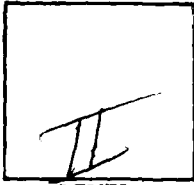


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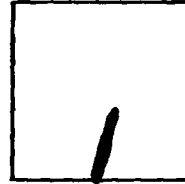
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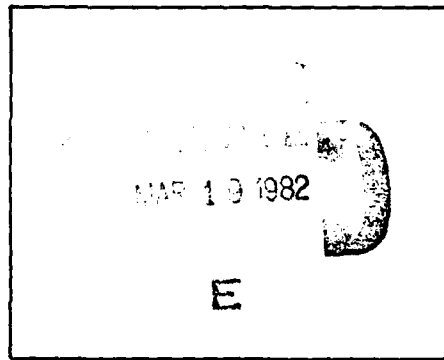
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E-TR-33-MS

MX SITING INVESTIGATION
GRAVITY SURVEY - MULESHOE VALLEY
NEVADA

Prepared for:

U. S. Department of the Air Force
Ballistic Missile Office (BMO)
Norton Air Force Base, California 92409

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Long Beach, California 90807

14 September 1981

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FOREWORD

Methodology and Characterization studies during Fiscal Years 1977 and 1978 (FY 77 and 78) included gravity surveys in 10 valleys, five in Arizona, two in Nevada, two in New Mexico, and one in California. The gravity data were obtained for the purpose of estimating the gross structure and shape of the basins and the thickness of the valley fill. There was also the possibility of detecting shallow rock in areas between boring locations. Generalized interpretations from these surveys were included in Ertec Western's (formerly Fugro National, Inc.) Characterization reports (FN-TR-26a through e).

During the FY 77 surveys, measurements were made to form an approximate 1-mile grid over the study areas, and contour maps showing interpreted depth to bedrock were made. In FY 79, the decision was made to concentrate on verifying and refining suitable area boundaries. This decision resulted in a reduction in the gravity program. Instead of obtaining gravity data on a grid, the reduced program consisted of obtaining gravity measurements along profiles across the valleys where Verification studies were also performed.

The Defense Mapping Agency (DMA), St. Louis, was requested to provide gravity data from their library to supplement the gravity profiles. For Big Smoky, Hot Creek, and Big Sand Springs valleys, a sufficient density of library data was available to permit construction of interpreted contour maps instead of just two-dimensional cross sections.

In late summer of FY 79, supplementary funds became available to begin data reduction. At that time, inner zone terrain corrections were begun on the library data and the profiles from Big Smoky Valley, Nevada, and Butler and La Posa valleys, Arizona. The profile data from Whirlwind, Hamlin, Snake East, White River, Garden, and Coal valleys, Nevada, became available from the field in early October 1979.

A continuation of gravity interpretations was incorporated into the FY 80 and 81 programs, and the results are being summarized in a series of valley reports. Reports covering Nevada-Utah gravity studies are being numbered "E-TR-33-" followed by the abbreviation for the subject valley. In addition, more detailed reports of the results of FY 77 surveys in Dry Lake and Ralston valleys, Nevada, were prepared. Verification studies were continued in FY 80, and gravity studies were included in the program. DMA continued to obtain the field measurements, and there was a return to the grid pattern. The interpretation of the grid data allows the production of contour maps which are valuable in the deep basin structural analysis needed for computer modeling in the water resources program. The gravity

interpretations will also be useful in Nuclear Hardness and Survivability (NH&S) evaluations.

The basic decisions governing the gravity program are made by BMO following consultation with TRW, Inc., Ertec Western, and the DMA. Conduct of the gravity studies is a joint effort between DMA and Ertec Western. The field work, including planning, logistics, surveying, and meter operation is done by the Defense Mapping Agency Hydrographic/Topographic Center (DMAHTC), headquartered in Cheyenne, Wyoming. DMAHTC reduces the data to Simple Bouguer anomaly (see Section A1.4, Appendix A1.0). The Defense Mapping Agency Aerospace Center (DMAAC), St. Louis, Missouri, calculates outer zone terrain corrections.

Ertec Western provides DMA with schedules showing the valleys with the highest priorities. Ertec Western also recommended locations for the profiles in the FY 79 studies with the provision that they should follow existing roads or trails. Any required inner zone terrain corrections are calculated by Ertec Western prior to making geologic interpretations.

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1.0 INTRODUCTION

1.1 OBJECTIVE

Gravity data from Muleshoe Valley were studied for the purpose of making a geologic interpretation which includes estimates of the shape of the structural basin, the thickness of the alluvial fill, and the location of concealed faults. The estimates will be useful in modeling the dynamic response of ground motion in the basin and in evaluating ground-water resources.

1.2 LOCATION

Muleshoe Valley is located in east-central Nevada about 40 miles (64 km) west of the Utah border and 80 miles (129 km) south of the town of Ely by U.S. Route 93 (Figure 1).

Muleshoe Valley lies between Lake Valley and Cave Valley and opens southward into Dry Lake Valley. Muleshoe Valley is bounded on the west by the southern Schell Creek Range and on the east by the Fairview Range, Grassy Mountain, and Dutch John Mountain (Figure 2).

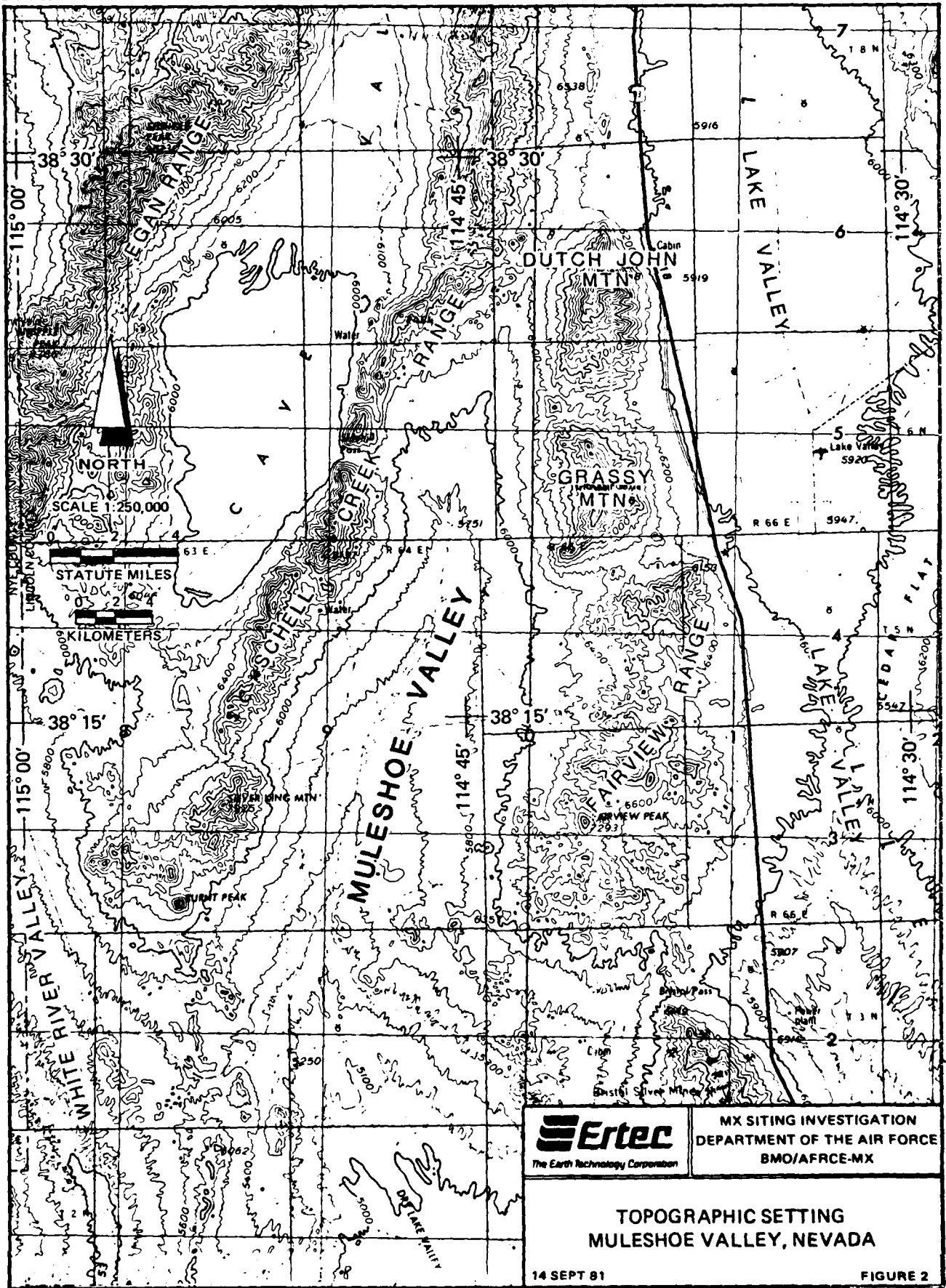
The area covered by this report lies between North latitudes 38°05' and 38°30' and West longitudes 114°35' and 114°50'.

1.3 SCOPE OF WORK

Five primary work elements were completed during this study.

They are:

1. Computation and merging of terrain corrections;
2. Synthesis of regional and valley-specific geological data;
3. Evaluation of the regional field and separation of the residual field;



4. Inverse modeling to estimate depth to bedrock; and
5. Interpretation of structural relationships.

The gravitational field within Muleshoe Valley was defined by data from 224 stations. The distribution of stations is shown in Drawing 1.0 and the station data are listed in Appendix A2.0. The Defense Mapping Agency Aerospace Center (DMAAC) supplied 86 gravity stations from its library, and 138 new gravity measurements were made by the Defense Mapping Agency Hydrographic-Topographic Center/Geodetic Survey Squadron (DMAHTC/GSS).

Muleshoe Valley and Cave Valley were studied together, with the results presented in separate reports. The rectangular region between North latitudes $38^{\circ}05'$ and $38^{\circ}45'$ and West longitudes $114^{\circ}35'$ and $115^{\circ}00'$ contains both of the valleys and surrounding mountains. All of the 522 gravity stations within this region were used to establish a common regional trend for the two valleys. Following separation of the residual field, the geologic modeling of the two valleys was done independently.

2.0 GRAVITY DATA REDUCTION

DMAHTC/GSS obtained the basic observations for the new stations and reduced them to Simple Bouguer Anomalies (SBA) as described in Appendix A1.0. Up to three levels of terrain corrections were applied to the new stations to convert the SBA to the Complete Bouguer Anomaly (CBA). Only the first two levels of terrain corrections described below were applied to the library stations.

First, the DMAAC, St. Louis, Missouri, used its library of digitized terrain data and a computer program to calculate corrections out to 104 miles (167 km) from each station. When the program could not calculate the terrain effects near a station, Ertec Western used a ring template to estimate the effect of terrain within approximately 3000 feet (914 m) of the station. The third level of terrain corrections was applied to those stations where relief of 10 feet (3 m) or more was observed within 130 feet (40 m). In these cases, the elevation differences were measured in the field at a distance of 130 feet (40 m) along six directions from the stations. These data were used by Ertec Western to calculate the effect of the very near relief.

The CBA values and principle facts for the Muleshoe Valley stations are listed in Appendix A2.0.

3.0 GEOLOGIC SUMMARY

Muleshoe Valley is a small valley in the eastern part of the Great Basin section of the Basin and Range physiographic province as described by Fenneman (1931).

Muleshoe Valley opens into Dry Lake Valley on the south. On the west and north, it is bounded by the southern Schell Creek Range. On the east, it is bounded by Dutch John Mountain, Grassy Mountain, and the Fairview Range (Figure 2). Narrow gaps on the north and east of the valley lead into Lake Valley.

The southern Schell Creek Range is primarily composed of carbonate and siliceous clastic rocks (dolomite, limestone, shale, and quartzite). At the northwest margin of the valley, these sedimentary rocks are overlain by Oligocene welded tuffs and are intruded by minor Tertiary rhyolite dikes and plugs (Ekren and others, 1977). Dutch John Mountain and Grassy Mountain are composed of carbonate and siliceous clastic rocks (limestone, shale, and quartzite) (Tschanz and Pampeyan, 1970); the Fairview Range is composed of middle-Tertiary ash-flow tuff, welded tuff, rhyolite lava, and basalt lava (Ekren and others, 1977). Isolated outcrops of carbonate and siliceous clastic rocks are found in the southern part of the Fairview Range.

The present topographic relief of Muleshoe Valley and the surrounding mountains is largely the result of late-Cenozoic extensional block faulting (Stewart, 1980). Surface data give little indication of the subsurface configuration of the valley.

The Schell Creek Range block on the west appears to be tilted down to the east. Repetition of Oligocene welded tuffs and the underlying Paleozoic rocks in the Dutch John-Grassy-Fairview Range across the valley suggests a major fault along the eastern margin of the valley, but minor faults may also occur beneath the western margin (Ertec, 1981a).

The valley-fill deposits are older alluvium and younger alluvium. The older alluvium is Quaternary in age (Tschanz and Pampeyan, 1970) and consists of nonindurated and partly indurated gravels, sands, and silts derived from the surrounding bedrock (Eakin, 1963). The younger alluvium is composed of reworked older alluvium and is found only in Coyote Wash, the valley's axial drainage channel (Tschanz and Pampeyan, 1970).

Aerial photograph analysis and geologic field reconnaissance (Ertec, 1981a) indicate only a few minor Pleistocene fault scarps breaking the valley-floor surface, although lineaments of unknown origin occur along the southeastern and southwestern valley margins. The Coyote Wash fault extends into Muleshoe Valley from the southwest just north of Silver King Mountain but does not appear to disturb the Quaternary alluvium. Aligned, faceted ridge spurs along the western flank of Dutch John and Grassy mountains suggest Quaternary faulting in the northeastern portion of the valley.

4.0 INTERPRETATION

The basis of interpretation in this report is the Complete Bouguer Anomaly (CBA). Complete Bouguer Anomaly contours and the gravity station locations are shown in Drawing 1.

The interpretation of irregularly spaced data is both difficult and inefficient. In order to simplify the interpretation, the CBA data were reduced to a set of values on the nodes of a regularly spaced grid. The value at each node was computed using a minimum curvature gridding program (Briggs, 1974; and Swain, 1976). Minimum curvature gridding is an iterative process, the purpose of which is to find the smoothest surface that fits the irregularly spaced data. This smooth surface is then used to interpolate between the existing data points. A 0.62-mile (1-km) grid spacing, which is slightly more dense than the average data spacing, was used throughout this analysis.

4.1 REGIONAL RESIDUAL SEPARATION

A fundamental difficulty in gravity interpretation is that the gravity expression of short wavelength, shallow structural features of interest are overlapped and obscured by long wavelength features occurring at all depths. The purpose of a regional-residual separation is to remove the effect of the longer wavelength structures so that the features of interest may be correctly interpreted.

In order to estimate the form and magnitude of the long wavelength contribution (regional), the CBA was continued upward

using a fast Fourier transform (FFT) and a frequency domain filter (Gunn, 1975). The data were continued upward to a height at which no correlation could be seen between the upward-continued CBA and the surface structure. This was at an altitude of 60,000 feet. The regional was then subtracted from the CBA and the resulting residual anomaly was further adjusted by a constant -8.0 mgal to make the zero residual contour approximately fit outcrops of Paleozoic carbonate rocks.

4.2 DENSITY SELECTION

The correct interpretation of the residual anomaly requires density values that are representative of the subsurface rock. In this analysis, only very generalized density information was available. Three borings were drilled approximately 100 feet (30 m) into the alluvium during Verification studies (Ertec, 1981b). The average density measured at the bottom of these borings was slightly less than 2.0 g/cm³. To account for compaction with depth (Woollard, 1962; and Grant and West, 1965) a density of 2.3 g/cm³ was assigned to the alluvium.

Basement rocks underlying the alluvium are assumed to be similar to rocks exposed in the nearby mountains. These consist of Tertiary volcanic and plutonic rocks and Paleozoic carbonate and siliceous sedimentary rocks. Published values for the density of the Paleozoic rocks typically range from 2.6 to 2.9 g/cm³. Carbonate rocks in the Paleozoic section are the most dense with some in Nevada and Utah having values near 2.8 g/cm³. The siliceous clastic sediments generally have densities ranging

between 2.6 and 2.7 g/cm³. Densities representative of the Tertiary volcanic rocks range from 2.0 to 2.5 g/cm³ for tuffaceous material, depending on the degree of welding, compaction, and alteration; from 2.3 to 2.6 g/cm³ for andesite and rhyolite flows; and from 2.6 to 2.7 g/cm³ for plutonic rocks.

4.3 MODELING

Modeling was accomplished using three computer programs. Two of these programs compute the gravitational effect of two- and three-dimensional bodies (Talwani and others, 1959; and Plouff, 1975). The third program calculates an inverse three-dimensional solution (Cordell, 1970). The two forward modeling programs were used to augment the inverse modeling program because the inverse program is capable of handling only a single density contrast; whereas, there are several density contrasts that contribute to the form of the residual anomaly.

A contour map showing the thickness of alluvial fill, based on the results of the inverse program, is shown in Drawing 2. The density contrast between alluvium and bedrock used in this analysis was -0.5 g/cm³. There is very little independent information with which to compare this interpretation. One well, 4N-64E-7dc, drilled to a depth of 1150 feet (351 m) did not penetrate bedrock. Its location is noted in Drawing 2.

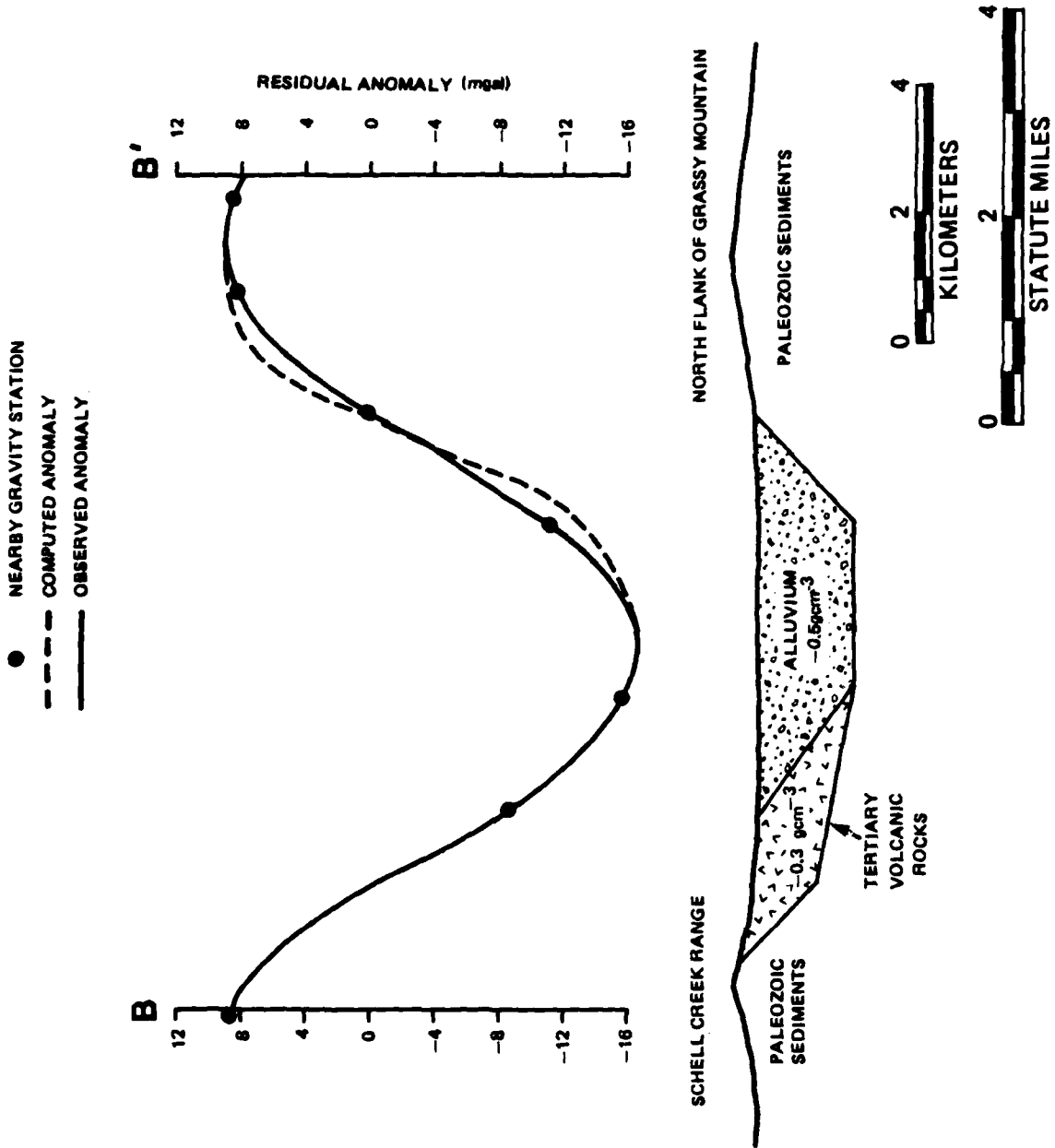
Two gravity profiles were selected for forward modeling with the Talwani program using two density contrasts: -0.5 g/cm³ for alluvium and -0.3 g/cm³ for volcanic rock. Profile B-B'

crosses the narrow northern part of the valley and Profile A-A' crosses the deepest part of the valley (Drawing 1). The interpretations resulting from forward modeling are shown in Figures 3A and 3B.

There are three principal sources of error in this analysis. First, because there is no detailed study of the true densities of the rocks, we have had to rely on estimates. Second, the inverse modeling program, upon which most of the interpretation is based, is capable of handling only a single density contrast; whereas, there are probably several density contrasts that contribute to the residual anomaly. Third, the distribution of gravity data in the mountains is not uniform, leaving areas in which the interpretation is based on interpolated trends of the data.

4.4 DISCUSSION OF RESULTS

The interpreted structure of Muleshoe Valley is shown on the contour maps of depth to rock (Drawing 2). Cross-sectional views of the interpreted stratigraphy and structure are shown in Figures 4A and 4B. The interpretations are based on geological information from published reports, analysis of aerial photographs, and geological field reconnaissance as well as gravity data. For example, wherever there was sufficient gravity data, the placement of faults could be made by finding the maximum horizontal gravity gradients. However, in areas lacking detailed gravity data, placement of faults was guided by field reconnaissance and published geologic maps. Major faults on the drawing generally comprise zones of smaller faults.



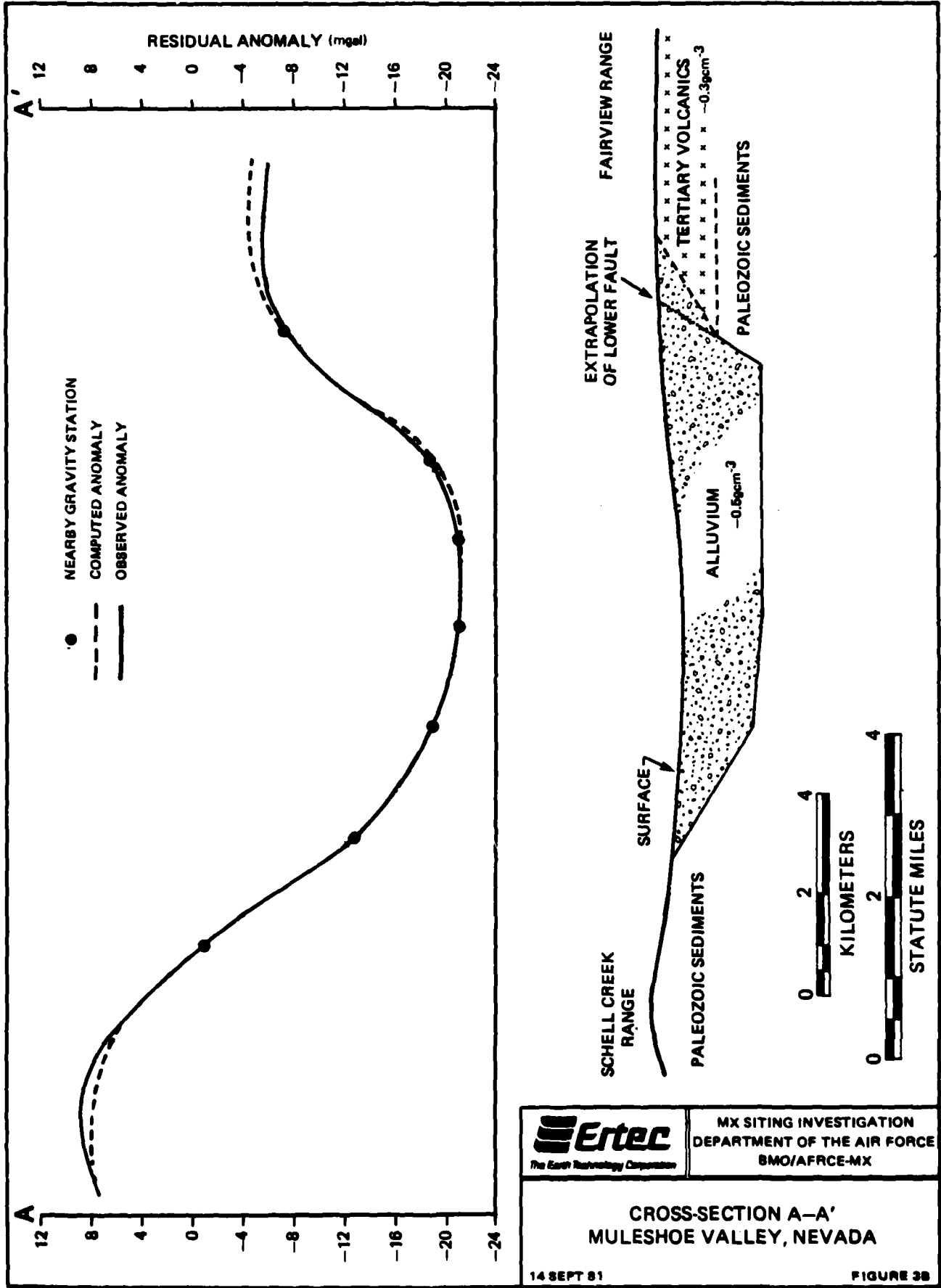
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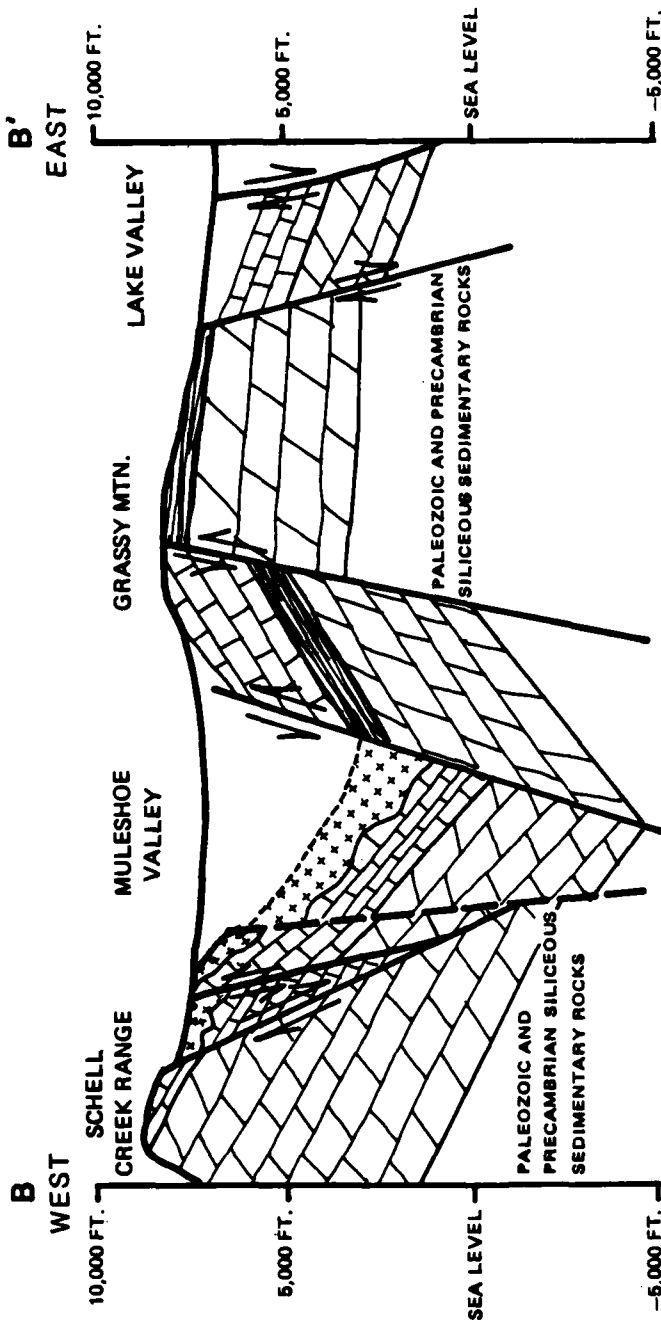
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CROSS-SECTION B-B'
MULESHOE VALLEY, NEVADA

14 SEPT 81






FIGURE 3A




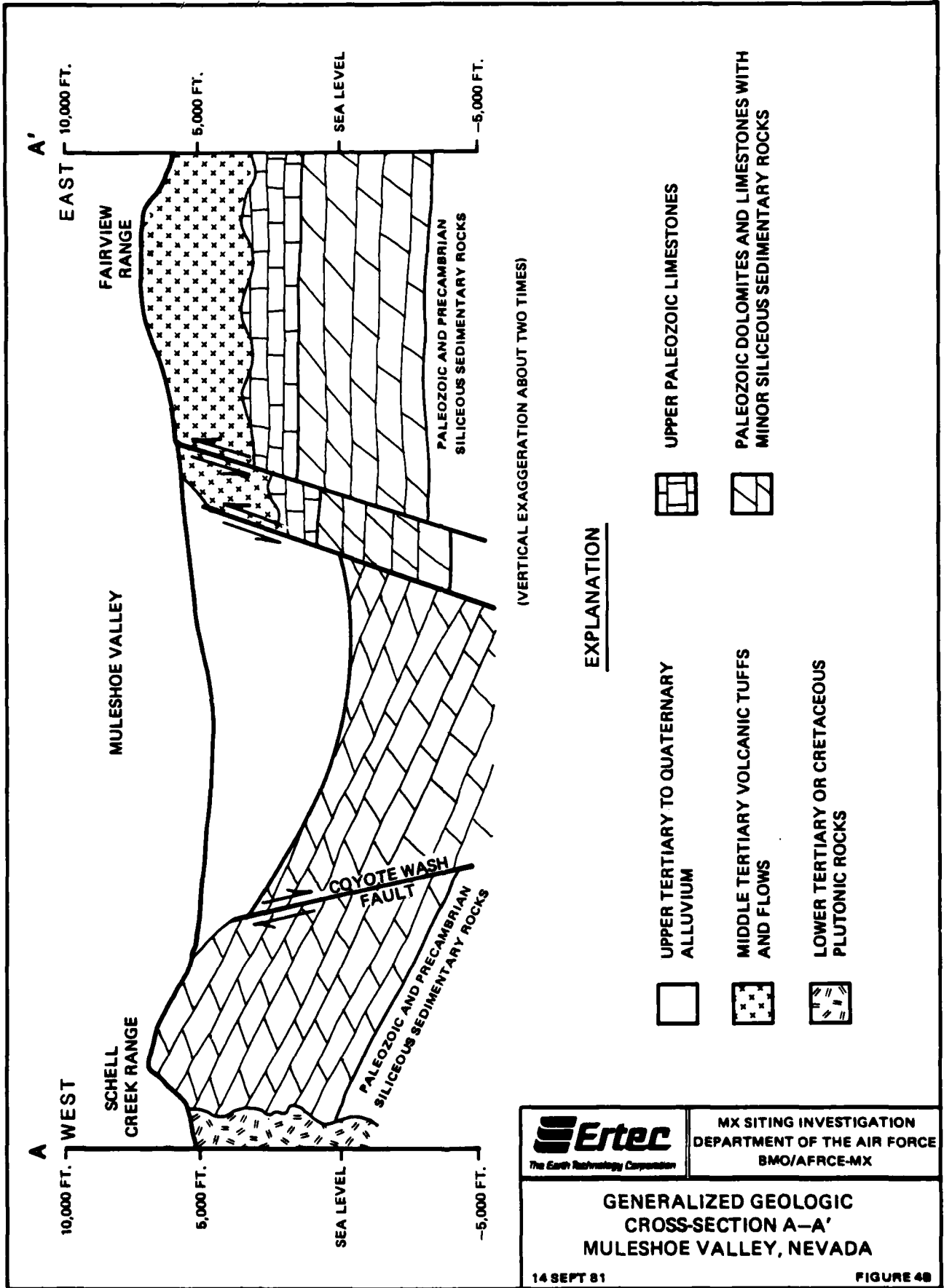


(VERTICAL EXAGGERATION ABOUT TWO TIMES)

EXPLANATION

-  UPPER TERTIARY TO QUATERNARY ALLUVIUM
-  MIDDLE TERTIARY VOLCANIC TUFFS AND FLOWS
-  UPPER PALEOZOIC LIMESTONES
-  PALEOZOIC SHALE
-  PALEOZOIC DOLOMITES AND LIMESTONES WITH MINOR SILICEOUS SEDIMENTARY ROCKS

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GENERALIZED GEOLOGIC CROSS-SECTION B-B' MULESHOE VALLEY, NEVADA	
14 SEPT 81	FIGURE 4A



In terms of its gravitational expression, Muleshoe Valley is not separated from Lake Valley as much as the topography would suggest. For example, the gravity low crossing the Fairview Range is a consequence of the low density of its volcanic materials. Muleshoe Valley is basically a graben that is wider at its southern end than at its northern end (Drawing 2 and Figure 4). The range-bounding fault on the east side of the valley appears to have more displacement than the corresponding fault on the west side of the valley.

Judging from the results of gravity modeling (Figure 3), the average slope of the base of the alluvium is about 30° on the west side of the valley. On the east side of the valley, the slope is between 45° and 60° in the north, and about 25° in the south. The 25° slope becomes steeper below the volcanic layer (Figure 3B).

The contour map of depth to rock (Drawing 2) shows that there is about 5000 feet (1524 m) of alluvium in the deepest parts of the valley. There may be as much as 3000 feet (914 m) of volcanic rock overlying Paleozoic rock in the Fairview Peak (Figure 2) area. (This estimate is based on a mass defect calculated by planimetric integration of the residual gravity anomaly [described by Grant and West, 1965] and an assumed density contrast of -0.35 g/cm^3 between volcanic rock and Paleozoic rock). It is primarily because of this substantial thickness of volcanic rock that the CBA contours defining Muleshoe Valley cut across the Fairview Range into Lake Valley.

5.0 CONCLUSION

The Complete Bouguer Anomaly data indicate that Muleshoe Valley is an irregular graben that is filled with as much as 5000 feet (1524 m) of alluvium. The deepest part of the graben is generally under the axis of the valley with the maximum depth south of the center. The northern part of the graben is less than half as wide as the southern part. On the west side of the valley, the interface between Paleozoic bedrock and alluvium has a dip of about 30° while the equivalent interface on the east side of the valley has a dip in excess of 50°. This, combined with the observation that the gravity low is displaced eastward with respect to the axis of the valley, indicates that the graben is tilted slightly to the east, as well as to the south. The southern end of the valley has a thin layer of alluvium and volcanic rocks overlying Paleozoic rocks. Apparently the southern end of the graben has no block-like edge but approaches the surface gradually by a stair-step series of unobservably small faults.

5.0 CONCLUSION

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REFERENCES

- Briggs, I. C., 1974, Machine contouring using minimum curvature, *Geophysics*, v. 39, no. 1, p. 39-48.
- Cordell, Lindreth, 1970, Iterative solution of three-dimensional gravity anomaly data, Geological Survey Computer Contribution No. 10, U.S. Geological Survey, Washington, D.C.
- Eakin, T. E., 1963, Ground-water appraisal of Pahranaagat and Pahroc Valleys, U.S. Geological Survey, Ground Water Resources-Reconnaissance Series, Report 21, p. 35.
- Ekren, E. B., Orkild, P. P., Sargent, K. A., and Dixon, C. L., 1977, Geologic map of Tertiary rocks, Lincoln County, Nevada, U.S. Geological Survey Map I-1041.
- Ertec Western, Inc., 1981a, MX siting investigation fault study summary of Muleshoe Valley, Nevada, unpublished.
- _____, 1981b, MX siting investigation verification study, Muleshoe Valley, Nevada, Volume II - Geotechnical Data.
- Fenneman, W. W., 1931, *Physiography of Western United States*, New York: McGraw-Hill, p. 534.
- Goguel, Jean, 1954, A universal table for the prediction of the lunar-solar correction in gravimetry (tidal gravity corrections); *geophysical prospecting*, v. II, supplement, March.
- Grant, F. S., and West, G. F., 1965, *Interpretation theory in applied geophysics*, New York: McGraw-Hill Book Co.
- Gunn, P. J., 1975, Linear transformations of gravity and magnetic fields, *Geophysical Prospecting*, v. 23, p. 300-312.
- Howard, E. L., 1978, Geologic map of the eastern Great Basin, Nevada and Utah, Denver, Colorado, Terra Scan Group Ltd., scale 1:250,000.
- Plouff, D., 1975, Derivation of formulas and FORTRAN programs to compute gravity anomalies of prisms, U.S. Geological Report No. USGA-GD-75-015.
- Stewart, J. H., 1980, *Geology of Nevada*, Nevada Bureau of Mines and Geology Special Publication 4, p. 136.
- Swain, C. J., 1976, A FORTRAN IV program for interpolating irregularly spaced data using the difference equations for minimum curvature, *Computers and Geosciences*, v. 1, p. 231-240.

Talwani, M., Worzel, J. L., and Landisman, M., 1959, Rapid computations for two-dimensional bodies with application to the Mendocino submarine fracture zone, *J. Geophysical Research*, v. 64, p. 49-59.

Tschanz, C. M., and Pampeyan, E. H., 1970, *Geology and mineral deposits of Lincoln County, Nevada*, Nevada Bureau of Mines Bulletin 73, p. 188.

Woollard, G. P., 1962, *The relation of gravity anomalies to surface elevation, crustal structure, and geology*, University of Wisconsin, Department of Geology, Geophysical and Polar Research Center, Madison, Wisconsin, Report 62-9.

APPENDIX A1.0

GENERAL PRINCIPLES OF THE
GRAVITY EXPLORATION METHOD

A1.0 GENERAL PRINCIPLES OF THE GRAVITY
EXPLORATION METHOD

A1.1 GENERAL

A gravity survey involves measuring the differences in the gravitational field between various points on the earth's surface. The gravity values are associated with the force which causes a 1 gm mass to be accelerated at 980 cm/sec^2 . This force is normally referred to as a 1 g force.

Even though in many applications the gravitational field at the earth's surface is assumed to be constant, small but distinguishable differences in gravity occur from point to point. The differences in gravity are caused by geometrical effects, such as differences in elevation and latitude, and by lateral variations in density within the earth. The lateral density variations are a result of changes in geologic conditions. For measurements at the surface of the earth, the largest factor influencing the pull of gravity is the density of all materials between the center of the earth and the point of measurement.

To detect changes produced by differing geological conditions, it is necessary to detect differences in the gravitational field as small as a few milligals. A milligal is equal to 0.001 cm/sec^2 or 0.00000102 g . To recognize changes due to geological conditions, the measurements are "corrected" to account for changes due to differences in elevation and latitude.

A difference in gravity between two points which is not caused by the effects of known geometrical differences, such as in

elevation, latitude, and surrounding terrain, is referred to as an "anomaly." The anomaly is the basic concept of the gravitational exploration method. If, instead of being an oblate spheroid characterized by complex density variations, the earth were made up of concentric, homogeneous shells, the gravitational field would be the same at all points on the surface of the earth. The complexities in the earth's shape and material distribution are the reason that the pull of gravity is not the same from place to place.

An anomaly reflects lateral differences in material densities. The gravitational attraction is smaller at a place underlain by relatively low density material than it is at a place underlain by a relatively high density material. The term "negative gravity anomaly" describes a situation in which the pull of gravity within a prescribed area is small compared to the area surrounding it. Low-density alluvial deposits in basins such as those in the Nevada-Utah region produce negative gravity anomalies in relation to the gravity values in the surrounding mountains which are formed by more dense rocks.

The objective of gravity exploration is to deduce the variations in geologic conditions that produce the gravity anomalies identified during a gravity survey.

A1.2 INSTRUMENTS

The gravity field data was measured with a LaCoste and Romberg Model D gravimeter. The sensing element of the gravimeter is a mass suspended by a zero-length spring. Deflections of the

mass from a null position are proportional to changes in gravitational attraction. These instruments are sealed and compensated for atmospheric pressure changes. They are maintained at a constant temperature by an internal heater element and thermostat. The absolute value of gravity is not measured directly by a gravimeter. It measures relative values of gravity between one point and the next. Gravitational differences as small as 0.01 milligal can be measured.

A1.3 FIELD PROCEDURES

The gravimeter readings were calibrated in terms of absolute gravity by taking readings twice daily at nearby USGS gravity base stations. Gravimeter readings fluctuate because of small time-related deviations due to the effect of earth tides and instrument drift. Field readings were corrected to account for these deviations. The magnitude of the tidal correction was calculated using an equation suggested by Goguel (1954):

$$C = P + N \cos \phi (\cos \phi + \sin \phi) + S \cos \phi (\cos \phi - \sin \phi)$$

where C is the tidal correction factor, P, N, and S are time-related variables, and ϕ is the latitude of the observation point. Tables giving the values of P, N, and S are published annually by the European Association of Exploration Geophysicists.

The meter drift correction was based on readings taken at a designated base station at the start and end of each day. Any difference between these two readings after they were corrected for tidal effects was considered to have been the result of

instrumental drift. It was assumed that this drift occurred at a uniform rate between the two readings. Corrections for drift were typically only a few hundredths of a milligal. Readings corrected for tidal effects and instrumental drift represented the observed gravity at each station. The observed gravity values represent the total gravitational pull of the entire earth at the measurement stations.

A1.4 DATA REDUCTION

Several corrections or reductions are made to the observed gravity to isolate the portion of the gravitational pull which is due to the crustal and near-surface materials. The gravity remaining after these reductions is called the "Bouguer Anomaly." Bouguer Anomaly values are the basis for geologic interpretation. To obtain the Bouguer Anomaly, the observed gravity is adjusted to the value it would have had if it had been measured at the geoid, a theoretically defined surface which approximates the surface of mean sea level. The difference between the "adjusted" observed gravity and the gravity at the geoid calculated for a theoretically homogeneous earth is the Bouguer Anomaly.

Four separate reductions, to account for four geometrical effects, are made to the observed gravity at each station to arrive at its Bouguer Anomaly value.

a. Free-Air Effect: Gravitational attraction varies inversely as the square of the distance from the center of the earth. Thus corrections must be applied for elevation. Observed

gravity levels are corrected for elevation using the normal vertical gradient of:

$$FA = -0.09406 \text{ mg/ft } (-0.3086 \text{ milligals/meter})$$

where FA is the free-air effect (the rate of change of gravity with distance from the center of the earth). The free-air correction is positive in sign since the correction is opposite the effect.

b. Bouguer Effect: Like the free-air effect, the Bouguer effect is a function of the elevation of the station, but it considers the influence of a slab of earth materials between the observation point on the surface of the earth and the corresponding point on the geoid (sea level). Normal practice, which is to assume that the density of the slab is 2.67 grams per cubic centimeter was followed in these studies. The Bouguer correction (B_C), which is opposite in sign to the free-air correction, was defined according to the following formula.

$$B_C = 0.01276 (2.67) h_f \text{ (milligals per foot)}$$

$$B_C = 0.04185 (2.67) h_m \text{ (milligals per meter)}$$

where h_f is the height above sea level in feet and h_m is the height in meters.

c. Latitude Effect: Points at different latitudes will have different "gravities" for two reasons. The earth (and the geoid) is spheroidal, or flattened at the poles. Since points at higher latitudes are closer to the center of the earth than points near the equator, the gravity at the higher latitudes is larger. As the earth spins, the centrifugal acceleration

causes a slight decrease in gravity. At the higher latitudes where the earth's radii are smaller, the centrifugal acceleration diminishes. The gravity formula for the Geodetic Reference System, 1967, gives the theoretical value of gravity at the geoid as a function of latitude. It is:

$$g = 978.0381 (1 + 0.0053204 \sin^2 \phi - 0.0000058 \sin^2 2\phi) \text{ gals}$$

where g is the theoretical acceleration of gravity and ϕ is the latitude in degrees. The positive term accounts for the spheroidal shape of the earth. The negative term adjusts for the centrifugal acceleration.

The previous two corrections (free air and Bouguer) have adjusted the observed gravity to the value it would have had at the geoid (sea level). The theoretical value at the geoid for the latitude of the station is then subtracted from the adjusted observed gravity. The remainder is called the Simple Bouguer Anomaly (SBA). Most of this gravity represents the effect of material beneath the station, but part of it may be due to irregularities in terrain (upper part of the Bouguer slab) away from the station.

d. Terrain Effect: Topographic relief around the station has a negative effect on the gravitational force at the station. A nearby hill has upward gravitational pull and a nearby valley contributes less downward attraction than a nearby material would have. Therefore, the corrections are always positive. Corrections are made to the SBA when the terrain effects were 0.1 milligal or larger. Terrain corrected Bouguer values are

called the Complete Bouguer Anomaly (CBA). When the CBA is obtained, the reduction of gravity at individual measurement points (stations) is complete.

A1.5 INTERPRETATION

To interpret the gravity data, the portion of the CBA that might be caused by the light-weight, basin-fill material must be separated from that caused by the heavier bedrock material which forms the surrounding mountains and presumably the basin floor. The first step is to create a regional field. A regional field is an estimation of the values the CBA would have had if the light-weight sediments (the anomaly) had not been there. Since the valley-fill sediments are absent at the stations read in the mountains, one approach is to use the CBA values at bedrock stations as the basis for constructing a second order polynomial surface to represent a regional field over the valley.

Where there are insufficient bedrock stations to define a satisfactory regional trend, another approach is to estimate the regional by the process of upward continuation of the CBA field.

In Potential Theory, a field normal to a surface, regardless of its actual source, may be considered as originating in an areal distribution of mass on that surface. If the field strength is known the surface density of mass (grams per square centimeter) can be calculated. The observed gravity field at the surface of the earth approximately fulfills the requirements of this theory: thus the observed (Bouguer anomaly) field can be used

to compute a surficial distribution of mass which would reproduce the field, and most importantly, account for the gravity field anywhere above the surface of observation. On this basis, the Bouguer anomaly field is readily "continued" to level surfaces above the ground.

An important property of such "upward continuation" is that the resultant field with increasing altitudes of continuation, changes more with respect to shallow sources than it does with respect to deeper sources. The anomalous parts of the field ascribed to shallow density distribution tend to vanish as the continuation is carried upward whereas the field produced by deeper sources changes only slightly, so that upward continuations produce "regional"-type fields.

The difference between the CBA and the regional field is called the "residual" field or residual anomaly. The residual field is the interpreter's estimation of the gravitational effect of the geologic anomaly. The zero value of the residual anomaly is not exactly at the rock outcrop line but at some distance on the "rock" side of the contact. The reason for this is found in the explanation of the terrain effect. There is a component of gravitational attraction from material which is not directly beneath a point.

If the "regional" is well chosen, the magnitude of the residual anomaly is a function of the thickness of the anomalous (fill) material and the density contrast. The density contrast is the difference in density between the alluvial and bedrock material.

If this contrast were known, an accurate calculation of the thickness could be made. In most cases, the densities are not well known and they also vary within the study area. In these cases, it is necessary to use typical densities for materials similar to those in the study area.

If the selected average density contrast is smaller than the actual density contrast, the computed depth to bedrock will be greater than the actual depth and vice-versa. The computed depth is inversely proportional to the density contrast. A ten percent error in density contrast produces a ten percent error in computed depth. An iterative computer program is used to calculate a subsurface model which will yield a gravitational field to match (approximately) the residual gravity anomaly.

The second vertical derivative (SVD) of gravitational field is used to aid the interpreter in evaluating the subsurface mass distribution. Once the CBA field has been projected onto a uniform grid system, its SVD at the grid nodes is readily computed. In accordance with Laplace's Equation in Free Space, the negative of the second vertical derivative is equal to the sums of the second derivatives in the x-direction and in the y-direction. The second vertical derivative is an indication of the curvature of the Bouguer anomaly field. In particular the zero-value of the SVD indicates the inflection in the field as it changes from "concave-upward" (algebraically negative SVD) to "convex-upward" (algebraically positive SVD). In a general way the zero SVD falls on the tightest contours of the field and

where contours are nearly parallel its location can be established by eye. However, where contours diverge, converge, or change direction this is not always so readily done. The zero SVD contour line may be an indicator of a line of faulting, the pinchout of a stratum, truncation of a stratum at an unconformity or merely a marked change in shape or in density of a geologic unit.

E-TR-33-MS

APPENDIX A2.0
MULESHOE VALLEY, NEVADA
GRAVITY DATA

MULESHOE VALLEY GRAVITY DATA

STATION IDENT.	LAT. DEG MIN	LONG. DEG MIN	ELEV. +CODE	TER-COR. IN/OUT	NORTH UTM	EAST UTM	OBSV GRAV	THEO GRAV	FAA -----	CBA +1000
0260	38 515	114387561690T	0	216421775	70647143554199963	1610	80796			
0262	38 520	114407455889T	0	138421777	70356146709199970	-680	80388			
0263	38 531	114414254629T	0	132421795	70256147609199986	-990	80512			
0282	38 564	114427452881T	0	136421851	70061149557200035	-730	81366			
0266	38 575	114392958560T	0	187421884	70565145175200051	200	80427			
0275	38 593	114402556280T	0	145421914	70424146313200077	-820	80125			
0278	38 602	114417554062T	0	128421925	70204147913200090	-1320	80368			
5029	38 622	114413254511T	0	137421963	70266147426200120	-1410	80137			
0294	38 623	114449152500T	0	121421952	69741149611200121	-1120	81091			
0274	38 627	114406255600T	0	135421975	70368146706200127	-1120	80055			
1914	38 640	114489250951T	0	92421970	69155149925200146	-2290	80422			
0157	38 658	114478351161T	0	96422007	69313149692200172	-2350	80296			
0273	38 664	114399056660T	0	149422046	70472145711200181	-1170	79659			
0154	38 675	114467551801T	0	101422042	69470149198200197	-2270	80161			
0297	38 697	114423755000T	0	122422080	70110147246200215	-1230	80132			
0158	38 692	114456152530T	0	107422078	69636149307200222	-1500	80687			
0295	38 705	114443353930T	0	123422106	69822148584200241	-920	80803			
0296	38 710	114432055230T	0	129422119	69987147473200248	-820	80469			
7412	38 710	114441153930T	0	127422116	69854148611200248	-900	80837			
0276	38 711	114412156220T	0	121422128	70278146571200250	-790	80151			
0156	38 721	114476051709T	0	105422124	69344149708200265	-1910	80555			
0271	38 730	114391858281T	0	149422171	70574145057200278	-390	79879			
0270	38 740	114382459911T	0	183422193	70711144838200292	900	80653			
7411	38 750	114380061529T	0	150422212	70745143580200307	1140	80320			
0008	38 751	114450555479T	0	113422189	69715147515200308	-600	80593			
0781	38 761	114430357749T	0	126422214	70010146652200323	650	81086			
0780	38 762	114438556181T	0	107422213	69890146994200325	-480	80467			
0782	38 771	114422556270T	0	124422236	70123146360200333	-1040	79884			
0790	38 794	114396857621T	0	145422288	70498145868200371	-300	80195			
0784	38 807	114442158560T	0	147422295	69835145467200390	160	80337			
0005	38 808	114485256339T	0	197422282	69206147010200392	-380	80597			
0783	38 815	114427456870T	0	126422315	70050146386200402	-520	80216			
0785	38 819	114449157021T	0	122422315	69733146525200408	-240	80432			
0004	38 833	114499751509T	0	119422323	68993149882200428	-2090	80459			
0793	38 834	114387858471T	0	202422365	70627146123200430	690	80952			
0009	38 861	114455056650T	0	115422391	69645146898200469	-280	80515			
0786	38 873	114435257470T	0	122422420	69933146343200487	-80	80442			
0010	38 890	114467256001T	0	112422421	69465147321200497	-500	80512			
0800	38 892	114427660341T	0	149422458	70043144701200515	940	80519			
0011	38 892	114477355840T	0	111422440	69317147348200515	-640	80431			
0012	38 899	114489852510T	0	105422449	69135149678200525	-1450	80745			

MULESHOE VALLEY GRAVITY DATA

STATION IDENT.	LAT. DEG MIN	LONG. DEG MIN	ELEV. +CODE	TER-COR. IN/OUT	NORTH UTM	EAST UTM	OBSV GRAV	THEO GRAV	FAA	CBA +1000
0013	38 914	11449995	1860T	0	132422473	6898614991	1200547	-1850	80592	
0787	38 922	11444455	6850T	0	135422507	6979514668	9200559	-390	80355	
0018	38 939	11449345	5850T	0	152422521	6908014726	2200584	-780	80322	
0020	38 944	11446695	8261T	0	131422540	6946714591	9200591	120	80391	
0021	38 946	11445585	9341T	0	175422547	6962914518	7200594	410	80355	
0811	38 957	11437936	1329T	0	123422596	7074614434	1200610	1410	80633	
0804	38 963	11443236	1171T	0	163422587	6997114469	4200619	1610	80923	
0019	38 966	11448375	4692T	0	94422575	6922114830	5200623	-870	80574	
0022	38 971	11445725	8471T	0	119422630	6960714589	9200660	230	80419	
0803	38 997	11444006	0732T	0	150422647	6985714478	2200669	1230	80680	
0026	381003	11449825	2251T	0	128422638	6900714974	0200677	-1780	80518	
7286	381005	11436216	0246T	0	117422691	7099514383	1200680	-170	79397	
0805	381005	11442636	6581T	0	386422667	7005714111	1200680	3050	80736	
0024	381013	11447525	4491T	0	128422665	6934314839	8200692	-1030	80508	
0023	381024	11446585	6280T	0	97422688	6948014701	5200708	-750	80147	
0025	381034	11448765	6319T	0	111422699	6916114686	5200723	-880	80021	
7442	381055	11443686	0397T	0	140422756	6990214549	8200754	1550	81110	
0819	381068	11444255	9531T	0	125422778	6981814572	2200773	940	80775	
0032	381074	11447625	6969T	0	114422777	6932514644	9200781	-740	79944	
0034	381079	11445705	7110T	0	103422793	6960614719	8200789	120	80763	
0033	381084	11446725	5341T	0	100422799	6945614784	8200796	-890	80340	
0031	381070	11449575	2710T	0	126422800	6904014900	4200805	-2210	79936	
0820	381149	11444396	1680T	0	161422927	6979414378	4200891	910	80041	
1582	381230	11435306	0961T	0	122423110	7111714370	2201010	30	79382	
7440	381410	11448385	3960T	0	132423396	6920014795	1201274	-2560	79172	
1210	381418	11435306	1821T	0	115423458	7110814351	6201285	380	79425	
7049	381430	11435506	1837T	0	119423480	7107814360	1201303	460	79499	
7441	381515	11442206	0843T	0	121423612	7009714442	8201427	220	79611	
1581	381530	11435406	2359T	0	114423665	7108814324	2201449	440	79314	
1580	381730	11435906	2139T	0	132424033	7100514380	2201742	500	79462	
7443	381775	11449606	2487T	0	191424067	6900614505	9201808	2020	80921	
7424	381825	11441905	9551T	0	195424186	7012614522	8201882	-630	79255	
7032	381916	11436306	1568T	0	141424375	7093814429	1202015	180	79341	
1209	381932	11436426	1624T	0	140424405	7092014425	7202039	180	79320	
7423	382020	11443355	7799T	0	273424542	6990614605	5202168	-1730	78823	
1579	382060	11436906	0679T	0	127424640	7084414496	2202226	-190	79247	
7031	382304	11437385	9760T	0	125425089	7076214608	3202584	-280	79455	
1578	382330	11438106	0259T	0	145425227	7065414630	2202696	280	79885	
7283	382535	11443386	2221T	0	169425494	6987814471	9202923	320	79289	
7419	382545	11448656	0312T	0	135425494	6911114733	4202938	1120	80705	
7220	382550	11448656	0285T	0	134425503	6911014733	5202945	1090	80684	

MULESHOE VALLEY GRAVITY DATA

STATION IDENT	LAT. DEG MIN	LONG. DEG MIN	ELEV. +CODE	TER-COR. IN/OUT	NORTH UTM	EAST UTM	OBSV GRAV	THEO GRAV	FAA	CBA +1000
1208	382590	114382	259892T	0	230425615	706261465	25203004		-140	79670
1577	382720	114386	059419T	0	330425854	705651470	72203195		-230	79840
7030	382782	114386	259846T	0	204425969	705591467	72203286		-210	79574
1589	382900	114387	060141T	0	134426187	705421466	52203459		-230	79394
LV0124	382875	114357	15915Y	0	97426152	709781444	75203422		-3278	76645
LV0162	382686	114401	28860V	463635	425785	703451272	01203145		7461	80923
LV0164	382468	114381	15998C	0	180425390	706481464	94202825		120	79843
LV0173	381791	114379	47866V	561888	424138	707051328	02201832		5013	80128
LV0201	381690	114356	4227Y	0	123423960	710451435	19201684		443	79328
LV0163	382590	114381	56007C	0	408425615	706361464	18203004		-49	79871
LV0262	381250	114363	56538Y	7	260423143	709631411	18201039		1615	79583
CAV073	382274	114486	17406C	322236	424993	691281373	44202540		4514	81522
CAV075	382399	114487	560420T	0	241425224	691021471	25202724		1267	80901
CAV086	382676	114456	07383V	38	951425747	695491376	67203130		4030	79838
CAV093	382030	114497	8114C	673245	424539	690141318	91202182		6086	81724
CAV085	382461	114476	97048C	94	573425342	692541395	56202814		3080	79708
LV0156	382896	114438	27574C	16	591426160	697971370	75203453		4914	79688
LV0165	382428	114407	07048C	6	347425306	702731414	07202766		4979	81293
LV0170	382090	114403	18613C	122214	424682	703451289	58202270		7766	80615
LV0282	38 539	114364	57907V	261624	421828	709821316	83199998		6114	80795
CAV070	382764	114499	26010S	0	114425895	689161453	14203259		-1381	78234
CAV074	382317	114491	760321T	0	176425071	690451468	61202603		1030	80632
CAV076	382482	114493	35981S	0	117425375	690141459	82202845		-573	79145
CAV077	382572	114495	05990S	0	109425541	689861455	49202977		-1053	78626
CAV078	382668	114497	16001C	0	111425719	689861454	34203118		-1205	78438
CAV079	382789	114488	6018S	0	121425945	690691454	51203296		-1205	78390
CAV080	382895	114490	36030S	0	127426140	690401452	73203451		-1426	78134
CAV081	382983	114486	86043S	0	138426304	690871455	88203581		-1118	78409
CAV082	382885	114476	260420T	0	154426127	692451460	84203437		-487	79059
CAV083	382678	114480	36022C	0	142425742	691951470	48203133		592	80195
CAV084	382547	114484	460312T	0	157425499	691411473	21202941		1142	80729
CAV087	382764	114464	26321C	0	218425907	694251457	71203259		2004	80663
CAV088	382856	114462	06401B	0	248426096	694531452	79203409		2116	80532
CAV091	382196	114485	061089T	0	225424849	691481466	57202426		1728	81117
CAV092	382164	114495	960200T	0	345424786	689901470	56202379		1336	81148
LV0123	382970	114369	45921Y	0	103426323	707941447	49203562		-3087	76821
LV0129	382727	114367	75919Y	0	114425874	708311452	36203205		-2261	77665
LV0136	382618	114363	15921Y	0	109425674	709031454	14203045		-1905	78010
LV0137	382501	114359	75927Y	0	104425459	709581454	54202873		-1636	78252
LV0144	382421	114357	5929Y	0	103425312	709941453	13202756		-1641	78240
LV0146	382324	114358	45936Y	0	104425132	709851453	89202613		-1357	78501

MULESHOE VALLEY GRAVITY DATA

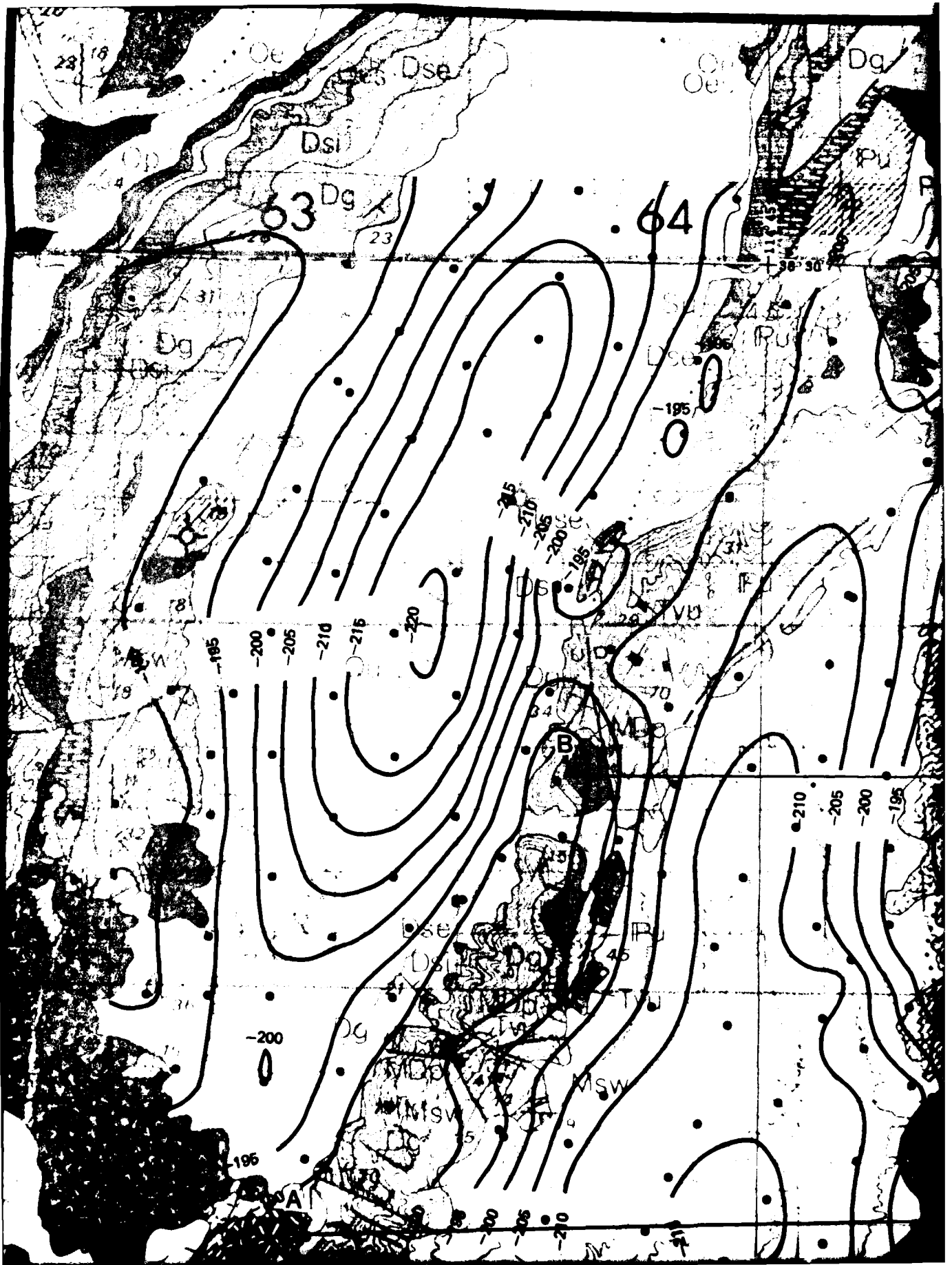
STATION IDENT.	LAT. DEG MIN	LONG. DEG MIN	ELEV. +CODE	TER-COR. IN/OUT	NORTH UTM	EAST UTM	OBSV GRAV	THEO GRAV	FAA	CBA +1000
LV0153	382979	1143884	5998B	0	122426332	70518146686203575			-438	79227
LV0154	382927	1144004	6097C	0	153426231	70346146165203499			50	79408
LV0155	382930	1144103	6220C	0	164426233	70202145257203503			296	79245
LV0157	382794	1144171	6316C	0	212425979	70109144777203303			919	79589
LV0158	382846	1144046	6246B	0	209426080	70288145570203380			977	79883
LV0159	382855	1143932	6056C	0	185426101	70454146520203393			124	79654
LV0160	382849	1143766	5937Y	0	120426096	70695145896203384			-1611	78259
LV0161	382734	1143843	5938C	0	240425880	70589146831203215			-498	79489
LV0166	382377	1143898	6146C	0	183425218	70526146614202691			1768	80989
LV0167	382326	1144027	6445Y	0	206425119	70340144953202616			2997	81221
LV0168	382278	1143787	6048C	0	138425039	70692145970202546			346	79856
LV0169	382172	1143877	6375B	0	235424840	70566144535202391			2145	80637
LV0171	382041	1143836	6407B	0	218424599	70632143833202198			1937	80302
LV0172	381931	1143771	62441T	0	171424435	70731144516202066			1218	80092
LV0174	382240	1143648	59550T	0	111424974	70896145699202490			-744	79056
LV0175	382240	1143541	59531T	0	100424978	71052145135202490			-1327	78469
LV0180	382129	1143541	5960Y	0	103424773	71057145148202327			-1086	78689
LV0181	382131	1143697	6023Y	0	126424771	70830145488202330			-156	79427
LV0182	382046	1143676	6075Y	0	122424614	70865144673202206			-355	79046
LV0188	381999	1143526	5997S	0	105424533	71085144303202137			-1392	78259
LV0189	381910	1143637	61621T	0	142424364	70928144232202006			221	79346
LV0202	381574	1143557	62359T	0	116423746	71061143241201514			420	79266
LV0254	381460	1143550	6190S	0	117423535	71076143556201347			468	79473
LV0296	382327	1143716	5957Y	0	120425132	70793146043202618			-510	79293
MSV076	381303	1144439	5847V	23	114423212	69787145469201117			-618	79577
MSV075	381149	1144439	6168S	201	159422927	69794143805200891			966	80289
CAV086	382676	1144560	7383C	238	634425747	69549137676203130			4040	79730
CAV090	382946	1144466	8269S	405	1503426250	69673132228203526			6539	80244
CAV093	382030	1144947	8114C	974	2209424539	69014131888202182			6084	81592
MSV013	382314	1144120	6652S	10	198425093	70205143740202599			3751	81271
MSV014	382161	1144047	8396S	599	2016424813	70319130486202374			7146	81124
MSV021	381935	1144165	7396S	609	874424391	70157136904202043			4477	80734
MSV015	381801	1143942	7418S	22	820424151	70489136569201847			4546	80087
MSV016	381545	1143959	6877C	16	254423677	70476139408201471			2665	79480
MSV052	381849	1144961	7190C	27	787424204	69001138329201917			4088	80379
MSV080	381213	1144109	7239S	22	698423057	70272136185200985			3337	79367
CAV085	382461	1144769	7048C	0	576425342	69254139553202814			3077	79614
CAV091	382196	1144850	61089T	0	219424849	69148146653202426			1724	81107
MSV001	382271	1144651	6039V	0	114424995	69434145796202536			97	79614
MSV002	382330	1144664	6150C	0	131425196	69410145525202696			712	79867
MSV003	382423	1144540	6202C	0	125425280	69589144875202759			488	79460

MULESHOE VALLEY GRAVITY DATA

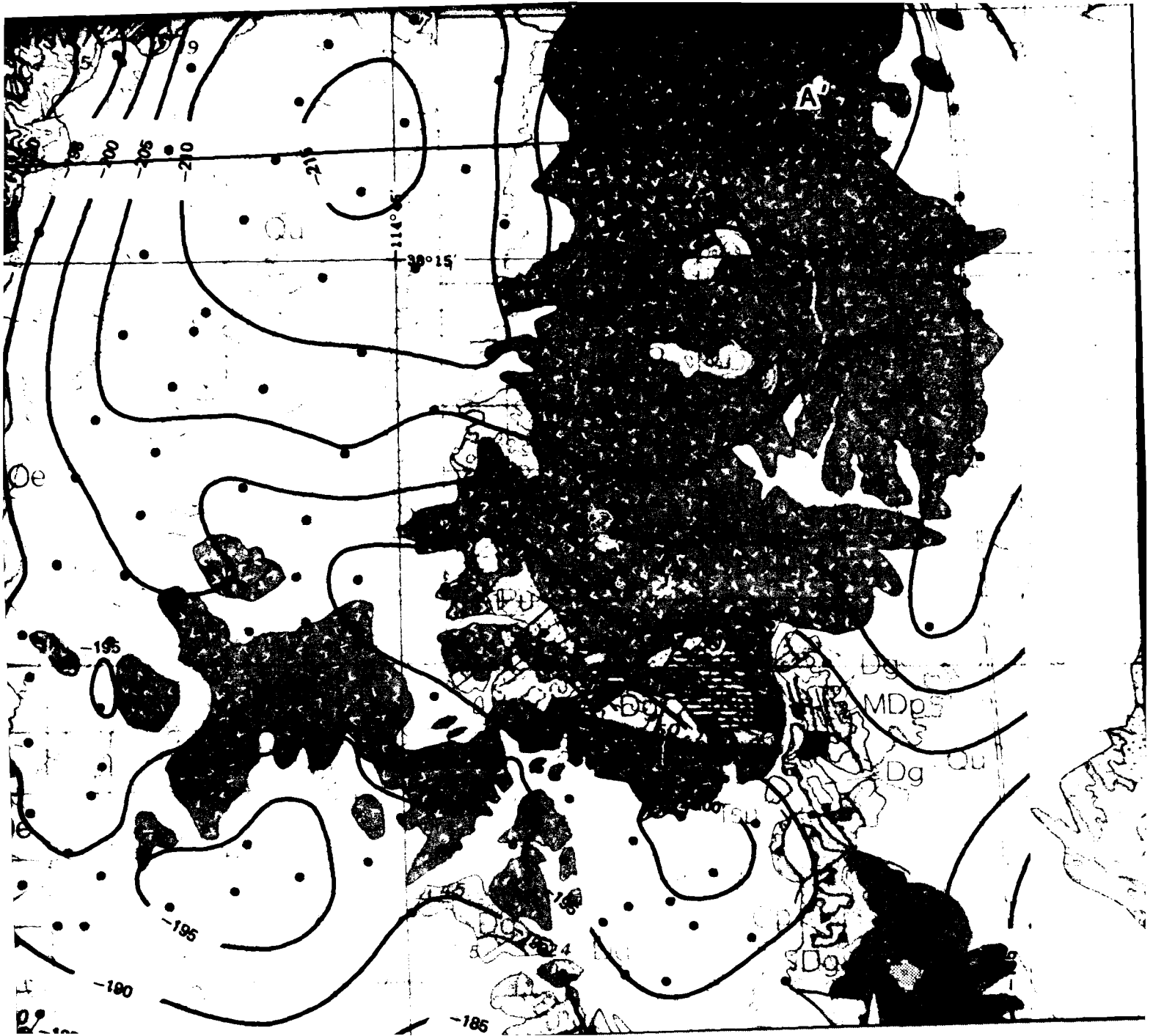
STATION IDENT.	LAT. DEG MIN	LONG. DEG MIN	ELEV. +CODE	TER-COR. IN/OUT	NORTH UTM	EAST UTM	OBSV GRAV	THEO GRAV	FAA	CBA +1000
MSV004	382295	1144515	6002C	0	114425044	69631145270202571			-812	78831
MSV005	382308	1144390	6088S	0	136425073	69813145024202590			-267	79104
MSV006	382441	1144379	6134S	0	145425319	69823144979202785			-74	79149
MSV007	382538	1144346	6230S	0	167425500	69866144692202928			401	79319
MSV008	382655	1144272	6396S	0	203425719	69968144182203099			1282	79670
MSV009	382725	1144209	6445S	0	220425850	70057144036203202			1494	79732
MSV011	382416	1144273	6338S	0	171425277	69978144376202748			1281	79835
MSV012	382285	1144276	6356C	0	179425034	69980144494202556			1760	80260
MSV017	381547	1144098	6255S	0	139423676	70273143308201474			706	79510
MSV018	381728	1144139	6310S	0	158424009	70205143858201740			1508	80144
MSV019	381830	114412462300T		0	175424198	70222144744201889			1491	80418
MSV020	381874	114404664071T		0	203424282	70334143725201954			2074	80425
MSV022	382182	1144269	6390V	0	211424844	69995144371202405			2108	80524
MSV023	382073	1144268	6410S	0	214424642	70001144215202245			2301	80652
MSV024	381954	1144233	6347S	0	224424423	70057144075202071			1742	80318
MSV025	381843	114422861519T		0	161424218	70070144788201908			781	79959
MSV026	381645	1144218	6146B	0	124423852	70094144522201618			749	79911
MSV027	381547	1144318	5981C	0	111423668	69952144473201474			-710	79001
MSV028	381620	1144382	5899V	0	108423800	69856144698201581			-1364	78624
MSV029	381735	1144323	5964V	0	115424015	69936144709201750			-909	78864
MSV030	381894	114431260289T		0	154424310	69945144992201983			-248	79343
MSV031	381937	114438759550T		0	131424386	69834145203202046			-796	79024
MSV032	382069	1144398	6013S	0	140424630	69812145439202239			-207	79424
MSV033	382211	1144437	5997S	0	131424891	69749145132202448			-873	78804
MSV034	381817	1144465	5842V	0	112424162	69726145292201870			-1595	78591
MSV035	381680	1144483	5798S	0	104423908	69706145184201669			-1917	78412
MSV036	381592	1144556	5722V	0	103423742	69603145524201540			-2163	78424
MSV037	381557	114475255121T		0	117423671	69319147018201489			-2595	78722
MSV038	381635	114469855571T		0	112423817	69394146756201603			-2547	78611
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MSV040	381786	1144607	5647C	0	114424099	69520146435201825			-2243	78611
MSV041	381926	1144556	5722C	0	120424360	69588146193202030			-1984	78620
MSV042	382038	1144579	5805C	0	119424566	69550146109202194			-1450	78869
MSV043	382139	1144530	5905C	0	116424755	69616145386202342			-1380	78595
MSV044	382143	114467459130T		0	131424757	69406146282202348			-415	79548
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MSV046	381951	1144691	5982C	0	143424402	69390145419202066			-347	79393
MSV047	381825	1144775	5915C	0	143424166	69273145347201882			-865	79104
MSV048	381757	1144836	5892C	0	139424038	69187145413201782			-916	79127
MSV049	381650	1144876	5782C	0	125423838	69134145983201625			-1224	79180
MSV050	381514	1144919	5595C	0	119423585	69077146995201426			-1774	79262

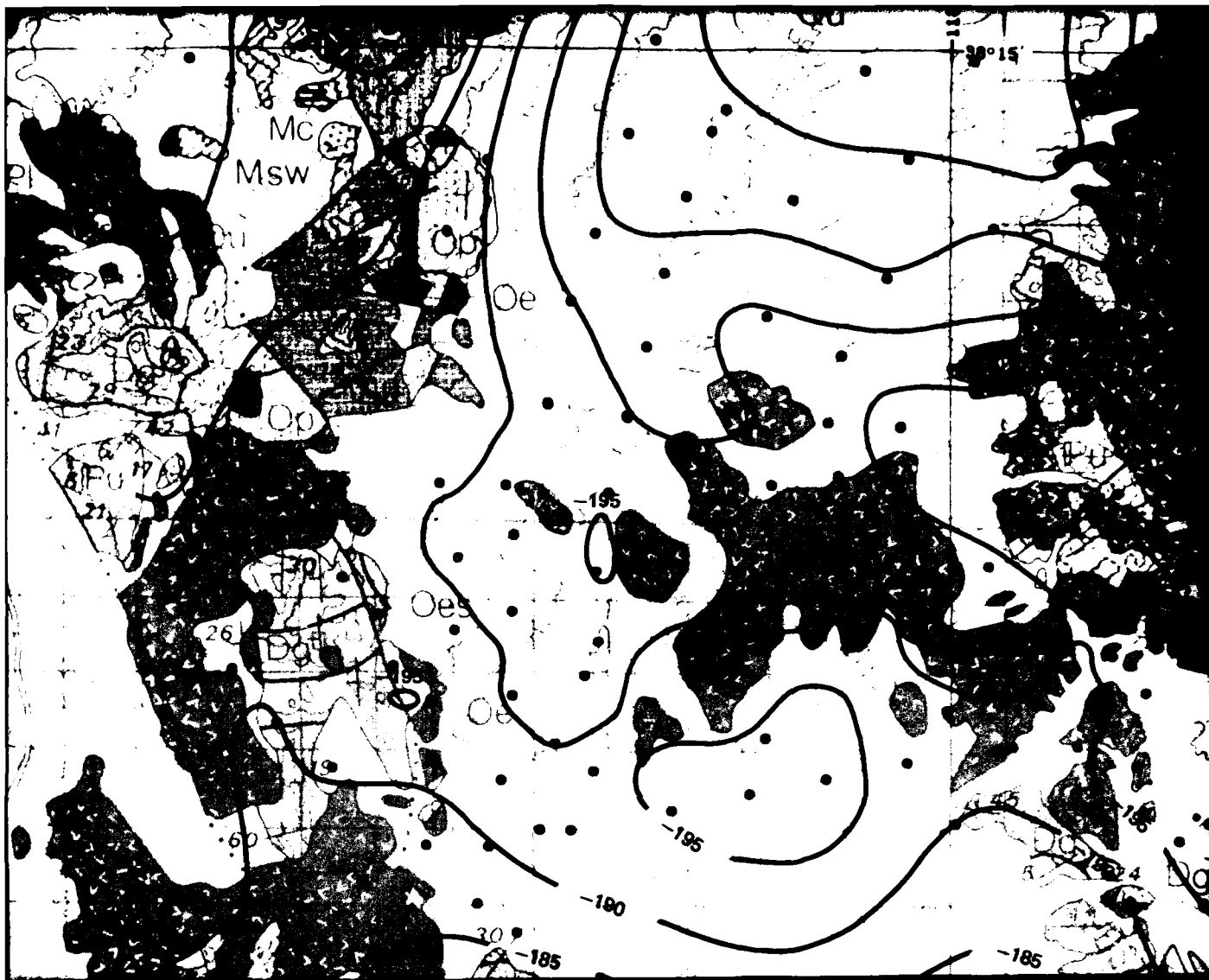
MULESHOE VALLEY GRAVITY DATA

STATION IDENT.	LAT. DEG MIN	LONG. DEG MIN	ELEV. +CODE	TER-COR. IN/OUT	NORTH UTM	EAST UTM	OBSV GRAV	THEO GRAV	FAA	CBA +1000
MSV051	381765	1144951	6110C	0	174424049	69019145440201794			1152	80487
MSV061	381407	1144956	5552Y	0	115423386	69028147159201269			-1858	79321
MSV062	381435	1144818	5468Y	0	118423443	69228147432201310			-2416	79052
MSV063	381338	1144874	53970T	0	128423261	69150147931201168			-2444	79276
MSV064	381252	1144904	53599T	0	118423101	69110148495201042			-2103	79734
MSV065	381169	1144930	53271T	0	122422947	69076148650200920			-2137	79816
MSV066	381039	1144957	52710T	0	124422798	69040149016200803			-2181	79965
MSV067	381074	1144762	5697Y	0	113422777	69325146467200781			-697	79985
MSV068	381203	1144760	5453V	0	104423016	69323148344200970			-1306	80199
MSV069	381334	1144723	5516V	0	99423259	69371147241201162			-2007	79278
MSV070	381479	1144622	5649V	0	97423531	69512146112201375			-2096	78734
MSV071	381330	1144559	5692V	0	100423350	69608146031201230			-1628	79058
MSV072	381247	1144589	5632V	0	108423103	69570146596201035			-1433	79466
MSV073	381160	1144654	5539V	0	104422940	69479147758200907			-1020	80192
MSV074	381030	1144570	5711Y	0	105422795	69606147218200790			177	80803
MSV077	381490	1144467	5806Y	0	102423557	69737145184201391			-1563	78736
MSV078	381376	1144347	6005Y	0	116423350	69918143795201224			-911	78723
MSV079	381460	1144208	6107S	0	131423511	70117144240201347			372	79674
MSV081	382521	1144111	7834C	0	1120425477	70209135887202903			6725	81126









EXPLANATION

FAULTS SHOWN ON GEOLOGIC BASE MAP:

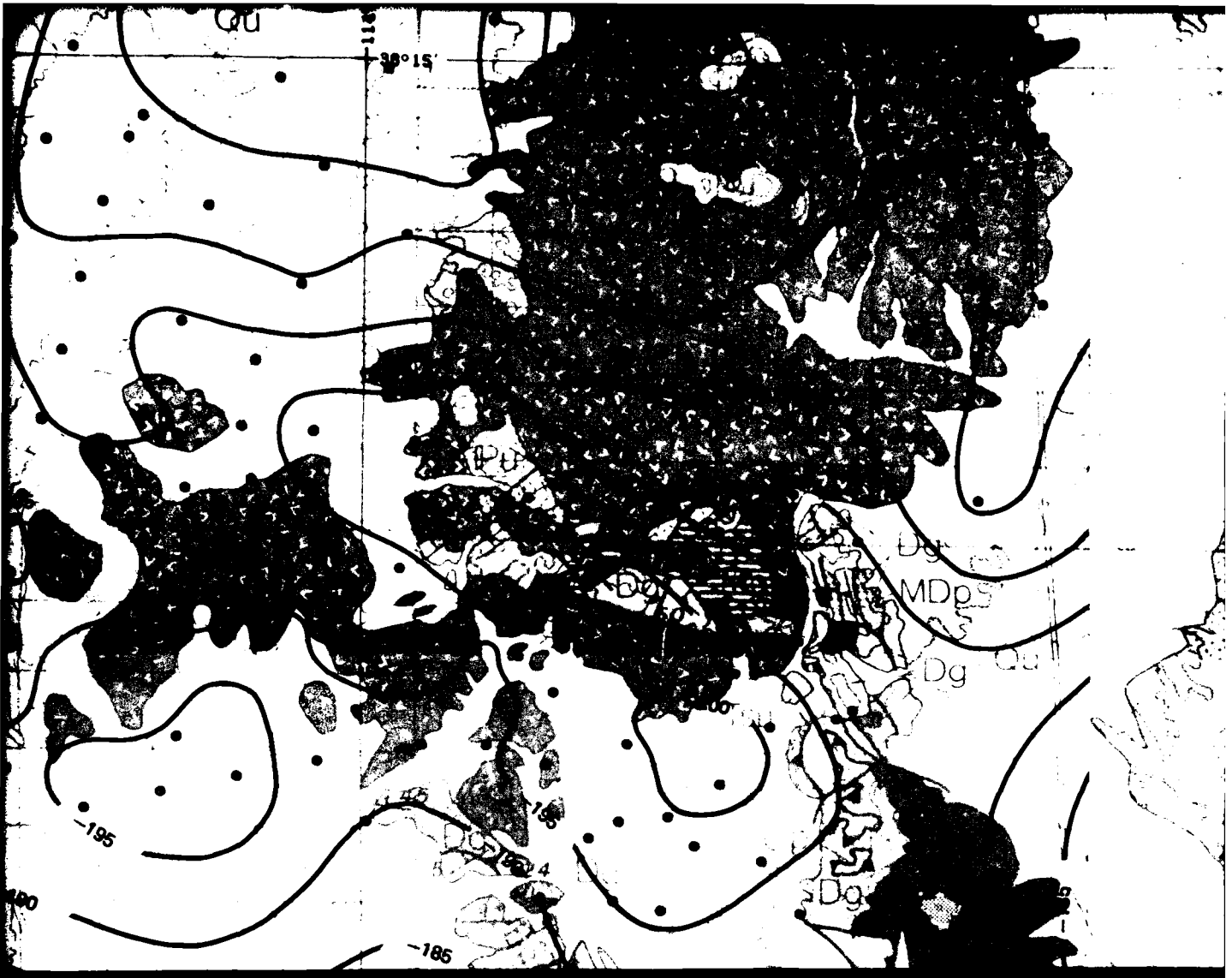
- NORMAL
- - - INFERRED
- CONCEALED
- ALLUVIAL MATERIAL
- ▒ ROCK (ALL PATTERNS)
- GRAVITY STATIONS




LOCATION OF PROFILE

CONTOUR INTERVAL = 5 MILLIGALS

GEOLOGIC BASE MAP: E.L.Howard (197

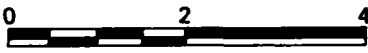


MAP:

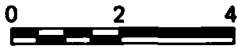


NORTH


SCALE 1: 125,000



STATUTE MILES




KILOMETERS

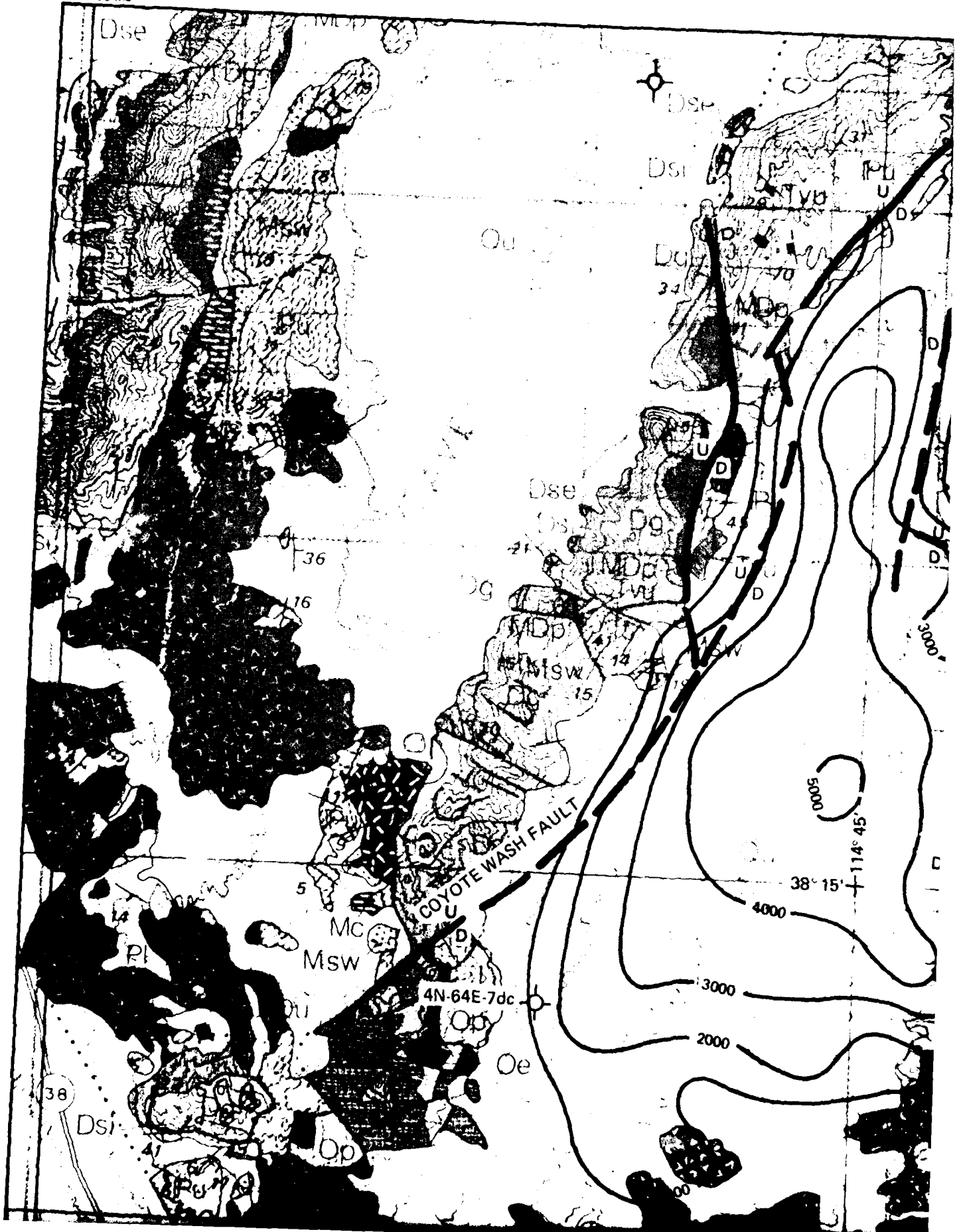


LOCATION OF PROFILE

CONTOUR INTERVAL = 5 MILLIGALS

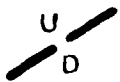
GEOLOGIC BASE MAP: E.L.Howard (1978)

 <p><i>The Earth Technology Corporation</i></p>	<p>MX SITING INVESTIGATION DEPARTMENT OF THE AIR FORCE BMO/AFRCE-MX</p>
<p>COMPLETE BOUGUER ANOMALY CONTOURS MULESHOE VALLEY, NEVADA</p>	
14 SEPT 81	DRAWING 1





EXPLANATION



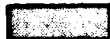
FAULTS INFERRED FROM GRAVITY DATA



FAULTS SHOWN ON GEOLOGIC BASE MAP



ALLUVIAL MATERIAL



ROCK (ALL PATTERNS)

CONTOUR INTERVAL - 1000 FT.

DEPTH CALCULATIONS BASED ON DENSITY CONTRAST OF $0.5g\ cm^{-3}$



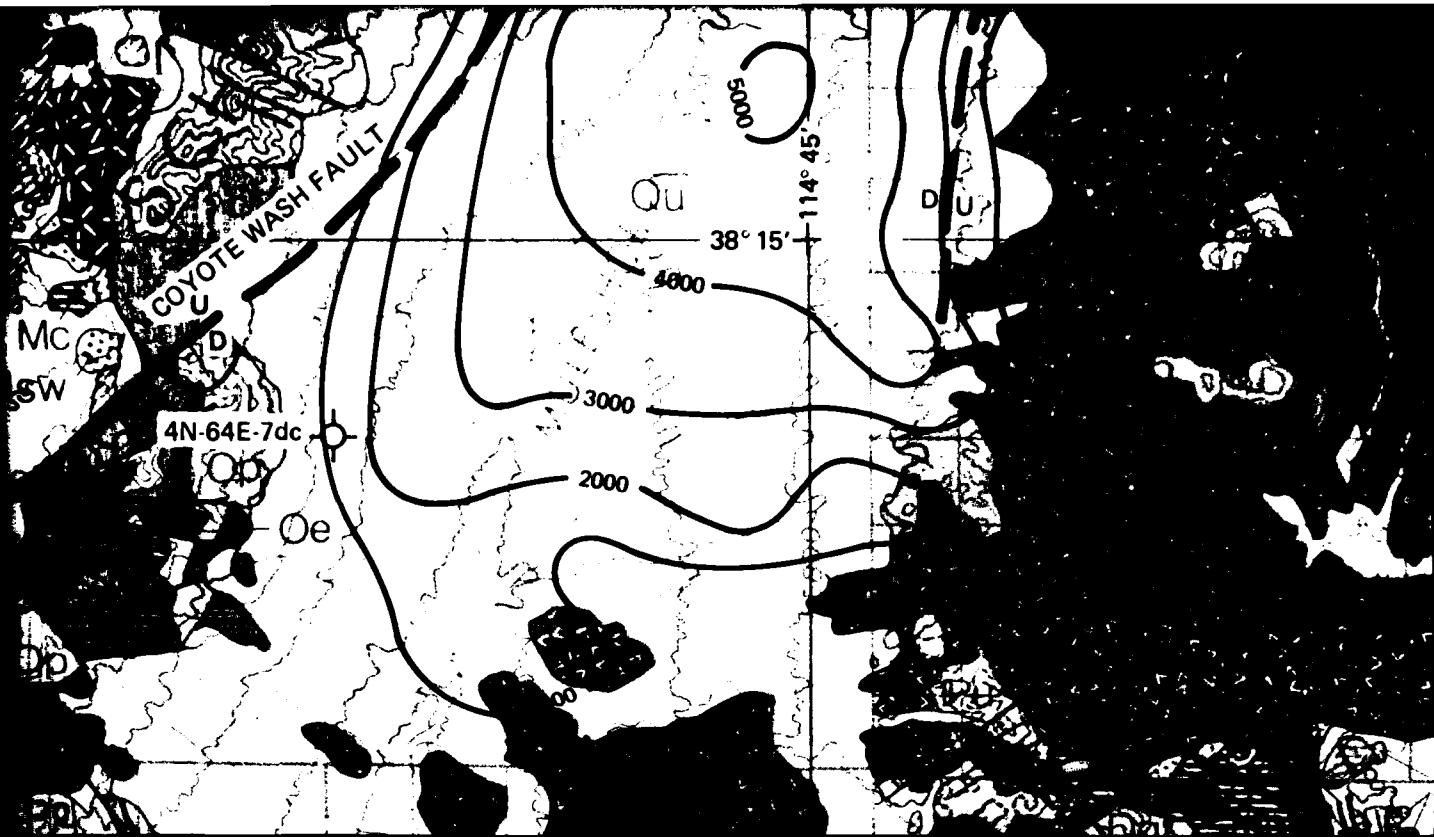
WELL

GEOLOGIC BASE MAP: E. L. Howard (1978)



INTER
MUI

14 SEPT 81



ATION

RRD FROM
TA

N ON
SE MAP

TERIAL

TTERNS)

ERVAL = 1000 FT.

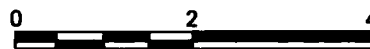
LATIONS BASED ON
TRAST OF $-0.5g\text{ cm}^{-3}$

SE MAP: E. L. Howard (1978)



NORTH

SCALE 1: 125,000



STATUTE MILES



KILOMETERS



MX SITING INVESTIGATION
DEPARTMENT OF THE AIR FORCE
BMO/AFRC-MX

DEPTH TO ROCK
INTERPRETED FROM GRAVITY DATA
MULESHOE VALLEY, NEVADA

14 SEPT 81

DRAWING 2

4

