

Polar Ionospheric Imaging at Storm Time

Ms Ping Yin and Dr Cathryn Mitchell

Department of Electronic and Electrical Engineering
University of Bath
BA2 7AY
UNITED KINGDOM

p.yin@bath.ac.uk / eescnm@bath.ac.uk

Dr Gary Bust

ARL University of Texas
USA

gbust@arlut.utexas.edu

ABSTRACT

During periods of major geomagnetic storms the polar ionosphere becomes very variable. Ground-based measurements, such as NIMS, GPS or ionosonde data are too sparse to do tomographic imaging, however, combination of these data sources as well as other available measurements, such as satellite-based data, may have the possibility to monitor the structure of the disturbed polar ionosphere. In this paper, LEO-based GPS data onboard CHAMP as well as ground-based GPS and ionosonde observations are input into a four dimensional tomographic algorithm – MIDAS (Multi-Instrument Data Analysis System) to image the disturbed ionosphere at Alaska and Greenland as well as over Europe for the major storm in October 2003. In contrast, electron density images produced by another independent method – IDA3D (Ionospheric Data Assimilation Three Dimensional), which assimilates primarily NIMS data and other data source, are involved to perform comparisons. A general good agreement can be obtained between them. As a result, the addition of LEO-based GPS data presents a great potential in polar ionospheric imaging.

1.0 INTRODUCTION

Tomographic imaging is an established method to study the ionosphere [for example see 1, 2]. Because the GPS coverage at the higher latitudes is relatively sparse due to the orbital inclination of the GPS constellation, it is difficult to image the ionosphere there with tomographic techniques using only ground-based GPS data. With the development of the radio-occultation technique in ionospheric imaging [3], however, the characteristics of the polar ionosphere at the storm-time are expected to be represented better with occultation data. Radio-occultation measurements observed by Low earth orbiters (LEO) have been adopted to improve ionospheric tomography [4]. In this study, we use both navigation and occultation measurements onboard the LEO CHAMP to make the inversion.

Using the CHAMP navigation and occultation data and some available ionosonde data, in conjunction with a 4D tomographic method (MIDAS) [5], images of electron density over Greenland and Alaska and the high latitudes over Europe are generated for periods of the October 2003 storm. A comparison between different data sources for imaging, such as GPS/CHAMP/ionosonde measurements, is made. Experimental results of using radio occultation data for imaging the disturbed ionosphere have been discussed previously by Yin et al [6].

Yin, P.; Mitchell, C.; Bust, G. (2006) Polar Ionospheric Imaging at Storm Time. In *Characterising the Ionosphere* (pp. 3-1 – 3-12). Meeting Proceedings RTO-MP-IST-056, Paper 3. Neuilly-sur-Seine, France: RTO. Available from: <http://www.rto.nato.int/abstracts.asp>.

Report Documentation Page				Form Approved OMB No. 0704-0188	
Public reporting burden for the collection of information is estimated to average 1 hour per response, including the time for reviewing instructions, searching existing data sources, gathering and maintaining the data needed, and completing and reviewing the collection of information. Send comments regarding this burden estimate or any other aspect of this collection of information, including suggestions for reducing this burden, to Washington Headquarters Services, Directorate for Information Operations and Reports, 1215 Jefferson Davis Highway, Suite 1204, Arlington VA 22202-4302. Respondents should be aware that notwithstanding any other provision of law, no person shall be subject to a penalty for failing to comply with a collection of information if it does not display a currently valid OMB control number.					
1. REPORT DATE 01 JUN 2006		2. REPORT TYPE N/A		3. DATES COVERED -	
4. TITLE AND SUBTITLE Polar Ionospheric Imaging at Storm Time				5a. CONTRACT NUMBER	
				5b. GRANT NUMBER	
				5c. PROGRAM ELEMENT NUMBER	
6. AUTHOR(S)				5d. PROJECT NUMBER	
				5e. TASK NUMBER	
				5f. WORK UNIT NUMBER	
7. PERFORMING ORGANIZATION NAME(S) AND ADDRESS(ES) Department of Electronic and Electrical Engineering University of Bath BA2 7AY UNITED KINGDOM				8. PERFORMING ORGANIZATION REPORT NUMBER	
9. SPONSORING/MONITORING AGENCY NAME(S) AND ADDRESS(ES)				10. SPONSOR/MONITOR'S ACRONYM(S)	
				11. SPONSOR/MONITOR'S REPORT NUMBER(S)	
12. DISTRIBUTION/AVAILABILITY STATEMENT Approved for public release, distribution unlimited					
13. SUPPLEMENTARY NOTES See also ADM002065., The original document contains color images.					
14. ABSTRACT					
15. SUBJECT TERMS					
16. SECURITY CLASSIFICATION OF:			17. LIMITATION OF ABSTRACT UU	18. NUMBER OF PAGES 41	19a. NAME OF RESPONSIBLE PERSON
a. REPORT unclassified	b. ABSTRACT unclassified	c. THIS PAGE unclassified			

Polar Ionospheric Imaging at Storm Time

For verification purposes, images in the same region are produced by another tomographic algorithm - Ionospheric Data Assimilation Three Dimensional (IDA3D) [7] with NIMS data, ground-based and space-based GPS measurements. NIMS data are excellent for producing 2-D images of the ionosphere but are limited by the number of suitable satellite passes. By comparing between images reconstructed with GPS, GOS and GIOS (see their representation in the 4.0 section) results, the advantages of using occultation/navigation data over the polar region are clearly identified.

2.0 METHOD

Since the ionosphere is dispersive in nature, either the code or the phase delay in the GPS signals is proportional to the inverse of the squared frequency and directly proportional to the change in TEC. For dual-frequency GPS receivers, the differential phase shift is formed from linear combinations of carrier phases at L1 and L2, as shown in the following equation,

$$\phi_{L1} - \phi_{L2} = 40.3 \times \left(\frac{1}{f_{L2}^2} - \frac{1}{f_{L1}^2} \right) \times TEC + \varepsilon_{\phi}$$

in which Φ_{L1} and Φ_{L2} are the phase delays of L1 and L2 respectively, f_{L1} and f_{L2} are GPS two carrier frequencies, TEC is the relative slant TEC values and ε_{ϕ} is the residual errors including integer ambiguities in addition to multi-path and inter-frequency biases. Therefore, calibration is needed if accurate GPS TEC will be used to image the ionosphere due to the residual errors. In this study, ground-based and space-based GPS data are firstly bias-corrected to remove most of biases including the satellite-receiver inter-frequency biases. Then calibrated GPS data are re-input into MIDAS to image the disturbed ionosphere for the 29 and 30 October 2003 storm.

As is well described in [5], the MIDAS inversion algorithm which is used in this study requires a priori assumptions to image the ionosphere. For the storm conditions, an extended range of Chapman profiles with a range of scale heights and peak heights is set to construct the background ionosphere during storms. Then as in the standard method, spherical harmonics are applied concerning the horizontal variations, and EOFs (Empirical Orthonormal Functions) are generated to represent the vertical distribution of ionization.

Finally, the reconstructed inversion results, for example, the electron concentration distribution and total electron content (TEC) variation against latitudes/longitudes will be shown hourly in forms of maps with 1° (Latitude) $\times$ 4° (Longitude) $\times$ 50 km (Altitude) grid voxels extending from 80 to 1080 km above the earth.

3.0 MEASUREMENTS

The observational inputs to MIDAS include ground-based GPS measurements (phase and time delay) that are available from SOPAC (Scripps Orbit and Permanent Array Center). In this study, ground GPS receivers are selected from 160°W to 40°E in longitude and 50°N to 80°N in latitude. Space-based GPS measurements and occultation data onboard LEO/CHAMP are supplied by the ISDC (Information System and Data Center, GFZ) and the ground ionosondes data are available from the SPIDR (Space Physics Interactive Data Resource). Table 1 lists the available ionosondes for this storm.

Table 1: Ionosondes locations

Ionosonde code	Station name	Geographic Latitude	Geographic Longitude
THJ77	Thule/Qaanaaq	77.5°N	69.2°W
KS759	King Salmon	58.4°N	156.4°W
NQJ61	Narssarssuaq	61.2°N	45.4°W
JR055	Juliusruh/Rugen	54.6°N	13.4°E
TR169	Tromso	69.7°N	19.0°E
FF051	Fairford	51.7°N	1.5°W

The data used are:

Ground-based GPS receivers observations with 30 seconds data interval; Space-based GPS receiver onboard CHAMP satellite, that is, GPS code and carriers phase data records including occultation data with 1Hz sample rate and GPS RINEX data with 10 sec sample rate. These data were processed for the period of 28 to 31 October 2003, here we focus on the 29th and 30th that are the main phases of this storm.

4.0 RESULTS AND DISCUSSIONS

The results description as shown in Table 2 explains the notation used for the different data combinations. GPS is used to represent the inversion images from ground-based GPS data alone, GOS represents the reconstructed images from the combination of ground-based GPS and CHAMP occultation and navigation data, while GIOS stands for the inversion results from ground-based GPS, CHAMP (occultation and navigation) and ionosonde data.

Table 2: Inversion results representation in terms of their input data

Results	Ground-based GPS	CHAMP observations	Ionosondes
GPS	×		
GOS	×	×	
GIOS	×	×	×

4.1 29 October 2003 Storm

For the storm on 29th October, the focus is put on the analysis of the addition of CHAMP data, which is investigated from the two aspects, that is, with a CHAMP pass and without a CHAMP pass in that area. The areas of interest are concentrated on the higher latitudes, in particular, Greenland and Alaska.

Hourly TEC maps reconstructed by CHAMP navigation/occultation data over the northern hemisphere on 29 October 2003 is shown in Figure 1, from which the CHAMP traces can be identified. For example, The CHAMP passed over Alaska and Greenland at 0900, 1100, 2000, 2200 UT and 0300, 0500, 1400, 1600 UT, respectively.

Polar Ionospheric Imaging at Storm Time

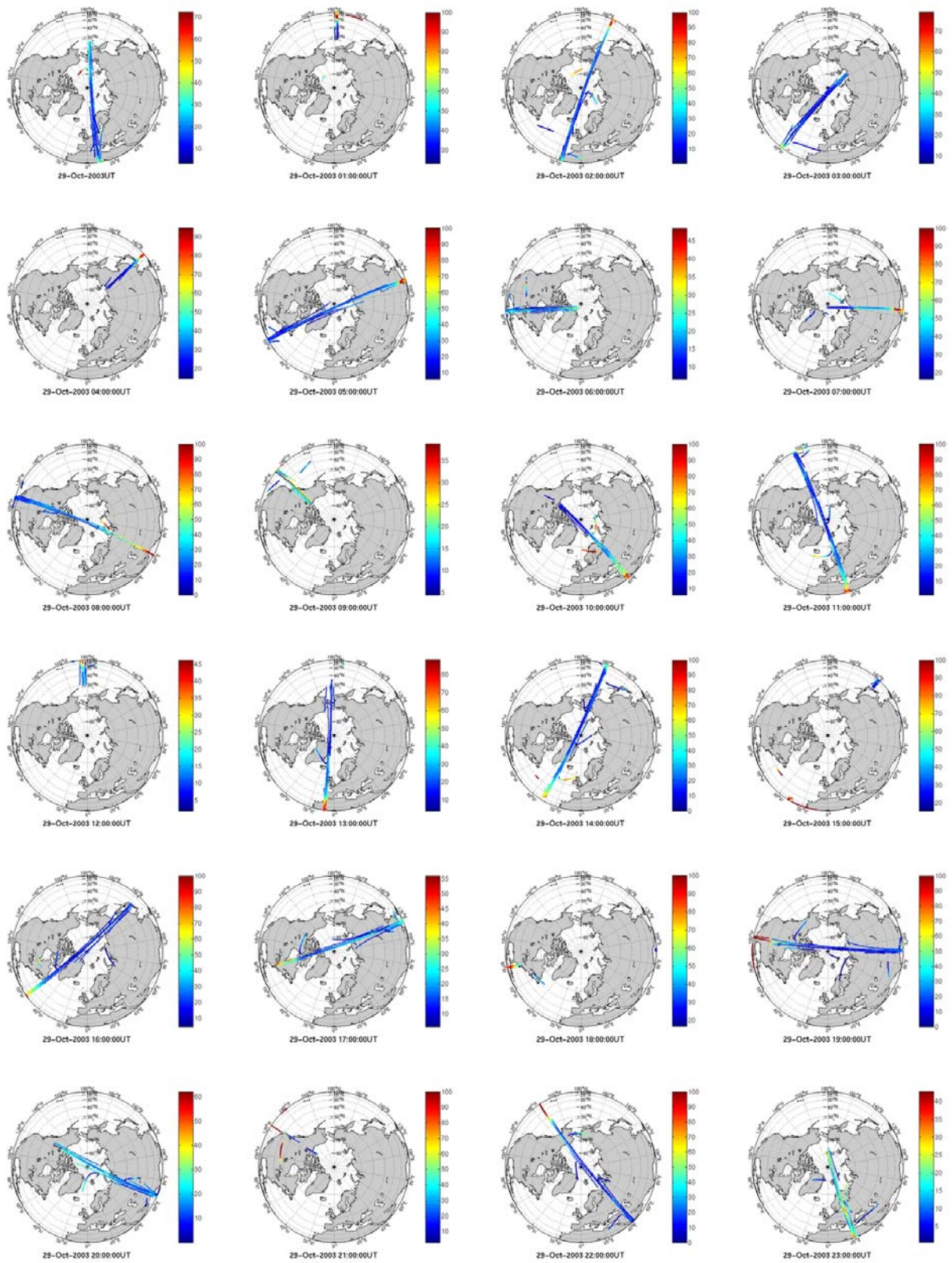


Figure 1: Hourly CHAMP TEC maps over the northern hemisphere on 29 October 2003

4.1.1 Without CHAMP Passes

The effect of additional CHAMP data is considered when there was no CHAMP pass over that region. As can be seen from CHAMP passes over the north hemisphere on 29 October in Figure 1, the images over Alaska at 0700 UT and those over Greenland at 2200 UT are selected to present the results when there is no CHAMP passes. The latitude range of the images is from 50°N to 80°N over Alaska, while 60°N to 90°N were selected for Greenland.

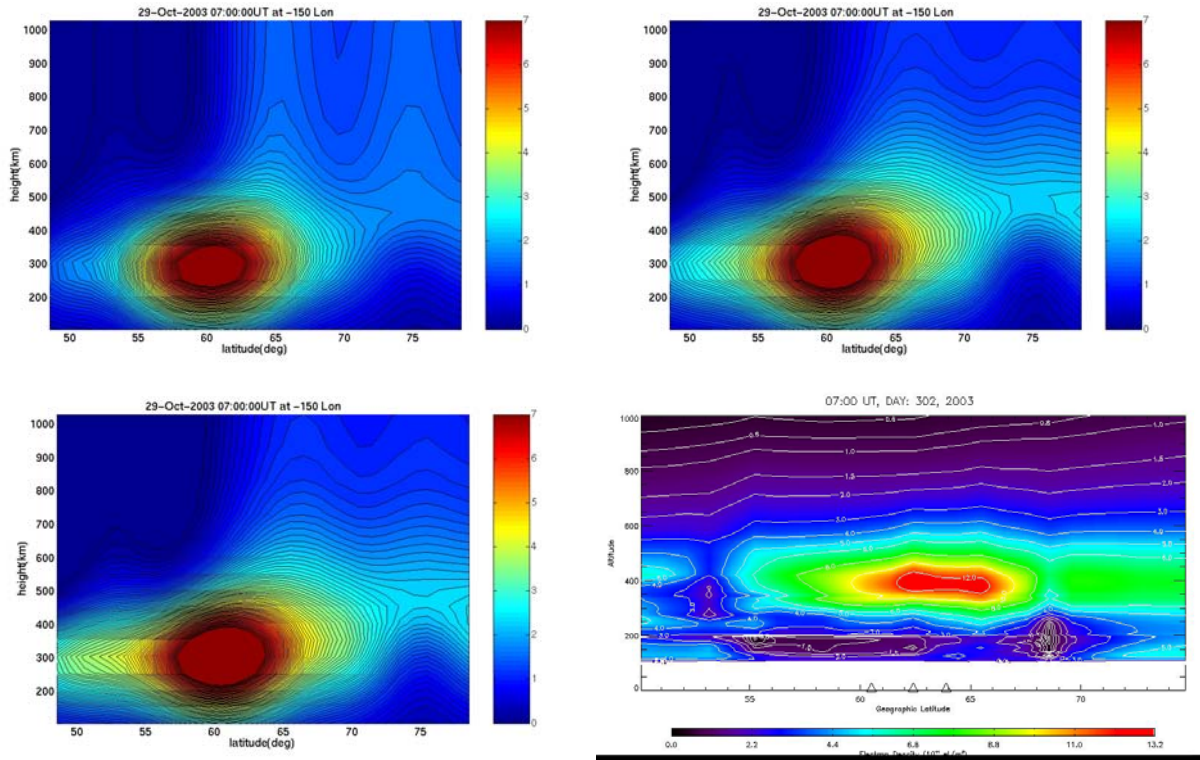


Figure 2: Electron density images over Alaska at 0700 UT on 29 October (top left: GPS results; top right: GOS results; bottom left: GIOS results; bottom right: NIMS results)

When there is no CHAMP pass over the region, the addition of occultation and navigation data has little effect on the final images in terms of the vertical profile, but the electron density shows some difference, indicating that the data input from another region can have an effect over the entire image. The images over Alaska at 0700 UT as shown in Figure 2 which derived from calibrated GPS, GOS and GIOS results respectively, are nearly the same at all latitudes. They all see the electron density enhancement around 60°N, which is also observed in the NIMS image as shown in the bottom right panel in Figure 2, in which triangles on the X-axis stand for the location of NIMS receivers.

Furthermore, all of the images over Greenland at 2200 UT from GPS, GOS and GIOS results in Figure 3 present the ionospheric uplift to 700 km at 78°N. However, it is apparently not effected by the contribution of CHAMP occultation and navigation data since there is no difference in the region above 75°N among the GPS/GOS/GIOS reconstructed images. This is interesting because it indicates that the ground-based GPS data contains information about the vertical profile of the electron density. This may be more the case at high latitudes where the rays are predominantly at low elevations.

Polar Ionospheric Imaging at Storm Time

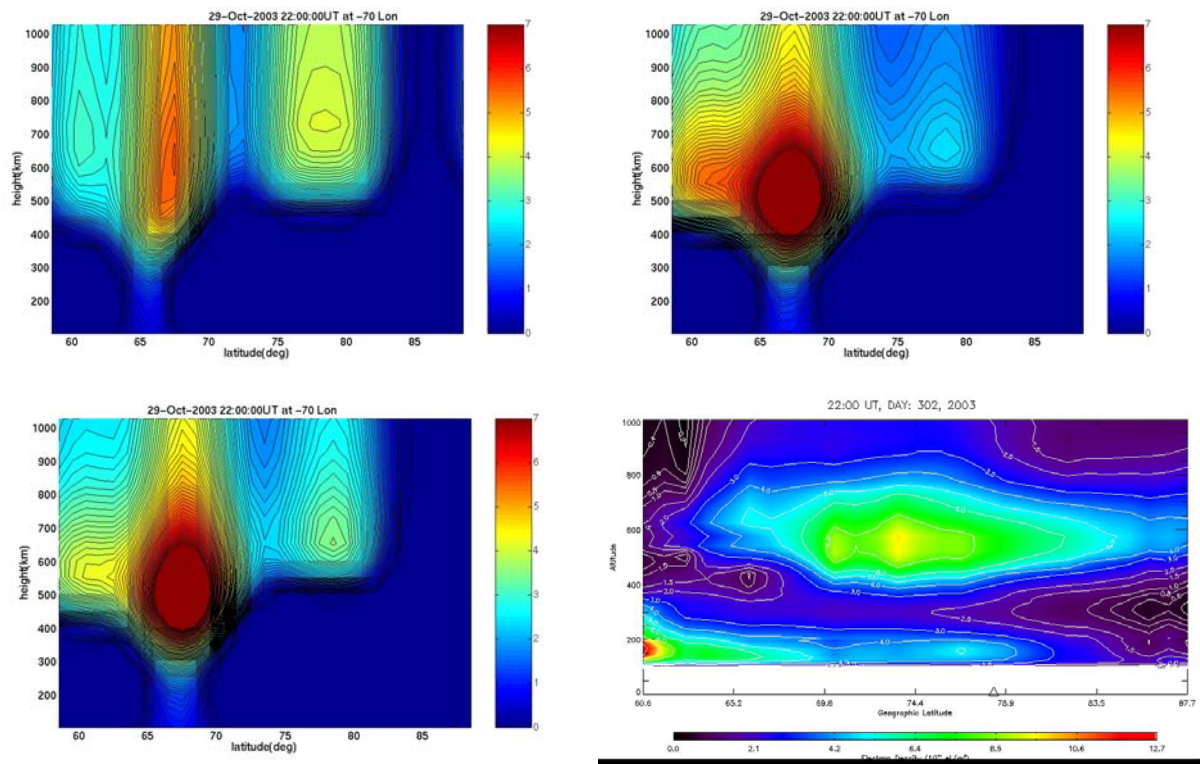


Figure 3: Electron density images over Greenland at 2200 UT on 29 October (top left: GPS results; top right: GOS results; bottom left: GIOS results; bottom right: NIMS results)

4.1.2 With CHAMP Passes

As can be seen from CHAMP passes over the north hemisphere on 29 October in Figure 1, the images over Greenland at 1400 UT and those over Alaska at 1100 UT are selected to present the advantage of CHAMP measurements.

Figure 4 illustrates electron density distribution produced from calibrated GPS, GOS and GIOS inverted results, respectively. Electron concentration is plotted as a function of geographic latitudes and altitude at 1100 UT on 29 October over Alaska. The bottom right panel in Figure 4 shows the electron density image at 1115 UT over Alaska produced by mainly NIMS inverted results, which is used for verification purpose.

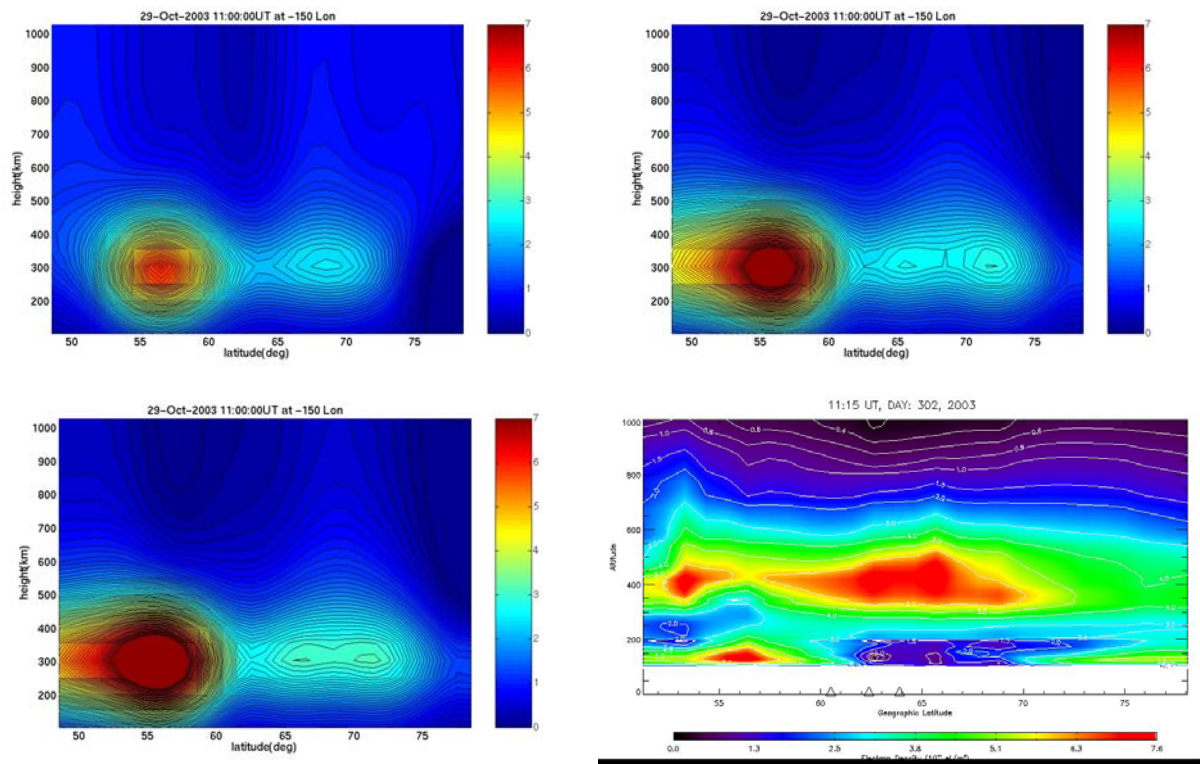


Figure 4: Electron density images over Alaska at 1100 UT on 29 October (top left: GPS results; top right: GOS results; bottom left: GIOS results; bottom right: NIMS results)

From comparisons among the above images, we can see that starting at 53°N there is an electron density peak clearly seen from the NIMS image in Figure 4, which is also shown in GOS and GIOS images (see Figure 4). However this peak moves to higher latitude (above 55°N) in the reconstructed GPS image (see the top left panel). Besides, further north there is another electron density peak around 65°N, which is clearly seen in the NIMS image. Again from the top right and bottom left panels, we can observe this peak at exactly 65°N, but in the GPS reconstructed image, this peak moved to 68°N. Hence, a conclusion can be made that the GOS/GIOS reconstructed images represent the ionospheric distribution more accurately than GPS reconstructed images.

Polar Ionospheric Imaging at Storm Time

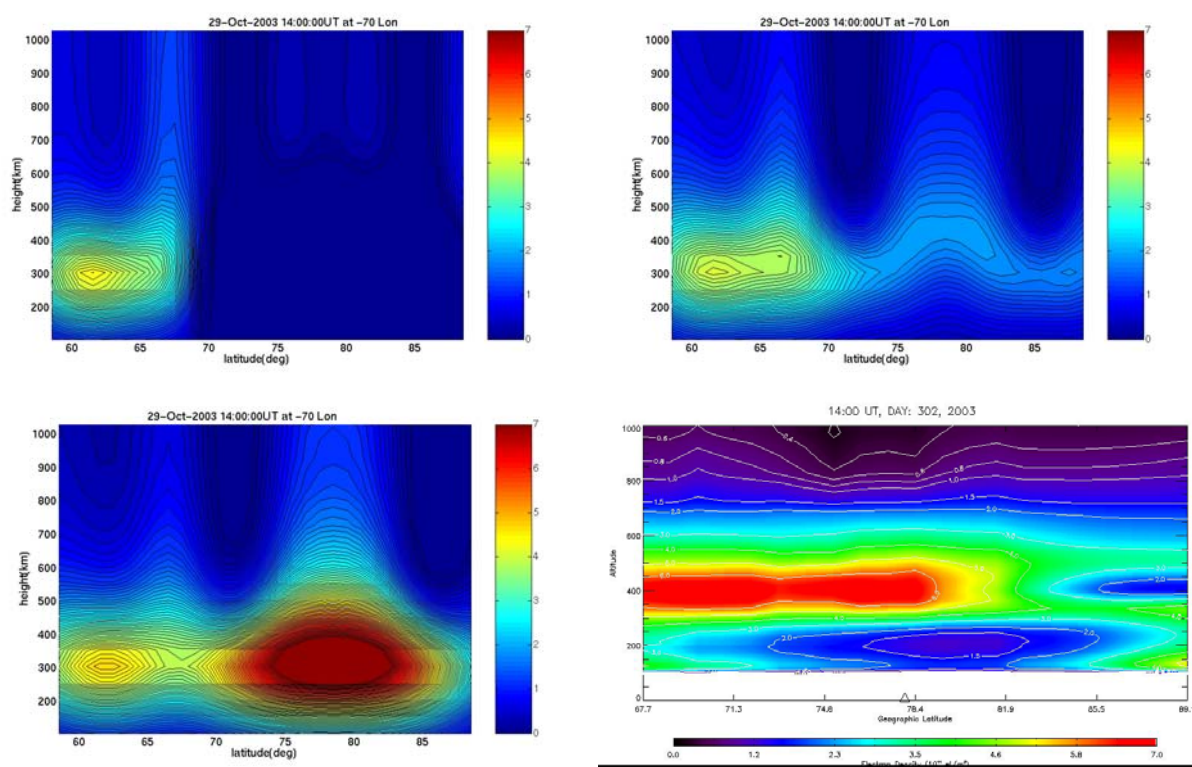


Figure 5: Electron density images over Greenland at 1400 UT on 29 October (top left: GPS results; top right: GOS results; bottom left: GIOS results; bottom right: NIMS results)

Furthermore, GOS/GIOS results are much better in regions above 70°N where GPS ground measurements are in lack, which we can tell from the following images over Greenland. Figure 5 illustrates electron density distribution produced from calibrated GPS, GOS and GIOS inverted results, respectively. They are plotted at 1400 UT on 29 October over Greenland. In addition, as a reference, the bottom right panel in Figure 5 also shows the image of electron density from NIMS results at the same time over Greenland.

As compared to the NIMS image, the GPS reconstructed image shown in the top left panel does not exactly reproduce the distribution of electron density above 70°N latitudes. This can be proved by the occurrence of electron density peak around 78°N in the NIMS image, which is not seen in the GPS reconstructed image. However, those GOS/GIOS images (the top right and bottom left panels) produced with the assimilation of CHAMP occultation/navigation data into ground GPS data greatly improved the ionospheric images of electron density in the higher latitudes where CHAMP passed. They both show the density peak except for a bit different peak height. The peak height derived from the GOS results (about 400 km) is similar to that from the NIMS results. The addition of ionosonde measurements may make the electron density greater, particularly the peak density.

4.2 30 October 2003 Storm

For the storm on the 30th, the area of interest is focused on the electron density images at Svalbard in the late evening, where the simultaneous EISCAT scan can be obtained from the Madrigal database.

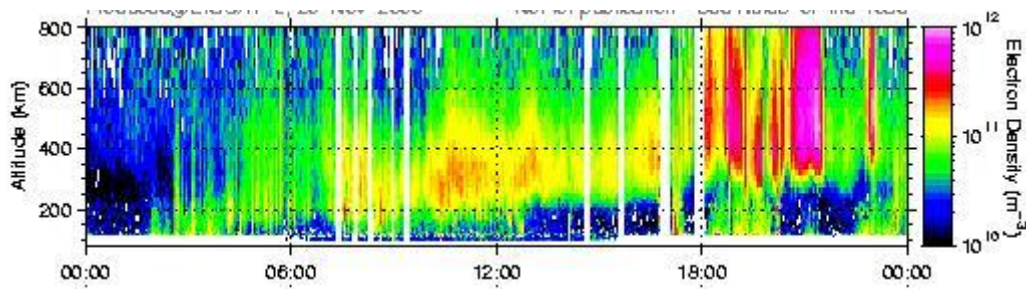


Figure 6: EISCAT scan at Svalbard on 30 October 2003

The electron density image as a function of altitude and time (UT) obtained with the EISCAT radar at Svalbard (78°N and 16°E) is shown in Figure 6. It can be clearly seen that electron densities are increased after 1800 UT on 30 October. In particular, the F layer height elevated above 500 km with the enhancement in electron densities around 2100 UT. In contrast, Figure 7 shows calibrated GIOS reconstructed images at 1900 - 2200 UT on 30 October. Here the images were generated by calibrated GIOS results in that the ESR is located at the high latitude where the ionosphere is represented well by calibrated GIOS results. The image at 2100 UT coincides with the ESR scan very well, where they both present the uplifts of the F-layer in conjunction with the electron density increase. The uplift in the peak height was also observed over the USA for the July 2000 storm [8].

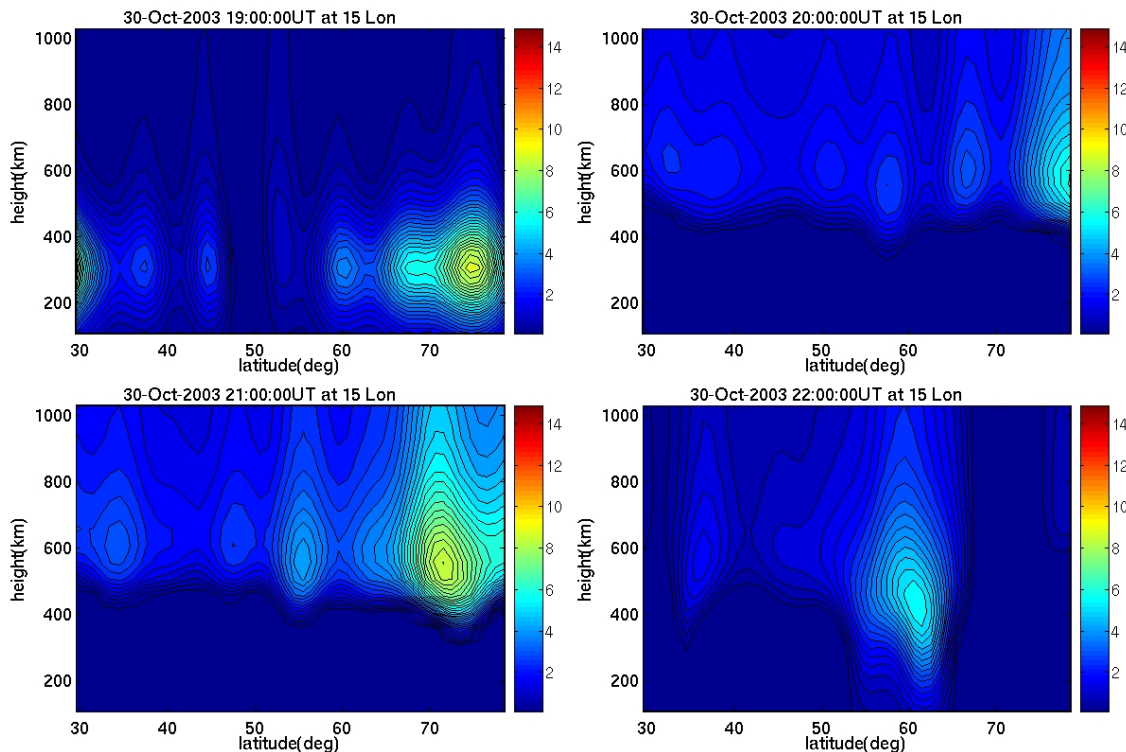


Figure 7: Images reconstructed from GIOS results at 15°E for 1900 - 2200 UT on 30 October

5.0 SUMMARY

The 29 October 2003 storm was chosen to compare GPS/GOS/GIOS reconstructed images by means of the verification of NIMS images over the polar region. The results indicate that GOS/GIOS reconstructed

Polar Ionospheric Imaging at Storm Time

images show good agreement with the NIMS images when CHAMP passed that area. The addition of CHAMP occultation data can improve the imaging of the detailed structure in the disturbed ionosphere over the polar region because of the shortage of the ground-based GPS receivers there.

Ionosondes are also sparsely distributed at the high latitude regions. There is only one available over Greenland (NQJ61 with 61.2°N and 45.4°W) on 29 October, while another one over Alaska (KS759 with 58.4°N and 156.4°W) was unavailable for the selected times on 29 October. This accounts for the similar images between GOS/GIOS results over Alaska, however, those images over Greenland have a difference in electron density. Thus, ionosonde data has some effect on the distribution of the electron density, in particular, increasing the peak density.

In summary, the images derived from GOS/GIOS results resemble the NIMS images much better than those from GPS results, which are lower resolution and less accurate at higher latitudes due to the shortage of ground receivers. In terms of the peak height, there seems no big difference among GPS, GOS and GIOS results. This may be due to the predomination of low-elevation rays from ground-based GPS there. The addition of CHAMP data into the tomographic algorithm greatly improved the details in ionospheric images over the polar region, e.g., from GOS images over Greenland, the electron density enhancement can be seen over the auroral and polar regions (above 65°N), but there is little information from GPS reconstructed images due to the lack of GPS receivers there.

Dramatic uplifts in the F-layer at the high latitudes over Europe for the periods of 2000 – 2100 UT are captured by the reconstructed images, which are validated by the correspondent ESR scan. Meanwhile, the abnormal enhancement in electron-density in the local evening was also observed by the inverted images and the ESR scan, which may be attributed to the plasma convection from the mainland USA [9].

6.0 ACKNOWLEDGEMENT

The authors are grateful to SOPAC for the provision of GPS data, SPIDR for the ionosonde data, ISDC/WDZ for the CHAMP data and the Madrigal database for the EISCAT scan. Support is also acknowledged from the UK PPARC and EPSRC.

7.0 REFERENCES

- [1] Fremouw E. J., Secan J. A. and Howe B. M. (1992). *Application of Stochastic inverse theory to ionospheric tomography*. Radio Science, 27 (5), 721-732.
- [2] Bernhardt, P. A., et al. (1998). *Two dimensional mapping of the plasma density in the upper atmosphere with computerized ionospheric tomography (CIT)*. Physics of Plasmas, 5, 2010-2021.
- [3] Hajj G. A., Ibañez-Meier R., Kursinski E. R. and Romans L. J. (1994). *Imaging the ionosphere with the Global Positioning System*. Int. J. Imag. Syst. Technol. 5, 174.
- [4] Rius A., Ruffini G. and Cucurull L. (1997). *Improving the vertical resolution of ionospheric tomography with GPS occultations*. Geophys. Res. Lett. 24 (18), 2291-2294.
- [5] Mitchell, C. N. and Spencer P. S. (2003). *A three dimensional time-dependent algorithm for ionospheric imaging using GPS*. Ann. Geophysice-ITALY, 46 (4), 687.
- [6] Yin, P. and Mitchell C. N. (2005). *Use of radio-occultation data for ionospheric imaging during the April 2002 disturbances*. GPS Solutions, 9 (2), 156-163.

- [7] Bust G. S., Garner T. W., Gaussiran II T. L. (2004). *Ionospheric Data Assimilation Three-Dimensional (IDA3D): A global, multisensor, electron density specification algorithm*. J. Geophys. Res., 109, A11312, doi: 10.1029/2003JA010234.
- [8] Yin, P., Mitchell C. N., Spencer P. S. and Foster J. C. (2004). *Ionospheric electron concentration imaging using GPS over the USA during the storm of July 2000*. Geophys. Res. Lett., 31, L12806, doi: 10.1029/2004GL19899.
- [9] Mitchell C. N., Alfonsi L., De Franceschi G., et al. (2005). *GPS TEC and scintillation measurements from the polar ionosphere during the October 2003 storm*. Geophys. Res. Lett., 32 (12), Art. No. L12S03.

Polar Ionospheric Imaging at Storm Time



Polar ionospheric imaging at storm time

Ping Yin (1), Cathryn Mitchell (1) and Gary Bust (2)

(1) University of Bath, UK

(2) ARL at University of Texas, USA

Outline

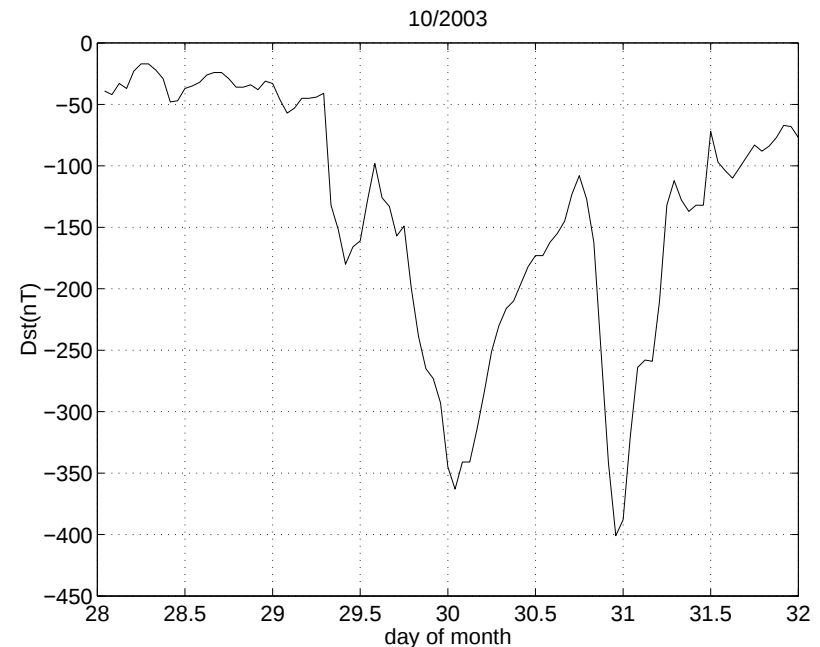
- ☐ Introduction
- ☐ Techniques
- ☐ Data sources
- ☐ Images for the 29th October 2003 storm event
- ☐ Summary
- ☐ Future work

Introduction

- ☐ Ionospheric storm description;
- ☐ Geomagnetic activity indices;
- ☐ Comparisons of electron-density images between quiet-time and storm-time.

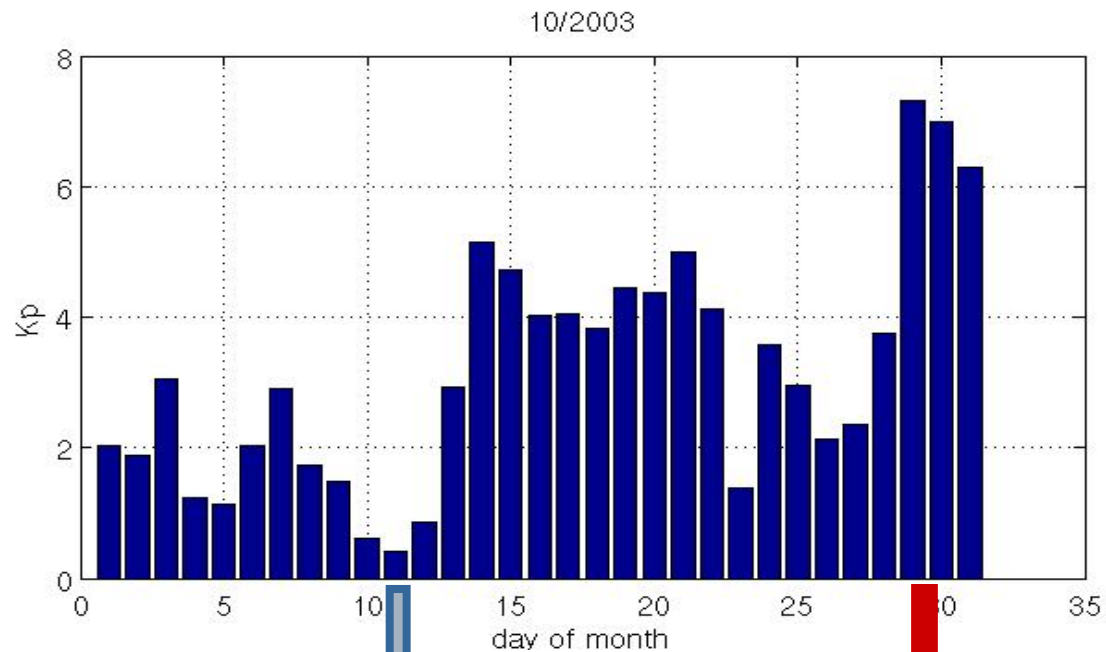
What is an ionospheric storm?

- ❑ An extreme form of space weather;
- ❑ Driven by highly variable solar and magnetosphere energy inputs to the Earth's upper atmosphere;
- ❑ The structure and dynamics of the disturbed ionosphere may be severely changed;
- ❑ Magnetic activity indices indicate the intensity of a storm, such as Dst, Kp and AE etc.



Dst variations on 28th – 31st October 2003

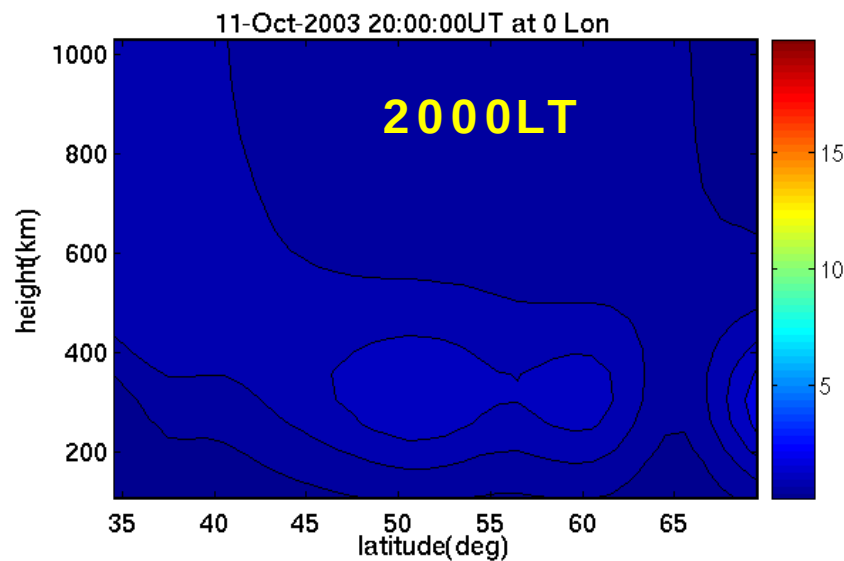
Kp index in October 2003



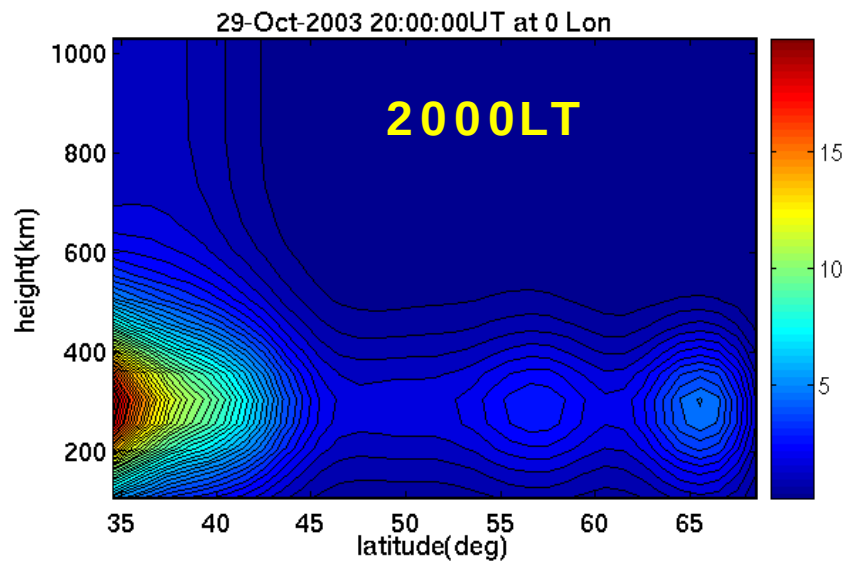
Quiet day on 11th

Storm day on 29th and 30th

Image comparison over Europe

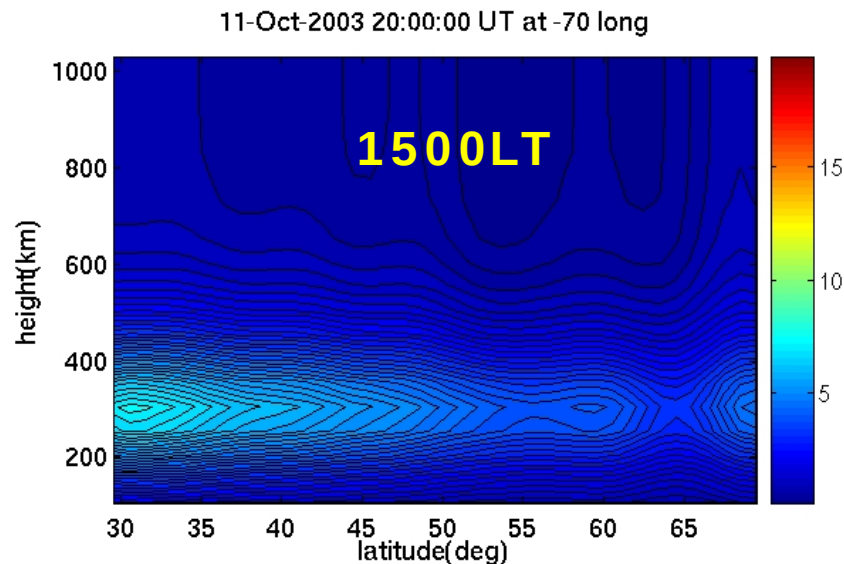


❑ Entire ionosphere depleted in the evening at quiet time

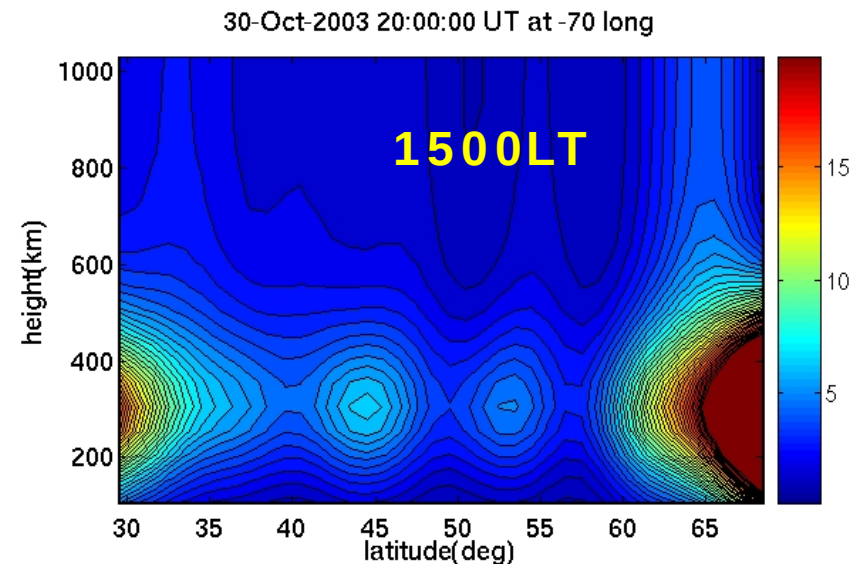


❑ Enhancement in electron-density with a steep wall of ionisation at lower latitudes

Image comparison over USA



☐ Stratified ionospheric structure in latitude with two minimums in electron-density at layer peak near 55°N and 65°N at quiet time



☐ Electron density highly increased at lower and higher latitudes, and narrow troughs at 40°N, 49°N and 57°N at storm time

Technique — GPS imaging

- GPS dual-frequency observations from both ground-based and LEO-based GPS receivers

$$\phi_{L1} - \phi_{L2} = 40.3 \times \left(\frac{1}{f_{L2}^2} - \frac{1}{f_{L1}^2} \right) \times TEC + \varepsilon_{\phi} + amb$$

Where, ϕ_{L1} and ϕ_{L2} are carrier phases at L1 and L2 frequency respectively, TEC is total electron content, amb is phase ambiguity, and ε_{ϕ} is the residual errors.

Technique — tomography

- ❑ Four-dimensional ionospheric tomography--MIDAS to reconstruct the disturbed ionosphere in latitude/longitude/altitude with time evolution;
- ❑ Spherical harmonics is applied concerning the horizontal variations and EOFs are generated to represent the vertical distribution of ionization;
- ❑ Capable of assimilating multiple data sources into inversions.

Data sources

☐ Inversion:

- Ground-based dual frequency GPS measurements from SOPAC
- LEO/CHAMP GPS navigation/occultation data from ISDC, GFZ
- Available ionosonde data from SPIDR

☐ Verification:

- NIMS observations

Tomographic imaging results

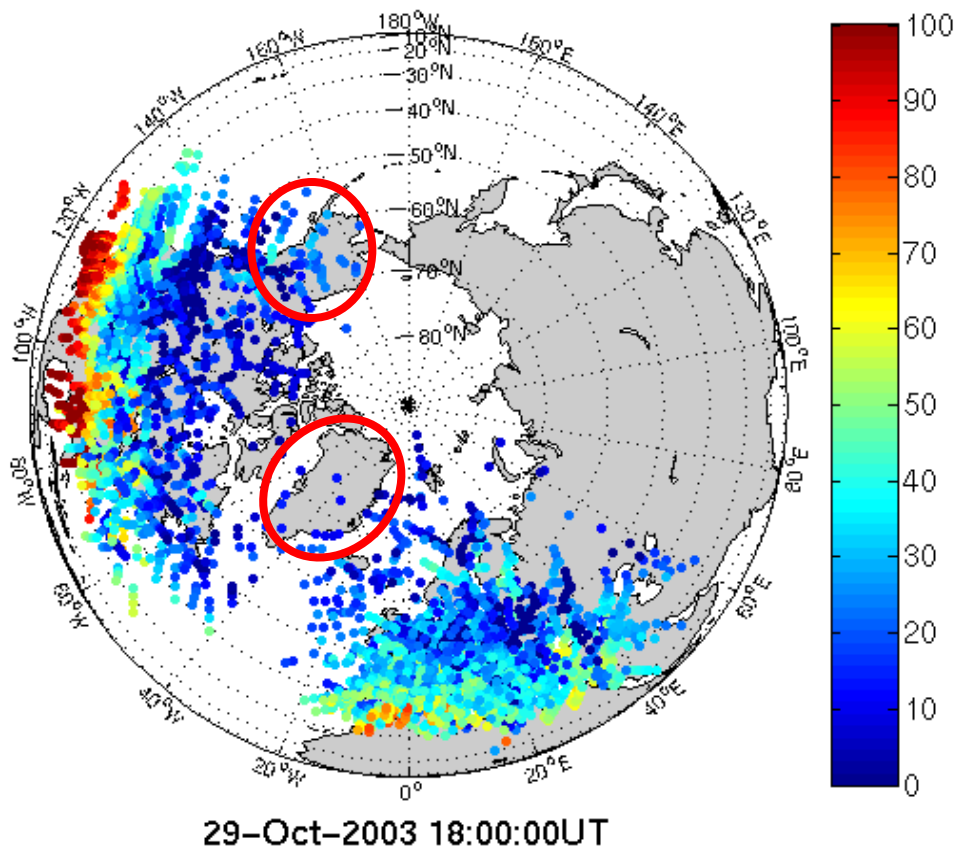
Results	Ground-based GPS	CHAMP observations	Ionosondes
GPS	√		
GOS	√	√	
GIOS	√	√	√

Table 1 Inversion results representation in terms of their input data



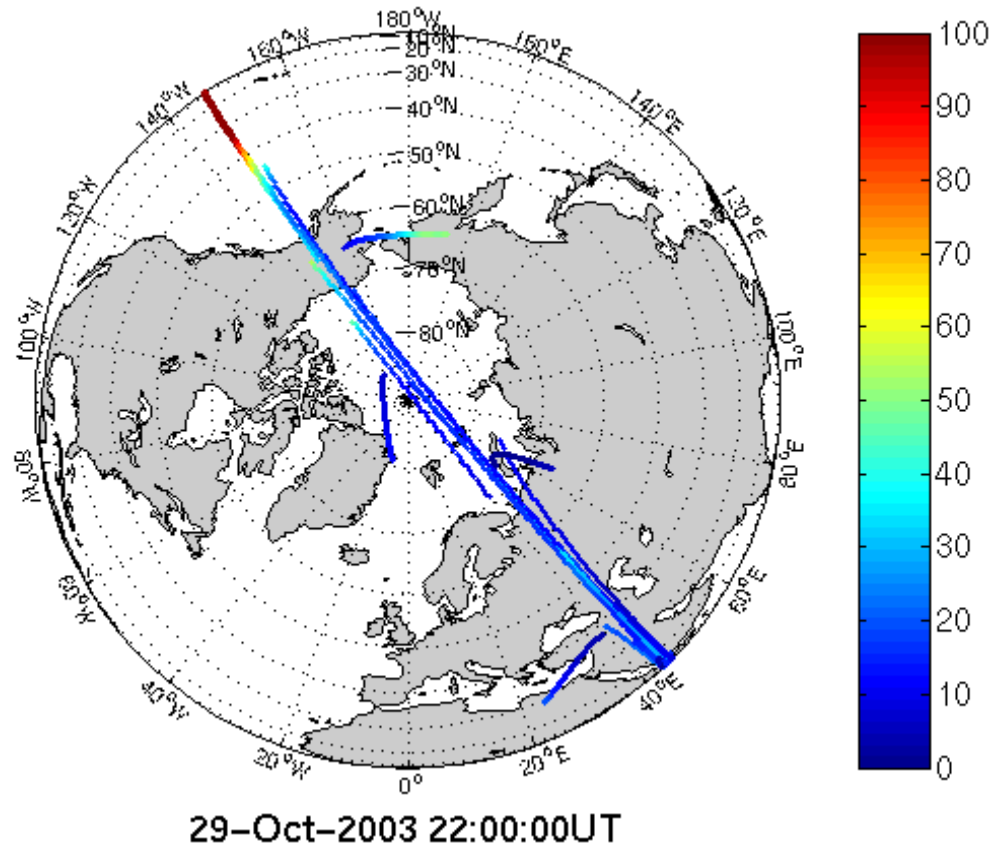
29th October 2003 storm

GPS TEC maps over North Hemisphere on 29th Oct 2003



- Strong TEC gradients were seen over the southern USA in the morning sector, while lower TEC was observed over Europe in the late afternoon sector;
- GPS TEC measurements in the polar regions, such as **Alaska and Greenland**, are sparse due to the lack of GPS ground receivers.

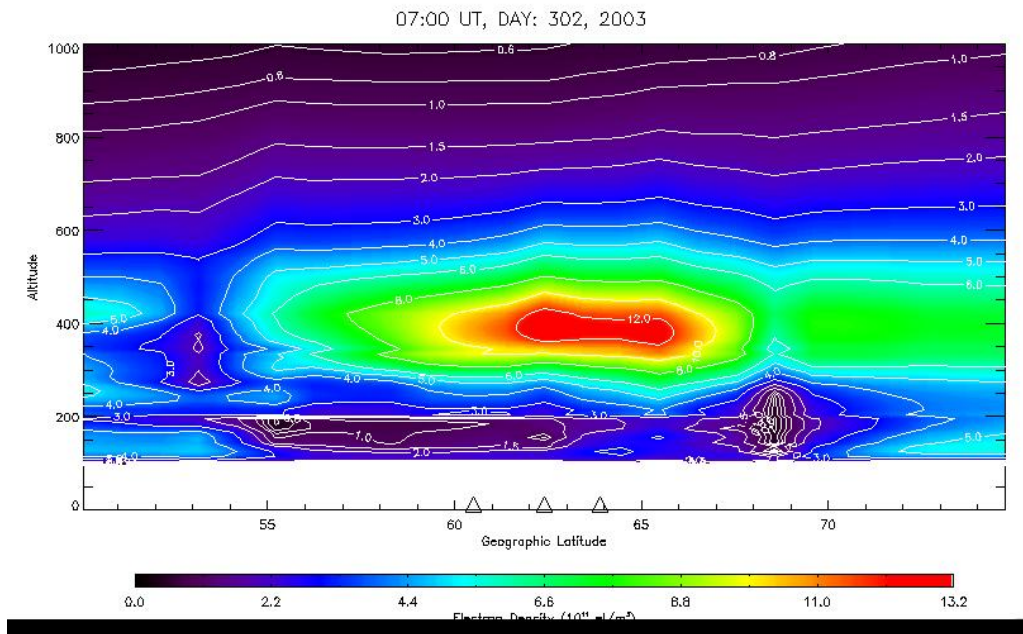
CHAMP traces



- TEC maps inverted by CHAMP GPS navigation and occultation data
- CHAMP passed over **Alaska** at 0900, 1100, 2000 and 2200 UT on 29th October; while over **Greenland** at 0300, 0500, 1400 and 1600 UT on 29th October

Without CHAMP trace over Alaska

□ At 07:00 UT (NIMS image)

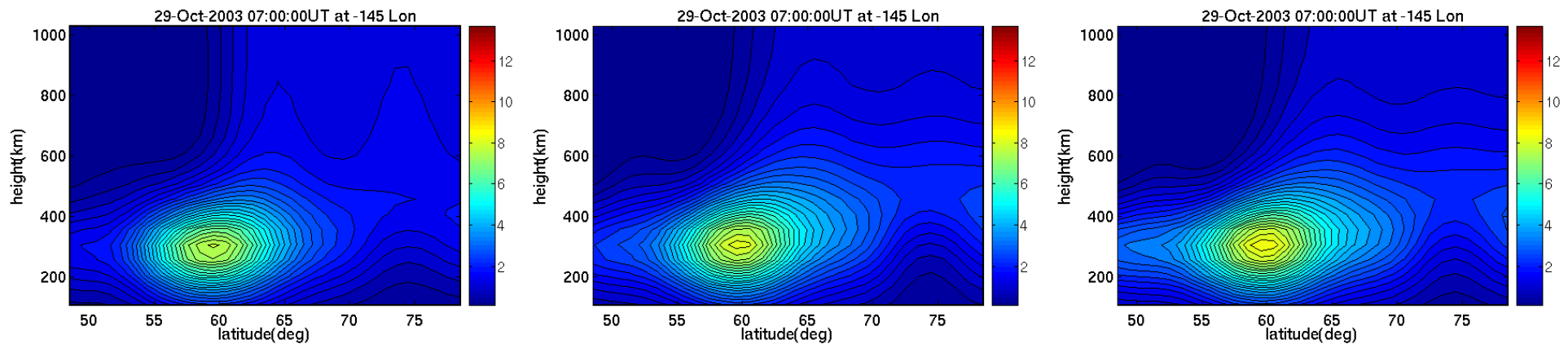


□ Three NIMS receivers over Alaska to make reconstruction with another independent inversion method

□ Enhancement in electron-density between 55°N - 68°N

Without CHAMP trace over Alaska

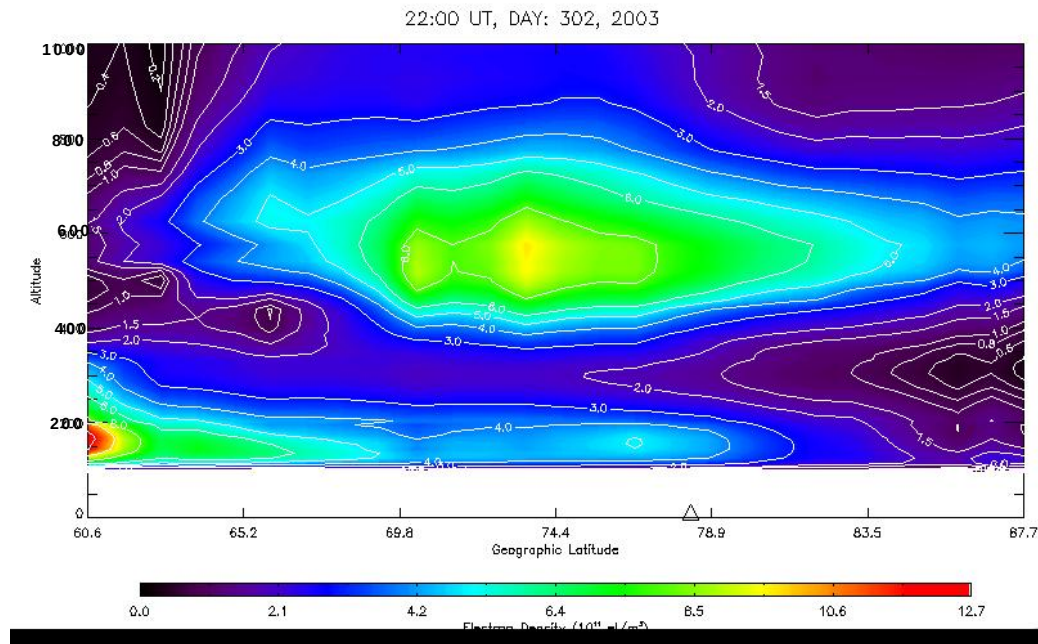
□ At 07:00 UT (GPS, GOS and GIOS images)



□ Increases in electron density are seen between 55°N – 65°N in all three images

Without CHAMP trace over Greenland

□ At 22:00 UT (NIMS image)

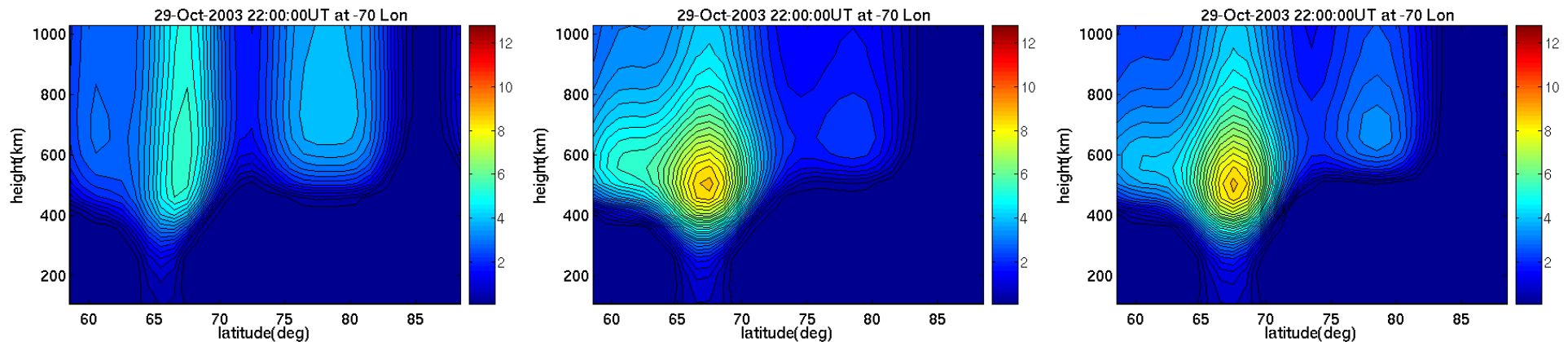


□ NIMS data from 1 Receiver and other data source to perform inversion using another tomographic method

□ The ionosphere was elevated to **600 km** near 70°N to 80°N with electron density increases

Without CHAMP trace over Greenland

□ At 22:00 UT (GPS, GOS & GIOS images)



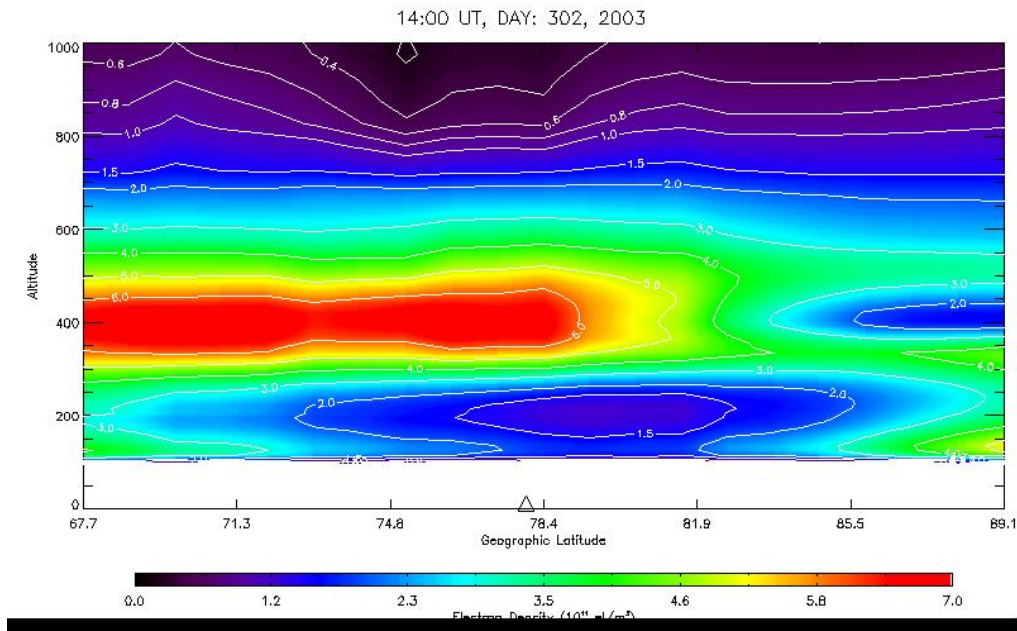
□ The F-layer uplifts to **>600km** between 70°N - 80°N were observed by all three images with different electron densities

Discussions

- When there is no CHAMP trace over the region, the addition of occultation and navigation data has little effect on the final images in terms of the vertical profile, but the electron density shows some difference, indicating that the data input from another region can have an effect over the entire image.
- Ground-based GPS data contains information about the vertical profile of the electron density. This may be more the case at high latitudes where the rays are predominantly at low elevations.

With CHAMP trace over Greenland

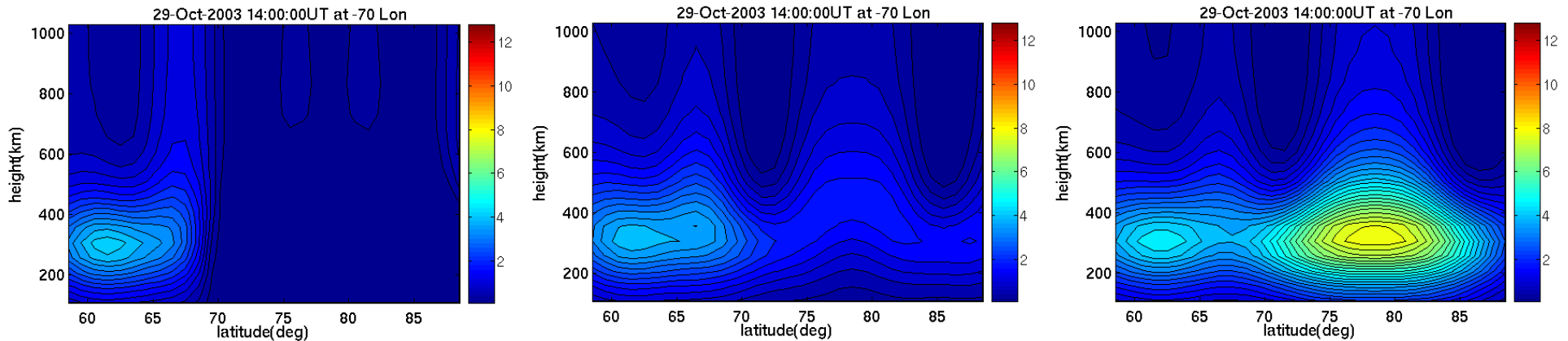
□ At 14:00 UT (NIMS image)



□ Electron densities keep the maximum in latitude from 80°N to the south with the layer height of 400km

With CHAMP trace over Greenland

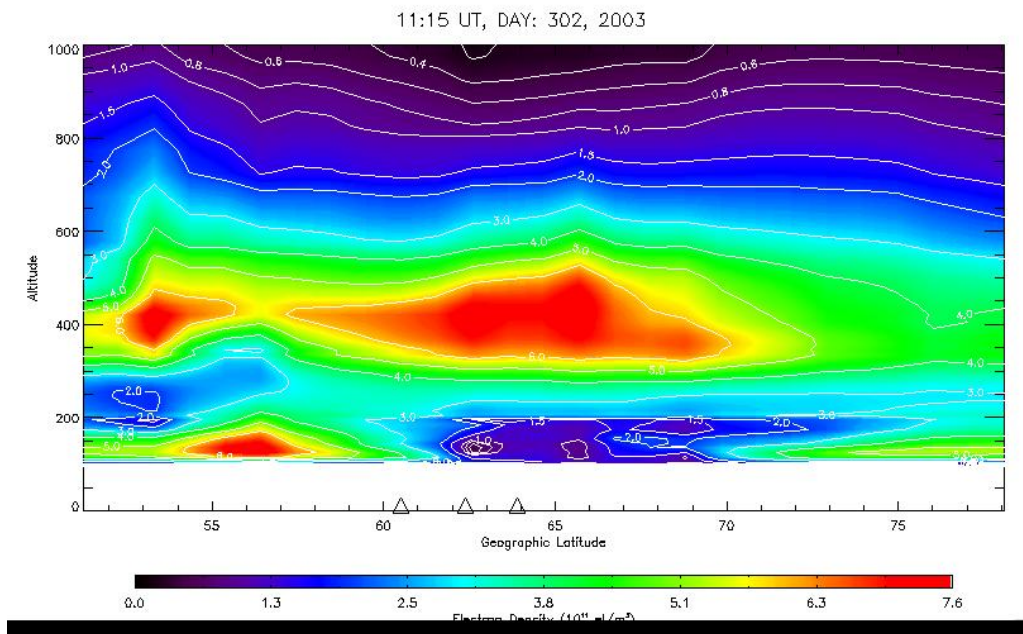
□ At 14:00 UT (GPS, GOS and GIOS image)



□ Electron-density maximum around 75°N - 80°N was reconstructed by GOS and GIOS inversions but GPS inversion

With CHAMP trace over Alaska

□ At 11:00 UT (NIMS image)

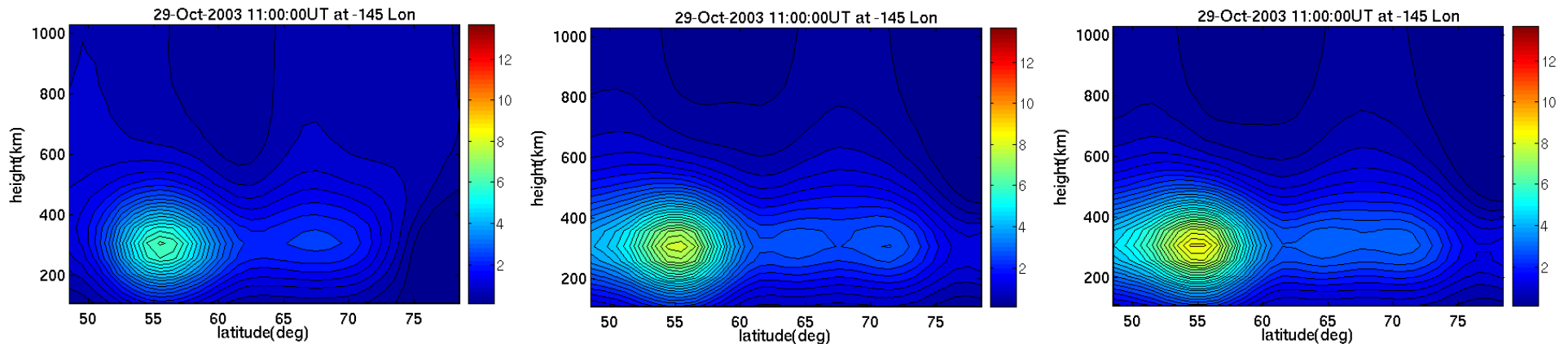


□ Starting at 53°N there is an electron density peak clearly seen from the NIMS image

□ Further north there is another electron density peak around 65°N

With CHAMP trace over Alaska

□ 11:00 UT (GPS, GOS and GIOS images)

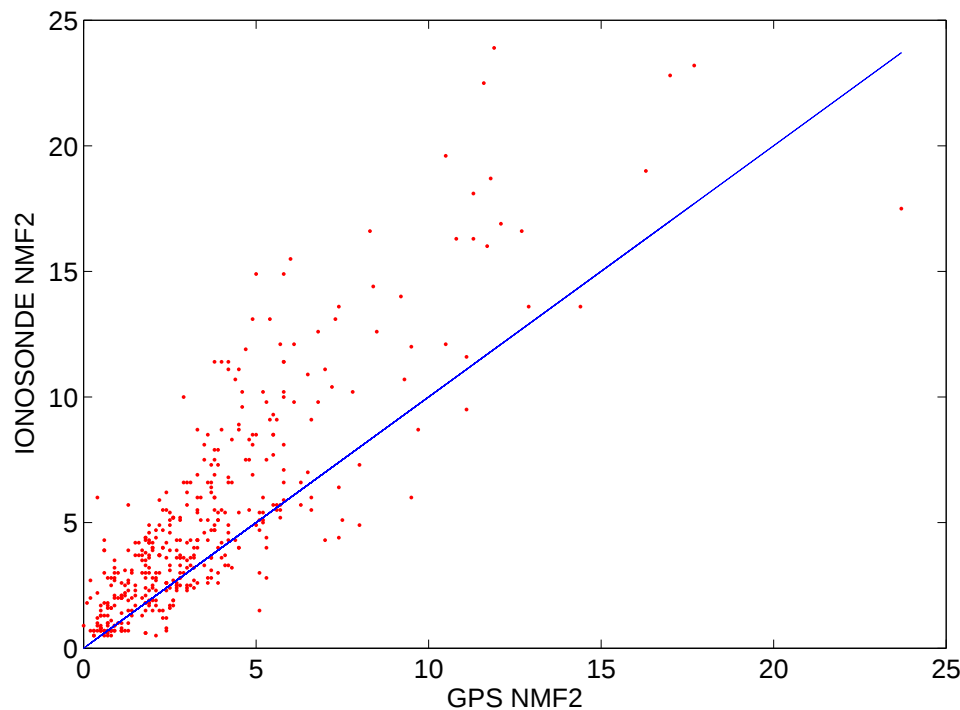


□ Two peak structures are observed in GOS/GIOS inverted images but not obvious in the GPS image, especially the further north one

Discussions

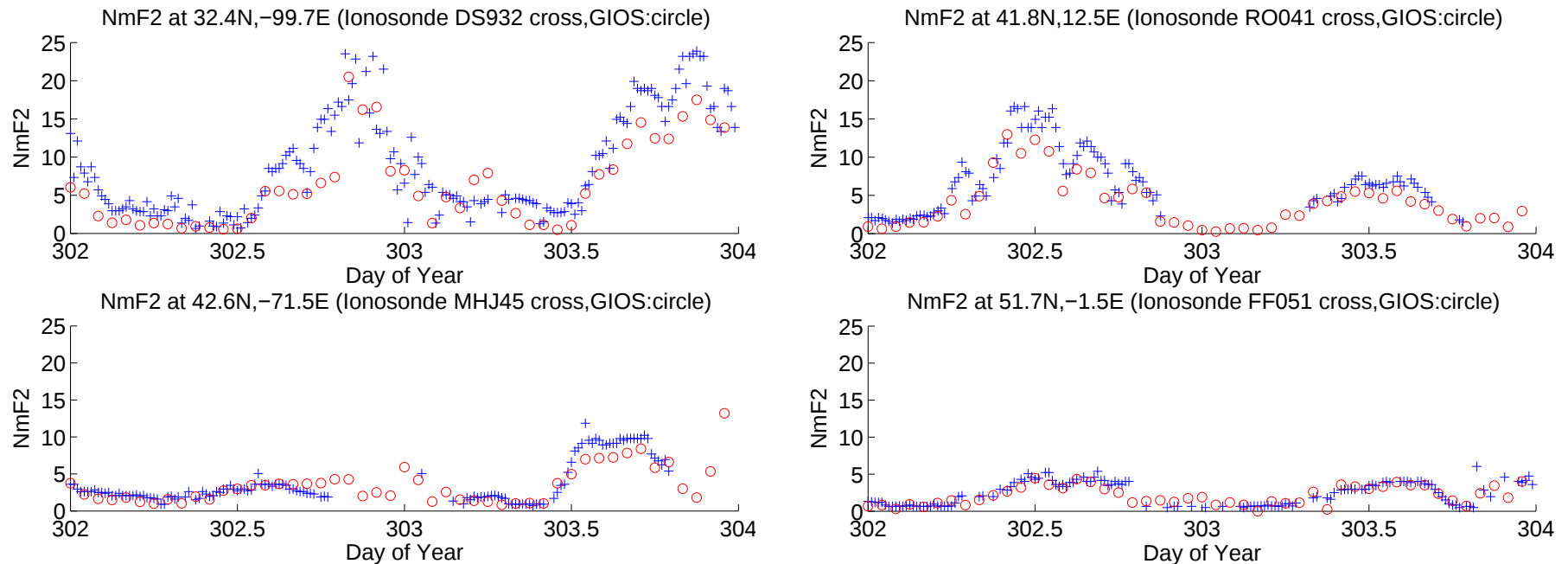
- ❑ CHAMP provides a wealth of information complementing ground GPS data;
- ❑ GOS/GIOS reconstructed images represent the ionospheric distribution more accurately than GPS reconstructed images;
- ❑ The addition of ionosonde measurements may make the electron density greater, particularly the peak density.

GPS underestimate ionosonde NmF2?



- ❑ Assume auto-scaled ionosonde data are correct; then
- ❑ About 74% inverted NmF2 from only ground based GPS data underestimate the F2 layer peak density measured by ionosondes at storm time;
- ❑ The mean under-estimation error is about 21%.

GIOS vs. ionosonde in NmF2



❑ GIOS greatly improved the peak density in the F2 layer of the disturbed ionosphere with a correlation of 90% to ionosondes

Summary

- ❑ GPS in conjunction with tomographic technique open up the possibility of imaging the disturbed ionosphere;
- ❑ The addition of CHAMP data is demonstrated the potential of imaging the polar ionosphere due to the lack of ground GPS receivers there;
- ❑ The assimilation of ionosonde data can greatly improve the layer peak density provided that they are accurate.

Future work

- ☐ Model study is required by the interpretation of tomographic observations and understanding the underlying mechanisms.
- ☐ Radio occultation data from COSMIC/FORMOSAT-3 that launched this April will be input into the inversion.

Acknowledgements

- ☐ Sponsorship from the UK PPARC & EPSRC
- ☐ GPS data from SOPAC
- ☐ Ionosonde data from SPIDR
- ☐ CHAMP data from ISDC/GFZ