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Hemispheric Asymmetries Detected From Satellite Ionization Gauge Measurements

JOSEPH P. McISAAC



25 February 1980

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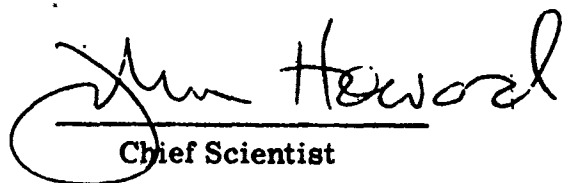
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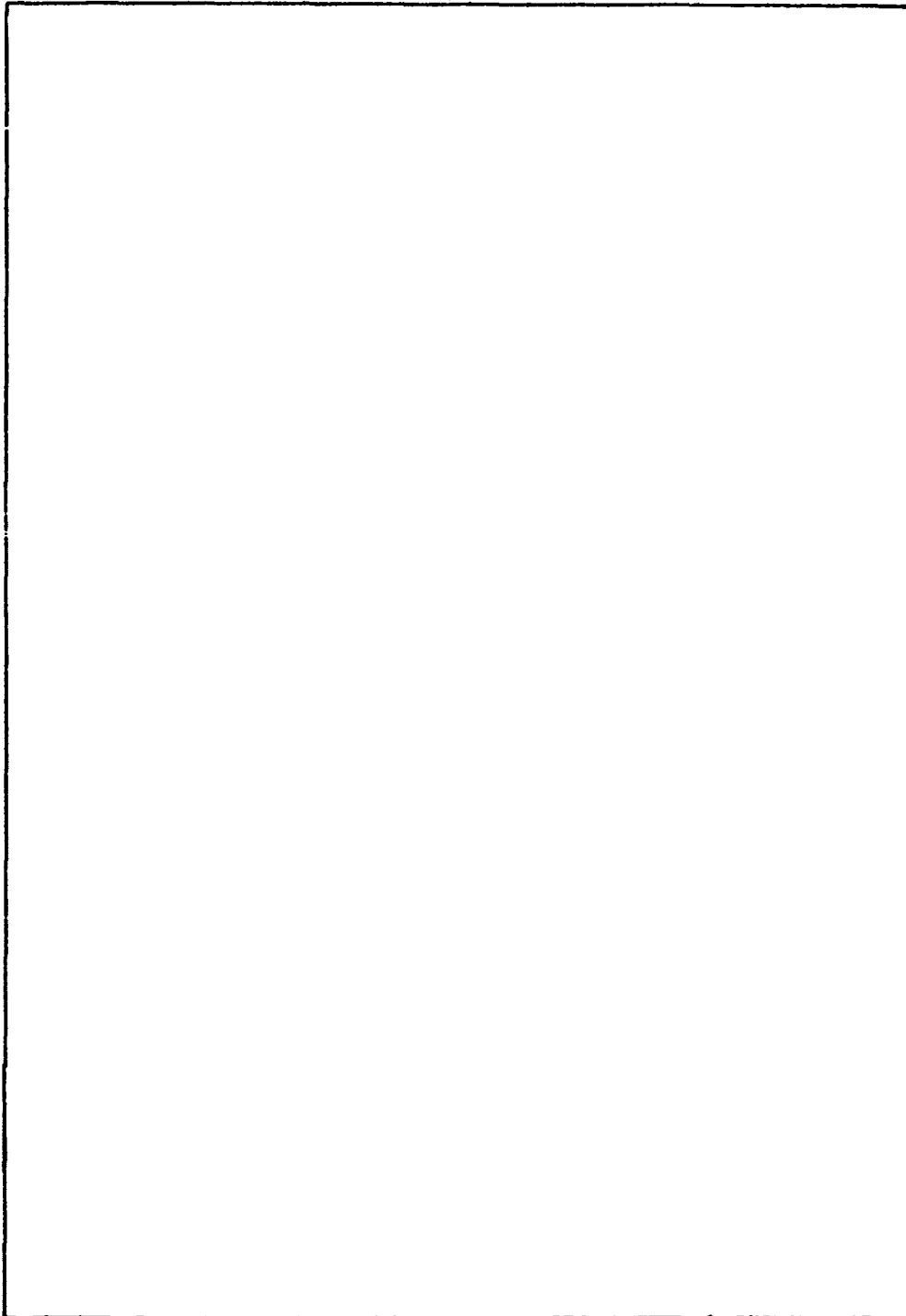
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Preface

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Hemispheric Asymmetries Detected From Satellite Ionization Gauge Measurements

1. INTRODUCTION

The ionization gauge experiment successfully flown onboard the Air Force S3-1 satellite resulted in the accumulation of an extensive file of neutral density measurements taken over the period extending from November 1974 to May 1975. This report is a first attempt at performing an analysis of the total data file gathered as a result of the S3-1 flight. The satellite orbit, experiment details, reduction procedures including computer flow charts, and representative results have been previously presented and described in detail in earlier reports.^{1, 2} The statistical analysis performed here is similar to that done by Forbes et al³ in which the statistical parameter analyzed is the ratio R between the measured density and a model density. Frequency distribution plots of R together with mean values and standard deviations are given for different levels of geomagnetic activity. The three levels of geomagnetic activity used are characterized as quiet, moderate, and

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1. McIsaac, J. P., Champion, K. S. W., McInerney, R. E., and Delorey, D. (1976) Ionization Gauge Measurements of Atmospheric Density From a Low Altitude Satellite, AFGL-TR-76-0113, AD A032373.
2. McIsaac, J. P., McInerney, R. E., and Delorey, D. (1978) Satellite Ionization Gauge Measurements of Atmospheric Density, AFGL-TR-78-0201, AD A061613.
3. Forbes, J. M., Marcos, F. A., and Gillette, D. F. (1979) An Evaluation of Thermospheric Models, AFGL-TR-79-0140, AD B036638.

disturbed levels. Comparisons of north and south hemispheric ratios are constructed over a range of different altitudes and latitudes for the levels of geomagnetic activity given above. The ratio analysis provides a more meaningful comparison of the relative hemispheric behavior of atmospheric density than would a comparison of the density measurements alone. This is so because the S3-1 satellite measurements were obtained over relatively short spatial intervals, that is, the satellite's data acquisition duty cycle routinely involved the taking of one perigee pass every alternate revolution. A perigee pass consisted of a 20-min tape recording centered around the time of perigee occurrence. The net result was a very large accumulation of measurements under diverse conditions of local time, latitude, longitude, altitude, season, and varied geophysical conditions of solar flux and geomagnetic activity. Figure 1 provides a time history of the satellite's perigee locations in terms of altitude, local time, and latitude. In this report the Jacchia 1971 model⁴ is used to form the ratio parameters, where the model values are adjusted to the appropriate conditions pertaining to the time of each measurement. An advantage in using the J71 model for forming the ratio parameter is that the latitudinal variation of the diurnal bulge in the Jacchia model is symmetrical in both hemispheres wherein the maximum density migrates north and south of the equator in phase with the sub-solar point. A disadvantage to the use of any model in determining the ratio parameter is that whatever model is employed, inherent in the ratios are the inadequacies of the modelling technique as well as unmodelled real atmospheric differences.

2. STATISTICAL PROCEDURE

The three parameters derived from the data base are the mean ratio $\bar{R}$ (mean of the ratios, data/J71), the standard deviation σ_R and N the total number of measurements. The expressions used to determine $\bar{R}$ and σ_R are as follows:

$$\bar{R} = \frac{1}{N} \sum_{m=1}^N R_m \quad (1)$$

and

$$\sigma_R = \left[\frac{1}{N-1} \sum_{m=1}^N (R_m - \bar{R})^2 \right]^{1/2} \quad (2)$$

-
4. Jacchia, L. G. (1971) Revised Static Models of the Thermosphere and Exosphere With Empirical Temperature Profiles, Spec. Rpt. 332, Smithsonian Astrophys. Observatory, Cambridge, MA.

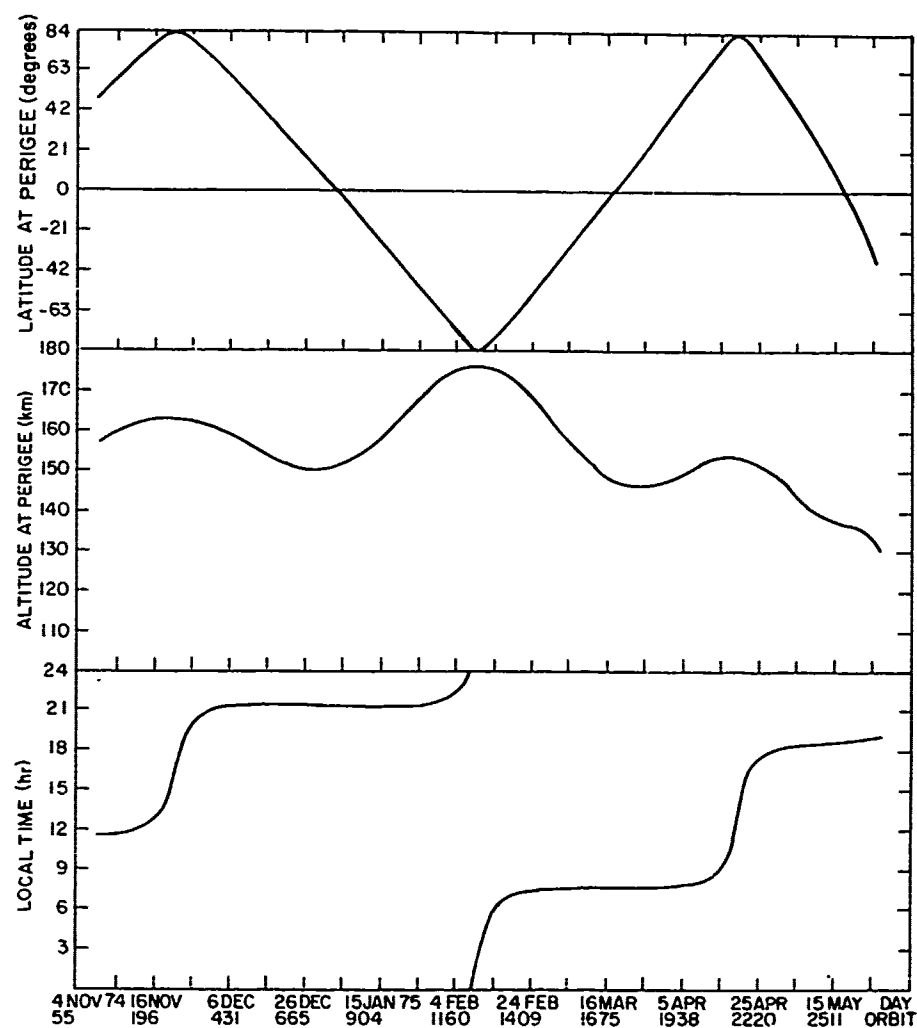


Figure 1. Satellite Position in Terms of Perigee Location, With Local Time, Altitude and Latitude of Perigee Given as a Function of Date and Orbit Number

The data base is divided into different combinations of parameters to form data sets. Within each set, the total number of points (ratios) are sorted in terms of their occurrence and partitioned into fixed intervals to form frequency histograms. For each histogram the three parameters $\bar{R}$, σ_R and N are computed and a normal distribution curve constructed which is then superimposed over the histogram. The expression used in deriving the normal distribution curve⁵ is:

5. Freeman, H. (1963) Introduction to Statistical Inference, p 140, Addison-Wesley Publishing Co.

$$Y = \frac{i N}{\sigma_R \sqrt{2\pi}} e^{-\frac{(R - \bar{R})^2}{2\sigma_R^2}} \quad (3)$$

where Y, the ordinate, is the number of occurrences per interval i. The normal curve superimposition although interesting is not too significant. Its worth is greatest when used for model evaluations. Some insight, however, is gained from histogram examination as one can determine how results are distributed about their means and detect asymmetric distributions. The real value of the statistical treatment is derived from the employment of the mean ratios and the variation of these ratios under diverse combinations of solar, geophysical and global conditions obtainable from an extensive data base. Used in this manner the statistical analysis provides an excellent technique for investigating average atmospheric behavior and examining systematic variations.

3. DATA

The data base partitioning used in this report is shown in Figure 2 and the respective sample size for each bin category is shown in Figure 3. As seen in Figure 2, the base is first separated into the two central divisions of northern and southern hemispheric data followed by a further division into three sub-levels. The first sub-level separates hemispheric data according to the degree of geomagnetic activity; quiet, moderate, or disturbed. Or more precisely, the geomagnetic division is determined by the value of the K_p planetary index (6 hr lag). The range limits for the activity levels used are:

Quiet activity (low):	$0 \leq K_p \leq 3$
Moderate activity (moderate):	$3 < K_p \leq 5$
Disturbed activity (high):	$5 < K_p \leq 9$

The second and third sub-level breakdowns are by altitude and latitude respectively. These breakdowns are determined in a straightforward manner reflecting a compromise between adequate statistical sample sizes and a desire to avoid an excessive number of computer runs needed to compute numerical values and to run off plotting routines. The partitioning as given in Figure 2 involved the generation of 135 frequency histograms, a portion of which are shown in the appendix. In addition to the above, various bins were combined to provide larger bins resulting in more histogram computations. An example of this is shown in Table 1 where combinations of several bins were constructed in order to develop a more meaningful comparison between gauge results shown here and those obtained from another technique.

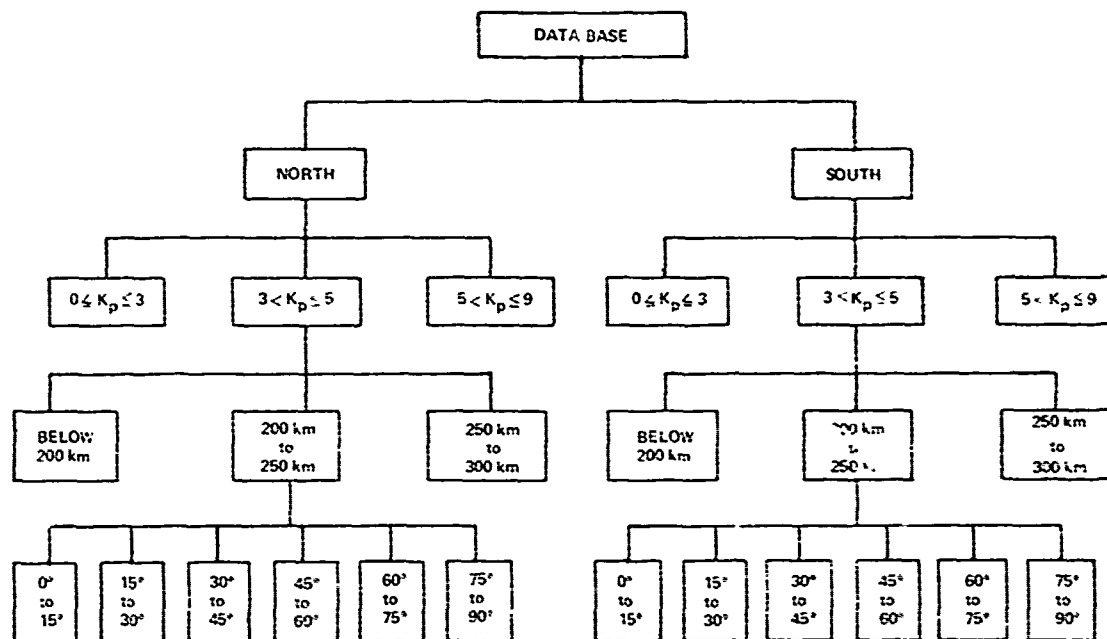


Figure 2. Partition Diagram Outlining Data Base Division

Histograms with normal curves superimposed are included in the appendix together with legend data consisting of values of $\bar{R}$, σ_R , N and the normalized Y value as well as the parameters that define the bin confines. Only a portion of the total number of histograms generated as a result of the bin partitioning has been included in the appendix. However, that portion is representative of the full histogram set. Figures, which are discussed in detail in the following section, representing mean behavior of neutral density over the globe, were developed that show results as a function of altitude under different conditions of geomagnetic activity and over several altitude regions.

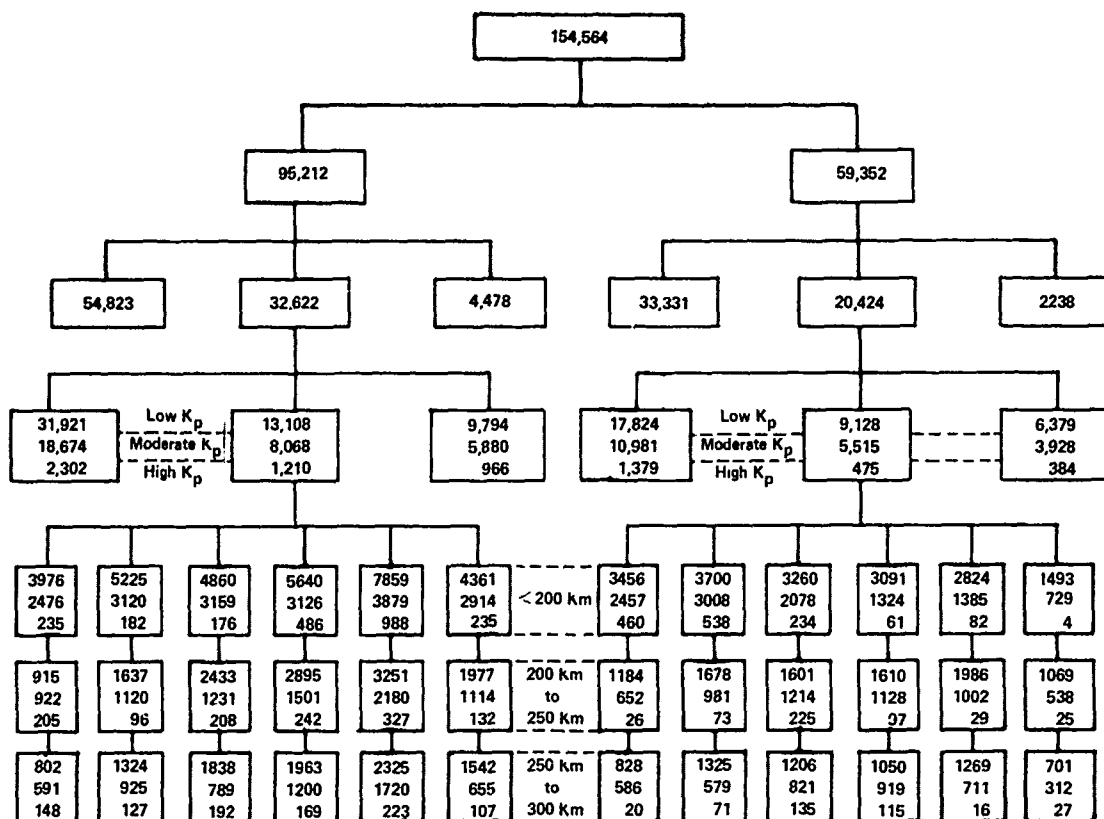


Figure 3. Sample Sizes Resulting From Data Base Partition Given in Figure 2.

4. DISCUSSION

Figure 4 shows differences between mean ratios at corresponding latitudes in the north and south hemispheres plotted as a function of latitude. The ratio differences are given for three levels of geomagnetic conditions; quiet, moderate, and disturbed, and for three altitude regions. Among the prominent features seen in this set of curves, perhaps most outstanding is the total absence of positive differences below 200 km over all latitudes. Only at higher altitudes and in regions about the equator do we observe positive ratio differences. The shapes of the difference curves above 200 km, also appear to develop into similar basic patterns that tend to be accentuated as the level of geomagnetic increases. This structural pattern is manifested in greater detail in the later figures.

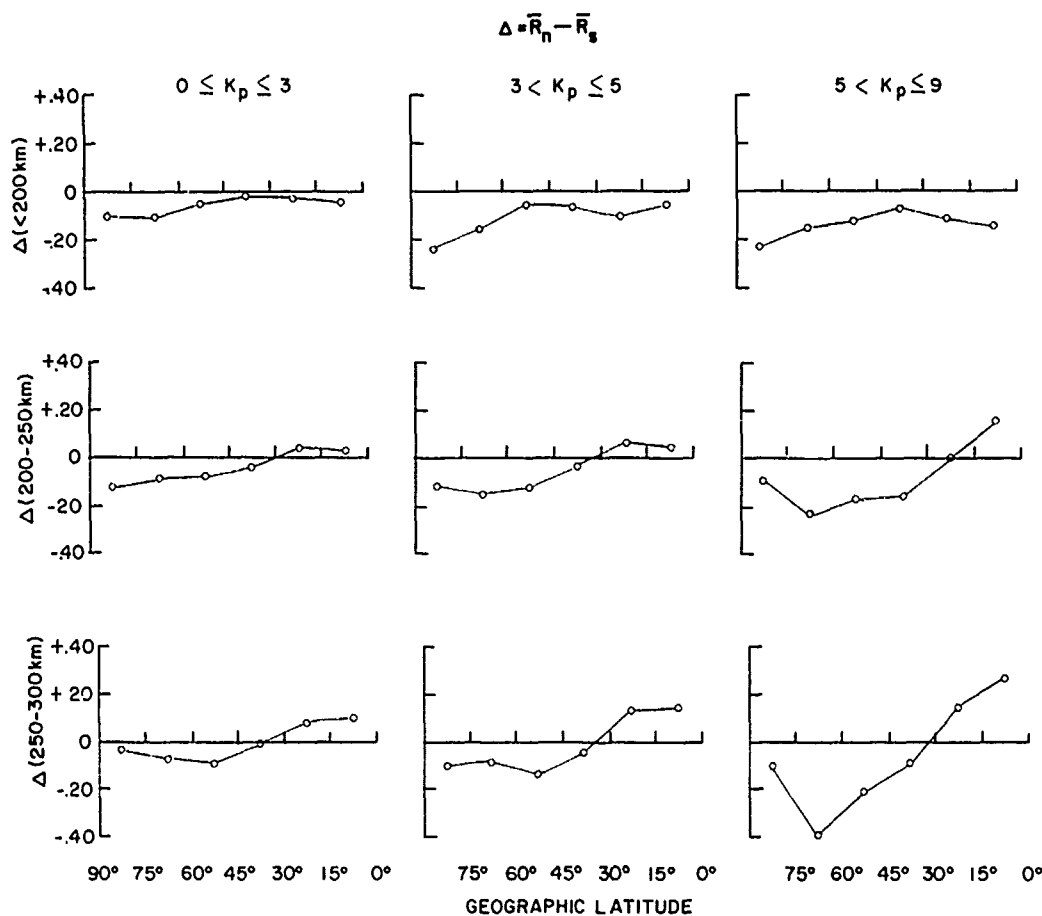


Figure 4. Differences in Average Density Ratio vs Latitude for Three Geomagnetic Groups. Top left to right: low, moderate and high activity. Top for altitudes less than 200 km, middle for altitudes 200 to 250 km and bottom for altitudes 250 to 300 km

The overall negative differences seen in Figure 4 indicate that atmospheric density over the period of this study was greater over the southern hemisphere, which is in agreement with what was found by Barlier et al.^{6,7} His examination of over 12,000 values obtained from drag data of satellites indicated that densities in the southern hemisphere were systematically larger than or equal to the corresponding values in the northern hemisphere. However, in terms of atomic oxygen

6. Barlier, F., Perret, D., and Jaek, C. (1973) Structure of the neutral atmosphere between 150 and 500 km, Space Research XIII:349-355.
7. Barlier, F., Bauer, P., Jaek, C., Thuillier, G., and Kockharts, G. (1974) North-south asymmetries in the thermosphere during the last maximum of the solar cycle, J. Geophys. Res. 79:5273-5285.

concentrations, it⁷ was shown that at 200 km the concentrations tended to accumulate over the cold hemisphere (winter) at the lower latitudes where seasonal effects did not dominate, which also is not in disagreement with what we observe in Figure 4.

It is interesting to note that the Jacchia 71 model treats hemispheric variations in a uniform manner, that is, the diurnal maximum migrates symmetrically about the equator in phase with the declination angle. In forming our ratios the normalization of the density measurement to the J71 model value should, in practice, compensate somewhat for the seasonal effects due to the migration of the diurnal bulge across the equator. As the southern hemisphere measurements in our data base were taken almost entirely during the summer season, the appropriate southern hemispheric J71 model values will, all other factors being equal, be greater than the J71 northern hemispheric values (spring/fall); consequently, from the model viewpoint the north hemispheric ratios should be larger than south hemispheric ratios.

Figure 5 illustrates global latitudinal structure more clearly; shown here are ratios of measured densities to model densities given as a function of north and south latitudes for three altitude regions and different levels of geomagnetic activity. In the low altitude (< 200 km) curve set, northern ratios are fairly constant while southern ratios increase in the direction of the pole. Again, southern ratios are seen to be considerably larger than northern ratios. There appears to be little noticeable effect from increasing geomagnetic activity at the lower altitudes. For the higher altitude curve sets, two prominent features stand out; the enhancement in the southern auroral region during periods when $K_p > 5$ and the north-south bulge/trough structure located about the equator. It is also interesting to note that for the most part the ratio magnitudes appear to be inversely proportional to the level of geomagnetic activity indicating the J71 model overcompensates for geomagnetic effects. The auroral region enhancement is present only in the southern hemisphere auroral region where, unlike the data below 200 km, the density is seen to decrease poleward of the auroral region. This same effect was observed by Forbes et al⁸ from accelerometer data taken onboard the NASA AE-C satellite, however, in that report north and south latitudes were not separated. This particular feature found here and in the AE-C results was deduced in both cases from rather small statistical data samples, therefore the observations regarding the auroral region feature; enhancement and subsequent poleward drop-off, should be viewed as inconclusive. Rice and Sharp⁹ have reported observations

8. Forbes, J. M., Marcos, F. A., and Champion, K. S. W. (1977) Lower thermosphere response to geomagnetic activity, Space Res. XVIII:173-176.
9. Rice, C. J., and Sharp, L. R. (1977) Neutral Atmospheric Waves in the Thermosphere and Tropospheric Weather System, SAMSO-TR-77-98.

from the ion gauge measurements performed onboard the AE-C spacecraft wherein they observed an asymmetric auroral region feature consisting of a southern latitude region (40° to 68° S) in which there was a high occurrence of wavelike structural features not seen in the corresponding northern latitude region.

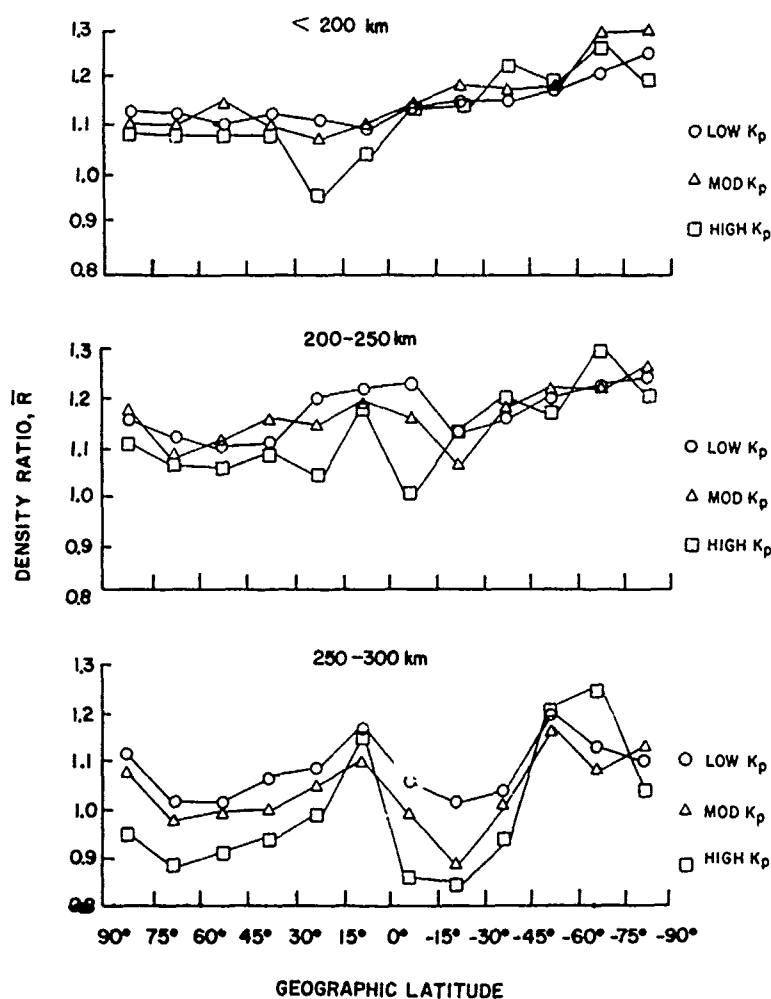


Figure 5. Average Density Ratio vs Geographic Latitude for Three Altitude Regions < 200 km, 200 to 250 km, and 250 to 300 km. Open circles are low K_p values, triangles are for moderate K_p values and squares are for high K_p values

The other prominent structure observation apparent in Figure 5 is the equatorial asymmetry, a north-south, bulge/trough formation located about the equator. This bulge/trough equatorial feature is highly visible at the higher altitudes but seems to disappear below 200 km. Certain aspects or characteristics of this structure are more evident in Figure 6 where ratio curves are grouped together in sets according to the level of geomagnetic activity. Foremost, is the phase relationship between ratios below 200 km to those above 200 km. The equatorial structure below 200 km is clearly out of phase with the structure above 200 km. Below 200 km at northern equatorial latitudes there is a trough in place of the bulge seen at altitudes above 200 km. In a like manner, the deep southern latitude trough at high altitudes is reversed at low altitudes and although one does not observe a well defined bulge at the southern latitudes below 200 km, one can observe a shape that suggests a bulge merged with other effects. A second characteristic of the equatorial structure apparent in both Figures 5 and 6 is that the minimum of the southern latitude trough appears to drift towards the equator during periods of increase geomagnetic activity ($K_p > 5$).

Equatorial located structure of a similar nature has been reported elsewhere^{10, 11} and, although no explanations were given by the authors, in each case the structure was characterized as persistent as opposed to being of a transient nature. The persistent nature of the structure described here requires a non-transient type mechanism be invoked if a satisfactory explanation is to be found. Recall that the ratio values are mean values representative of data obtained over a six month period. One possible explanation would involve solar tidal variations of density. The north-south pattern observed in our data set is suggestive of tidal mode propagation into the thermosphere with antisymmetric, semidiurnal modes dominant. Semidiurnal tidal effects have been theoretically examined by Hong and Lindzen¹² and large equatorial latitudinal variations, particularly at solar cycle minimum, have been predicted. Our data base lacks adequate local time coverage to properly evaluate tidal oscillations as an explanation. However, although not shown in this report, the data given in Figures 5 and 6 have been further constrained to periods of constant local time (± 1 hr). The two local time periods were at 0700 and 2200 hours. Under this added constraint of constant local time the observed structural features remained basically unchanged indicating that the variations in the data are not due to differences in local time.

10. Philbrick, C. R., and McIsaac, J. P. (1972) Measurements of atmospheric composition near 400 km, Space Research XVI:743-749.

11. Ching, B.K., and Carter, V. L. (1974) Ion gauge measurements of latitudinal density variations at night, Geophys. Res. Lett. 1:93-96.

12. Hong, S. S., and Lindzen, R. S. (1976) Solar semidiurnal tide in the thermosphere, J. Atmos. Sci. 33:135-153.

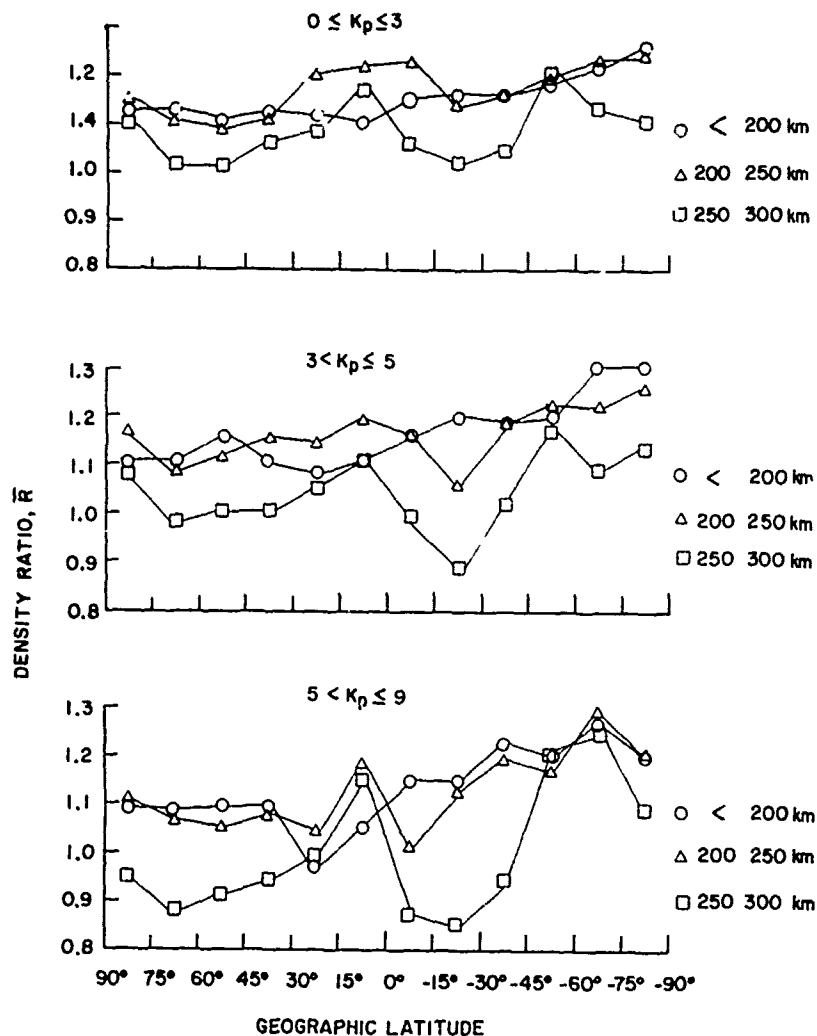


Figure 6. Average Density Ratio vs Latitude for Three Levels of Geomagnetic Activity as Shown. Altitude symbols are circles < 200 km, triangles 200 to 250 km and squares 250 to 300 km

A point that requires investigation is the absence of the structural features discussed above from the accelerometer results¹³ obtained onboard the same S3-1 spacecraft. In an attempt to ascertain the reason for this discrepancy the gauge ratio parameters given in this report were recomputed to coincide with the

13. Marcos, F.A., McInerney, R.E., and Fioretti, R.W. (1978) Variability of the Lower Thermosphere Determined From Satellite Accelerometer Data, AFGL-TR-78-0134, AD A058-982.

partitioning used by Marcos et al in their processing of accelerometer results. This new partitioning of the gauge data involved combining the two altitude bins of < 200 km and 200 to 250 km to form a single bin extending from perigee to 250 km and expanding the latitude spans by 15° to form 30° bins. Putting results from both techniques in the same format provides a means for directly comparing results. This is done in Table 1 where results from both data sets are listed. Also, the data shown in Table 1 are plotted as a function of latitude (Figure 7). Referring to Figure 7, one readily observes that the figure shows two pairs of curves for two levels of geomagnetic activity in which results from each technique are compared. The upper and lower curve sets can be referred to as representing high and low geomagnetic, respectively, although the K_p limits for both sets are slightly different. Marcos et al¹³ omitted data in the K_p range from $2.6 \leq K_p \leq 3.6$ in order to emphasize the differences between the high and low K_p data groups. Also, in that study, there was no separate geomagnetic activity category that could be characterized as moderate, hence, a comparison of results for moderate geomagnetic conditions is omitted. Upon examination of Figure 7 one can observe that the agreement between results obtained from both techniques is generally good, particularly in the equatorial region where the structure reported in this report has been observed. The region of greatest discrepancy is in the high southern latitudes which is also the region where statistics are poorest. Accelerometer results also indicate higher southern hemispheric densities. One fact made immediately evident from Figure 7 is that when the ionization gauge results are configured in the format used in the accelerometer study, the results reveal little of the structure reported here; hence, indicating no general disagreement between the two techniques. Two reasons are offered for the lack of structure seen in Figure 7, namely, the bin limits were too broad (latitude and altitude) and, certainly a major factor, the altitude dependent phase reversal behavior reported here. By grouping data from below 200 km with data above 200 km, the result is a tendency to smear or negate the structural differences between the two altitude regions.

Table 1. Comparison of Accelerometer and Ionization Gauge Results Obtained Onboard the S3-1 Satellite

Ionization Gauge Results				Accelerometer Results			
Latitude Range	K _p Range	Number of Points	Mean Value	Latitude Range	K _p Range	Number of Points	Mean Value
90°N - 60°N	0 ≤ K _p ≤ 3	17,448	1.13	90°N - 60°N	0 ≤ K _p ≤ 2.6	3697	1.00
60°N - 30°N	0 ≤ K _p ≤ 3	15,828	1.12	60°N - 30°N	0 ≤ K _p ≤ 2.6	2177	1.02
30°N - 0°	0 ≤ K _p ≤ 3	11,753	1.13	30°N - 0°	0 ≤ K _p ≤ 2.6	1204	1.11
0° - 30°S	0 ≤ K _p ≤ 3	10,018	1.15	0° - 30°S	0 ≤ K _p ≤ 2.6	1529	1.14
30°S - 60°S	0 ≤ K _p ≤ 3	9,562	1.17	30°S - 60°S	0 ≤ K _p ≤ 2.6	1717	1.09
60°S - 90°S	0 ≤ K _p ≤ 3	7,372	1.23	60°S - 90°S	0 ≤ K _p ≤ 2.6	988	1.05
90°N - 60°N	5 < K _p ≤ 9	1682	1.09	90°N - 60°N	4 < K _p ≤ 9	2163	1.03
60°N - 30°N	5 < K _p ≤ 9	1112	1.08	60°N - 30°N	4 < K _p ≤ 9	1945	1.02
30°N - 0°	5 < K _p ≤ 9	718	1.08	30°N - 0°	4 < K _p ≤ 9	1424	1.07
0° - 30°S	5 < K _p ≤ 9	1097	1.15	0° - 30°S	4 < K _p ≤ 9	1180	1.10
30°S - 60°S	5 < K _p ≤ 9	617	1.20	30°S - 60°S	4 < K _p ≤ 9	1165	1.05
60°S - 90°S	5 < K _p ≤ 9	140	1.25	60°S - 90°S	4 < K _p ≤ 9	396	1.11

Altitude range is from 250 km to perigee.

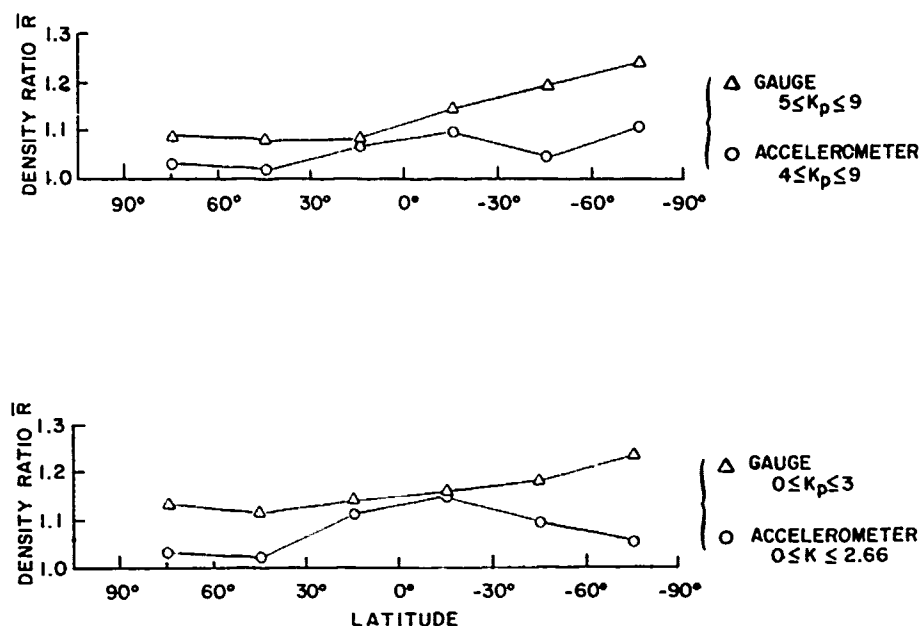


Figure 7. Ratios vs Latitude for Two S3-1 Satellite Experiments. Triangles are ionization gauge values, circles are accelerometer values

5. SUMMARY AND CONCLUSION

A statistical analysis of thermospheric density measurements, in excess of 150,000 values, described in this report has revealed hemispheric asymmetries. In general, at least for the period of this study, southern hemispheric densities are found to be greater relative to the Jacchia 71 model than northern hemisphere densities. Density enhancements are observed in the southern auroral region under disturbed geomagnetic conditions ($K_p > 5$) that are not seen in the corresponding north auroral region. A non-transient, equatorial-located, structural feature is observed. This structure which is asymmetric about the equator, is seen to intensify with increasing altitude and geomagnetic activity. A tentative explanation is advanced that attributes the equatorial behavior to atmospheric tidal effects. Further analysis of these results is planned to investigate seasonal influence upon the observed equatorial structure and determine the fine spatial resolution, as well.

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References

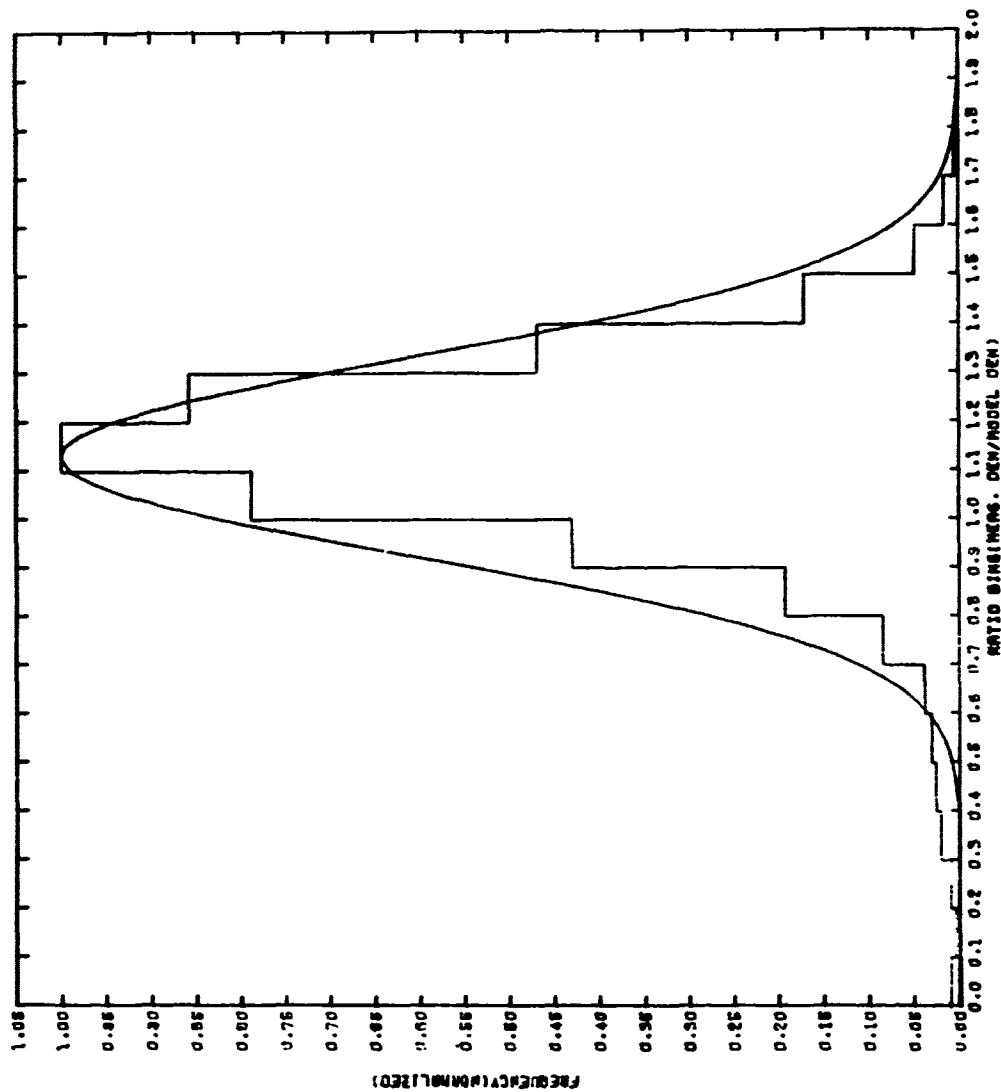
12. Hong, S.S., and Lindzen, R.S. (1976) Solar semidiurnal tide in the thermosphere, J. Atmos. Sci. 33:135-153.
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Appendix A

Histograms of Frequency Distributions of Ratios of Measurements/Jacchia 71 Model

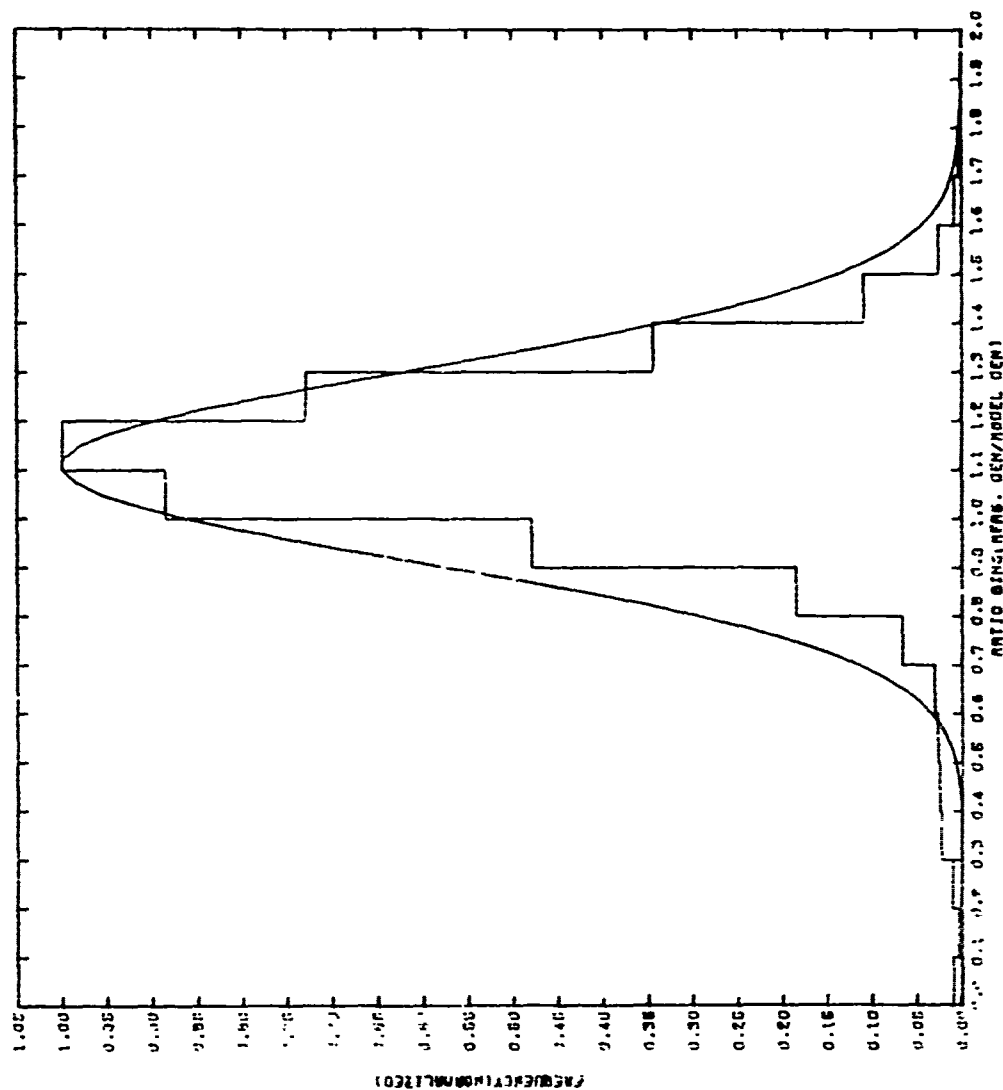
Number	Altitude Range	Latitude Range	K_p Range
A1	Perigee to 300 km	90°N to 90°S	0 to 9
A2	"	90°N to 0°	"
A3	"	90°S to 0°	"
A4	"	90°N to 0°	5 to 9
A5	"	90°S to 0°	"
A6	250 Km to 300 km	90°N to 0°	3 to 5
A7	"	90°S to 0°	"
A8	"	15°N to 0°	0 to 3
A9	"	15°S to 0°	"
A10	"	30°N to 15°N	5 to 9
A11	"	30°S to 15°S	"
A12	"	60°N to 45°N	3 to 5
A13	"	60°S to 45°S	"
A14	200 Km to 250 km	15°N to 0°	"
A15	"	15°S to 0°	"
A16	"	75°N to 60°N	5 to 9
A17	"	75°S to 60°S	"
A18	"	90°N to 75°N	3 to 5
A19	"	90°S to 75°S	"
A20	Perigee to 200 km	30°N to 15°N	5 to 9
A21	"	30°S to 15°S	"
A22	"	45°N to 30°N	3 to 5
A23	"	45°S to 30°S	"
A24	"	75°N to 60°N	0 to 3
A25	"	75°S to 60°S	"
A26	"	90°N to 75°N	3 to 5
A27	"	90°S to 75°S	"

100 DATA C.0- 8.0
 KP LINES 2.0000
 LAT/LN -20.00 40.00
 BATH MEASUREMENTS
 MAX Y VAL 50000.00
 MEAN VALUE 1.13
 ST. DEV. 2.00E-01
 TOTAL POINTS 154504

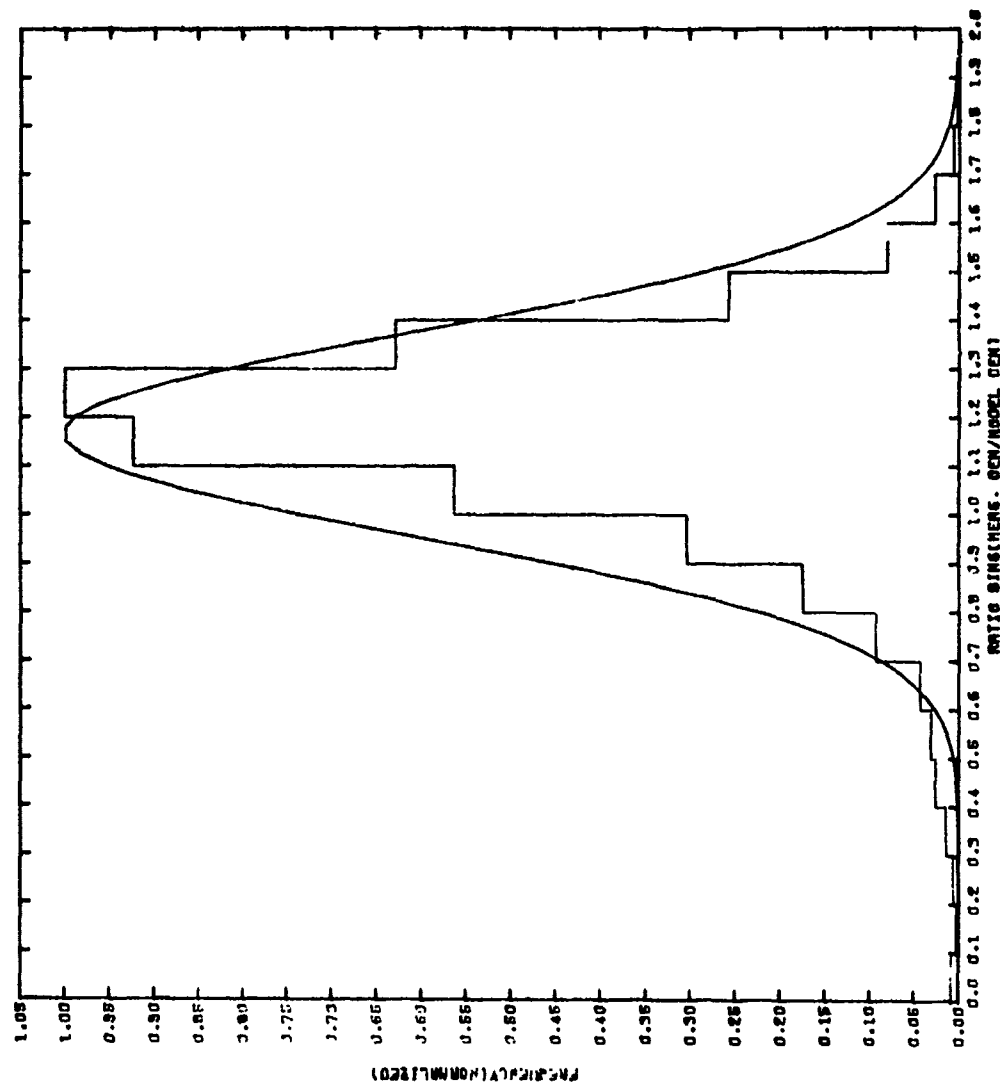


Number A1

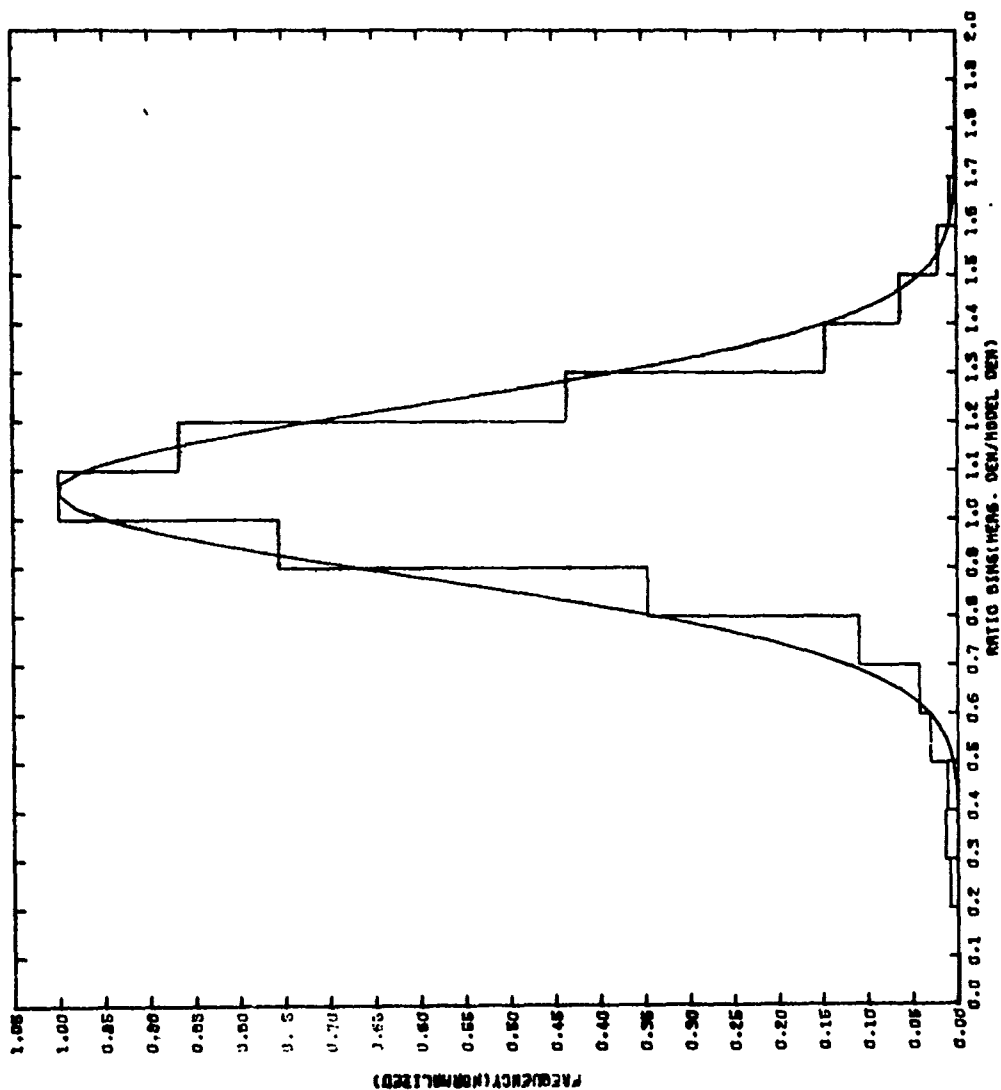
100 DATA
 KP LINE 0.0-0.1
 ALTITUDE 0.0-300
 LAT LINE 0.0-80
 NORTHERN HEMISPHERE
 MAX Y VAL 23166.00
 MEAN VAL 1.11
 ST. DEV. 1.97E-01
 TOTAL POINTS 81283



Number A2

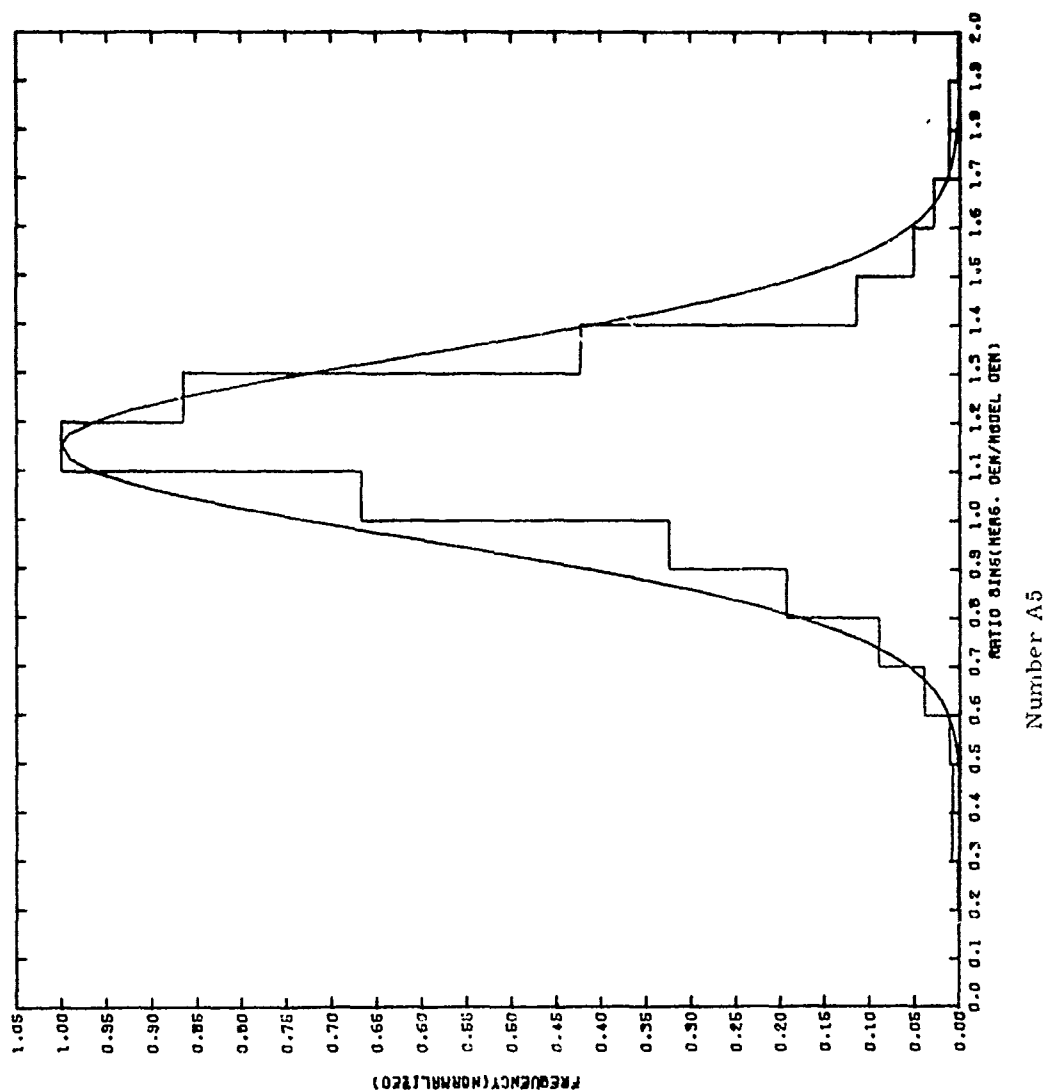


Number A3

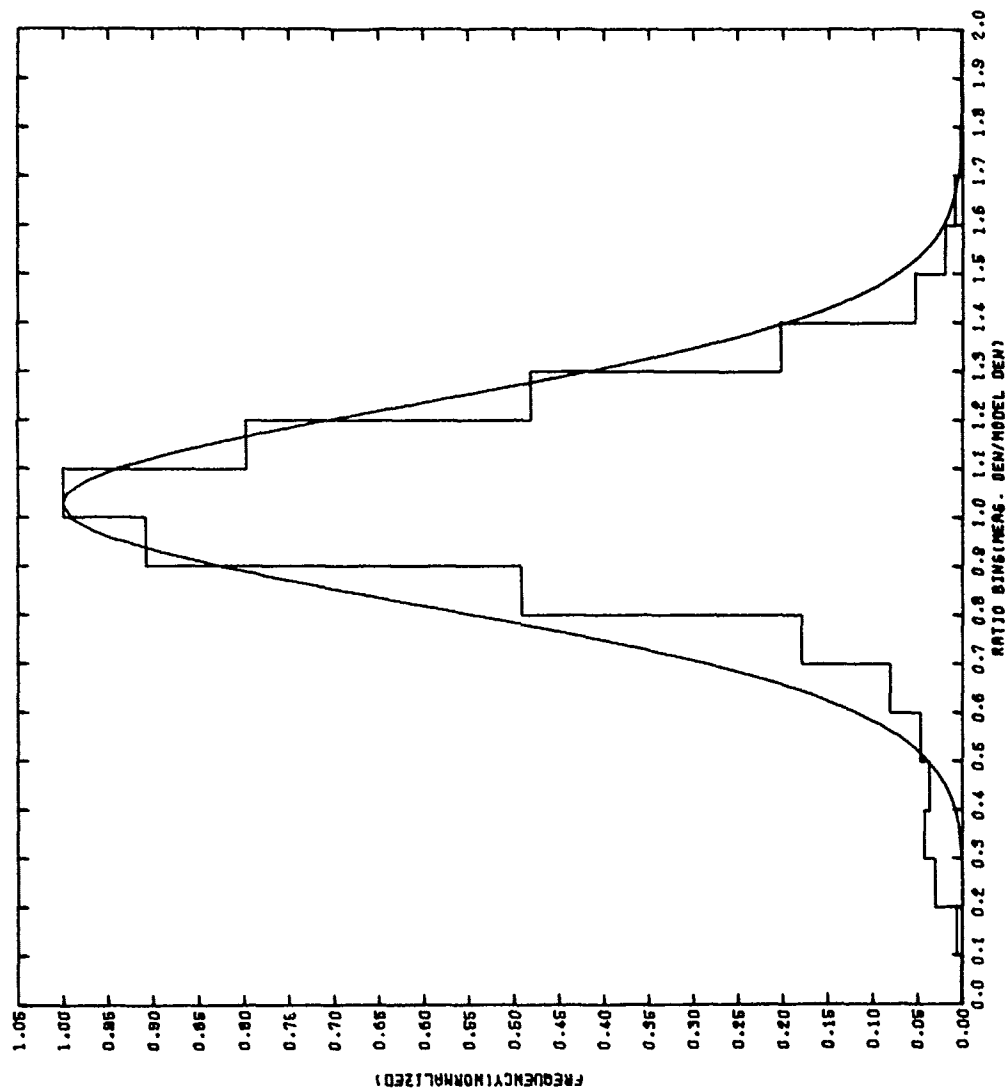


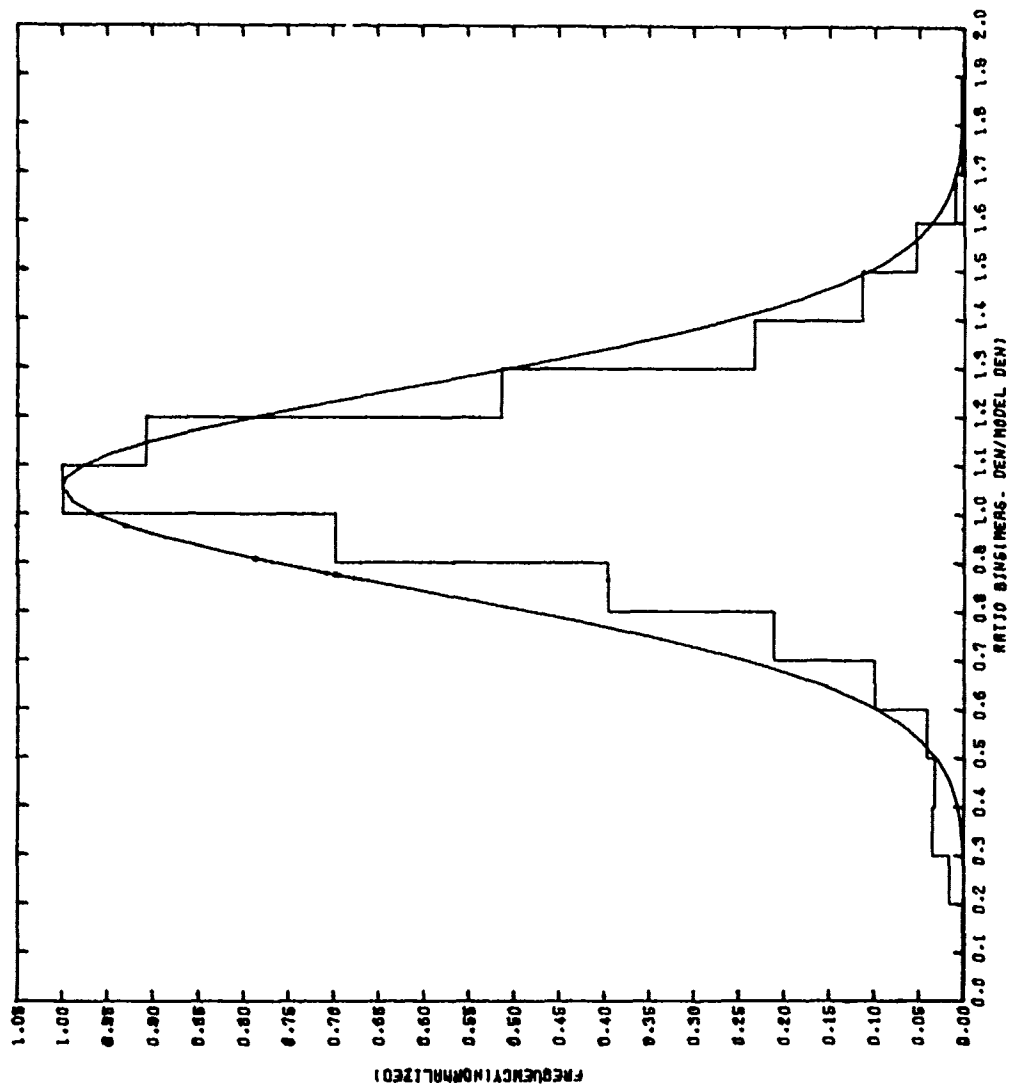
Number A4

IOB DATA
 KP LIMITS 5.0- 9.0
 ALTITUDE 0.0- 500.
 LAT LON 0.0- 90.
 NORTHERN HEMISPHERE
 MAX Y VAL 1159.00
 MEAN VALUE 1.00
 ST. DEV. 1.00E-11
 TOTAL POINTS= 4478



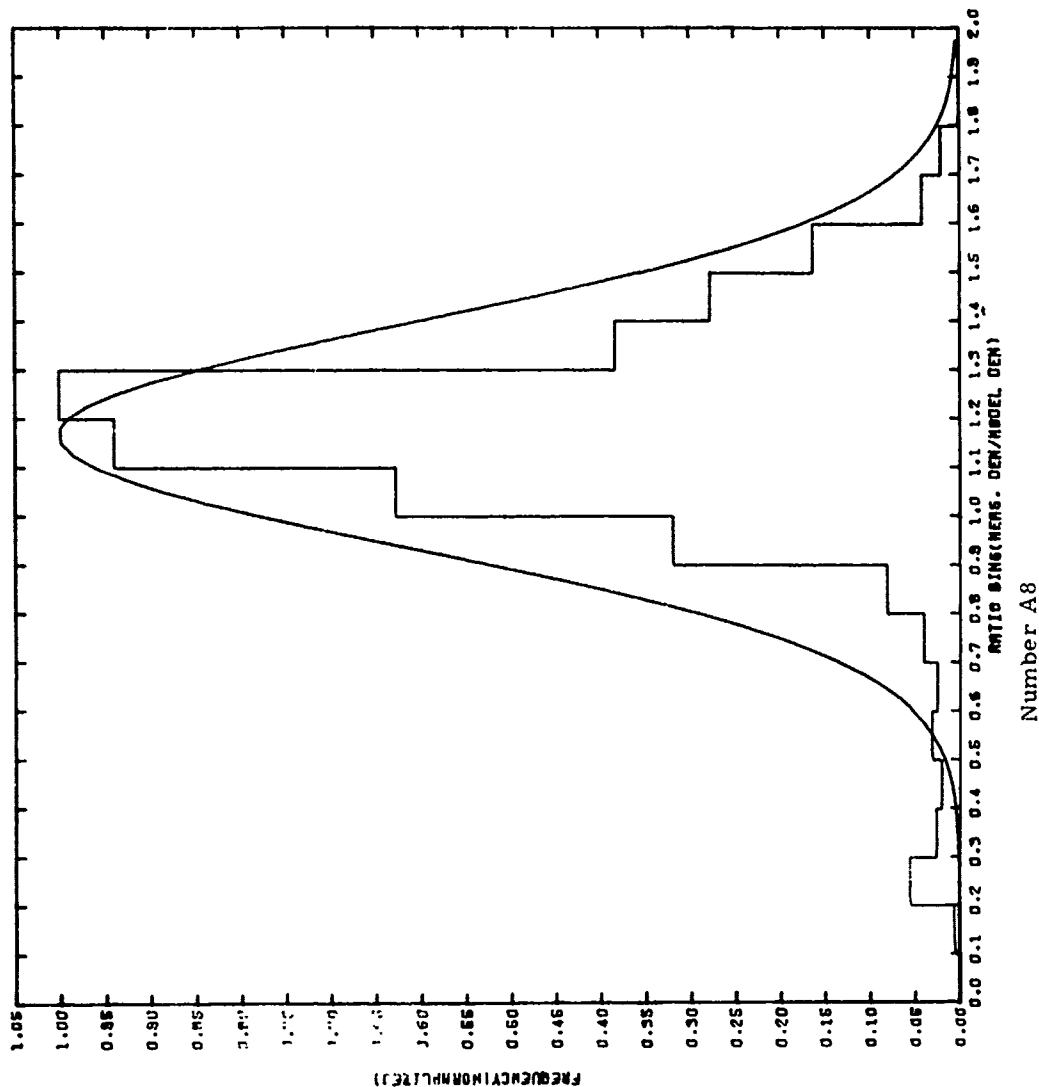
IDG DATA
 KP LIMITS 3.0- 5.0
 ALT LIM 250.- 300.
 LAT LIM 0.- 80.
 NORTHERN HEMISPHERE
 MAX Y VAL 1339.00
 MEAN VALUE 1.03
 ST. DEV. 2.07E-01
 TOTAL POINTS= 6880

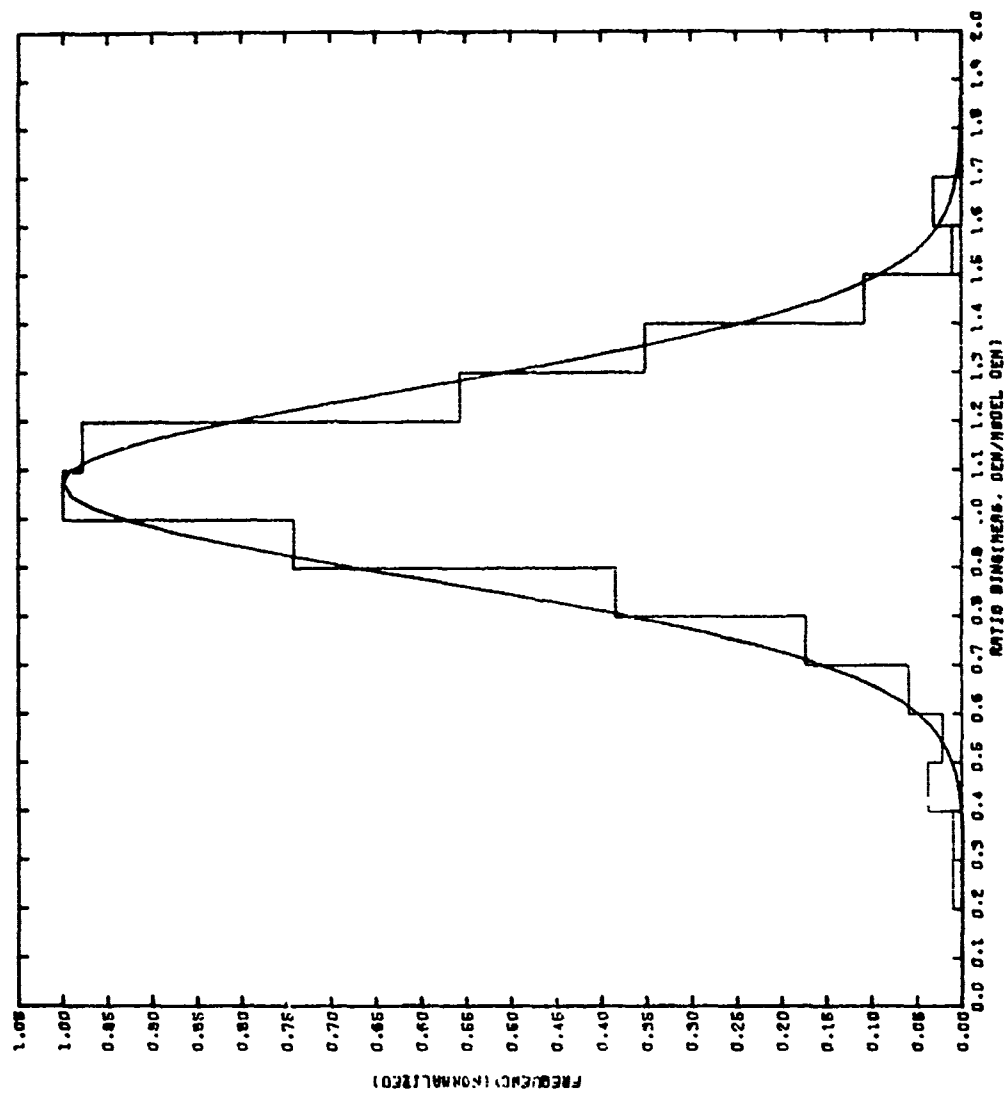




Number A7

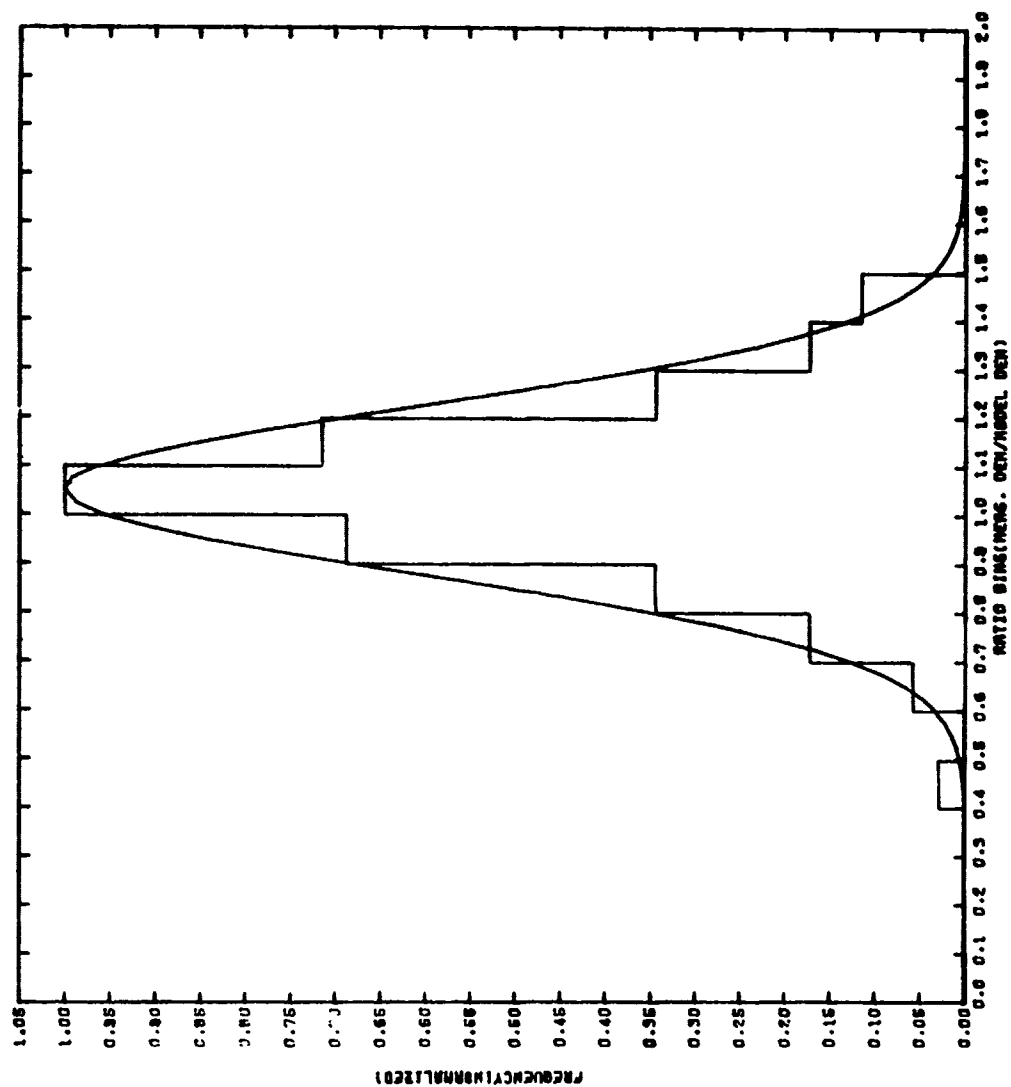
100 DATA 3.3-3.3
 KPLIMIT 3.3-3.3
 PLIMIT 3.3-3.3
 LPLIMIT 3.3-3.3
 NORMINAL MEAN 1.17
 MEAN VALUE 1.17
 ST. DEV. 0.332-01
 TOTAL POINTS= 802



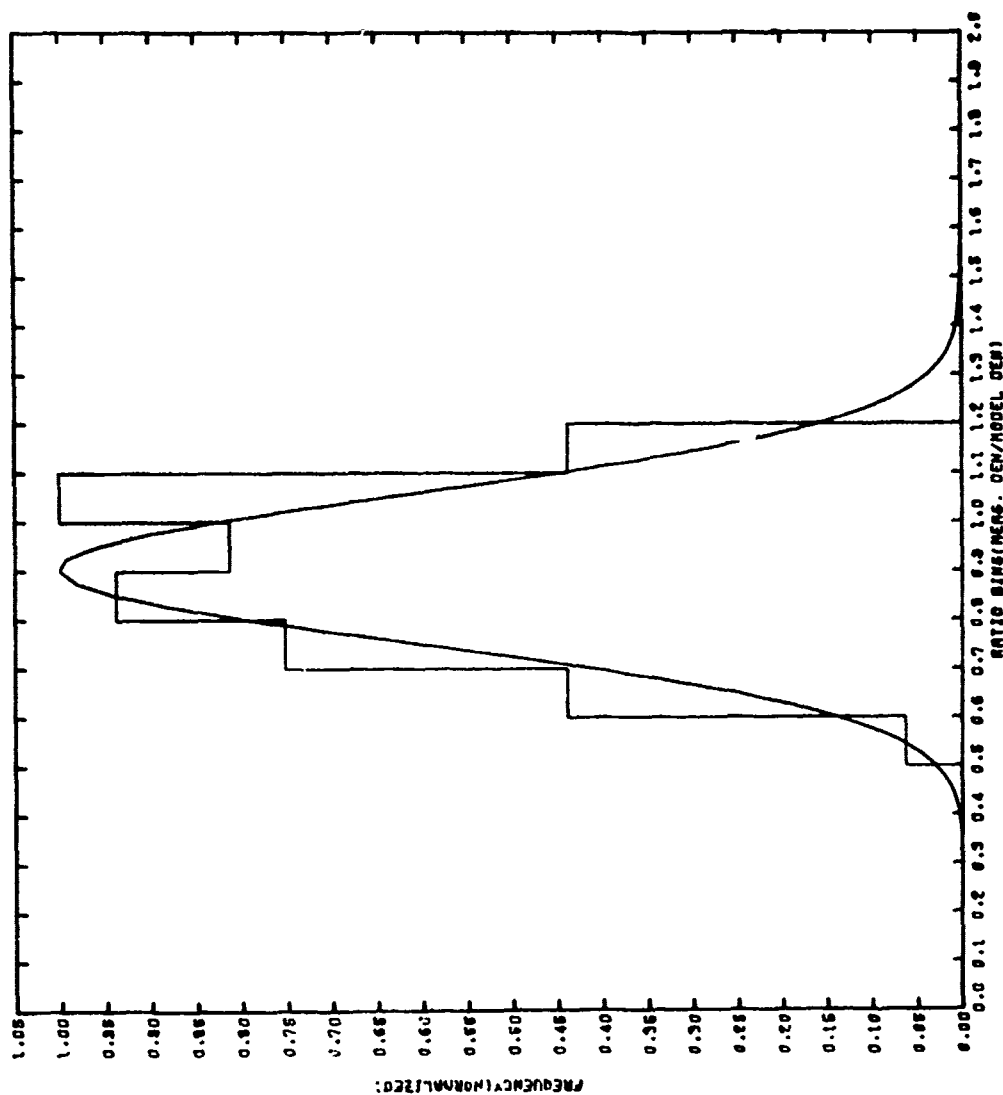


Number A9

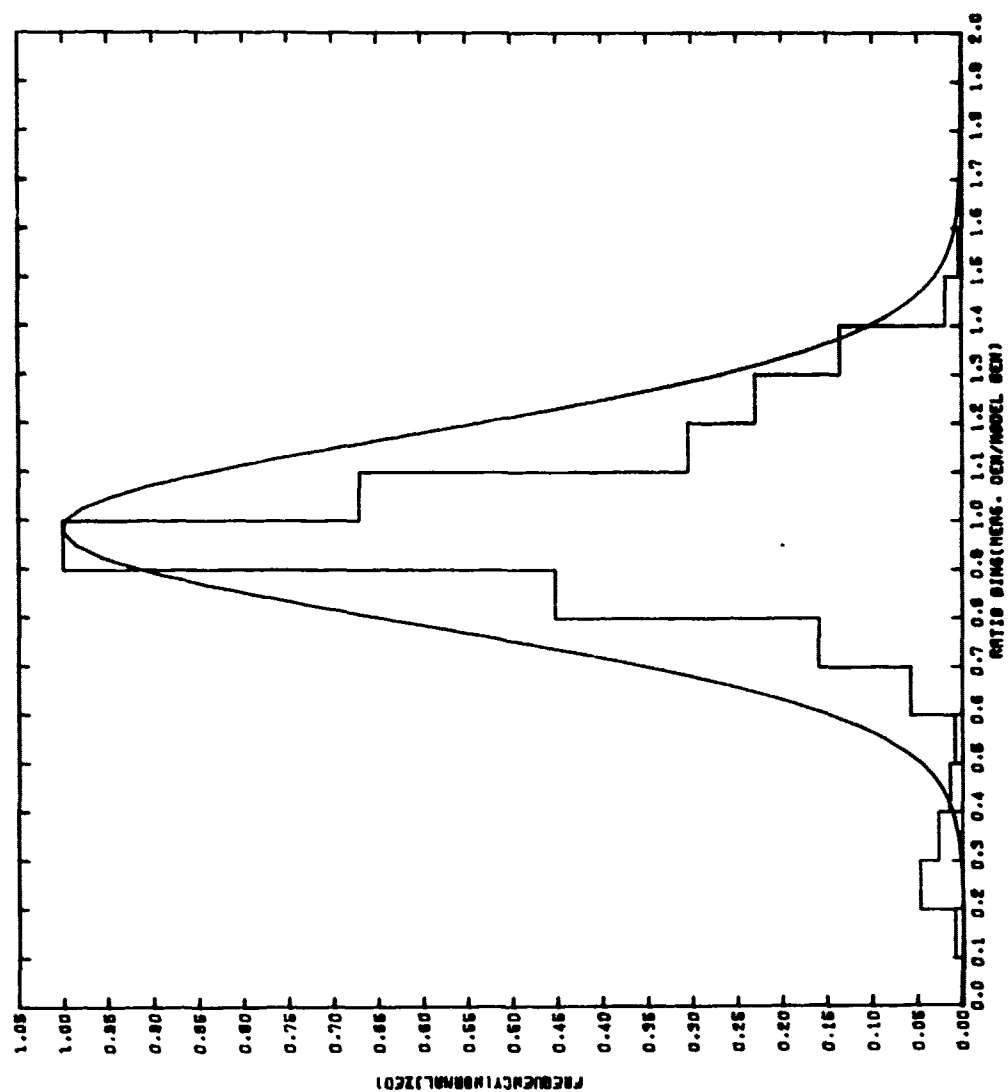
100 DATA
 NP LIMITS 8.0- 8.0
 ALTITUDE 250.- 300.
 LAT LJA 15.- 30.
 NORTHERN HEMISPHERE
 MAG Y VAL 35.00
 MEAN VALUE 1.08
 ST. DEV. 1.75E-01
 TOTAL POINTS= 127



Number A10



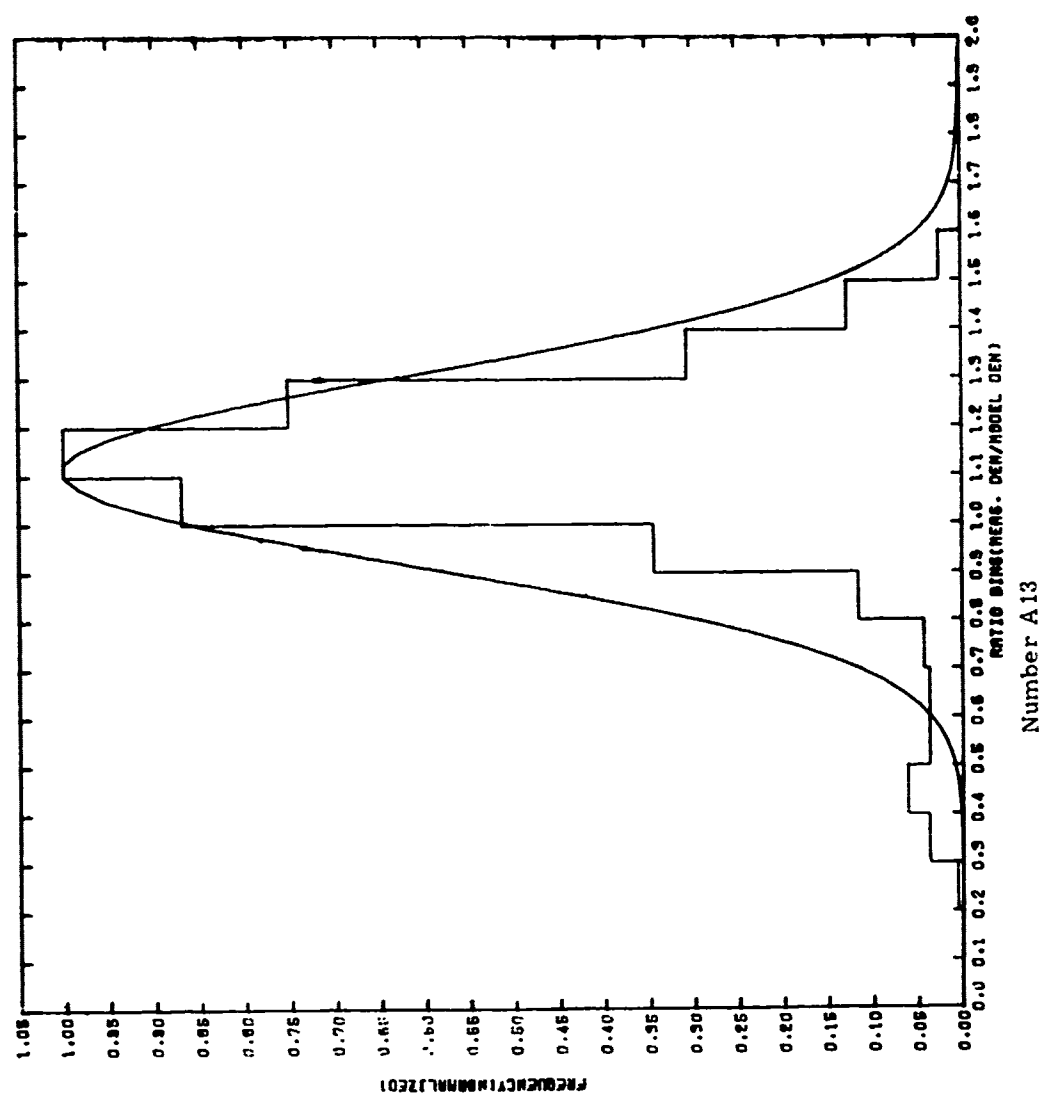
Number A11

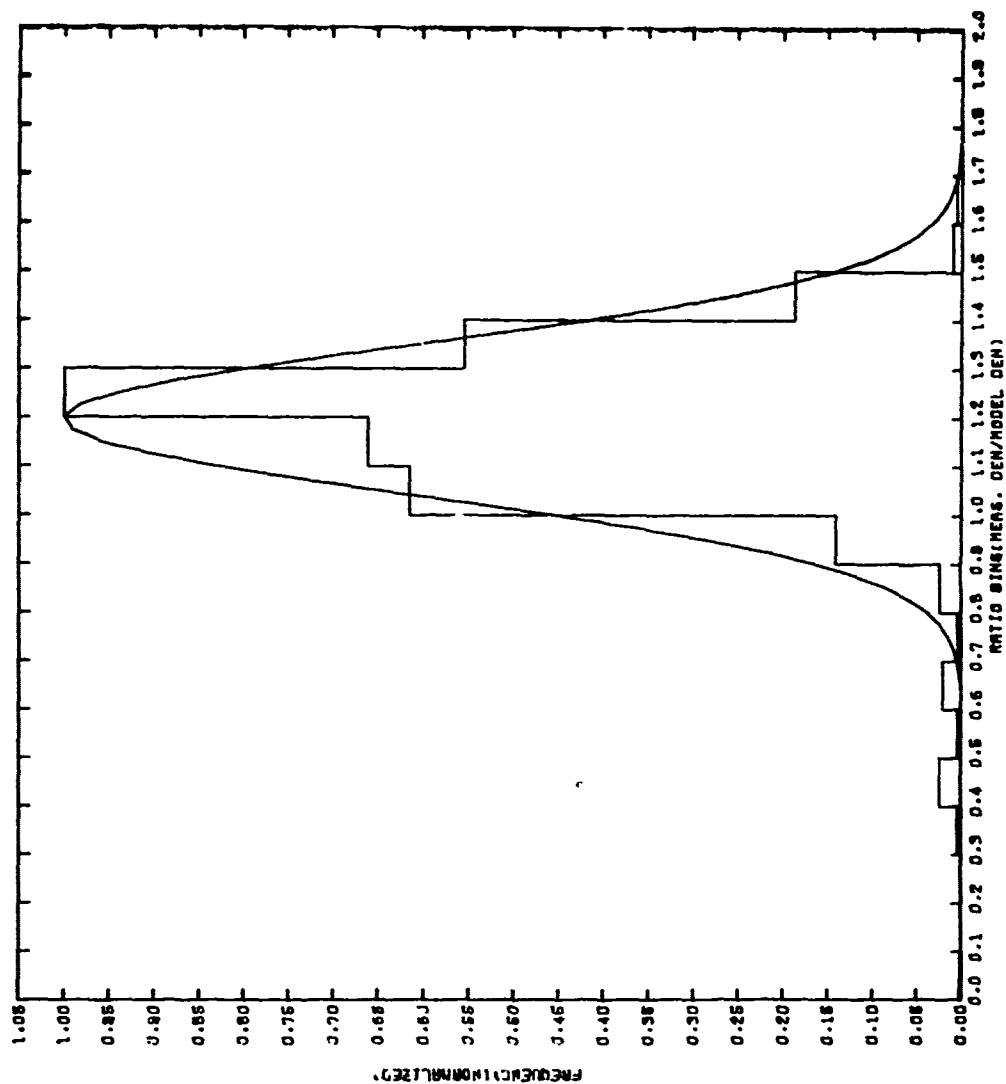


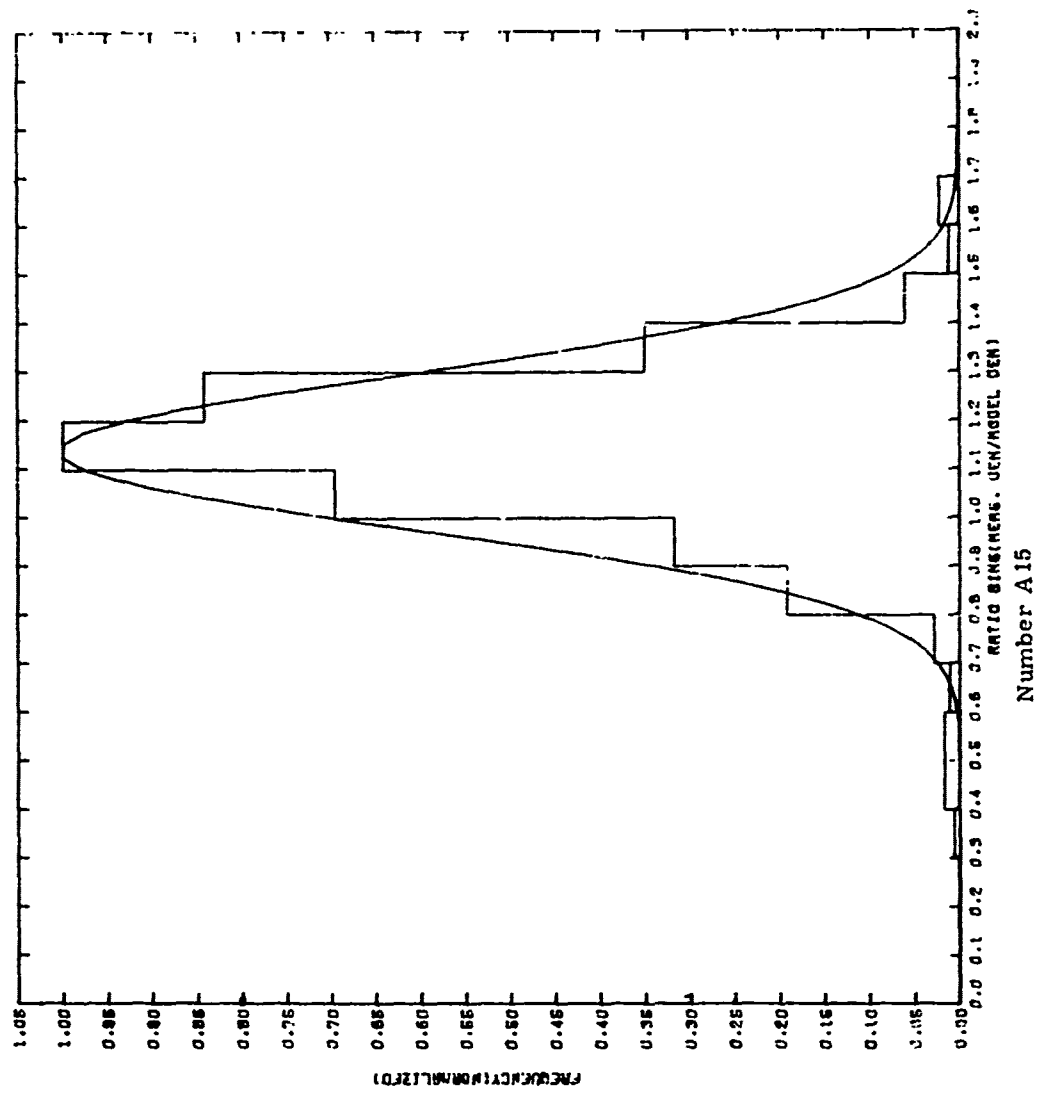
Number A12

100 DATA
KP LIMITS 3.0- 8.0
ALT LIM 250.- 500.
LAT LIM 45.- 60.
NORTHERN HEMISPHERE
MAX Y VAL 385.00
MEAN VALUE .85
ST. DEV. 1.85E-01
TOTAL POINTS= 1200

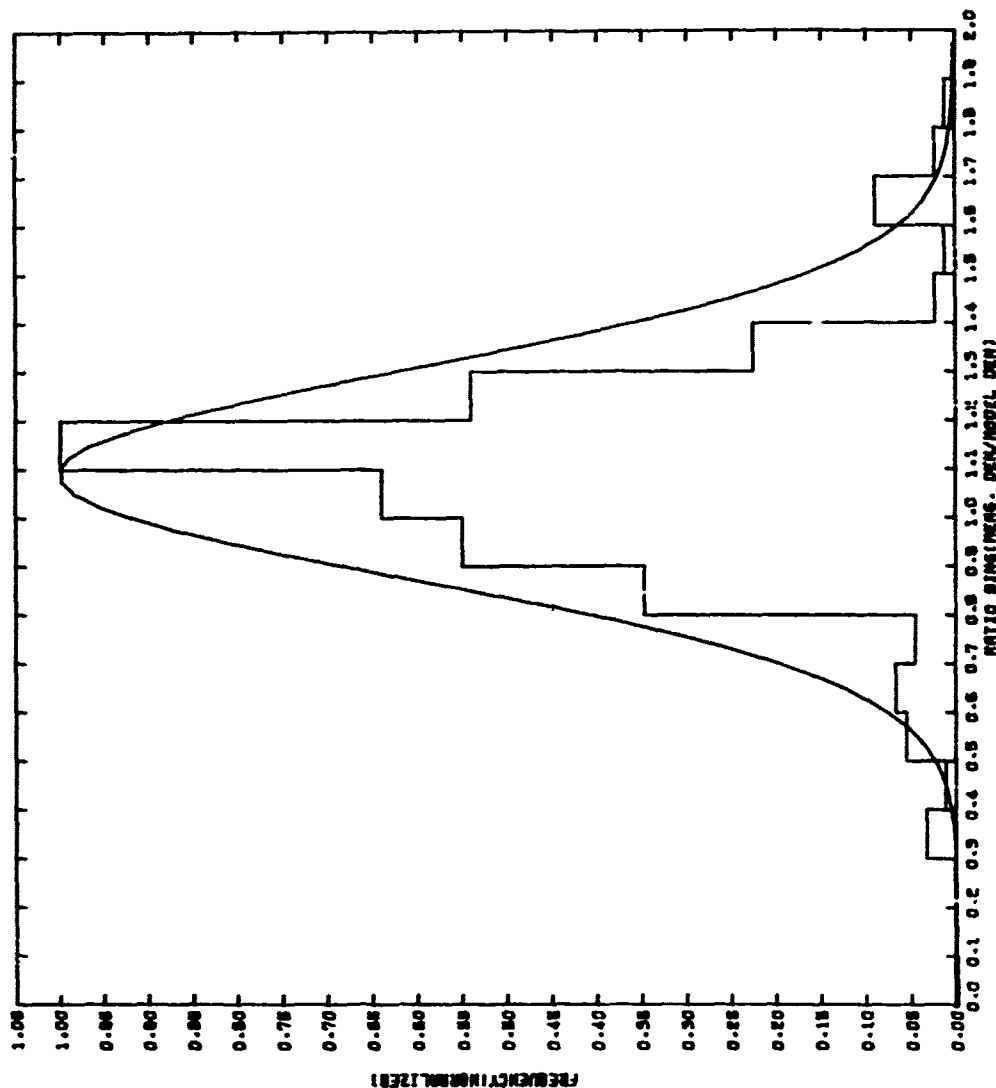
100 DATA
 KP LIMITE 3.0- 8.0
 ALT LIM 250.- 500.
 LAT LIM -60.- -45.
 SOUTHERN HEMISPHERE
 MAX Y VAL 248.00
 MEAN VALUE 1.11
 ST. DEV. 1.08E-01
 TOTAL POINTS= 819

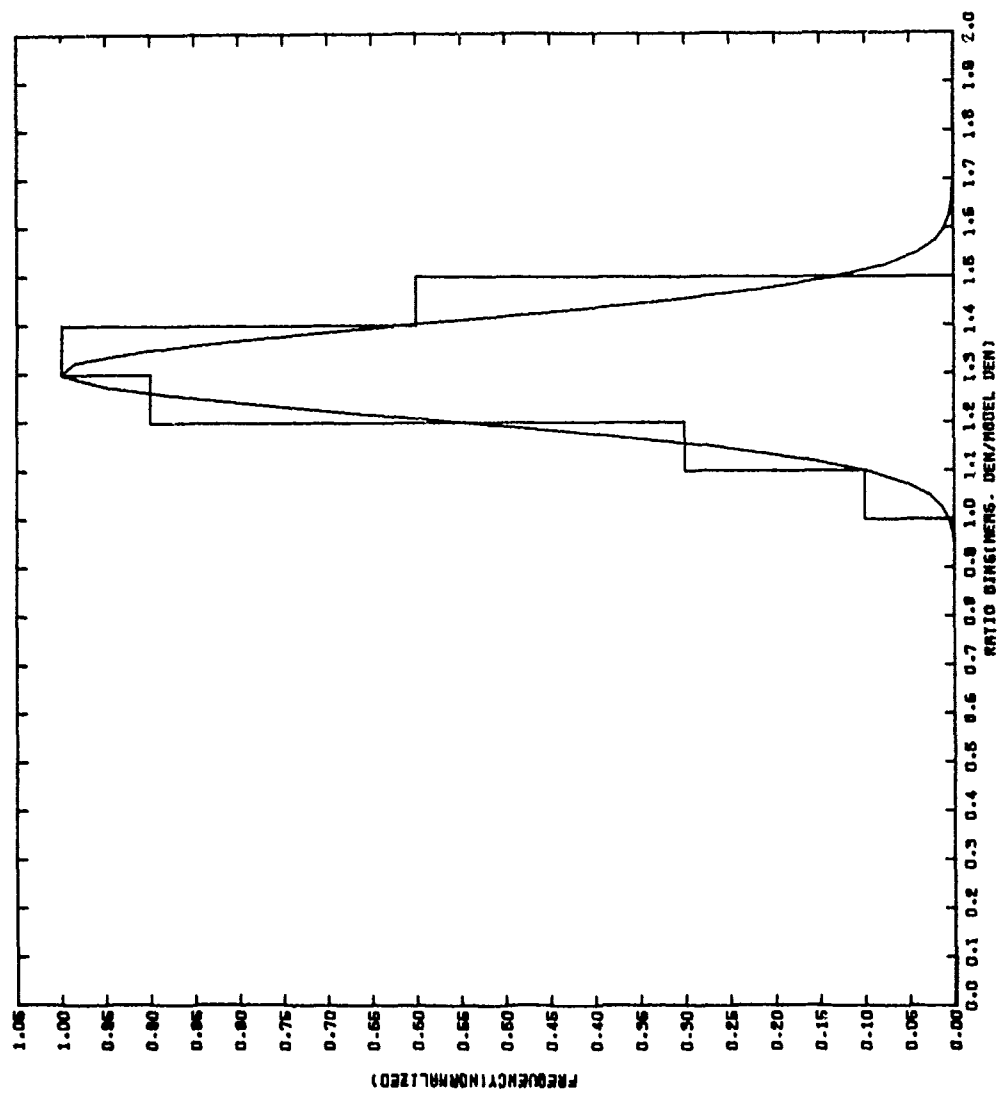




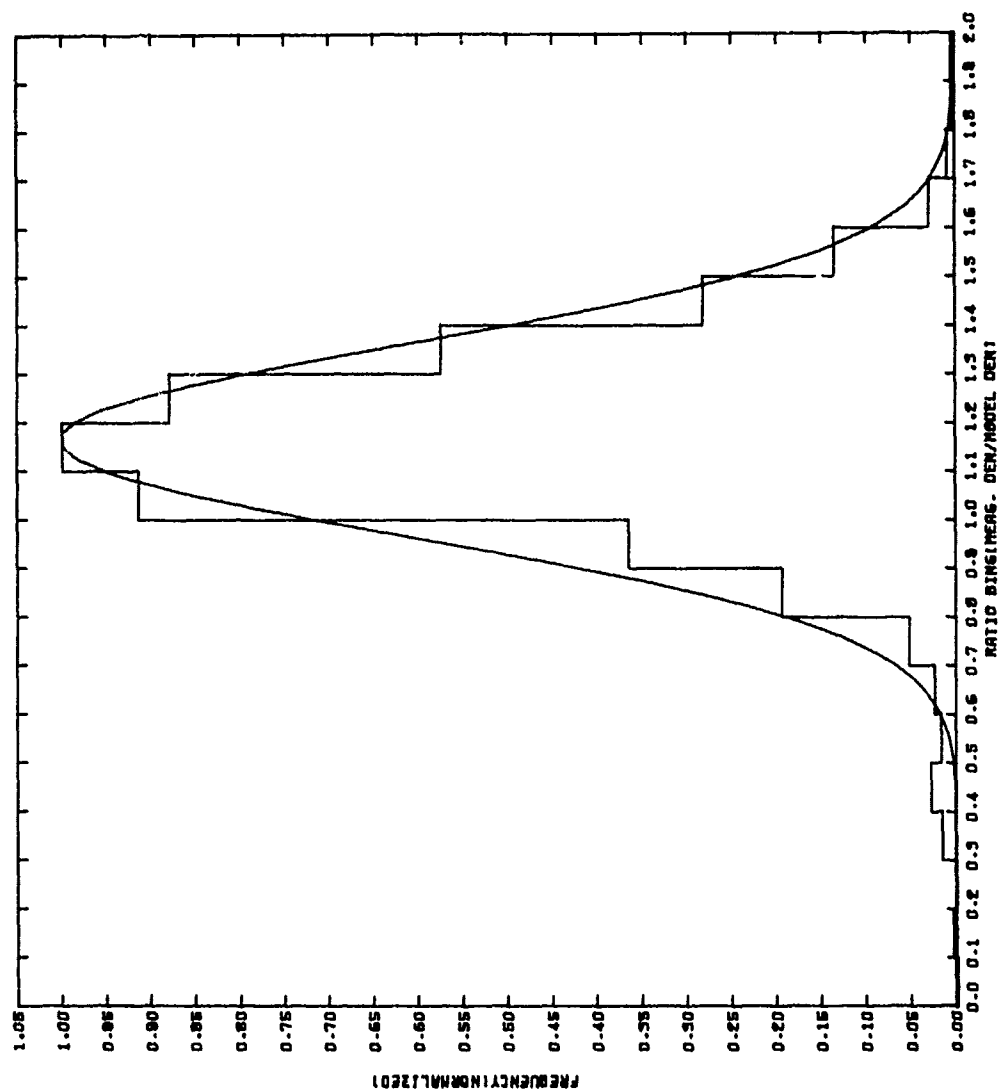


1.1, 1.14
 WY LIMITS 3.0- 5.0
 ALTITUDE 200.0- 250.0
 LAT LIM -15.0- 0.0
 COASTLINE MEMPHIS
 MAX Y VAL 195.00
 MEAN VALUE 1.14
 ST. DEV. 1.81E-01
 TOTAL POINTS= 652

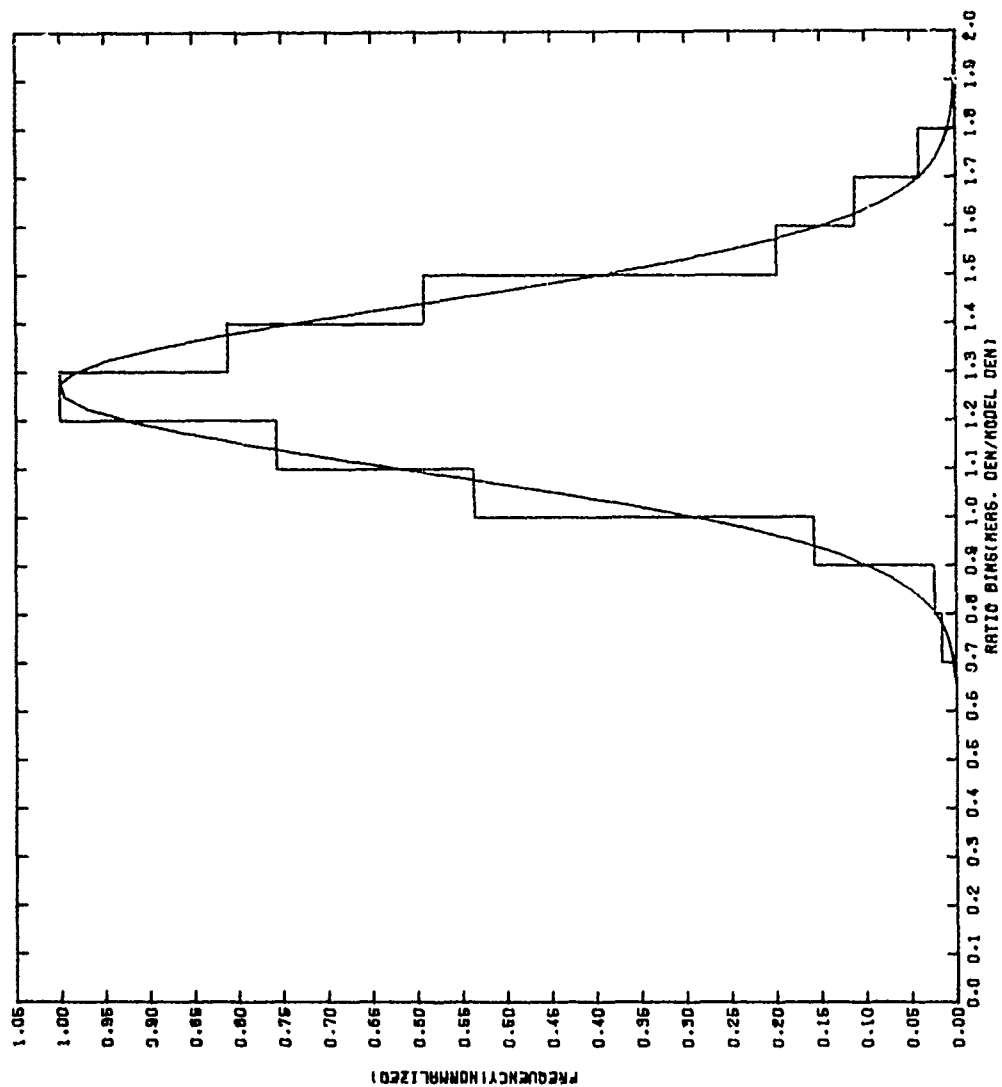




Number A17



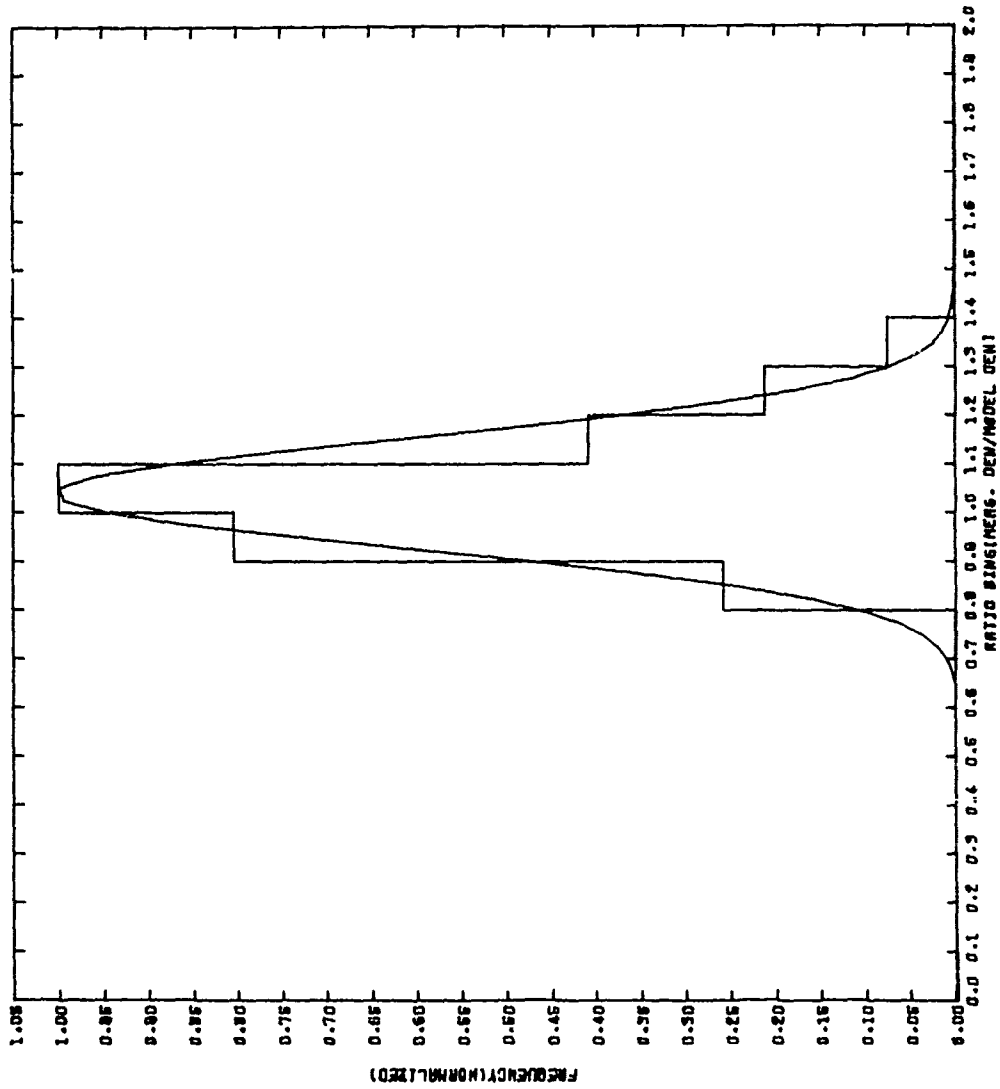
Number A18



Number A19

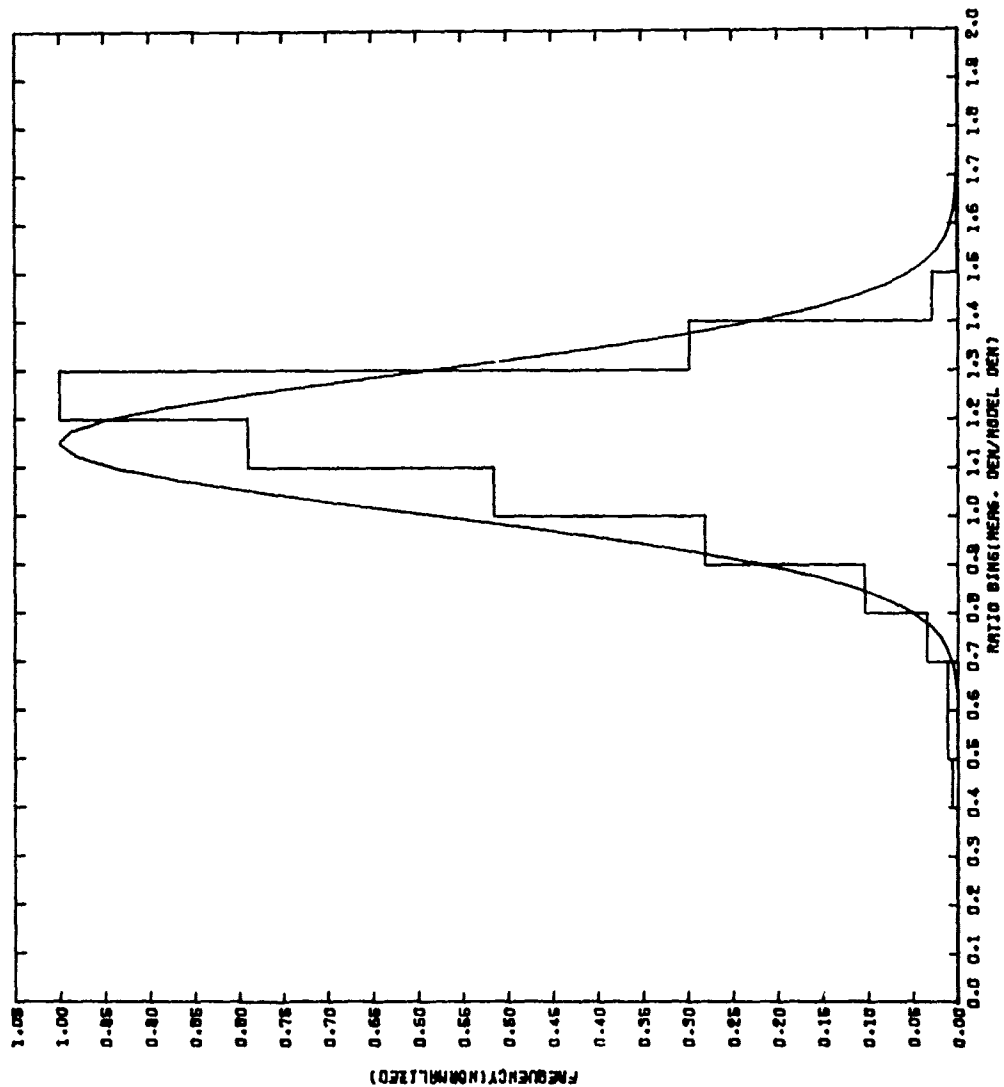
100 DATA
 KP LIMITS 3.0- 5.0
 ALT LN 200.- 250.
 LAT LN -30.- -75.
 SOUTHERN HEMISPHERE
 MAX Y VAL 127.00
 MEAN VALUE 1.27
 ST. DEV. 1.70E-01
 TOTAL POINTS= 536

100 DATA 5.0- 9.0
 MP LIMITS 0.0- 800.
 LAT LIM 15.0- 30.
 NORTHERN HEMISPHERE
 MAX Y VAL 66.00
 MEAN VALUE 1.04
 ST. DEV. 1.13E-01
 TOTAL POINTS= 102

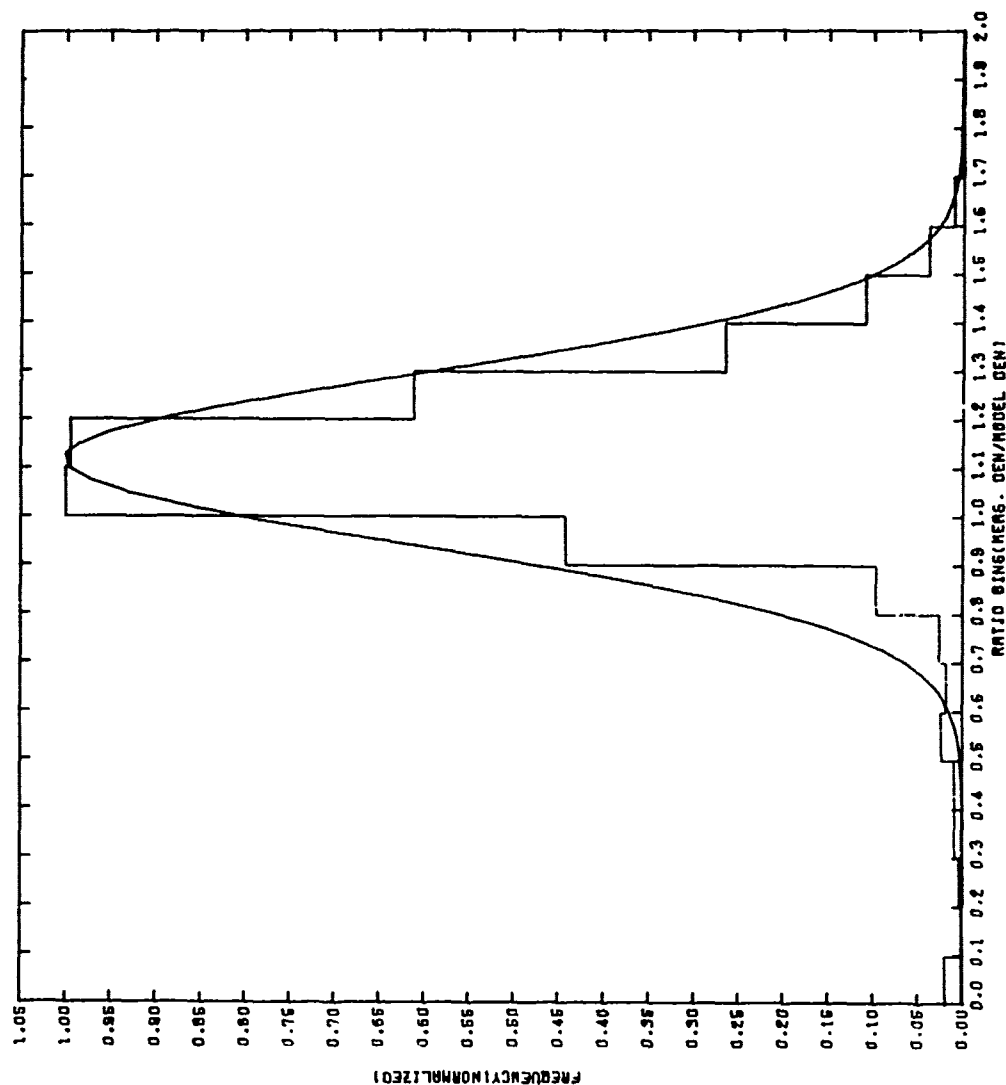


Number A20

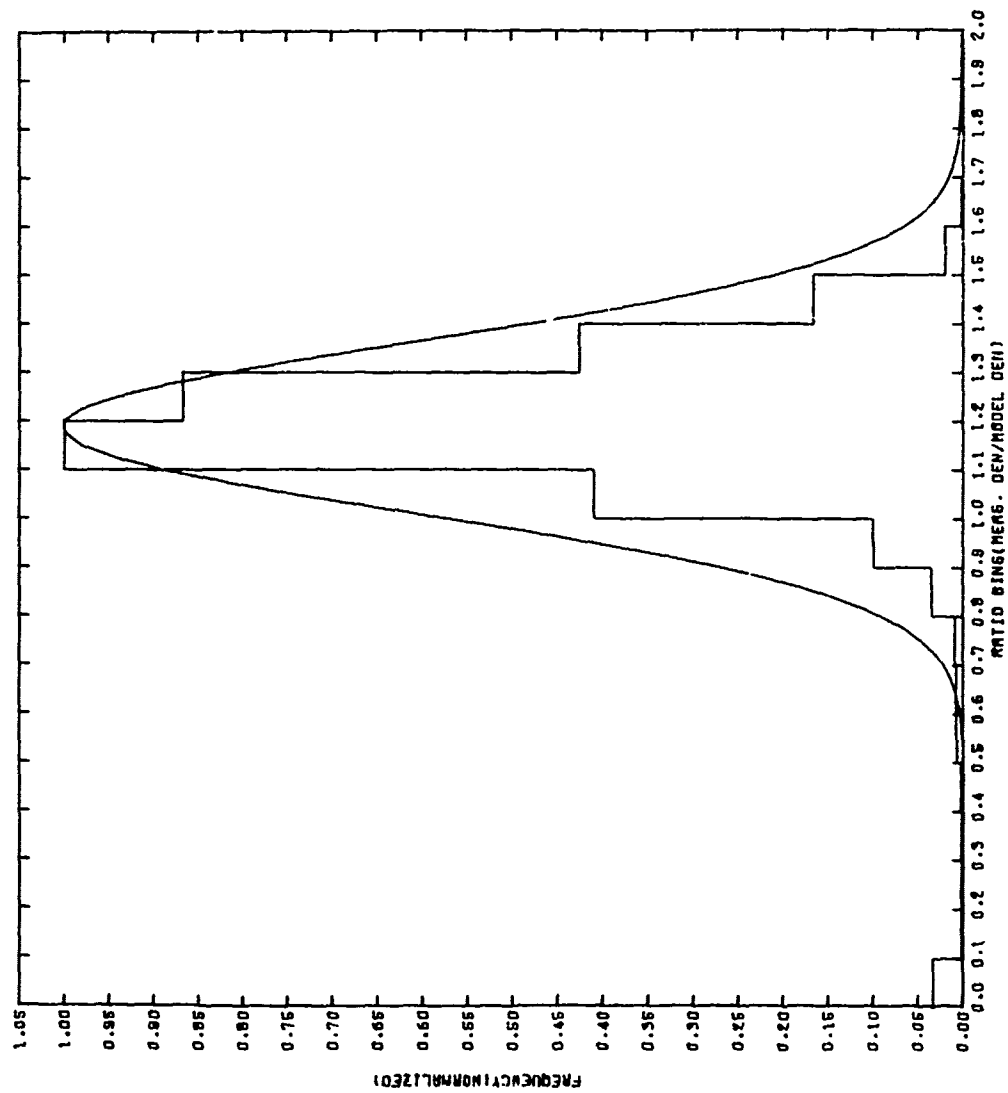
100 DATA 5.0- 8.0
 NP LIMITS 0.0- 200.
 LAT LIN -30.- -15.
 SOUTHERN HEMISPHERE
 MAX Y VAL 175.00
 MEAN VALUE 1.16
 ST. DEV. 1.44E-01
 TOTAL POINTS= 636



Number A21

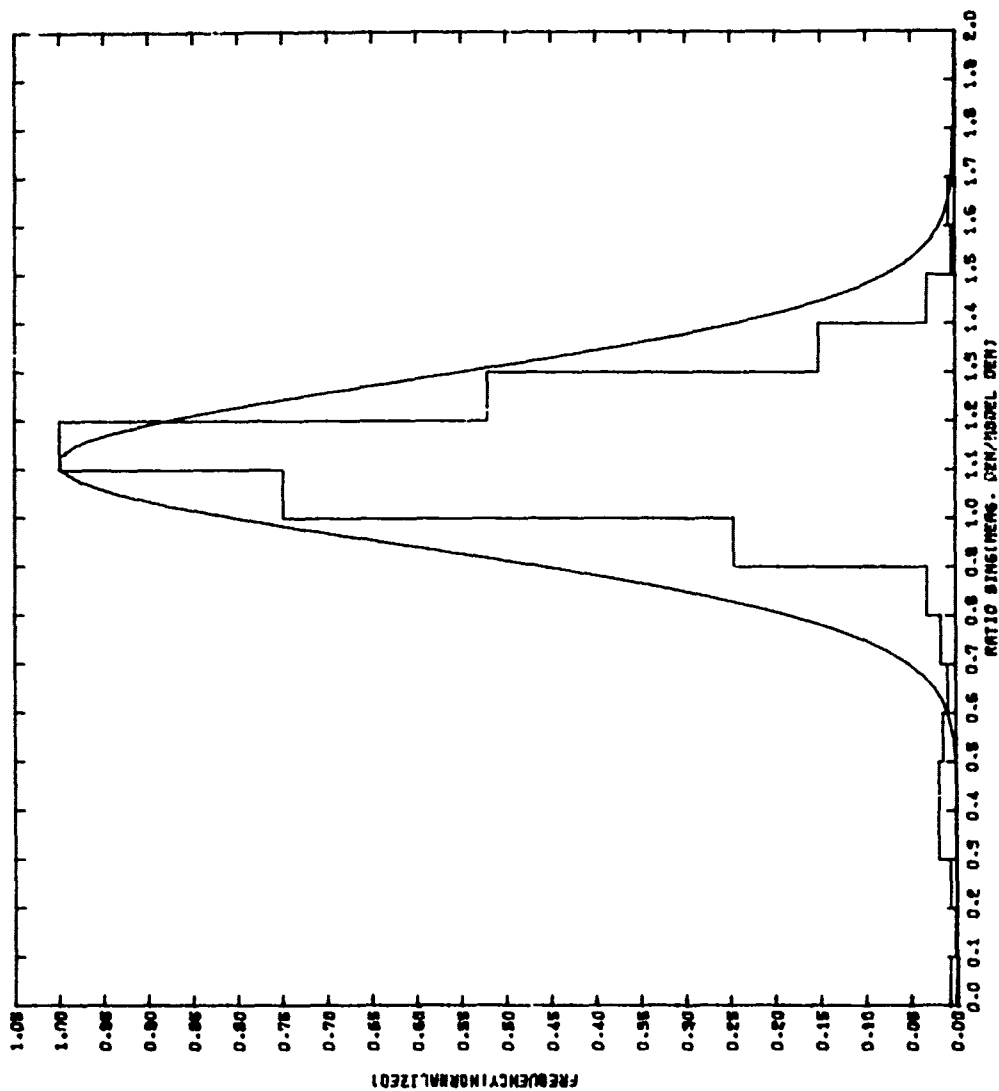


Number A22



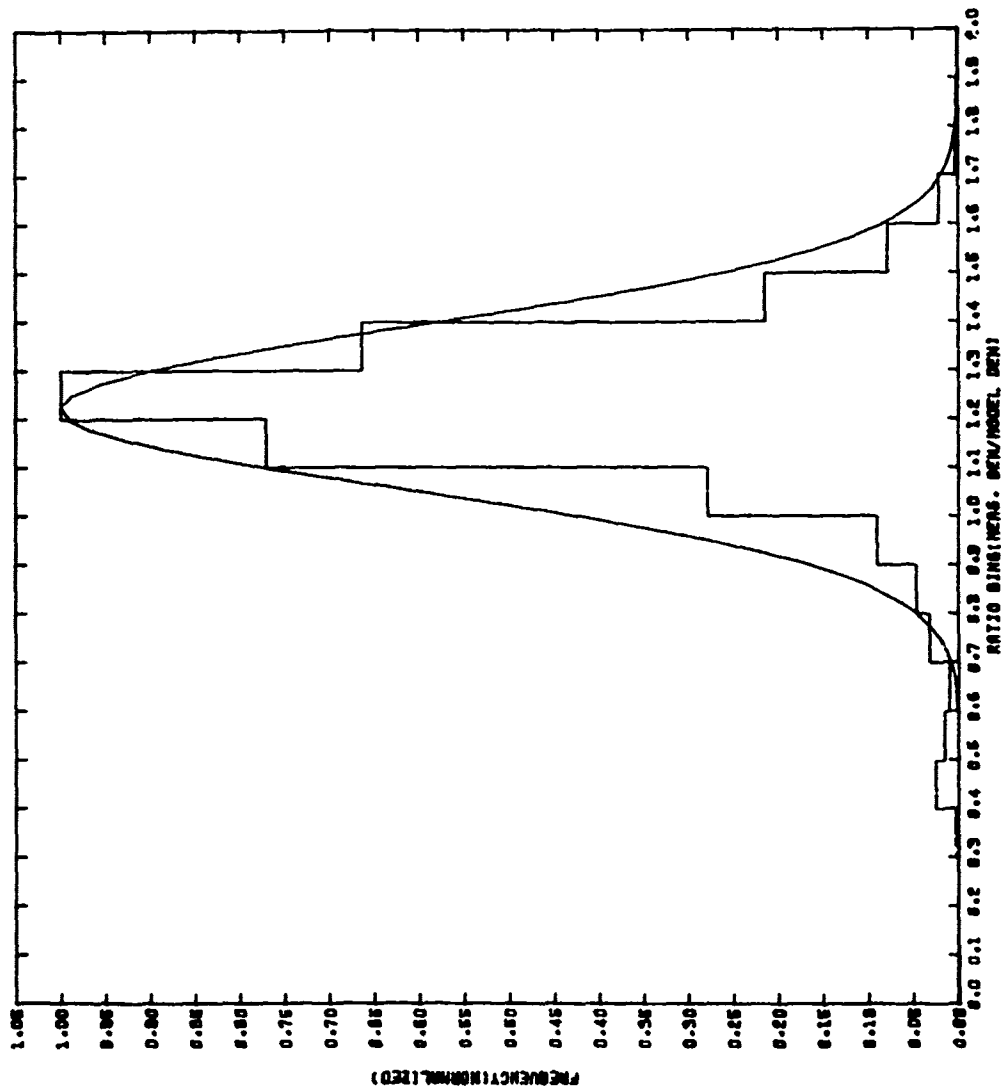
Number A23

JOB DATA 0.0- 5.0
 KP LIMITS 0.- 200.
 LAT LIN 60.- 75.
 NORTHERN HEMISPHERE
 MAX Y VAL 2750.00
 MEAN VALUE 1.11
 ST. DEV. 1.71E-01
 TOTAL POINTS= 7888

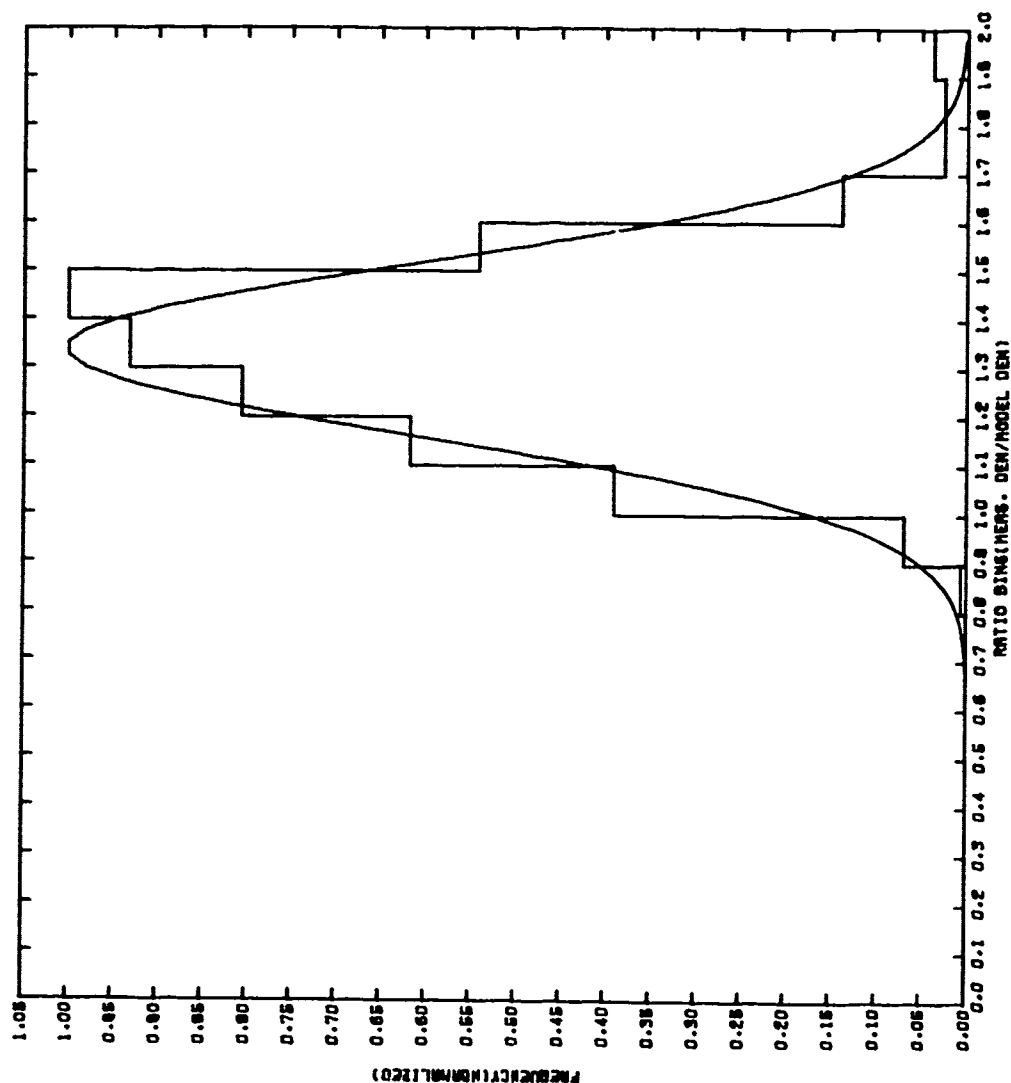


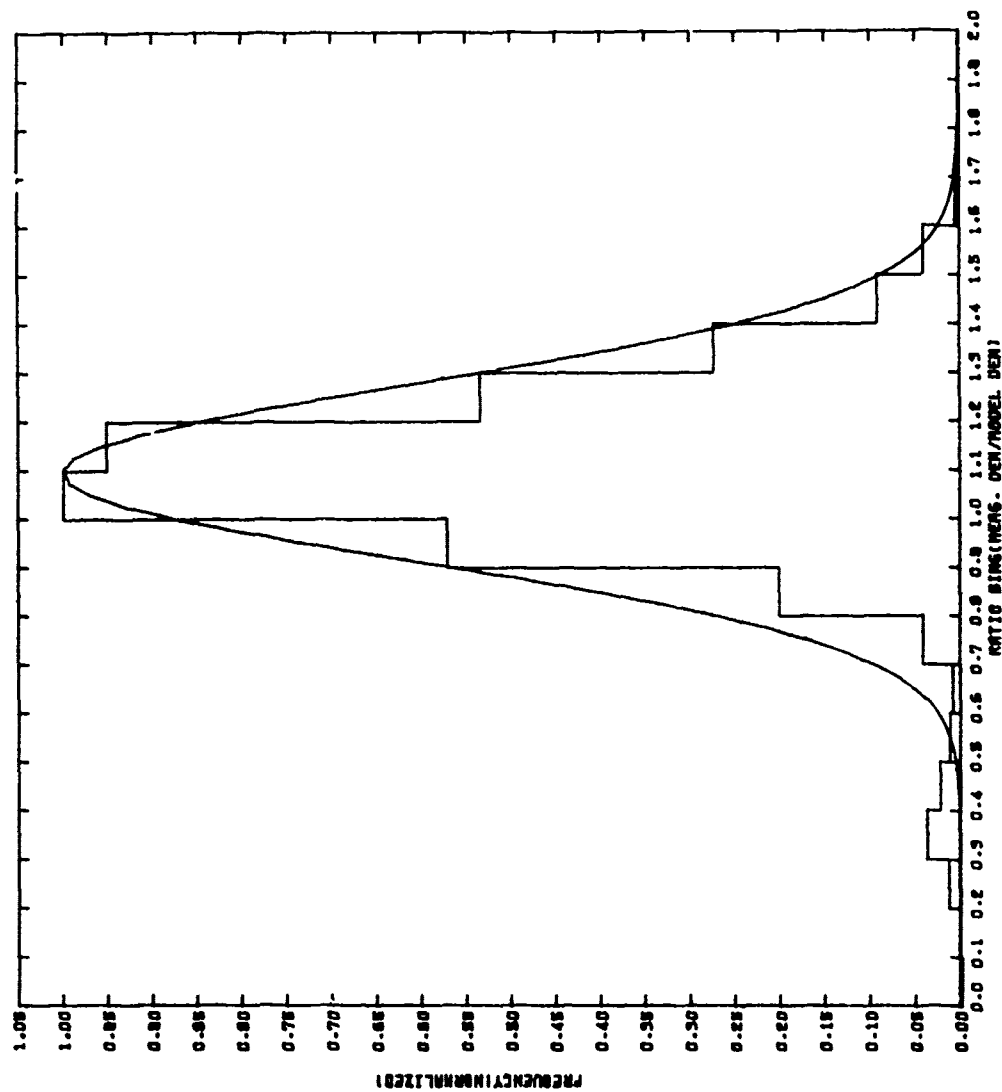
Number A24

100 DATA
 NP LIMITS 0.0- 3.0
 ALTITUDE 0.- 800.
 LAT LIM -75.- -80.
 SOUTHERN HEMISPHERE
 MAX Y VAL 569.00
 MEAN VALUE 1.32
 ST. DEV. 1.70E-01
 TOTAL POINTS= 5924



Number A25





Number A27