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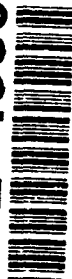
DEVELOPMENT OF MESOSCALE MODEL PERFORMANCE  
EVALUATION PROGRAMS

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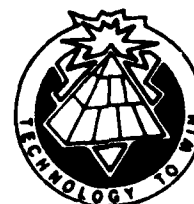
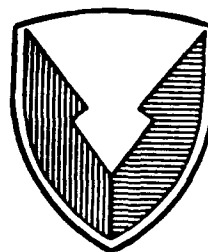
Teizi Henmi  
Roger Vega

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## 1. INTRODUCTION

In 1992, the U.S. Army will sponsor the workshop on mesoscale model technology exchange. The workshop will be handled by the MESOMET Panel. The objectives of the workshop are the following:\*

- to identify state-of-the-art technology of mesoscale modeling;
- to inform the scientific community about the availability of Project WIND (wind in non-uniform domains) data; and
- to serve as a forum for mesoscale modeling technology exchange.

The U.S. Army Atmospheric Sciences Laboratory (ASL) will be assigned to examine the subset of the outputs that are generated by eight different mesoscale models.

In the last several months, several computer programs to examine and display the outputs of the mesoscale model to compare with observations, have been developed by ASL using the output of a mesoscale model HOTMAC (High Order Turbulence Model for Atmospheric Circulation) (Yamada and Bunker, 1989) and Project WIND Phase I, Julian days 178-179, data (Cionco, 1990). Programs developed consist of those displaying horizontal, vertical, and temporal distributions of meteorological parameters--simulated and observed.

This report describes the methods used to examine model output and show examples of graphic display. Detailed results of comparisons between model simulation and observation will be described in the near future. The computer programs are developed by using the FORTRAN language with the National Center for Atmospheric Research (NCAR) GKS-Compatible Graphic System. The programs are on the HP 9000/840 computer at ASL.

## 2. PROJECT WIND DATA AND MODEL SIMULATION

Data used to develop the program was from Project WIND Phase I, covering 24 h from 0900 l.s.t. of day 178. These measurements consisted of upper-air sounding data at five locations every 2 h, with a few missing data, and 21 surface station data. Details of the data set are presented in Cionco (1990).

Model simulation was conducted over terrain as shown in figure 1.\*\* Latitude and longitude of the southwestern corner of the domain are  $39^{\circ} 11' 04.4''$  N and  $122^{\circ} 59' 58''$  W. The terrain heights were represented by grids of 81 by 81 with a unit grid distance of 2.5 km. The highest and lowest grid points are 2477 and 12 m, respectively, above sea level. In the figure, numbers represent the locations of upper-air stations. Meteorological parameters were calculated at every other grid point (40 by 40) for 16 vertical layers, using HOTMAC. The model was initialized at 0900 l.s.t. of day 178 using sounding data taken at station 04, and simulation continued until 0800 l.s.t. of the next day.

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\*J. E. Harris and R. E. Meyers, 1991, Trip report on meeting of MESOVET Panel in Bruges, Belgium, 9-10 May 1991 (unpublished).

\*\*Figures are presented at the end of the text.

In figure 1b the locations of surface observations are marked. Surface data contained windspeed, wind direction, temperature, relative humidity, pressure, incoming solar radiation, and precipitation. During the 24-h period (between 0900 l.s.t. of day 178 and 0900 l.s.t. of day 179), no precipitation occurred. Therefore, precipitation data was not examined. Relative humidity data was converted to dew point by using empirical formulas.

### 3. DATA INTERPOLATION METHODS

#### 3.1 Upper-Air Data

Generally, most mesoscale models are formulated on terrain-following coordinates, and meteorological parameters are calculated at particular heights determined by model design. To compare model output with observation at a desired height, one must interpolate both model outputs and observed data to the height. Model output values computed at a grid point most adjacent to an upper-air sounding station were used for the comparison study.

In HOTMAC model, the following equation is used to define a terrain-following vertical coordinate.

$$z^* = \bar{H} \frac{z - z_g}{H - z_g} \quad (1)$$

where  $z^*$  and  $z$  are the transformed and Cartesian vertical coordinates, respectively;  $z_g$  is ground elevation above sea level;  $\bar{H}$  is the material surface top of the model; and  $H$  is the corresponding height in the coordinate. For simplicity,  $H$  is specified as

$$H = \bar{H} + z_{gmax} \quad (2)$$

where  $z_{gmax}$  is the maximum value of  $z_g$ . From equation (1), height above ground  $H_g$  can be given as

$$H_g = z - z_g = z^* \frac{\bar{H} + z_{gmax} - z_g}{\bar{H}} \quad (3)$$

The author applied both linear interpolation and cubic spline methods to interpolate the values of meteorological parameters (horizontal wind components, temperature, and dew point) at desired height above ground. Few differences were found in the results produced by the two methods. In the linear interpolation method, meteorological parameter  $\varphi$  at height  $z$  can be obtained by using the values at  $z_i$  and  $z_{i+1}$ , where  $z_i < z < z_{i+1}$ , as follows: (figure 2)

$$\varphi(z) = A \cdot \varphi_i + B \cdot \varphi_{i+1} \quad (4)$$

where



$$A = \frac{z_{i+1} - z}{z_{i+1} - z_i} \quad (5)$$

$$B = \frac{z - z_i}{z_{i+1} - z_i} \quad (6)$$

The equation for cubic spline interpolation is expressed as

$$\varphi(z) = A \cdot \varphi_i + B \cdot \varphi_{i+1} + C \cdot \varphi''_i + D \cdot \varphi''_{i+1} \quad (7)$$

where

$$C = \frac{1}{6} \cdot (A^3 - A) \cdot (z_{i+1} - z_i)^2 \quad (8)$$

$$D = \frac{1}{6} \cdot (B^3 - B) \cdot (z_{i+1} - z_i)^2 \quad (9)$$

and

$$\varphi'' = \frac{d^2 \varphi}{dz^2} \quad (10)$$

FORTTRAN programs of cubic spline interpolation described in Press et al. (1989) were used.

### 3.2 Surface Data

To reduce the value of a meteorological parameter at a surface station  $k$ , values at four grid points surrounding the station were used as

$$\varphi_k = \frac{\sum_i \varphi_i \frac{1}{r_i^2}}{\sum_i \frac{1}{r_i^2}} \quad (11)$$

where the subscript  $i$  represents a grid point,  $r_i$  is the distance between station location and grid point  $i$ , and  $\varphi$  is an arbitrary meteorological parameter.

## 4. SURFACE METEOROLOGICAL PARAMETERS

### 4.1 Horizontal Distributions of Gridded Data

Plotting horizontal distributions of gridded data of model output is important to examine distribution patterns of meteorological parameters over the model

domain and to intercompare between models. Scalar variables such as temperature and sensible heat flux are plotted by using contour-line drawing routines in the NCAR GKS-Compatible Graphic System. Horizontal wind vectors are plotted by using vector drawing routines. In the following figures, (a) and (b) represent daytime and nighttime conditions for 1500 l.s.t. of day 178 and 0300 l.s.t. of day 179, respectively.

Examples of contour-line plotting of scalar variables are shown in figures 3 and 4. Figures 3a and 3b show air temperature distributions at 10-m levels above ground. Contour lines are drawn with 5 °C intervals. Terrain contour lines are also drawn using thin lines. Figure 4 shows the distribution of sensible heat flux ( $W/m^2$ ). Sensible heat fluxes are upward (positive) throughout the model domain during the day, and downward (negative) during the night. Upward fluxes are contoured by solid lines and downward fluxes by broken lines.

Figure 1b shows that surface observation stations were in the center of the model domain. For easier comparison, the right side of figures 5a and 5b show the horizontal wind vectors computed by the model for the center area only; the left side shows the observed wind vectors. Upslope wind conditions during the day and downslope wind conditions during the night were well simulated and in good agreement with observations.

#### 4.2 Comparisons of Simulation with Surface Observation

The computer program developed for comparing simulation with surface observation takes the following steps:

- determines the locations of surface station in grid coordinate.
- creates time series arrays of meteorological parameters for points representing surface stations, using data at four grid points surrounding the surface stations.

The program is designed to plot both simulation and observation for any surface station desired.

Since the model output file contained hourly data at grid points, time series plotting of both simulated and observed data were also made hourly. Wind direction and windspeed (meters per second), temperatures (degrees Celsius) at 2- and 10-m levels, dew point (degrees Celsius), and downward shortwave radiation (watts per square meter) were plotted. As examples, plottings for stations S1 and C1 are shown, respectively, in figures 6 and 7, with thin lines representing simulation and thick lines representing observation. These figures show that shortwave radiation observation data and temperatures at either level were not available at some stations.

### 5. UPPER-AIR METEOROLOGICAL PARAMETERS

#### 5.1 Vertical Distributions at the Locations of Sounding Stations

Upper-air data were observed at five stations during the 24-h period. Data were taken at each station every 2 h, with some exceptions. The program developed to plot vertical distributions of meteorological parameters extracts data at grid

points closest to the stations and creates arrays of data for each station. The program was arranged so plotting could be made for a desired station and time.

Figure 8 shows the vertical distributions of wind direction and windspeed, x and y components of wind vector, temperature, and dew point for station 01 and 1300 l.s.t. of day 178. Thin and thick lines represent, respectively, simulation and observation. Figure 9 shows station 04 at 0100 of day 179.

## 5.2 Time-Series at Different Levels

For comparison between different models, it is convenient to have meteorological parameters at the same heights. Participants in the workshop of mesoscale model technology exchange are asked to produce time series of meteorological parameters at the following: standard heights 2, 10, 50, 100, 200, 400, 800 m above the ground and standard pressure levels of 850, 700, 500, and 300 mbar. The HOTMAC model computed variables at the following 16 levels of terrain-following coordinates: 0, 2, 6, 10, 14, 28, 114, 281, 530, 861, 1273, 1767, 2342, 3000, 3729, 4559. Thus, interpolation of variables from model height to standard height was necessary. As described in section 3, a cubic spline interpolation method was used for interpolation from model heights to standard heights. The linear interpolation method was used to interpolate to standard pressure level. In the HOTMAC model, pressure is a diagnostic variable.

The program extracts parameters for grid points most adjacent to upper-air stations and calculates the values at standard heights or pressure levels. Observed data were also interpolated by using cubic spline or linear interpolation methods. Time series arrays of meteorological parameters for different station locations and standard levels were generated; therefore, time series of plotting could be easily made for different stations and levels. Wind direction and windspeed, temperature, and dew point were plotted.

Figures 10a and 10b show the time series of wind direction and windspeed at seven different heights for station 03. Continuous lines are used for simulation and asterisks (\*) are used to plot the values of observation. Figures 11a and 11b show examples of the time series of temperature and dew point.

Time series for three different pressure levels (500, 700, and 850 mbar) are given in figures 12 and 13. Neither observation nor simulation was available at the 300-mbar level. Plots for station 03 were used for these figures.

## 6. EVALUATION OF MODEL PERFORMANCE

Visual comparisons of simulation with observations (as shown in the previous sections) are useful. However, model performance evaluated quantitatively is also desirable since it enables us to compare objectively one model to another model and to gain insight into the sources of error. For the present simulation, continuous data of wind direction, windspeed, temperature, and humidity were available throughout the 24-h period at 21 surface stations; and 4 or 5 upper-air sounding data were available every 2 h. These data were used to perform the following statistical evaluations.

### 6.1 Surface Data

Mean, standard deviation, root mean square errors (rmse), unbiased rmse, and agreement measure were calculated hourly. Willmott (1981, 1982) and Willmott et al. (1985) recommend the use of the above statistical parameters to quantitatively evaluate model performance. The following equations are definitions of these parameters:

a. mean

$$\bar{\varphi} = \frac{\sum_k \varphi(k)}{N} \quad (12)$$

where  $\varphi(k)$  is meteorological parameters at kth station, and N is the number of stations. Means for both simulation and observation were calculated. In a good agreement case, means for both should have similar values.

b. rmse (E) and unbiased rmse ( $E_{ub}$ )

$$E = \left\{ \frac{\sum_k [\varphi_m(k) - \varphi_o(k)]^2}{N} \right\}^{1/2} \quad (13)$$

$$E_{ub} = \left\{ \sum_k \frac{[(\varphi_m(k) - \bar{\varphi}_m) - (\varphi_o(k) - \bar{\varphi}_o)]^2}{N} \right\}^{1/2} \quad (14)$$

If there is a perfect agreement, E and  $E_{ub}$  are zero.

c. standard deviation

$$\sigma = \left\{ \sum_k \frac{[\varphi(k) - \bar{\varphi}]^2}{N} \right\}^{1/2} \quad (15)$$

In a good agreement case, the standard deviation for both simulation and observation should have similar values.

d. agreement measure

$$A = 1 - \frac{\sum_k (\varphi_m(k) - \varphi_o(k))^2}{\sum_k (|\varphi_m(k) - \bar{\varphi}_m| + |\varphi_o(k) - \bar{\varphi}_o|)^2} \quad (16)$$

This dimensionless index has a theoretical range of 1.0 (for perfect agreement) to 0.0 (for no agreement).

These statistical parameters were calculated hourly for wind, temperature, and dew point. Figure 14 shows the results for wind. Mean wind direction was calculated from the means of horizontal wind components. Thin lines represent simulation and thick lines represent observation in the top three portions of figure 14. In the rmse plotting,  $E_b$  was drawn using a thin line and  $E_{ub}$  by using a thick line. Agreement measure of windspeed was slightly lower during the day-time than during the night, as mean windspeed showed greater discrepancies during the day than during the night.

Figure 15 is a similar figure for temperature at the 10-m level. There is a good agreement between simulation and observation during the day, as can be seen in mean temperature and agreement measure. Agreement becomes poor during the night, as the standard deviation of observed temperature is much greater than that of simulation during the night, probably resulting from the model's incapability of representing localized effects. The rmse shows also that agreement between simulation and observation becomes poor during the night.

## 6.2 Upper-Air Data

So far, for upper-air data, the following two statistical parameters were calculated at different levels as a function of time to evaluate model performance. Different statistical parameters may need further consideration.

- a. average difference between observation and simulation

$$\overline{\delta\varphi(t)} = \frac{\sum_k |\varphi_{k,o}(t) - \varphi_{k,m}(t)|}{N} \quad (17)$$

where  $\varphi$  represents meteorological parameters, subscripts o and m are for observation and simulation, and N is the number of observations.  $\delta\varphi$  of horizontal wind components, speed, temperature, dew point, and pressure were calculated at different levels every 2 h when observation of upper-air was available. Figures 16 and 17 are for the 10- and 1000-m levels, respectively. At the 10-m level, the difference of temperature becomes greater during the night, as has been mentioned in section 6.1. On the other hand, temperature difference at the 1000-m level did not show great difference at night, probably because temperature at the 1000-m level was influenced very little by surface heating and cooling. The average difference of the x component of wind at the 1000-m level grew after several hours of simulation. In this simulation, the model was initialized at 0900 l.s.t. using upper-air sounding data, and no adjustment was made during simulation.

b. correlation coefficient of variance

$$r(t) = \frac{\sum_k (\delta\varphi'_{k,o} \cdot \delta\varphi'_{k,m})}{(\sum_k \delta\varphi_{k,o}^2)^{1/2} \cdot (\sum_k \delta\varphi_{k,m}^2)^{1/2}} \quad (18)$$

where

$$\delta\varphi_o(t + \Delta t) = \varphi_o(t + \Delta t) - \varphi_o(t) \quad (19)$$

$$\delta\varphi_m(t + \Delta t) = \varphi_m(t + \Delta t) - \varphi_m(t) \quad (20)$$

$$\delta\varphi_o' = \delta\varphi_o - \overline{\delta\varphi_o} \quad (21)$$

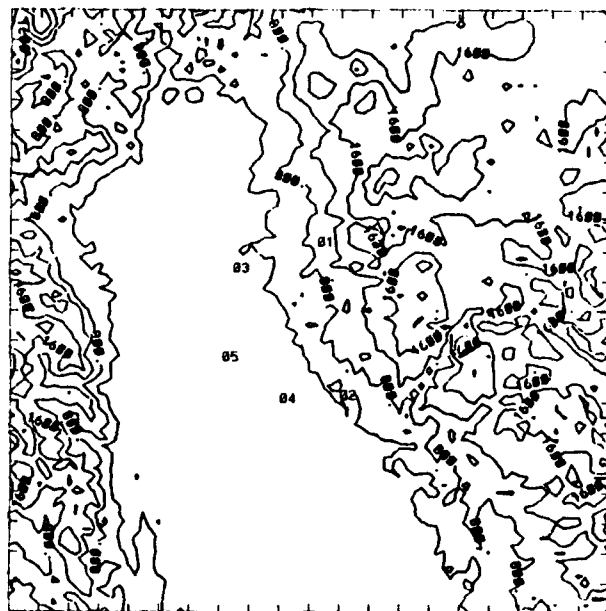
$$\delta\varphi_m' = \delta\varphi_m - \overline{\delta\varphi_m} \quad (22)$$

Here the overline denotes an average of an entire simulation period. The use of the correlation coefficient of variance was suggested by the MESOMET panel.

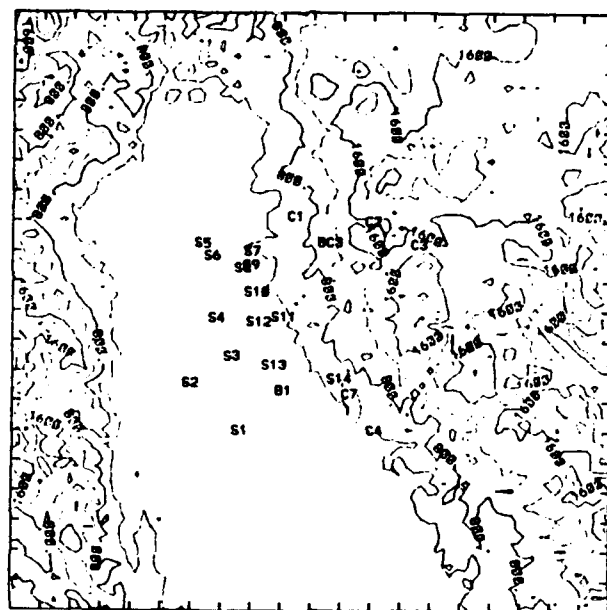
The coefficient  $r(t)$  was calculated for meteorological parameters including horizontal wind components, windspeed, temperature, dew point, and pressure at different levels. Figures 18 and 19 show the 10- and 1000-m levels, respectively. The values of  $r(t)$  vary considerably for all the meteorological parameters. Ideas on model performance are difficult to obtain from these figures. The correlation coefficient must be done carefully.

## 7. CONCLUDING REMARKS

This report describes and illustrates computer programs developed for a comparison between model simulation and observation by using Project WIND Phase I day 178 data and the HOTMAC model output. Temporal and spatial comparisons of simulation with observation can be made by using the program developed. Statistical parameters described in the report will become meaningful when different model simulations are compared with observations.



(a) Locations of upper-air sounding stations, marked with numerical numbers.



(b) Locations of surface stations.

Figure 1. Project WIND terrain map, with 400 m height contours.

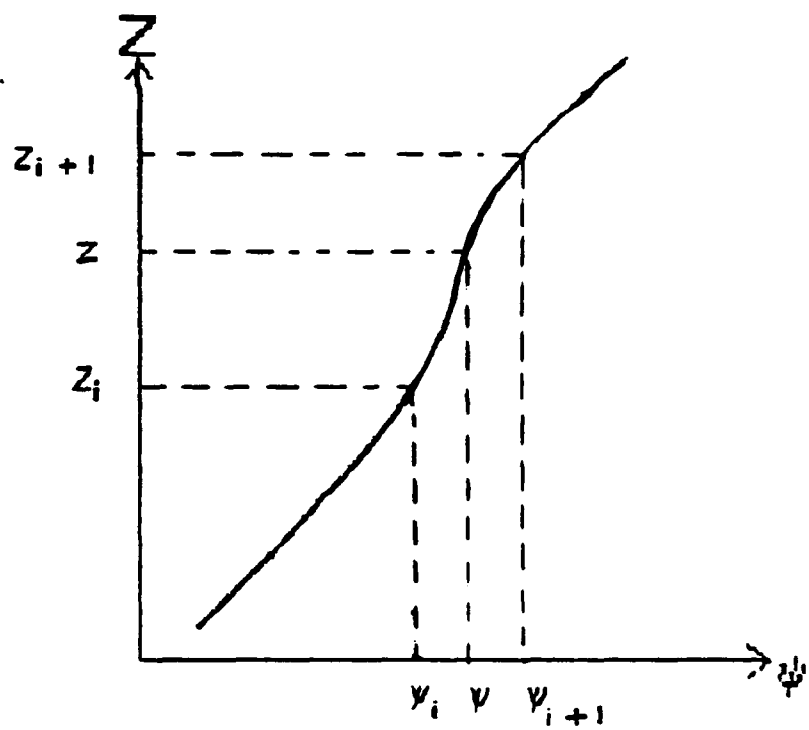
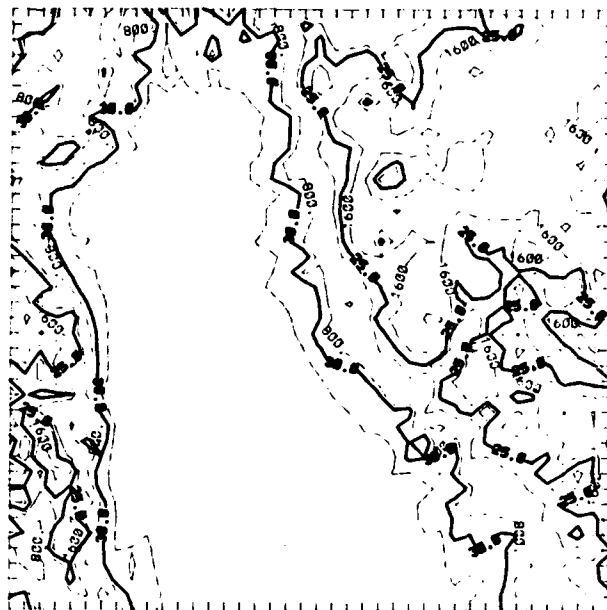
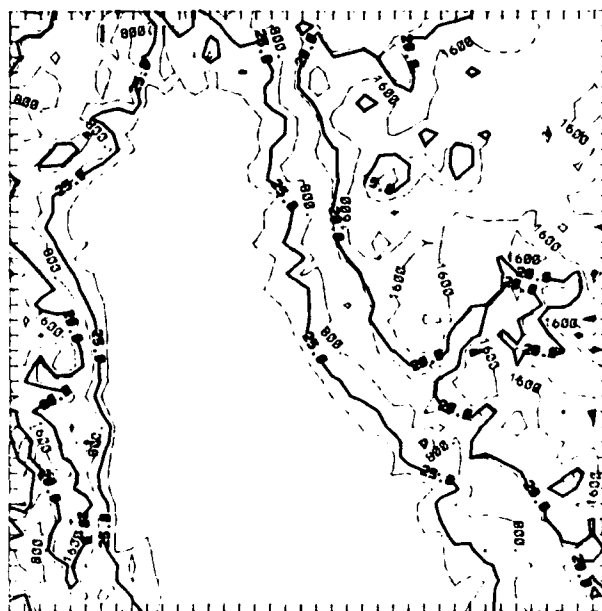


Figure 2. Interpolation method of vertical data.



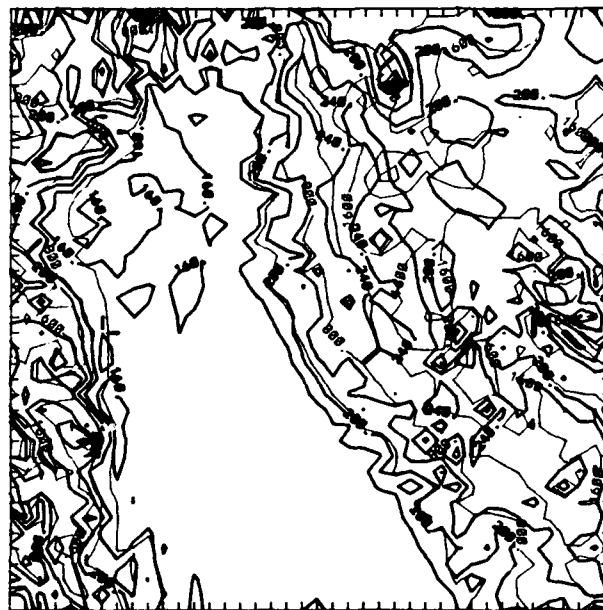


(a) 1500 l.s.t., day 178

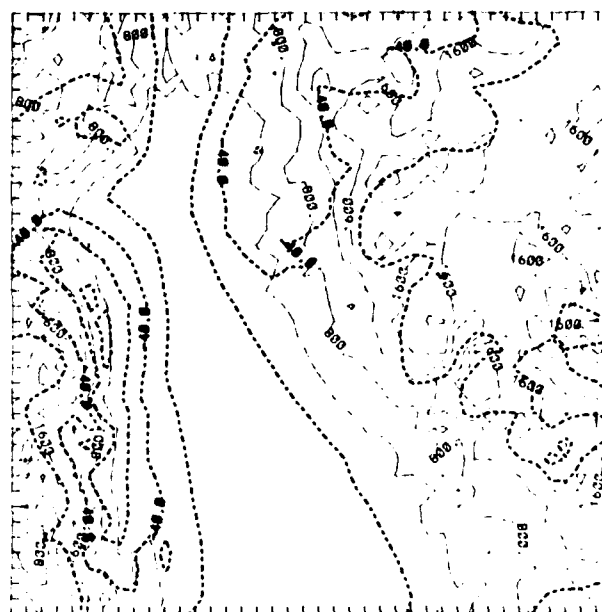


(b) 0300 l.s.t., day 179

Figure 3. Temperature distribution with 5 °C contours.

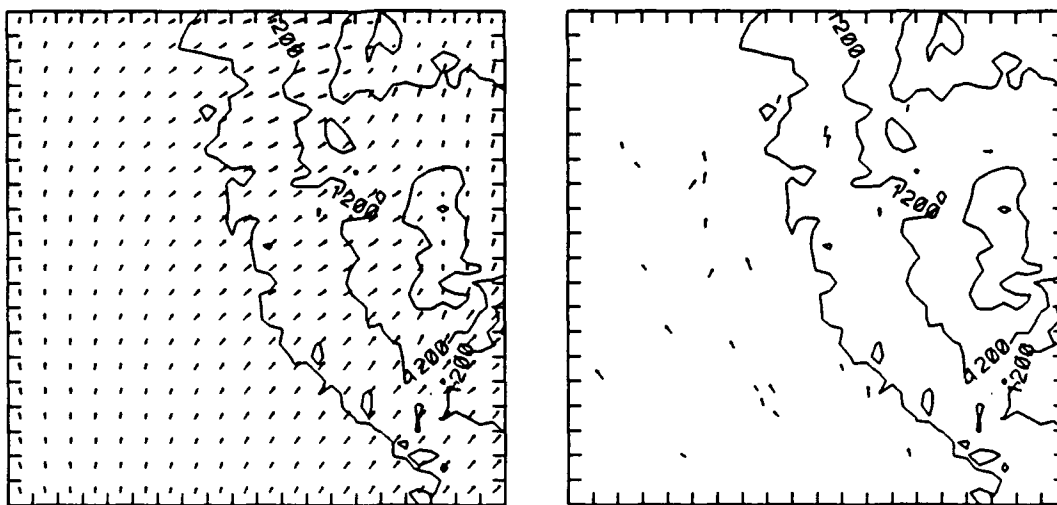


(a) 1500 l.s.t., day 178

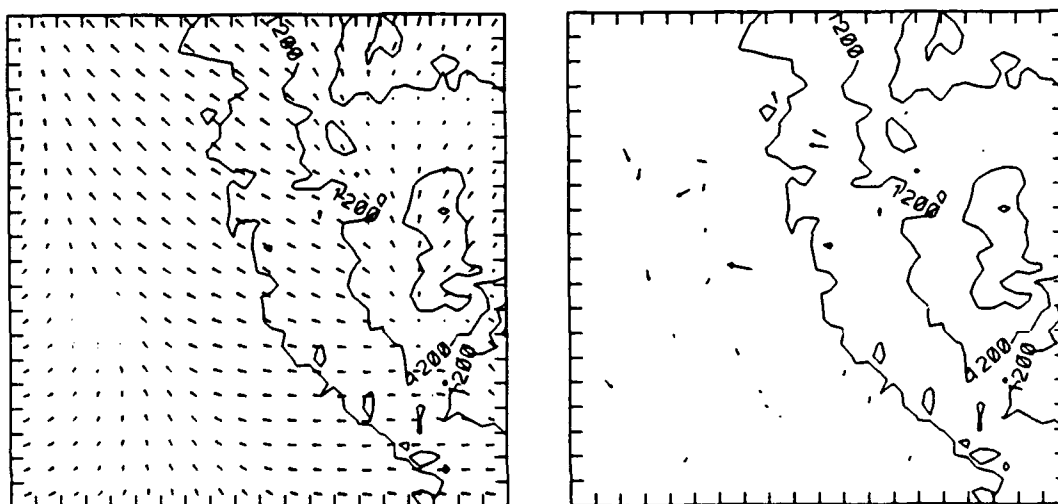


(b) 0300 l.s.t., day 179

Figure 4. Sensible heat flux distribution over terrain data--solid lines for upward flux and broken lines for downward flux.



(a) 1500 l.s.t., day 178



(b) 0300 l.s.t., day 179

Figure 5. Horizontal wind vector distributions--right side for simulation and left side for observation--maximum arrow = 7.35 m/s.

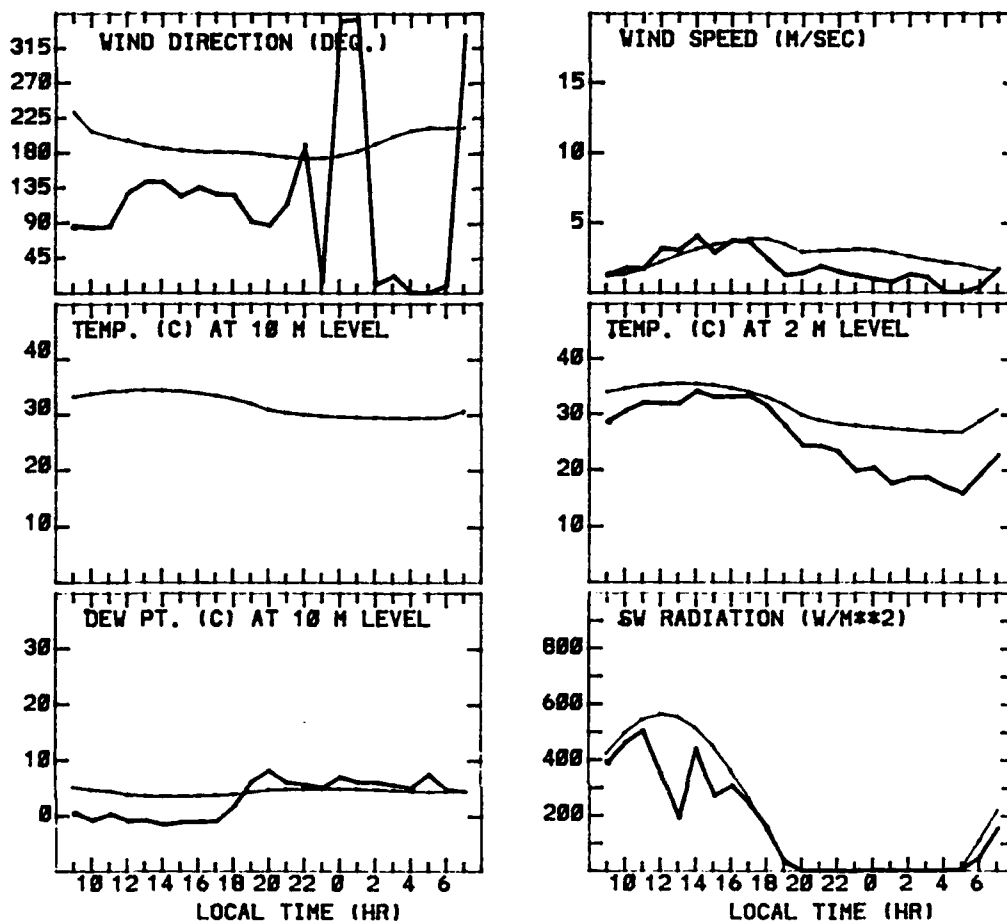


Figure 6. Comparison of surface data--thin lines for simulation and thick lines for observation, wind direction, and windspeed--temperatures at 2- and 10-m levels, dew point, and downward shortwave radiation as a function of time for station S1.

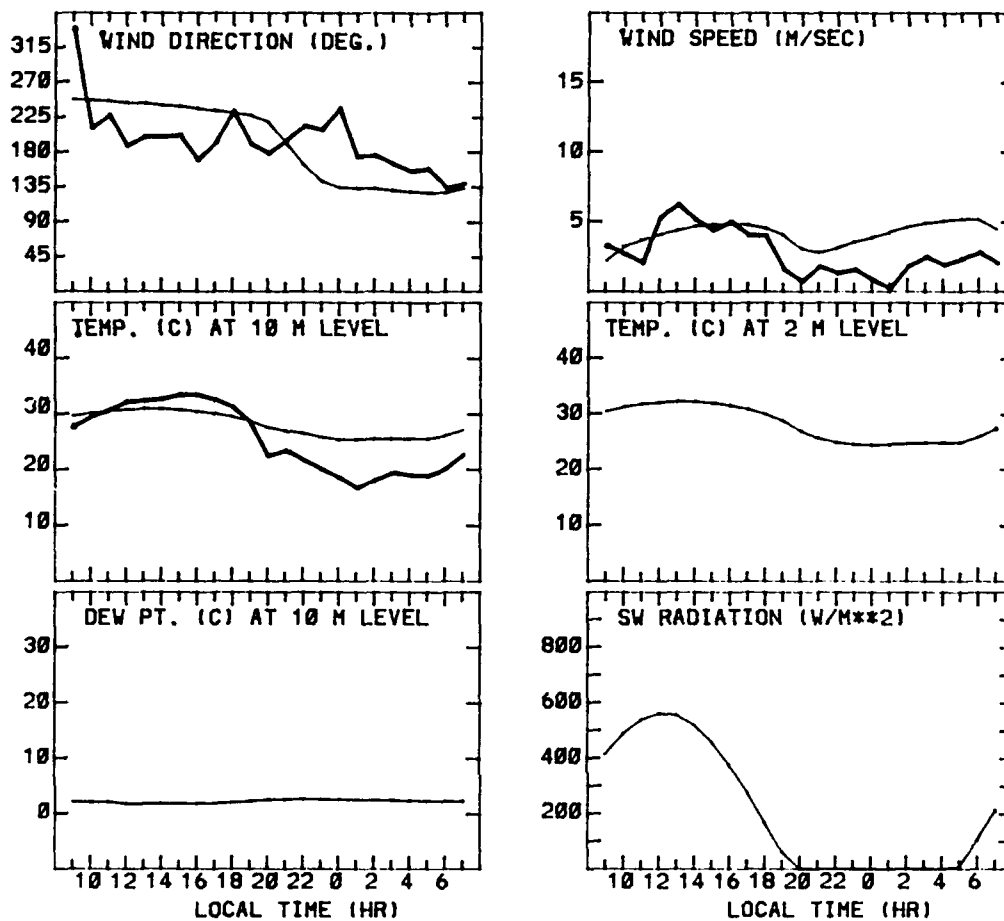


Figure 7. Comparison of surface data--thin lines for simulation and thick lines for observation, wind direction, and windspeed--temperatures at 2- and 10-m levels, dew point, and downward shortwave radiation as a function of time for station C1.

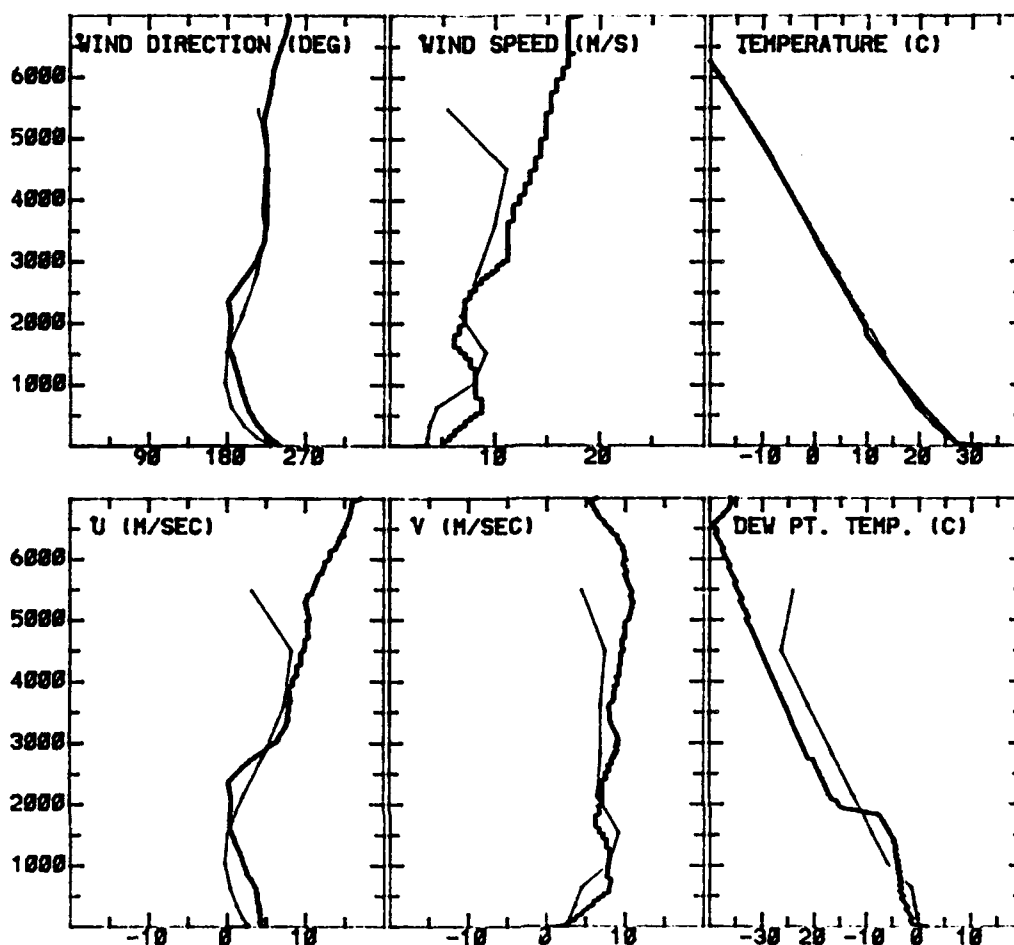


Figure 8. Vertical distributions of wind direction and windspeed, and horizontal components of wind vectors, temperature, and dew point--thin lines for simulation and thick lines for observation. Station = 0.1, 1300 l.s.t., day 178.

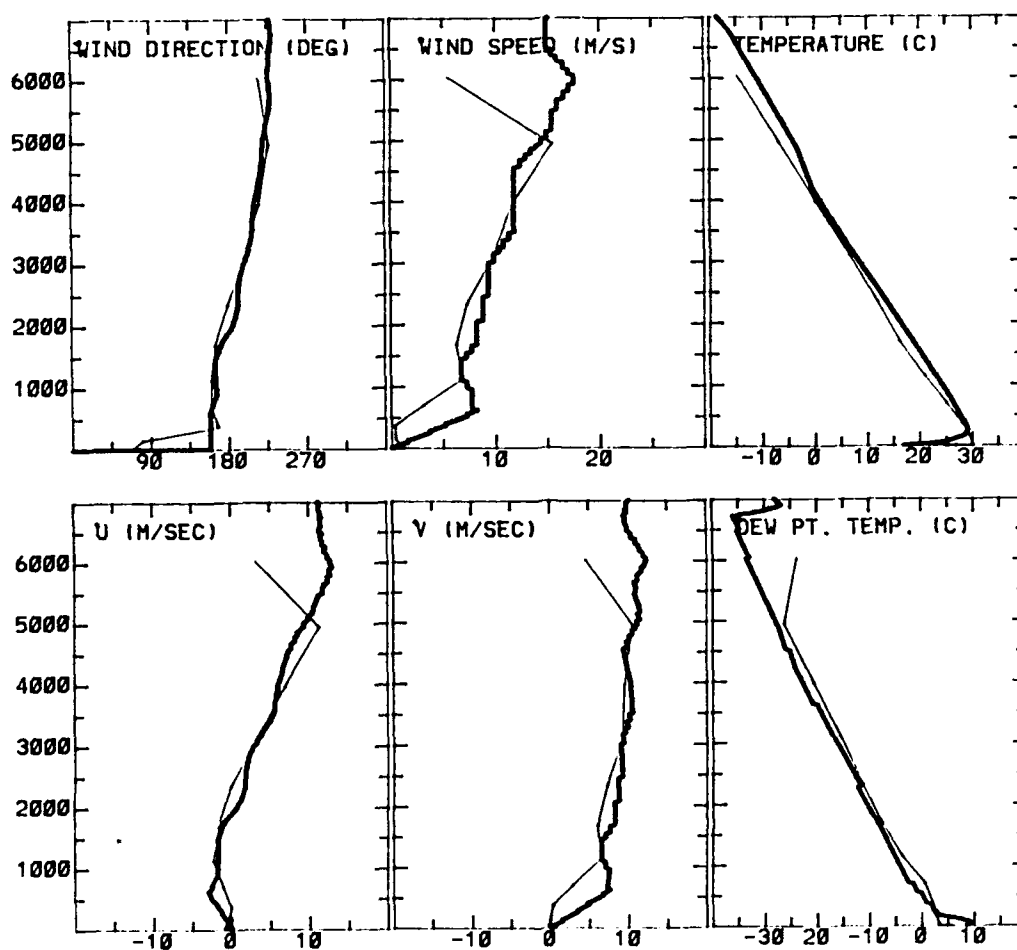


Figure 9. Vertical distributions of wind direction and windspeed, and horizontal components of wind vectors, temperature, and dew point--thin lines for simulation and thick lines for observation. Station = 04, 0100 l.s.t., day 179.

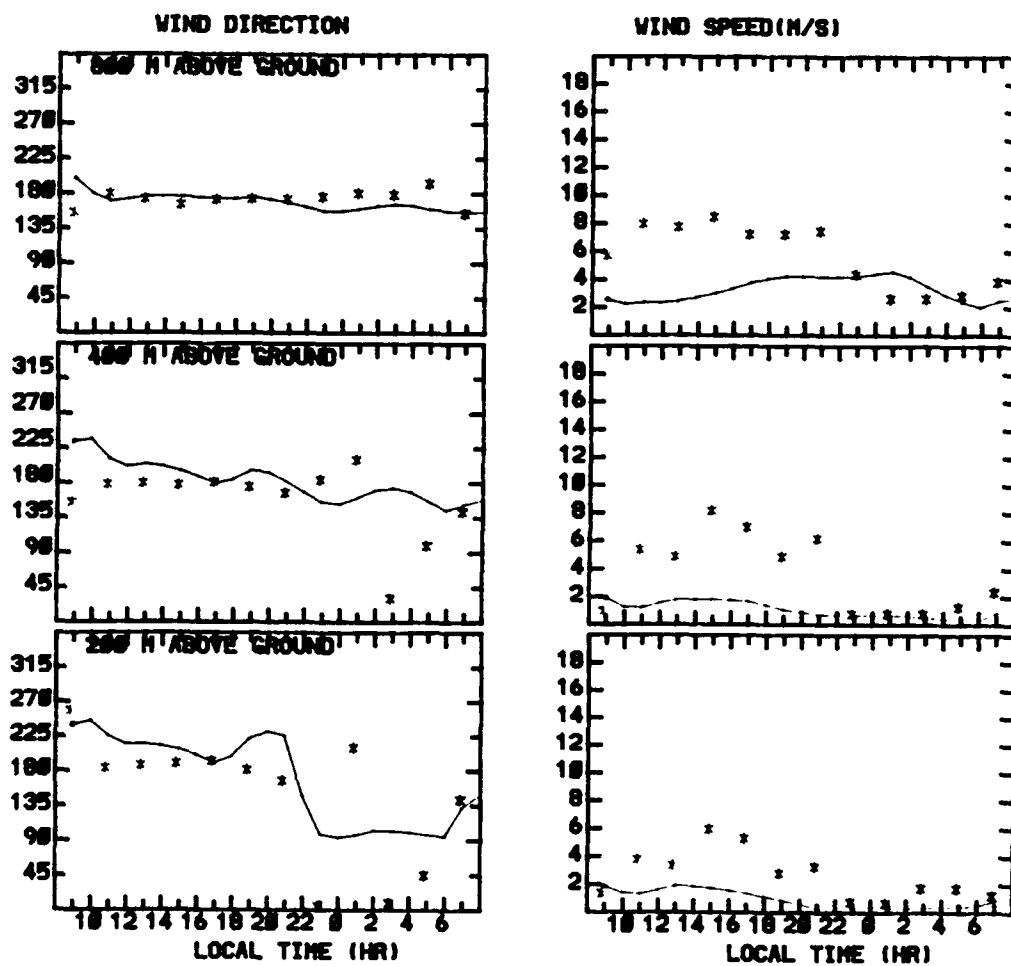


Figure 10a. Time series of wind direction and windspeed at 800-, 400-, and 200-m levels for station = 03.



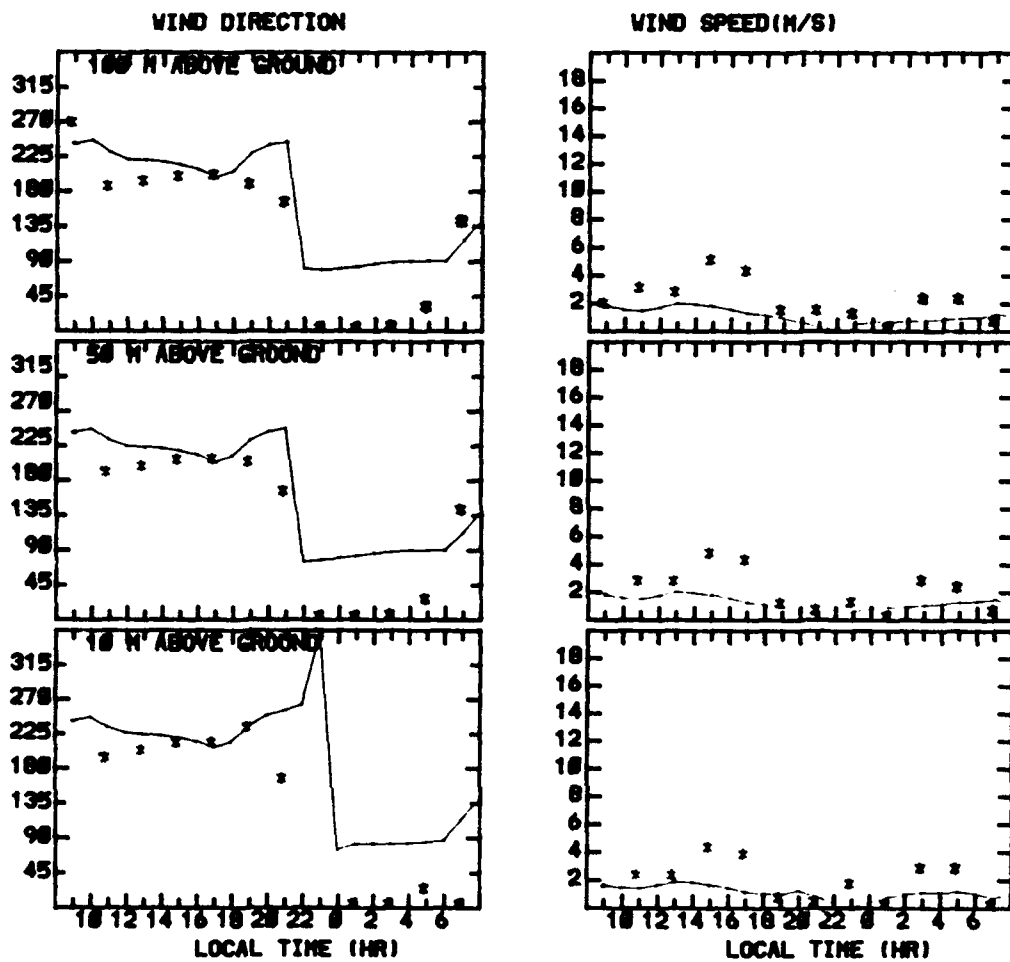


Figure 10b. Time series of wind direction and windspeed at 100-, 50-, and 10-m levels for station = 03.

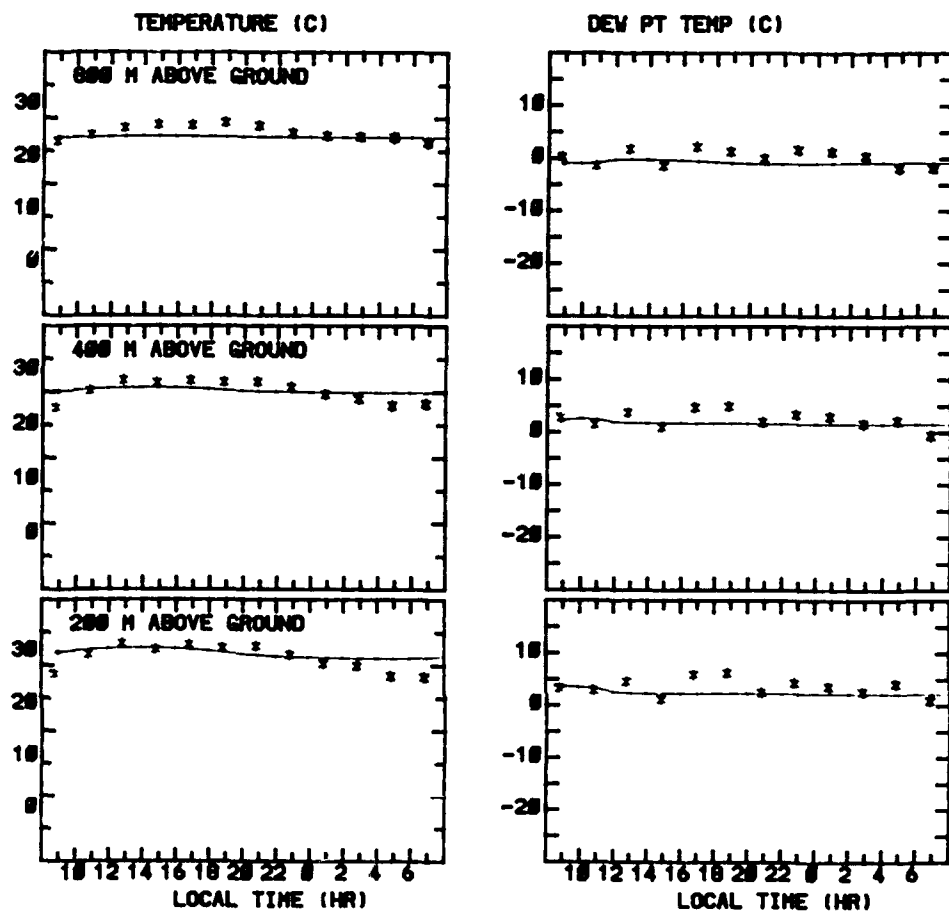


Figure 11a. Time series of temperature and dew point at 800-, 400-, and 200-m levels for station = 03.

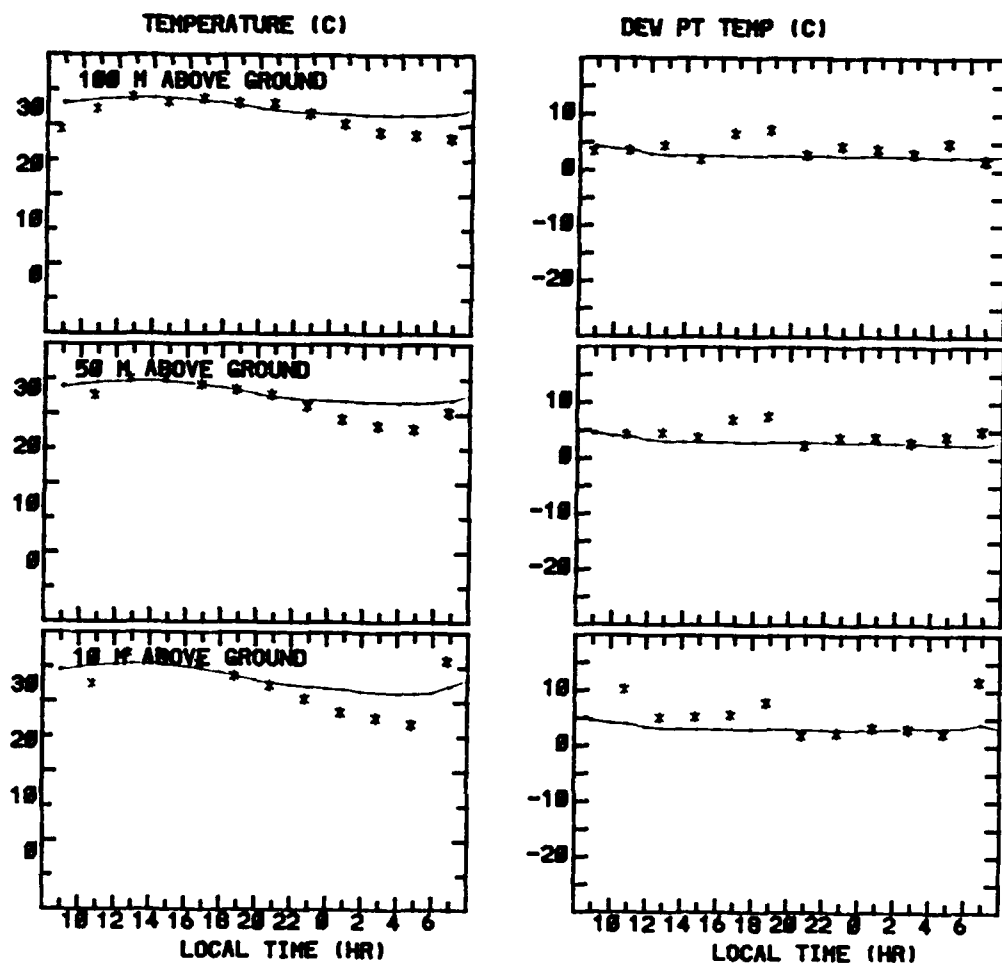


Figure 11b. Time series of temperature and dew point at 100-, 50-, and 10-m levels for station = 03.

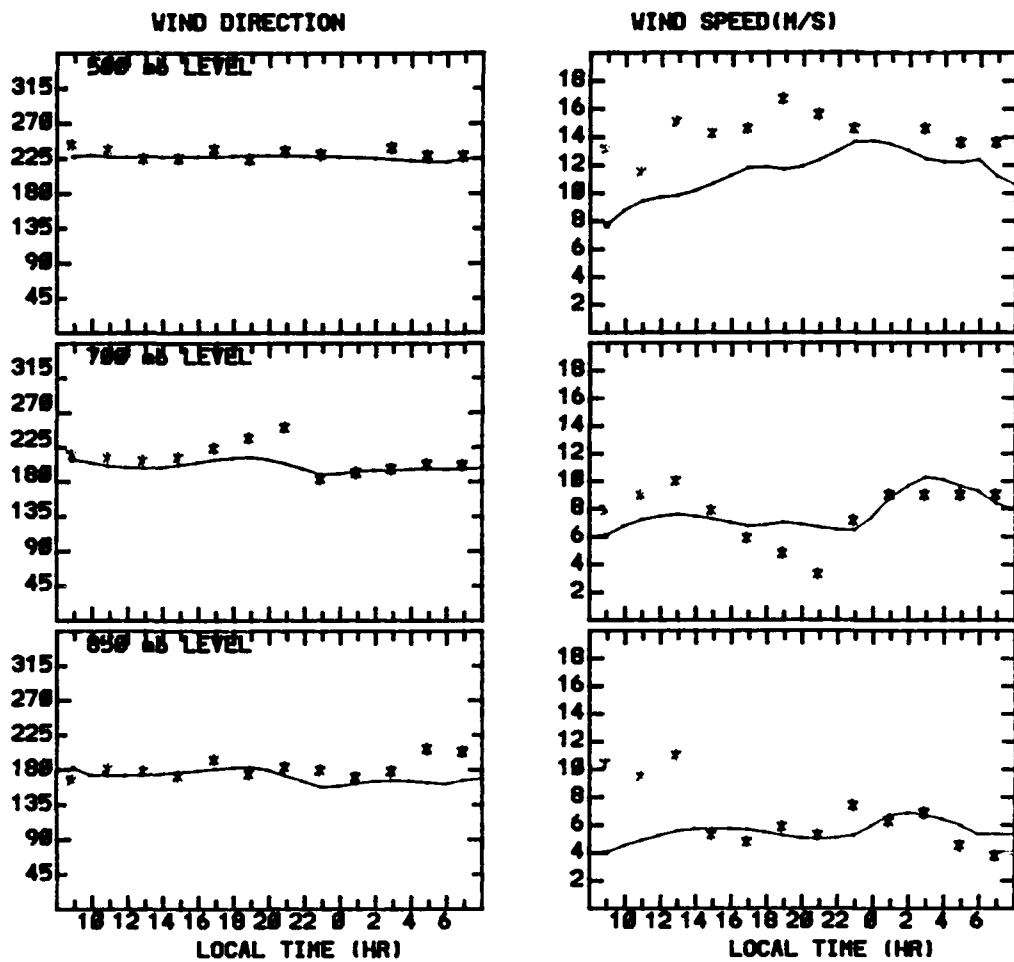


Figure 12. Time series of wind direction and windspeed at 500-, 700-, and 850-mbar levels for station = 03.

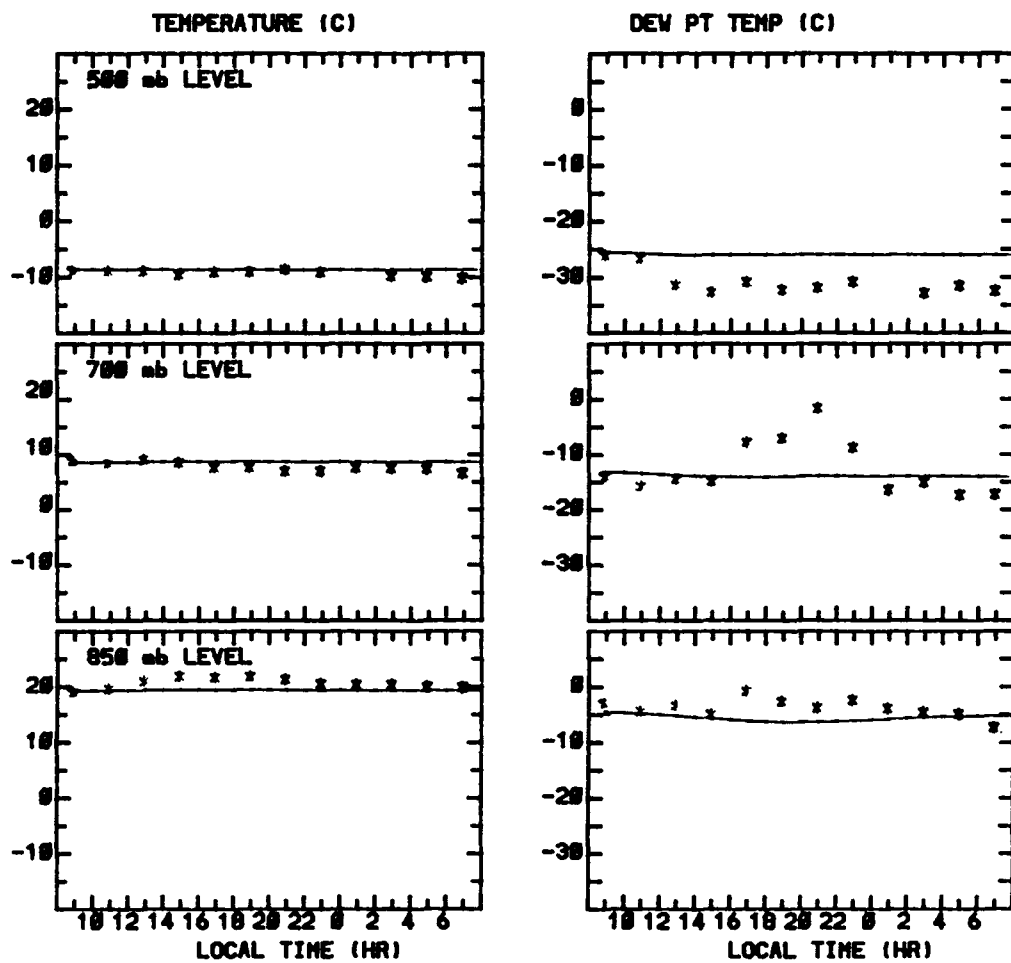


Figure 13. Time series of temperature and dew point at 500-, 700-, and 850-mbar levels for station = 03.

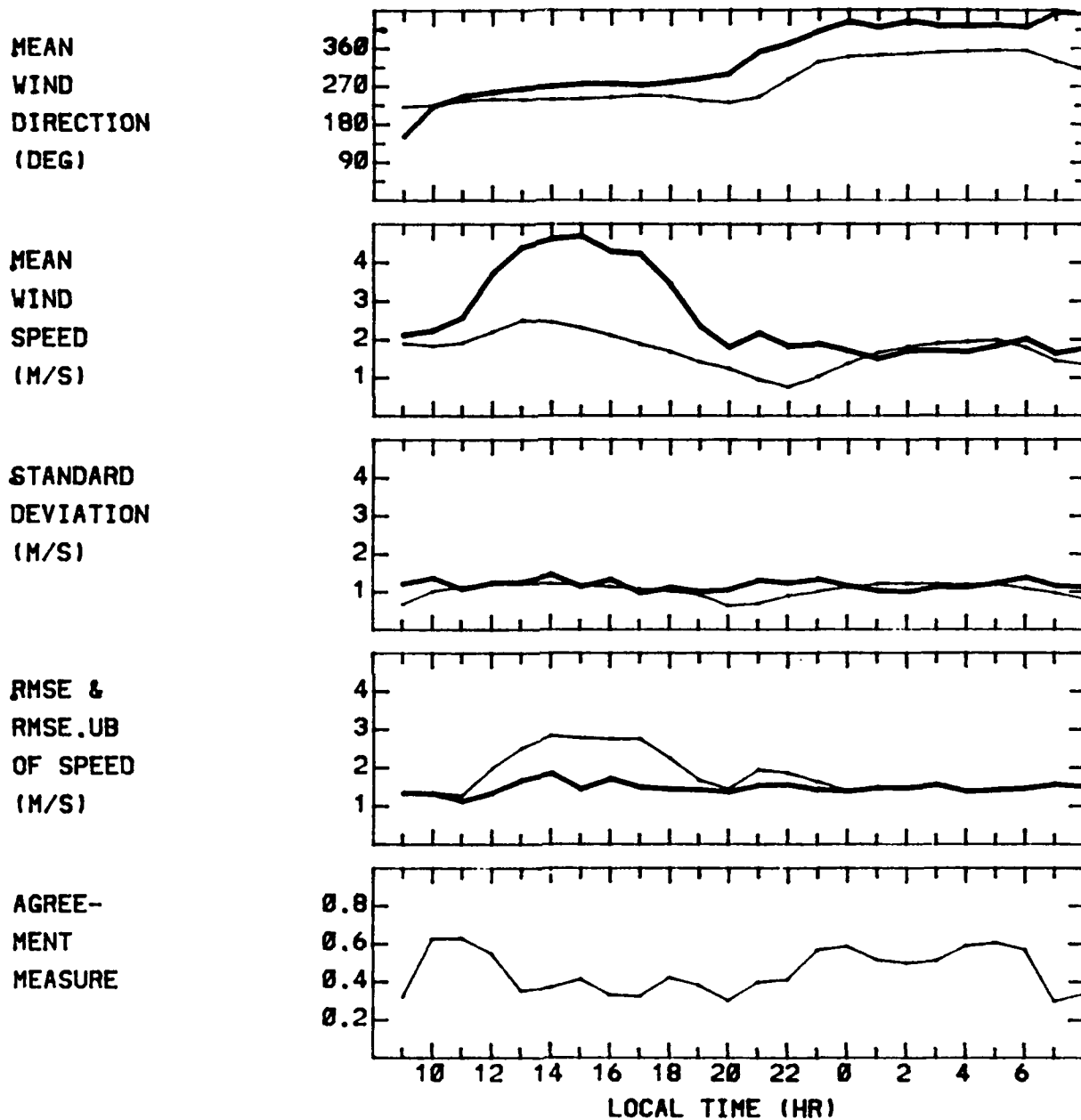
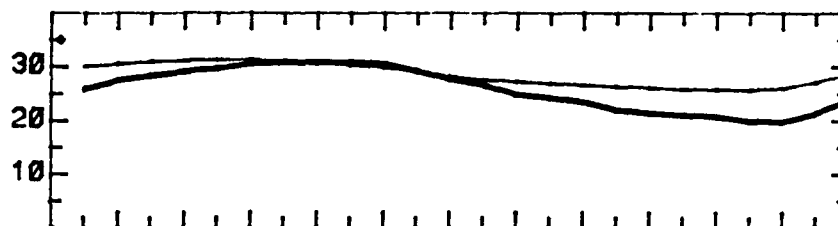


Figure 14. Time series of statistical parameters for surface wind data. Mean wind direction and windspeed, standard deviation, rmse, and agreement measure. In the top three portions of the figure, the thin lines are for simulation and the thick lines are for observation. In the fourth portion (rmse), the thin line is for  $E_b$  and the thick is for  $E_{ub}$ .

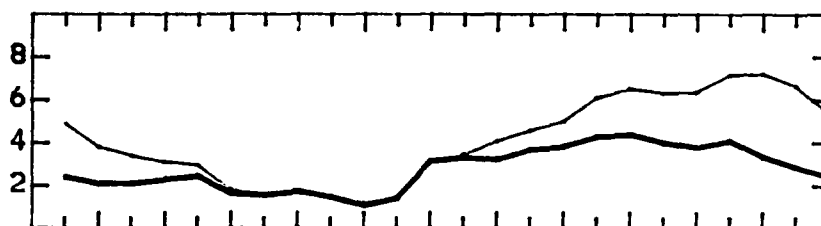
MEAN  
TEMPERATURE  
AT 10 M  
(C)



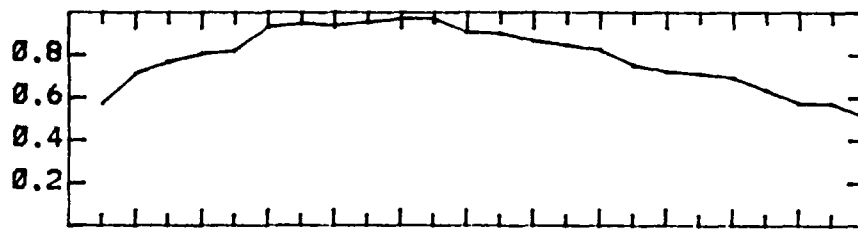
S.D. OF  
TEMPERATURE (C)  
AT 10 M



RMSE &  
RMSE.UB  
OF TEMP. (C)  
AT 10 M



AGREE-  
MENT  
MEASURE



LOCAL TIME (HR)

Figure 15. Time series of statistical parameters for surface (10-m level) temperature. In the top two portions of the figure, the thin lines are for simulation and the thick lines are for observation. In the third portion, the thin line is for  $E_b$  and the thick line is for  $E_{ub}$ .

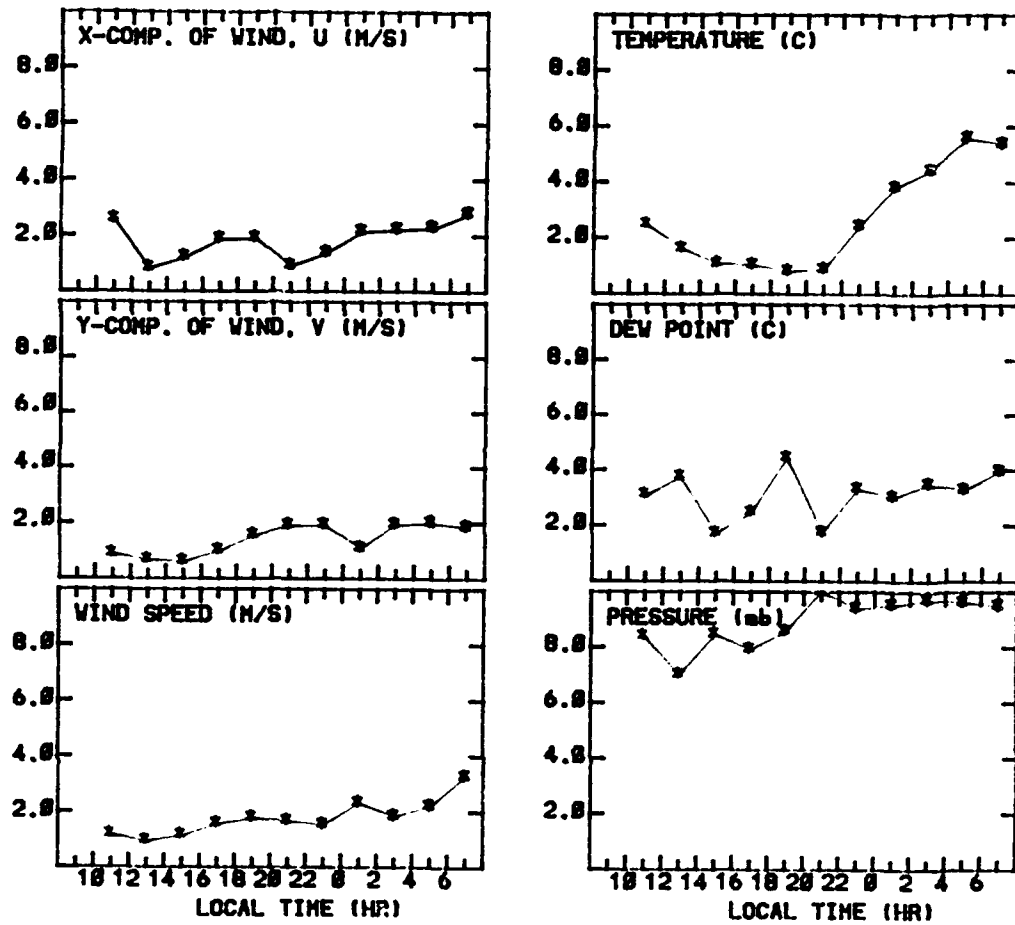


Figure 16. Time series of average differences between observation and simulation for horizontal wind vector components, speed, temperature, dew point, and pressure at the 10-m level.



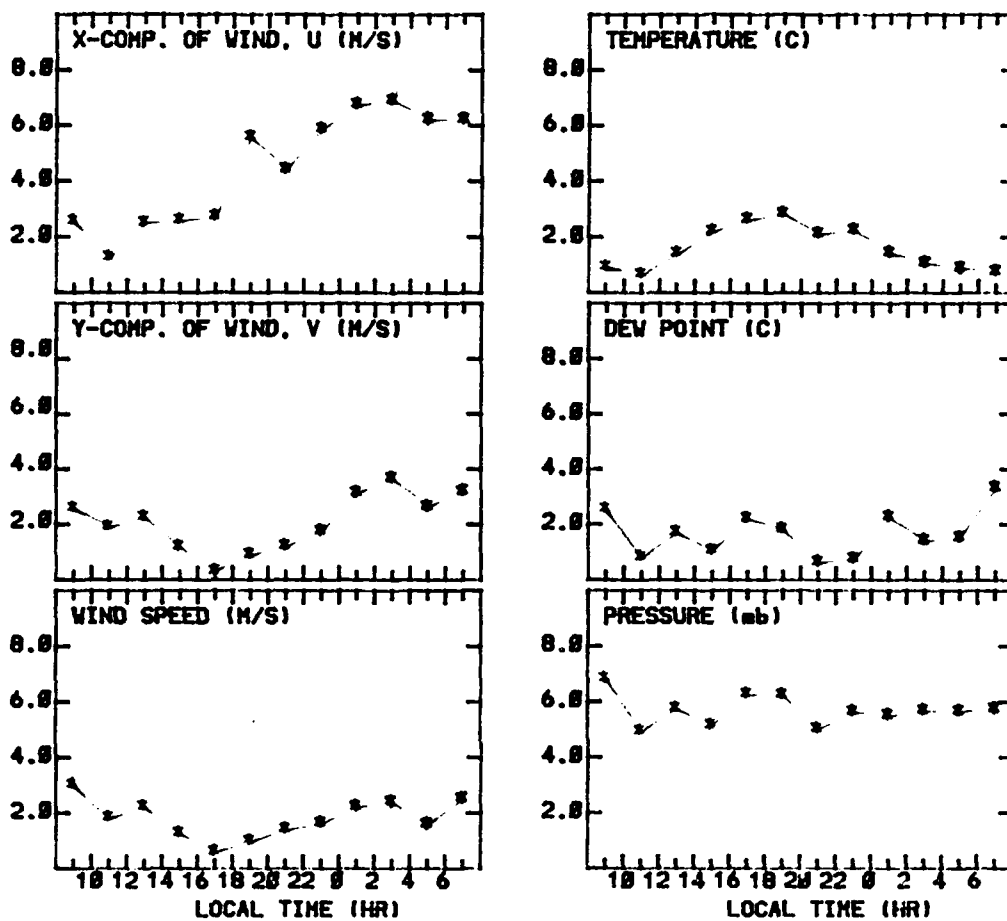


Figure 17. Time series of average differences between observation and simulation for horizontal wind vector components, speed, temperature, dew point, and pressure at the 1000-m level.

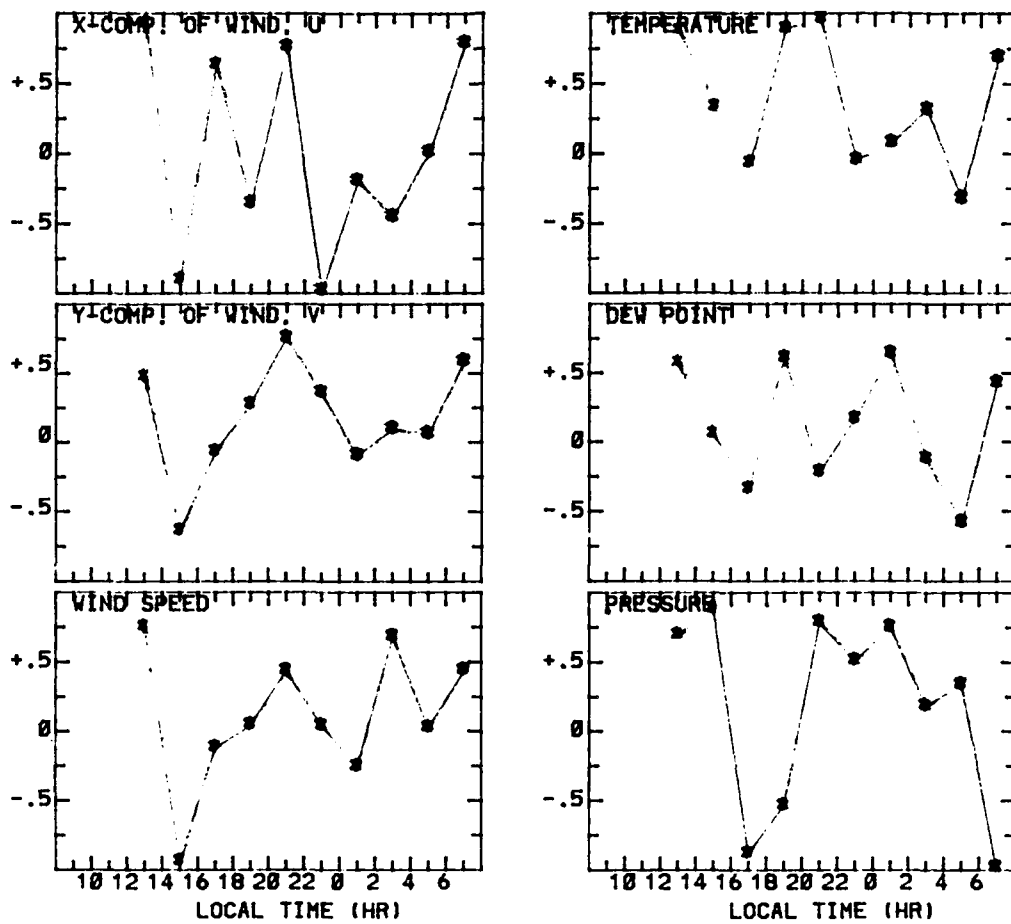


Figure 18. Time series of correlation coefficients of horizontal wind vector components, speed, temperature, dew point, and pressure for the 10-m level.

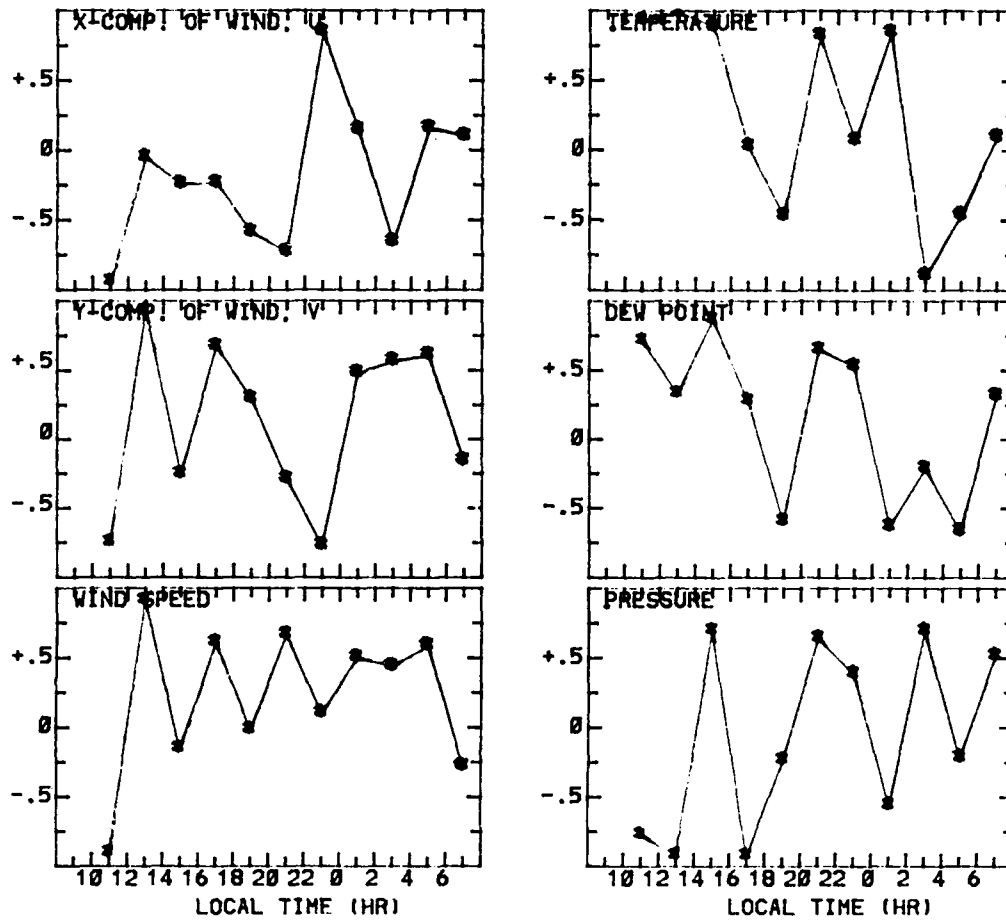


Figure 19. Time series of correlation coefficients of horizontal wind vector components, speed, temperature, dew point, and pressure for the 1000-m level.

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