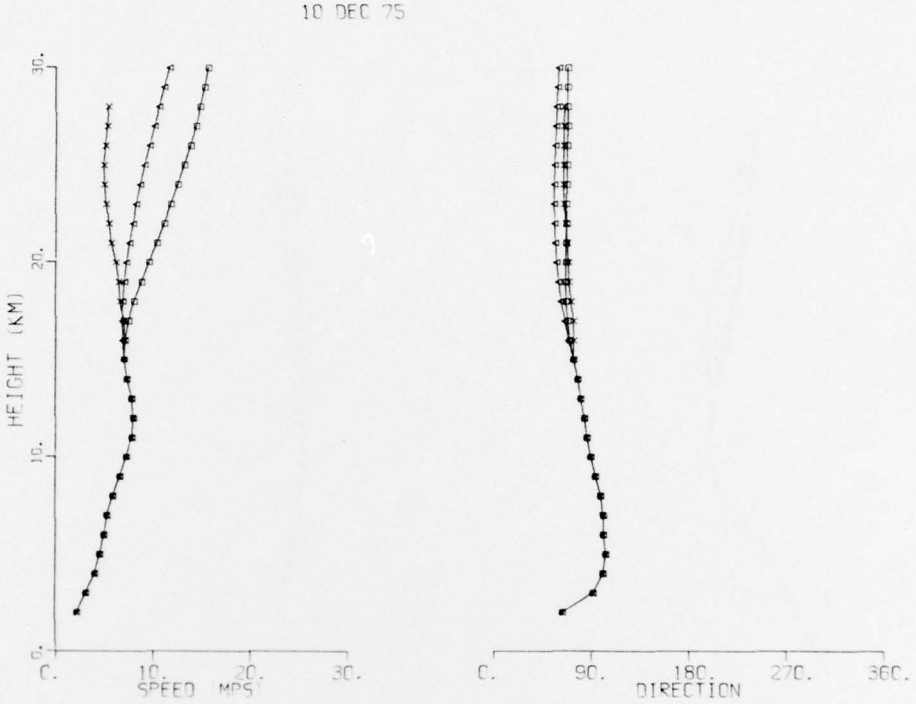

Figure A31. Satellite/radiosonde effective velocity comparison for 10 Dec 75.

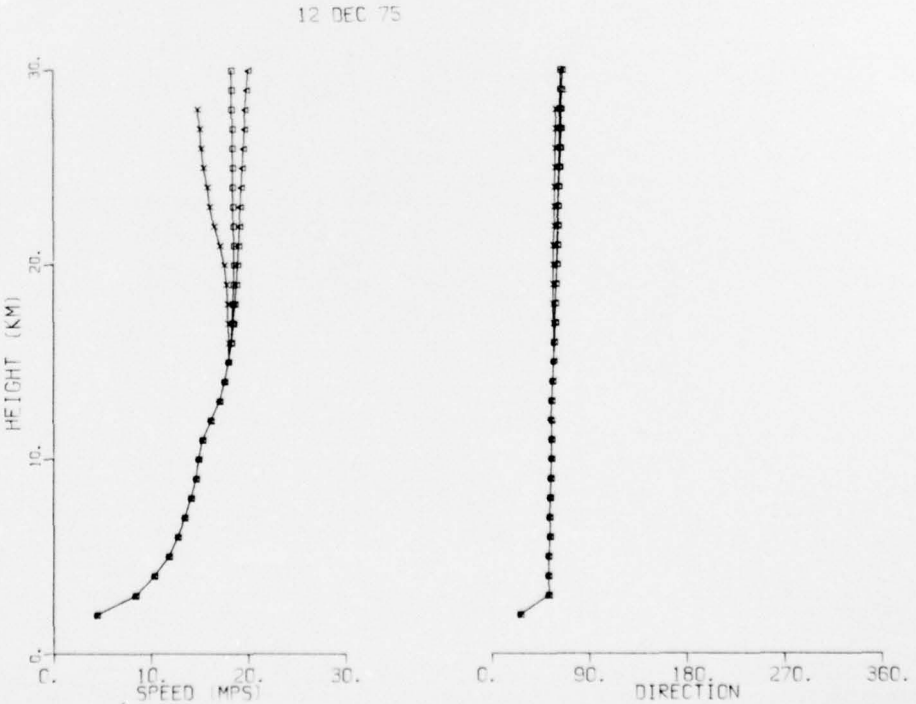

Figure A32. Satellite/radiosonde effective velocity comparison for 12 Dec 75.

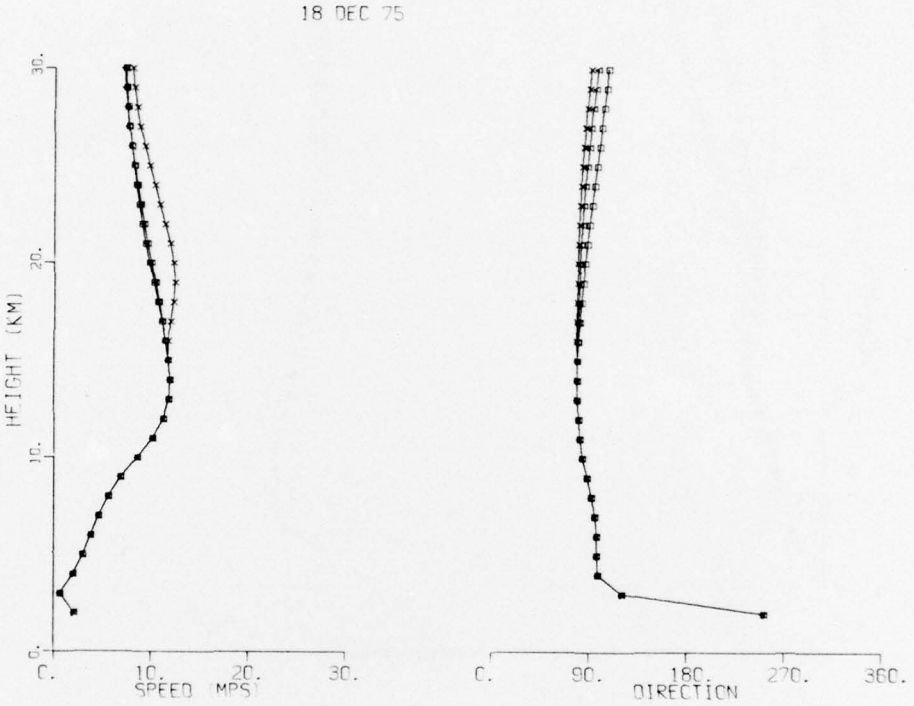

Figure A33. Satellite/radiosonde effective velocity comparison for 18 Dec 75.

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